

1 **Revision 4**

2

3 **Ca-Zn solid solutions in *C2/c* pyroxenes: synthesis, crystal structure**
4 **and implications for Zn geochemistry**

5

6

7 **CLAUDIA GORI¹, MARIO TRIBAUDINO^{1*}, LUCIANA MANTOVANI¹,**
8 **DAVIDE DELMONTE^{1,2}, FRANCESCO MEZZADRI^{2,3}, EDMONDO GILIOLI²**
9 **AND GIANLUCA CALESTANI^{2,3}**

10

11

12 ¹ Dipartimento di Fisica e Scienze della Terra “Macedonio Melloni”,
13 Parco Area delle Scienze, 157/A, 43124 Parma, Italy

14 ²IMEM-CNR, Istituto dei Materiali per l’Elettronica e il Magnetismo,
15 Parco Area delle Scienze, 37/A, 43124 Parma, Italy

16 ³Dipartimento di Chimica, Parco Area delle Scienze, 17/A, 43124 Parma, Italy

17

18 **ABSTRACT**

19 The effect of Zn substitution on a series of clinopyroxenes along the join $\text{CaZnSi}_2\text{O}_6$ -
20 $\text{Zn}_2\text{Si}_2\text{O}_6$ was studied. The pyroxenes were synthesized at $P = 4\text{-}5$ GPa and $T = 1000\text{-}1200^\circ\text{C}$ by a
21 multi anvil apparatus. SEM-EDS and XRD analysis showed complete solid solution; all of the
22 samples have the *C2/c* space group. No miscibility gap between clino and orthopyroxene nor phase
23 transition to the *P2₁/c* space group was found. Moreover, the cell volume of Ca-Zn pyroxenes
24 decreases less than expected from the decrease of the average cation size for the substitution of Zn
25 for Ca.

26 The crystal structures of three synthetic pyroxenes of composition $(\text{Ca}_{0.5}\text{Zn}_{0.5})\text{ZnSi}_2\text{O}_6$,
27 $(\text{Ca}_{0.3}\text{Zn}_{0.7})\text{ZnSi}_2\text{O}_6$ and $(\text{Ca}_{0.2}\text{Zn}_{0.8})\text{ZnSi}_2\text{O}_6$ were refined by single crystal X-ray diffraction ($R_{4\sigma}$
28 between 3 and 4.5%). It was observed that the Ca-Zn substitution occurs in the M2 polyhedron,
29 with a sub-site splitting of Zn in a position at approximately 0.7 Å from Ca. In this position, Zn
30 retains a highly distorted fourfold coordination; moreover, the tetrahedral chain configuration is
31 little changed along the series, and the M1 polyhedral size increases with Zn substitution in M2.

32 An implication of the present work is the interpretation of the partitioning of Zn between
33 clinopyroxene and melt. The distribution coefficients of Zn and Co are quite different in rocks of
34 the same composition, despite their very similar ionic radius, and the difference is related to the
35 preference of Zn for the M2 site, where Zn may find a suitable atomic coordination.

36 **Keywords:** clinopyroxene, high pressure, X-ray diffraction, trace elements, Zn.

37

38 *E-mail: mario.tribaudino@unipr.it

39

40

INTRODUCTION

41 The crystal structure of pyroxene, due to its inherent flexibility, provides an effective
42 playground for obtaining models relating the atomic structure to composition and P - T conditions
43 and for interpreting phase equilibria and solid solutions in minerals.

44 The formula unit of pyroxenes is $\text{M}_2\text{M}_1\text{T}_2\text{O}_6$, with M2 hosted in a distorted six-eightfold
45 polyhedron, M1 into a more regular octahedron and T in tetrahedral coordination (Fig. 1). In natural
46 pyroxenes, the M2 site is occupied by larger cations, such as Ca, Na, Mg, Fe^{2+} , and Li, the M1 site
47 is occupied by Mg, Fe^{2+} , Fe^{3+} , Al, and Cr, whereas the T site is occupied by only Si and Al. Further
48 substitutions are possible in synthetic pyroxenes: the M2 site can also be filled by Sr, Co, Ni, Mn
49 and Zn; the M1 site can be filled by Co, Ni, Cr, Mn, Ni, Zn, Sc, Ti, Cu, Ga and V; and Ge can fully
50 exchange with Si in the tetrahedral site (Ohashi et al. 1990; Redhammer and Roth 2004;
51 Redhammer et al. 2012; Redhammer and Tippelt 2014; Tribaudino et al. 2005a, b; Arlt and Angel

52 2000; Tribaudino et al. 2002; Heuer et al. 2005; Tribaudino et al. 2009; Thompson et al. 2005;
53 Hugh-Jones et al. 1994).

54 The above synthetic pyroxenes are a counterpart to natural ones. For instance, in a recent
55 investigation on the crystal structure, phase equilibria and phase transitions in synthetic pyroxenes
56 along the series $\text{CaCoSi}_2\text{O}_6$ - $\text{Co}_2\text{Si}_2\text{O}_6$ (Mantovani et al. 2013, 2014), Co-pyroxenes were compared
57 with natural Ca-Mg and Ca-Fe pyroxenes to study the effect of cation substitutions in the M2 site. It
58 was observed that in Ca-richer pyroxenes with $C2/c$ structure, the solid solution is ideal and volume
59 changes simply due to the different cation radii of the substituting cations; the apparent deviation
60 from the ideal behavior in Ca-poorer compositions is due to the negative excess volume for the
61 transition from the $C2/c$ to the $P2_1/c$ space group.

62 Moreover, it was observed that the $C2/c$ - $P2_1/c$ phase transition occurs at an average M2
63 cation radius between 0.86 to 0.88 Å.

64 A different behavior can be expected in Ca-Zn pyroxenes. Zn has a ionic radius very similar
65 to that of Co and Mg, i.e., 0.74 vs. 0.745 and 0.72 Å, respectively, but it enters differently in the M2
66 site. Co, Mg and Fe in the M2 site show a 4+2 sixfold coordination, whereas Zn exhibits a fourfold
67 coordination (Morimoto et al. 1975). Moreover, at room pressure and temperature, monoclinic
68 $\text{Zn}_2\text{Si}_2\text{O}_6$ has a $C2/c$ structure, whereas monoclinic $\text{Co}_2\text{Si}_2\text{O}_6$, $\text{Mg}_2\text{Si}_2\text{O}_6$ and $\text{Fe}_2\text{Si}_2\text{O}_6$ show a $P2_1/c$
69 space group (Morimoto et al. 1975; Ohashi et al. 1975; Morimoto et al. 1960; Sasaki and Takéuki
70 1981; Ohashi 1984). The fourfold coordination of Zn in the M2 cavity was explained by the
71 preference of Zn for covalent bonding and an sp^3 tetrahedral hybrid electronic configuration
72 (Morimoto et al. 1975).

73 Experimental data on Zn substitution in pyroxenes are scarce. Zn is a relatively common
74 element, but due to its low affinity for silicates, it occurs in rock forming silicates such as
75 pyroxenes as a trace element; actually, in natural pyroxenes, it is found as a major constituent in
76 only the rare mineral petedunnite (Essene and Peacor 1987). In petedunnite, the $\text{CaZnSi}_2\text{O}_6$
77 component is present in solid solution with johansennite ($\text{CaMnSi}_2\text{O}_6$).

78 Pure $\text{CaZnSi}_2\text{O}_6$ pyroxene was synthesized at $P > 0.8$ GPa (Huber et al. 2012); it has a $C2/c$
79 structure under room and high pressure conditions (Ohashi et al. 1996; Nestola et al. 2010). The Ca-
80 free orthorhombic $Pbca$ and monoclinic $C2/c$ polymorphs of the end member $\text{Zn}_2\text{Si}_2\text{O}_6$ were
81 synthesized by Syono et al. (1971), and their crystal structure was refined by Morimoto et al.
82 (1975).

83 At high pressure, the monoclinic $C2/c$ $\text{Zn}_2\text{Si}_2\text{O}_6$ polymorph shows two phase transitions
84 with pressure, one to $P2_1/c$ at $P = 1.92$ GPa and the other to a high pressure $C2/c$ structure at $P = 4.9$
85 GPa (Arlt and Angel 2000). The high pressure $C2/c$ structure has tetrahedral chain kinking and cell
86 parameters quite different from the room pressure $C2/c$ (Arlt and Angel 2000).

87 Orthorhombic Zn pyroxenes with mixed Ni, Co, Mg and Zn occupancy were synthesized
88 and refined by Syono et al. (1971) and Morimoto et al. (1975), showing that Zn partitions strongly
89 in the M2 site.

90 No data are available for intermediate compositions between $\text{CaZnSi}_2\text{O}_6$ and $\text{Zn}_2\text{Si}_2\text{O}_6$. The
91 relations between the $C2/c$ structures of the end members, one with Ca in eightfold and the other
92 with Zn in fourfold coordination, are not clear, nor is the extent of mutual solid solution. A specific
93 point of interest comes from the preference of Zn in the M2 polyhedron for a tetrahedral
94 coordination dictated by covalent bonding, whereas most cations substituting in the M2 polyhedron
95 of pyroxenes obey a simple ionic mechanism.

96 An investigation of Zn entrance in the M2 site also has an implication for the petrologic
97 behavior of Zn in pyroxenes as the structural behavior can provide a rationale for the geochemical
98 behavior of Zn in pyroxenes. The partitioning of Zn between clinopyroxene and melt was recently
99 proposed as a potential tracer for mantle source and an indicator for high temperature metamorphic
100 conditions (Le Roux et al. 2010; Huber et al. 2012).

101 We have therefore synthesized at high pressure a series of pyroxenes between $\text{CaZnSi}_2\text{O}_6$
102 and $\text{Zn}_2\text{Si}_2\text{O}_6$, finding that complete solid solution may be attained. Single crystal and powder
103 diffraction analysis detailed the structural mechanism for the Ca-Zn substitution; the implications

104 for the partitioning of Zn between pyroxene and melt, in comparison with other trace elements,
105 were eventually discussed.

106

107

EXPERIMENTAL METHODS

108

Synthesis

109

A set of Ca-Zn pyroxenes with composition between $\text{CaZnSi}_2\text{O}_6$ and $\text{Zn}_2\text{Si}_2\text{O}_6$ (Table 1)

110

was synthesized under high-pressure/high-temperature (HP/HT) conditions, using as a starting

111

material a stoichiometric mixture of CaCO_3 , ZnO and SiO_2 (Sigma-Aldrich > 99.99%). The mixtures

112

were previously annealed in an electric furnace for 12 hours at $T = 1000^\circ\text{C}$ and room pressure, to

113

obtain an assemblage of willemite (Zn_2SiO_4), hardystonite ($\text{Ca}_2\text{ZnSi}_2\text{O}_7$) and quartz. This

114

assemblage, finely ground, was put in a multi-anvil Walker-type press using Pt and Au capsules.

115

The capsules (30 μm -thick Pt or Au foil) were loaded into an octahedral MgO pressure cell and then

116

into the multi-anvil apparatus, in contact with a 10% Pt/Pt-Rh thermocouple for monitoring the

117

temperature during the reaction. The experiments were run at $P = 5$ GPa, and the capsules were

118

heated up to $T = 1200^\circ\text{C}$ at a rate of $25^\circ\text{C}/\text{min}$ for $(\text{Ca}_{0.8}\text{Zn}_{0.2})\text{ZnSi}_2\text{O}_6$, $(\text{Ca}_{0.5}\text{Zn}_{0.5})\text{ZnSi}_2\text{O}_6$,

119

$(\text{Ca}_{0.3}\text{Zn}_{0.7})\text{ZnSi}_2\text{O}_6$, and $(\text{Ca}_{0.2}\text{Zn}_{0.8})\text{ZnSi}_2\text{O}_6$ (hereafter referred to as $\text{Ca}_{0.8}\text{Zn}_{1.2}$, $\text{Ca}_{0.5}\text{Zn}_{1.5}$,

120

$\text{Ca}_{0.3}\text{Zn}_{1.7}$, and $\text{Ca}_{0.2}\text{Zn}_{1.8}$). For $\text{CaZnSi}_2\text{O}_6$ and $\text{Zn}_2\text{Si}_2\text{O}_6$ (hereafter referred to as Ca_1Zn_1 , Zn_2), the

121

runs were performed at $P = 4$ GPa and $T = 1000^\circ\text{C}$.

122

This condition was maintained for 3 hours of solid-state reaction, and then the capsules were

123

quenched to room temperature by switching off the heater. Finally, the pressure was slowly released

124

to prevent mechanical damage.

125

126

SEM-EDS and X-ray powder diffraction

127

A few grains of the run products were embedded in epoxy and polished for SEM-EDS

128

analysis using a Jeol 6400 SEM equipped with an Oxford EDS, operated at 20 kV. Electron

129

backscattered images were taken, showing that most of the sample was made by a phase with

130 pyroxene stoichiometry, with a little excess of silica. The crystal size varies from 10 to 100 μm ,
131 depending on the samples.

132 At least 10 analytical spots were collected on each run product. The chemical composition
133 confirmed the expected stoichiometry.

134 Powder X-ray diffraction was performed using a Bruker D2 PHASER diffractometer and a
135 Thermo ARL X'tra powder diffractometer equipped with a solid-state thermo electron detector,
136 operating with $\text{CuK}_{\alpha 1}$ radiation ($\lambda = 1.5406 \text{ \AA}$).

137 Rietveld analysis was performed to determine the unit cell parameters and the phase ratio,
138 using the GSAS-EXPGUI software package (Larson and Von Dreele 1994; Toby 2001). Starting
139 atomic parameters for the Ca-Zn pyroxenes were taken from Ohashi et al. (1996) for $\text{CaZnSi}_2\text{O}_6$
140 and from Morimoto et al. (1975) for $\text{Zn}_2\text{Si}_2\text{O}_6$. Rietveld analysis of powder patterns indicates that
141 clinopyroxene is the major phase, together with a small amount of coesite, less than 10%. Rietveld
142 refinement results are reported in Table 1.

143 **X-ray single-crystal refinement**

144 Single crystal X-ray diffraction data were collected on a Bruker AXS Smart diffractometer
145 equipped with an APEX II CCD area-detector, on selected crystals chosen for sharp optical
146 extinction, for the compositions $(\text{Ca}_{0.5}\text{Zn}_{0.5})\text{ZnSi}_2\text{O}_6$, $(\text{Ca}_{0.3}\text{Zn}_{0.7})\text{ZnSi}_2\text{O}_6$ and $(\text{Ca}_{0.2}\text{Zn}_{0.8})\text{ZnSi}_2\text{O}_6$.
147 The data collection was also attempted for other compositions, but suitable single crystals were not
148 found because of the highly twinned nature of the crystals.

149 Data collection conditions and parameters are reported in Table 2. The MoK_{α} radiation was
150 used, and the intensity data were corrected for absorption using the SADBS program (Sheldrick
151 1996). The $(\text{Ca}_{0.2}\text{Zn}_{0.8})\text{ZnSi}_2\text{O}_6$ crystals turned out to be systematically twinned on (100). The
152 refinement was performed on the intensities of one of the individuals after twin correction during
153 the data reduction.

154 Preliminary analysis of the reflection extinctions and intensities indicated a $C2/c$ space
155 group for all the samples.

156 Weighted structural anisotropic refinements were performed using the SHELX-97 program
157 (Sheldrick 1997), within the WinGX suite (Farrugia 1999) and the coordinates of $\text{CaZnSi}_2\text{O}_6$ by
158 Ohashi et al. (1996) and of $\text{Zn}_2\text{Si}_2\text{O}_6$ by Morimoto et al. (1975) as starting models. Full occupancy
159 of Zn and Si in the M1 and T sites was confirmed by site occupancy refinement. In the M2 site,
160 mixed occupancy of Zn and Ca were present and refined but without significant deviations from the
161 expected stoichiometry.

162 Similar to other pyroxenes where Ca (or Na) coexists with a smaller cation, we observed a
163 significant improvement in the agreement factor by assuming that in the M2 cavity, Ca and the
164 smaller cation (here Zn) are split between two subsites (M2 for Ca and M2' for Zn). In previous
165 works, the M2 and M2' subsites were located along the b axis, at $y/b \sim 0.30$ for Ca and 0.25 for the
166 smaller cation (Rossi et al. 1987; Tribaudino et al. 1989). In our case, the position for Zn was
167 refined further away from Ca, at $y/b \sim 0.23-0.24$, at the same position found for Zn in $C2/c$ $\text{Zn}_2\text{Si}_2\text{O}_6$
168 (Morimoto et al. 1975). The electron density map along the diad axis at the M2 site position clearly
169 shows the two split atoms (Fig. 1). Polyhedra volume and bond valence are calculated by IVTON
170 (Balic Zunic and Vickovic 1996).

171 Cell, data collection and refinement parameters, atomic coordinates, anisotropic
172 displacement parameters and a selection of bond lengths and angles are reported in Tables 2, 3 and
173 4.

174

175 RESULTS

176 Cell parameters

177 In Fig. 2, the cell parameters of Ca-Zn pyroxenes are shown. For comparison, the trend
178 observed in $\text{CaCoSi}_2\text{O}_6$ - $\text{Co}_2\text{Si}_2\text{O}_6$ pyroxenes, with Co replacing Zn, is reported (Mantovani et al.
179 2014); because the cation radius of Co is similar to that of Zn, the cell parameters are expected to
180 vary similarly. However, this is observed in only the Ca-rich end member, with no Zn in the M2
181 site (Fig. 2). As Zn enters the M2 polyhedron, we find significant differences: the a parameter

182 decreases little in Ca-Zn pyroxenes, much more than in Ca-Co ones; the b increases, but it is
183 constant in Ca-Co pyroxenes; c decreases in Ca-Co pyroxenes and increases in Ca-Zn ones, and the
184 β angle increases most in Ca-Zn pyroxenes. Volume changes are higher and non-linear in pyroxenes
185 of the series $\text{CaCoSi}_2\text{O}_6$ - $\text{Co}_2\text{Si}_2\text{O}_6$ and smaller and linear in Ca-Zn pyroxenes.

186 Similar differences can also be observed between Ca-Zn (this work), Ca-Mg (Tribaudino et
187 al. 1989) and Ca-Fe (Ohashi et al. 1975) pyroxenes.

188 The lower volume change in Ca-Zn pyroxenes can be in part ascribed to the absence of a
189 $C2/c$ - $P2_1/c$ phase transition at intermediate compositions, but the Ca-Zn and Ca-Co trends are also
190 different when both phases have $C2/c$ symmetry, i.e., when Ca is higher than 0.4 Ca atoms per
191 formula unit (a.p.f.u.).

192 The cell volume in Ca-Zn pyroxenes shows a linear change with composition, following an
193 ideal behavior (Fig. 2). The same was observed in Ca-Mg, Ca-Fe and Ca-Co $C2/c$ pyroxenes (Arlt
194 et al. 2000; Mantovani et al. 2014) where a deviation from the ideal behavior occurs for only the
195 $P2_1/c$ - $C2/c$ phase transition. The phase transition occurs with a volume decrease of the $P2_1/c$
196 structure with respect to the volume that would be observed in a $C2/c$ pyroxene with the same
197 composition; in the $P2_1/c$ field we observe linear changes in cell volume.

198

199 **M2 polyhedron**

200 In $\text{CaZnSi}_2\text{O}_6$ pyroxene, with only Ca in the M2 site, the M2-O bond distances (Fig. 3) are
201 very close to those observed in $\text{CaCoSi}_2\text{O}_6$ (Mantovani et al. 2013, 2014; Ghose et al. 1987), which
202 are, in turn, similar to those of diopside and hedenbergite (Bruno et al. 1982; Nestola et al. 2007)
203 indicating that the different cations in M1 have little effect on the Ca in the M2 site. However, when
204 Zn exchanges for Ca into the irregular eightfold polyhedron surrounding Ca atoms, which we can
205 call the M2 cavity (Fig. 1), the electron density map shows the coexistence of two distinct positions
206 for Ca and Zn. The two atomic positions split so much that an average structure is hardly significant
207 and, in our Zn-rich solid solution, we have to consider different distances for the Ca and the Zn

208 environments. In agreement with previous investigations (Rossi et al. 1987), we define M2 as the
209 polyhedron for Ca and M2' as that for the smaller cation (here Zn).

210 There are suggestions that oxygen also shows atom splitting as it was observed in synthetic
211 pyroxenes along the Di-En series by the analysis of the size of the U_{eq} vs. composition (Fig. 4)
212 (Bruno et al. 1982; Tribaudino et al. 1989; Tribaudino et al. 2003).

213 In Ca-Zn pyroxenes, similar to other *C2/c* pyroxenes where Ca exchanges with a smaller
214 cation, the U_{eq} for the O2 atoms are larger in the intermediate compositions than in the end
215 members and are generally larger than for other atoms of oxygen. Positional disorder explains the
216 increased size in intermediate compositions as the displacement parameters record an average of
217 local positions, which are different if the O2 is bonded to Zn in M2' or Ca in M2.

218 The M2'-O distances with the O2 and O1 atoms are close to 2.0 Å, whereas, as Ca is
219 exchanged by Zn, the M2'-O3 interatomic distances are greater than 3.0 Å. This results in an
220 irregular fourfold coordination of O1 and O2 atoms of oxygen around Zn (in the M2' site) (Fig. 1).
221 The same was found in $Zn_2Si_2O_6$ by Morimoto et al. (1975), and here, it is also observed in the
222 intermediate compositions: it is likely that Zn enters into fourfold coordination in the M2
223 polyhedron, even if present in a small amount.

224 Zn-O1 and Zn-O2 distances do not change with composition, and Zn-O3 distances become
225 even larger as zinc increases. However, from the geometrical point of view, we are far from a
226 regular tetrahedron: for instance, in $Ca_{0.3}Zn_{1.7}$, the O2-M2'-O2 angle is 169.7° and the other angles
227 of the O1 and O2 atoms of oxygen with M2' are between 90° and 100°, to compare with the ideal
228 109.4° predicted by tetrahedral coordination.

229 The fourfold coordination of Zn in M2' is specific to *C2/c* clinopyroxenes: in the
230 orthopyroxene and in the high pressure *P2₁/c* structure, Zn instead shows an asymmetric sixfold
231 coordination, with two out of four O3 oxygens in coordination (Morimoto et al. 1975; Arlt et al.
232 2000).

233 As Ca decreases, the bond distances with the furthestmost O3 atoms of oxygen (M2-
234 O3C2,D2) and those with the closer O3 (M2-O3C1,D1 distances) increase; the M2-O1 distance is
235 constant, and the M2-O2 distance decreases. Within the eightfold M2 cavity, Ca then shows a 4+2
236 sixfold coordination and Zn shows a fourfold one.

237 The contribution of the M2 cavity to the cell volume can be obtained by the volume of the
238 eightfold coordinated polyhedron, regardless of the actual coordination of the cation in the site. This
239 volume is compared in Fig. 5 with that of other pyroxenes where Ca exchanges with a smaller
240 cation. The volume of the M2 cavity changes in a way very similar to that of Ca-Co, Ca-Fe and Ca-
241 Mg pyroxenes, with the only exception in the end member $Zn_2Si_2O_6$, which has the same volume as
242 $Ca_{0.2}Zn_{1.8}Si_2O_6$.

243 This result is quite surprising and indicates that the volume of the M2 cavity changes almost
244 completely for the different cation radii of Zn and Ca, despite the different bonding of Zn and Co
245 in the M2 cavity (Fig. 5). This change is similar to what was found between Co and Ca. The
246 suggestion is that Ca, being the larger cation, accounts for the decrease in the size of the cavity, no
247 matter for the local coordination of Zn. The volume decrease in the M2 cavity of Ca-Co pyroxenes
248 explains the cell volume decrease almost completely; however, as discussed in the following
249 paragraph, in Zn pyroxenes, a cell volume increase driven by the M1 polyhedron opposes the
250 decrease in the M2 cavity and results in a lower volume change (Fig. 5).

251 **M1 polyhedron**

252 In the Ca-Zn series, Zn fully occupies the M1 site. The M1-O bond lengths vs. Ca are shown
253 in Fig. 6 and compared with those of Ca-Co pyroxenes. The two longer and smaller M1-O1
254 distances (M1-O1 A1,B1 and M1-O1 A2,B2, respectively) show a strong and opposite change with
255 Ca: the shorter distances decrease whereas the longer ones increase; the intermediate M1-O2
256 distances instead change slightly. The difference between shorter and longer distances also
257 increases in Ca-Co pyroxenes, but at a lower rate (Mantovani et al. 2013). For instance, the longer
258 M1-O1 distance increases by 0.064 Å along the series $CaCoSi_2O_6$ - $Co_2Si_2O_6$, but in Ca-Zn

259 pyroxenes, it increases by 0.174 Å. As a result, in $\text{Zn}_2\text{Si}_2\text{O}_6$ a 4+2 coordination is observed, with the
260 longer M1-O1B being less bonded. We also interpret this behavior as an attempt of Zn in the M1
261 site to approach the fourfold coordination, as in the M2 cavity. However, a transition to a fourfold
262 coordination in the M1 site is hindered because the M1 polyhedron is tighter and less flexible than
263 the M2 one; in the M2 cavity, which is more flexible and enables higher deformation, Zn finds its
264 own fourfold coordination.

265 Average bond distances and M1 polyhedral volume are very similar in $\text{CaZnSi}_2\text{O}_6$ and
266 $\text{CaCoSi}_2\text{O}_6$, although individual bond lengths are somewhat different. As shown in Fig. 7, the
267 polyhedral M1 volume is related to the M1 ionic radius in pyroxenes with only Ca in the M2 site,
268 suggesting that Zn in the M1 site shows an ionic behavior. In intermediate solid solutions, the
269 average M1-O distances increase with substitution of Zn for Ca in the M2 site, which is surprising
270 as the M1 site is always filled by Zn. In fact, the average bond distances (M1-O) are almost the
271 same for the whole series in Ca-Fe, Ca-Co and Ca-Mg pyroxenes (Fig. 8), where Fe, Co and Mg fill
272 the M1 polyhedron. The contribution of the increased size of the M1 polyhedron to the cell volume
273 explains why the M2 cavity in the case of Zn-pyroxenes does not account for the change in volume
274 as observed for the Co-pyroxenes: the Ca-Zn substitution in the M2 polyhedron decreases the cell
275 volume, but the increased M1 polyhedral size opposes such a decrease.

276

277 **T tetrahedron and tetrahedral chain**

278 Throughout the series, the volume and average bond distances in the Si filled tetrahedron do
279 not change significantly but the individual T-O bond distances and angles do show significant
280 changes with Ca. The bond distances of Si with the O3 atoms, i.e., the T-O bridging (T-O_{br}),
281 decrease with decreasing Ca, whereas the distances with O1 and O2 atoms (T-O non bridging, T-
282 O_{nbr}), slightly increase (Fig. 9). Moreover, the tetrahedral angle variation (TAV) decreases with
283 decreasing Ca, showing that the tetrahedron becomes more regular. The T-O_{br} distances decrease

284 for the lower electrostatic contribution from Ca and Zn as both the M2-O3 and M2'-O3 distances
285 increase with Zn, making the tetrahedral chains further from the cations.

286 The O3-O3 distance, which is a shared edge with the M2 polyhedron, increases as Ca
287 decreases, accounting for the strong increase in the *c* parameter. As in the M1 polyhedron, we have
288 therefore an effect of the M2-O distances on the arrangement of the tetrahedral bond distances. The
289 same rearrangement in Si-O bond lengths is also found in other pyroxenes when Ca is exchanged
290 by a smaller cation (Fig. 10a).

291 The geometry of the tetrahedral chain does not change much with Ca-Zn substitution: the
292 kinking angle decreases, following the same trend in *C2/c* Ca-Co, Ca-Mg and Ca-Fe pyroxenes
293 (Fig. 10b). Rather than a high temperature fully extended configuration similar to that found in
294 spodumene or high temperature pigeonite (Cameron et al. 1973), in the *C2/c* end member $Zn_2Si_2O_6$,
295 the tetrahedral chain shows intermediate kinking, close to that observed in diopside. The kinking of
296 the end member $Zn_2Si_2O_6$ is very close to that of $(Ca_{0.3}Zn_{0.7})ZnSi_2O_6$ and $(Ca_{0.2}Zn_{0.8})ZnSi_2O_6$.

297 The relationship between the kinking angle and the cell β angle (Thompson et al. 2005) is
298 not found here. The β angle increases much more in Ca-Zn than in Ca-Co pyroxenes (Fig. 2), but
299 the O3-O3-O3 kinking angle changes likewise (Fig. 10b), with the exception of the end member
300 $Zn_2Si_2O_6$. The strong increase in the β angle is described by the shift of the tetrahedral chains
301 (Benna et al. 1988; Tribaudino 1996; Mantovani et al. 2013) that accounts for the offset between
302 the two tetrahedral chains facing onto the M2 polyhedron along the *c* direction (Tribaudino et al.
303 1989). It was observed by Mantovani et al. (2013) that chain shift enables an increase in the bond
304 distance between M2 and the furthestmost O3, maintaining the *C2/c* symmetry; in Ca-Fe, Ca-Mg
305 and Ca-Co pyroxenes, it takes a maximum value of 0.9 Å, which was suggested to be a limiting
306 value before the transition to the *P2₁/c* symmetry. In Ca-Zn pyroxenes, higher shift can occur in the
307 *C2/c* symmetry, accounting for the higher values in the β angle (Fig. 10c).

308

309

310

311

DISCUSSION

312

Zn substitution and phase transition

313

314

315

316

The average ionic radius for six-coordinated M2 decreases from 1.0 to 0.74 Å along the series. Following the prediction of a critical cation radius at approximately 0.86-0.88 Å, a $C2/c$ - $P2_1/c$ phase transition should occur as Ca becomes less than 0.4-0.5 a.p.f.u., but the $C2/c$ symmetry is retained for any Zn substitution.

317

318

319

320

321

322

This effect reflects a different interaction of the cation in the M2 site with the tetrahedral chain. In Ca-Mg, Ca-Fe, Ca-Co and Mn-Mg pyroxenes, the phase transition occurs as the size of the average M2 cation decreases; then, in the $P2_1/c$ structure, the cation falls out from the special position along the diad axis and the tetrahedral chains are rearranged to alternatively decrease and increase the bond lengths with the M2 cation. Therefore, there is an active interaction between the M2 and O3 oxygen.

323

324

325

326

327

328

In the case of Ca-Zn pyroxenes, at least at room conditions, this interaction is not observed: the tetrahedral chain kinking does not change much along the series, and for Zn richer samples, it does not change at all. The tetrahedral chain kinking angle decreases only if the Ca content is large enough. For higher Zn content, the cation has no interaction with the tetrahedral chain and no significant kinking occurs. The reduced ionic radius for Ca-Zn exchange in M2 is compensated by the tetrahedral chain shift.

329

330

331

332

333

334

335

This effect does not occur simply because Zn lies in the M2' site at a position more displaced from the tetrahedral chain than Mg, Co or Fe (Tribaudino et al. 1989; Rossi et al. 1987; Mantovani et al. 2013; Ohashi et al. 1975; Nestola et al. 2007) and the bond distances with the O3 oxygen are longer (Fig. 3). A calculation on the charge on the O3 atom, which is, apart from Si, bonded by only the M2-M2' atoms (Fig. 1), shows only little difference between Ca-Co and Ca-Zn pyroxenes (Fig. 11). A likely explanation lies in the covalent character and the tetrahedral coordination of the Zn-O bonding, which does not involve the O3 atoms of the chains facing the

336 M2 polyhedron: the charge balance model, which assumes an ionic charge distribution, is not
337 correct for Zn in M2 in Zn pyroxenes. Bonding for Zn involves only the O1 and O2 atoms, and an
338 interaction with the O3 atoms is not present.

339 A switch from the above “covalent” behavior to a “ionic” one occurs at high pressure,
340 reducing the Zn-O3 distances and promoting an approach to the tetrahedral chains,. At high
341 pressure, an interaction of Zn with the O3 bridging atoms promotes the transition to the $P2_1/c$
342 symmetry, at $P= 1.9$ GPa in $Zn_2Si_2O_6$ (Arlt and Angel 2000). In the high pressure $P2_1/c$ structure, as
343 well as in the $Pbca$ structure of the orthopyroxene, which can be described as a twin at the unit cell
344 scale of the $P2_1/c$ structure, Zn is no longer fourfold coordinated (Morimoto et al. 1975; Arlt and
345 Angel 2000). Following the same rationale, we could also predict the transition to $P2_1/c$ at low
346 temperature, as a consequence of a supposed decrease in the volume of the M2 polyhedron and
347 approach of the tetrahedral chains to M2. This prediction is indeed supported by quantum
348 mechanical calculations at 0 K on monoclinic $Zn_2Si_2O_6$ (Karazhanov et al. 2009) but not by
349 experiments. Powder XRD down to 100 K by Galkin et al. (2007) in $Zn_2Si_2O_6$ does not show any
350 evidence of a transition to $P2_1/c$, and the transition, if any, should occur at even lower temperature.

351

352 **Solid solution in $C2/c$ pyroxenes**

353 A complete solid solution is obtained in Ca-Zn clinopyroxenes at a temperature of 1200°C;
354 under the same conditions where the solid solution was very limited in the Ca-Co and Ca-Mg
355 pyroxenes (Lindsley and Dixon 1976; Mantovani et al. 2014). In Ca-Co, as in Ca-Mg and Ca-Fe
356 pyroxenes, we have a miscibility gap separating a wider solid solution field of Ca-rich
357 clinopyroxenes from a narrow Ca-poor orthopyroxene field. Here, the orthopyroxene is not stable
358 and the two-pyroxene assemblage does not lead to a decrease in free energy with respect to the
359 single phase pyroxene.

360 Therefore, also in view of the ideal behavior of cell volumes, as long as the symmetry is
361 $C2/c$, complete solid solution can be expected.

362 The different coordination of Zn in the M2 cavity in *C2/c* and *Pbca* pyroxenes accounts for
363 the above behavior. In the M2 site of the *Pbca* $Zn_2Si_2O_6$ orthopyroxene, two of the O3 atoms of
364 oxygen are in coordination, although in an asymmetric position within the M2 cavity, and the
365 coordination in the M2 site can be described as 4+2 sixfold (Domeneghetti et al. 1995). The same
366 was observed in high pressure *P2₁/c* $Zn_2Si_2O_6$ (Arlt and Angel 2000). In *C2/c* clinopyroxene, the
367 shorter M2-O3 distances are out of bonding (~ 3.03 Å, Morimoto et al. 1975) and the coordination
368 of Zn in M2 is fourfold. Given the preference of Zn for a fourfold coordination, the higher stability
369 of the *C2/c* phase with respect to the *P2₁/c* and orthorhombic *Pbca* is not surprising (Syono et al.
370 1971).

371

372

IMPLICATIONS

373 In this work on synthetic pyroxenes, we have filled the M1 site with Zn and exchanged Zn
374 for Ca in the M2 site, which was possible because Ca does not enter the M1 site and any Ca will
375 buffer the M2 site. In natural petedunnite, we also have the M2 site completely filled by Ca and Zn
376 enters only the M1 site. However, Zn can enter both the M2 and the M1 sites if Ca does not buffer
377 the M2 site, as in clinopyroxenes from basalts and mantle peridotites.

378 Intracrystalline M2-M1 site preference for trace elements can be tested by the analysis of the
379 partition coefficients of the same elements between clinopyroxene and melt. Partitioning
380 coefficients are modified by the different chemical environments, but in a given sample, for a given
381 phase and in a given chemical site, they are related to the ionic radius of the partitioning element by
382 Onuma diagrams (Onuma et al. 1981). Onuma diagrams provide a curve showing the partitioning
383 coefficient vs. cation radius for cations in each structural site. Different curves are provided for
384 different structural sites. In general, cations with very similar ionic radii and charge, confined in a
385 given structural site, have very similar mineral-melt partition coefficients. Cations in different
386 structural sites may on the contrary have different crystal-melt partitioning even if they have similar
387 ionic size.

388 We could then expect that Zn will have a very similar partitioning coefficient to cations such
389 as Co and Mg, with the same valence and similar size if they are concentrated in the same structural
390 site. This similarity is not found.

391 In Fig. 12, the clinopyroxene-melt distribution coefficients for a series of divalent cations, as
392 reported for a Ca-poor Iherzolitic clinopyroxene by Adam and Green (2006), are plotted as a
393 function of ionic radius. The six- and eight-coordinated radii were reported (Shannon 1976),
394 respectively, for the M1 and M2 polyhedra. The Onuma diagram curves reported in Adam and
395 Green (2006) for the specific case were calculated from the lattice strain model of Blundy and
396 Wood (2003). In this model, the partition coefficients can be predicted by the polyhedral Young's
397 modulus and the ideal radius that the cation should have in a given polyhedron.

398 We observe that the clinopyroxene-melt partitioning coefficients for atoms such as Ca, Pb
399 and Mn, which are completely (Ca, Pb) or almost completely (Mn) partitioned in the M2 site, fall
400 on the curve calculated for the M2 site. Those for Co and Mg fall close to the curve calculated for
401 the M1 site, showing that their distribution is fully predicted, assuming that they enter the M1 site.
402 Zn instead shows a much lower clinopyroxene-melt distribution coefficient than expected if
403 concentrated in the octahedral M1 site, whereas, assuming that Zn enters the M2 site, the observed
404 partitioning agrees with predictions (Fig. 12).

405 The suggestion from crystal-melt partitioning is that Zn enters the M2 site. Although this
406 suggestion awaits experimental confirmation by an analysis of intracrystalline site partitioning
407 between the M2 and M1 sites in natural clinopyroxene, a strong preference was shown by Zn for
408 the M2 site in a solid solution of Co, Ni, Zn orthopyroxenes and in $\text{ZnMgSi}_2\text{O}_6$ synthetic
409 orthopyroxene (Morimoto et al. 1975; Ghose et al. 1975).

410 A likely crystal chemical explanation of the preference of Zn in the M2 site, in
411 clinopyroxenes and in orthopyroxenes, is that the M2 cavity is more flexible and better suited to
412 host the Zn cation in a fourfold sub-site than is the M1 polyhedron, which is a tighter cavity.
413 Incidentally, Ghose (1975) showed that there is a strong correlation between the intracrystalline

414 partition coefficients and ionic radii in the M1 site, and Zn is again an exception, being more
415 concentrated in the M2 site than expected by its ionic radius.

416 The crystal chemical considerations put forward by the present work may also have a
417 petrologic implication. The crystal chemical preference of Zn for the M2 site results in a different
418 (lower) partitioning of Zn in the clinopyroxene but also a higher partitioning in the melt. A melt
419 formed by clinopyroxene-rich source rocks will be then enriched in Zn. The enrichment of Zn in
420 basaltic melts resulting from partial melting of clinopyroxene-rich source rocks was used by Le
421 Roux et al. (2010, 2011) to constrain the origin of the observed magmas, to track the source rock of
422 the original magmas and to detect the mantle source heterogeneities.

423 The above petrologic implications derive from the peculiar bonding requirements of Zn in
424 pyroxenes that confine Zn to the M2 cavity, and the clinopyroxene-melt partitioning coefficient is
425 then lower than what we would have if Zn was present in the M1 octahedral coordination. The
426 empirical predictions by Onuma diagrams are respected, once it is assumed that the smaller Zn fits
427 into the larger M2 cavity.

428

429

430

ACKNOWLEDGMENTS

431 This work was supported by the Italian PRIN project "Spinelli e Pirosseni: nuovi materiali per
432 l'industria ceramica e l'elettronica" #2010EARRRZ_005. The authors thank Diego Gatta and two
433 anonymous reviewers for constructive comments.

434

435

436

437

438

439

440

441

442

443

444

445

REFERENCES CITED

446 Adam, J., and Green, T. (2006) Trace element partitioning between mica-and amphibole-bearing
447 garnet lherzolite and hydrous basanitic melt: 1. Experimental results and the investigation of
448 controls on partitioning behavior. *Contributions to Mineralogy and Petrology*, 152, 1-17.

449

450 Arlt, T., and Angel, R.J. (2000) Displacive phase transitions in *C*-centered clinopyroxenes:
451 spodumene, $\text{LiScSi}_2\text{O}_6$ and ZnSiO_3 . *Physics and Chemistry of Minerals*, 27, 719-731.

452

453 Arlt, T., Kunz, M., Stolz, J., Armbruster, T., and Angel, R.J. (2000) PTX data on *P21/c*-
454 clinopyroxenes and their displacive phase transitions. *Contributions to Mineralogy and Petrology*,
455 138, 35-45.

456

457 Balic Zunic, T., and Vickovic, I. (1996) IVTON: a program for the calculation of geometrical
458 aspects of crystal structures and some crystal chemical applications. *Journal of Applied*
459 *Crystallography*, 29, 305-306.

460

461 Benna, P., Tribaudino, M., and Bruno, E. (1988) Crystal structure of $\text{Di}_{50}\text{CaTs}_{50}$ synthetic
462 clinopyroxene ($\text{CaMg}_{0.5}\text{AlSi}_{1.5}\text{O}_6$). Crystal chemistry along the Di-CaTs join. *Mineralogy and*
463 *Petrology*, 38, 189-200.

464

- 465 Blundy, J., and Wood, B. (2003) Partitioning of trace elements between crystals and melts. Earth
466 and Planetary Science Letters, 210, 83-397.
467
- 468 Bruno, E., Carbonin, S., and Molin, G.M. (1982) Crystal structure of Ca-rich clinopyroxenes on the
469 CaMgSi₂O₆-Mg₂Si₂O₆ join. Tschermaks Mineralogische und Petrographische Mitteilungen, 29,
470 223-240.
471
- 472 Cameron, M., Sueno, S., Prewitt, C.T., and Papike, J.J. (1973) High-temperature crystal chemistry
473 of acmite, diopside, hedenbergite, jadeite, spodumene and ureyte. American Mineralogist, 58,
474 594-618.
475
- 476 Domeneghetti, M.C., Molin, G.M., Stimpfl, M., and Tribaudino, M. (1995) Orthopyroxene from the
477 Serra de Magé meteorite: structure refinement and estimation of *C2/c* pyroxene contributions to
478 apparent *Pbca* diffraction violations. American Mineralogist, 80, 923-929.
479
- 480 Essene, E.J., and Peacor, D.R. (1987) Petedunnite (CaZnSi₂O₆), a new zincian clinopyroxene from
481 Franklin, New Jersey, and phase equilibria for zincian pyroxenes. American Mineralogist, 72,
482 157-166.
483
- 484 Farrugia, L.J. (1999) WinGX suite for small-molecule single-crystal crystallography. Journal of
485 Applied Crystallography, 32, 837-838.
486
- 487 Galkin, V., Kuznetsov, G., and Turkin, A. (2007) Thermal expansion of ZnSiO₃ high-pressure
488 phases. Physics and Chemistry of Minerals, 34, 377-381.
489

- 490 Ghose, S., Wan, C., Okamura, P., Ohashi, H., and Weidner, J.R. (1975) Site preference and crystal-
491 chemistry of transition-metal ions in pyroxenes and olivines. *Acta Crystallographica*, A31, S76-35.
492
- 493 Ghose, S., Wan, C., and Okamura, F.P. (1987) Crystal structures of $\text{CaNiSi}_2\text{O}_6$ and $\text{CaCoSi}_2\text{O}_6$ and
494 some crystal-chemical relations in *C2/c* clinopyroxenes. *American Mineralogist*, 72, 375-381.
495
- 496 Heuer, M., Huber, A.L., Bromiley, G.D., Fehr, K.T., and Bente, K. (2005) Characterization of
497 synthetic hedenbergite ($\text{CaFeSi}_2\text{O}_6$)-petedunnite ($\text{CaZnSi}_2\text{O}_6$) solid solution series by X-ray single
498 crystal diffraction. *Physics and Chemistry of Minerals*, 32, 552-563.
499
- 500 Huber, A.L., Heuss-Aßbichler, S., Fehr, K.T., and Bromiley, G.D. (2012) Petedunnite ($\text{CaZnSi}_2\text{O}_6$):
501 Stability and phase relations in the system CaO-ZnO-SiO_2 . *American Mineralogist*, 97, 739-749.
502
- 503 Hugh-Jones, D.A., Woodland, A.B., and Angel, R.J. (1994) The structure of high-pressure *C2/c*
504 ferrosilite and crystal chemistry of high-pressure *C2/c* pyroxenes. *American Mineralogist*, 79,
505 1032-1041.
506
- 507 Karazhanov, S.Z., Ravindran, P., Vajeeston, P., Ulyashin, A.G., Fjellvåg, H., and Svensson, B.G.
508 (2009) Phase stability and pressure-induced structural transitions at zero temperature in ZnSiO_3 and
509 Zn_2SiO_4 . *Journal of Physics: Condensed Matter*, 21, 485801.
510
- 511 Larson, A.C., and Von Dreele, R.B. (1994) Gsas. General Structure Analysis System. LANSCE,
512 MS-H805, Los Alamos, New Mexico.
513

- 514 Le Roux, V., Lee, C.-T.A., and Turner, S.J. (2010) Zn/Fe systematics in mafic and ultramafic
515 systems: Implications for detecting major element heterogeneities in the Earth's mantle.
516 *Geochimica et Cosmochimica Acta*, 74, 2779-2796.
- 517 Le Roux, V., Dasgupta, R., and Lee, C.-T.A. (2011) Mineralogical heterogeneities in the Earth's
518 mantle: Constraints from Mn, Co, Ni and Zn partitioning during partial melting. *Earth and*
519 *Planetary Science Letters*, 307, 395-408.
- 520
- 521 Lindsley, D.H., and Dixon, S.A. (1976) Diopside-enstatite equilibria at 850 to 1400°C, 5 to 35 kb.
522 *American Journal of Science*, 276, 1285-1301.
- 523
- 524 Mantovani, L., Tribaudino, M., Mezzadri, F., Calestani, G., and Bromiley, G. (2013) The structure
525 of (Ca,Co)CoSi₂O₆ pyroxenes and the Ca-M²⁺ substitution in (Ca,M²⁺)M²⁺Si₂O₆ pyroxenes (M²⁺=
526 Co, Fe, Mg). *American Mineralogist*, 98, 1241-1252.
- 527
- 528 Mantovani, L., Tribaudino, M., Bertoni, G., Salviati, G., and Bromiley, G. (2014) Solid solutions
529 and phase transitions in (Ca,M²⁺)M²⁺Si₂O₆ pyroxenes (M²⁺= Co, Fe, Mg). *American Mineralogist*,
530 99, 704-711.
- 531
- 532 Morimoto, N., Appleman, D.E., and Evans, H.T. (1960) The crystal structures of clinoenstatite and
533 pigeonite. *Zeitschrift für Kristallographie*, 114, 120-147.
- 534
- 535 Morimoto, N., Nakajima, Y., Syono, Y., Akimoto, S., and Matsui, Y. (1975) Crystal structure of
536 Pyroxene-Type ZnSiO₃ and ZnMgSi₂O₆. *Acta Crystallographica*, B31, 1041-1049.
- 537

- 538 Nestola, F., Tribaudino, M., Boffa Ballaran, T., Liebske, C., and Bruno, M. (2007) The crystal
539 structure of pyroxenes along the jadeite-hedenbergite and jadeite-aegirine joins. American
540 Mineralogist, 92, 1492-1501.
541
- 542 Nestola, F., Ballaran Boffa, T., Angel, R.A., Zhao, J., and Ohashi, H. (2010) High-pressure
543 behavior of Ca/Na clinopyroxenes: The effect of divalent and trivalent 3d-transition elements.
544 American Mineralogist, 95, 832-838.
545
- 546 Ohashi, Y. (1984) Polysynthetically-twinned structures of enstatite and wollastonite. Physics and
547 Chemistry of Minerals, 10, 217-229.
548
- 549 Ohashi, Y., Burnham, C.W., and Finger, L.W. (1975) The effect of Ca-Fe substitution on the
550 clinopyroxene crystal structure. American Mineralogist, 60, 423-434.
551
- 552 Ohashi, H., Osawa, T. and Sato, A. (1990) Structures of Na(In,Sc)Si₂O₆ clinopyroxenes formed at 6
553 GPa pressure. Acta Crystallographica, B46, 742-747.
554
- 555 Ohashi, H., Osawa, T., Sato, A., and Tsukimura, K. (1996) Crystal structure of (Na, Ca)
556 (Sc,Zn)Si₂O₆ clinopyroxenes formed at 6 GPa pressure. Journal of Mineralogy, Petrology and
557 Economic Geology, 91, 21-27.
558
- 559 Onuma, N., Ninomiya, S., and Nagasawa, H. (1981) Mineral/groundmass partition coefficients for
560 nepheline, melilite, clinopyroxene and perovskite in melilite-nepheline basalt, Nyiragongo, Zaire.
561 Geochemical Journal, 15, 221-228.
562

- 563 Redhammer, G.J., and Roth, G. (2004) Structural changes upon the temperature dependent
564 $C2/c \rightarrow P2_1/c$ phase transition in $\text{LiMe}^{3+}\text{Si}_2\text{O}_6$ clinopyroxenes, $\text{Me} = \text{Cr, Ga, Fe, V, Sc}$ and In .
565 *Zeitschrift für Kristallographie*, 219, 585-605.
566
- 567 Redhammer, G.J., Nestola, F., and Miletich, R. (2012) Synthetic $\text{LiAlGe}_2\text{O}_6$: The first pyroxene
568 with $P2_1/n$ symmetry. *American Mineralogist*, 97, 1213-1218.
569
- 570 Redhammer, G.J., and Tippelt, G. (2014) Pyroxene-type compounds $\text{NaM}^{3+}\text{Ge}_2\text{O}_6$, with $M = \text{Ga,}$
571 Mn, Sc and In . *Acta Crystallographica*, 70, 852-857.
572
- 573 Robinson, K., Gibbs, G.V., and Ribbe, P.H. (1971) Quadratic elongation: A quantitative measure of
574 distortion in coordination polyhedra. *Science*, 172, 567-570.
575
- 576 Rossi, G., Oberti, R., Dal Negro, A., Molin, G.M., and Mellini, M. (1987) Residual electron density
577 of the M2 site in $C2/c$ clinopyroxenes relationship with bulk chemistry and sub-solidus evolution.
578 *Physics and Chemistry of Minerals*, 14, 514-520.
579
- 580 Sasaki S. and Takéuchi Y. (1981) Electron-density distributions of three orthopyroxenes, $\text{Mg}_2\text{Si}_2\text{O}_6$,
581 $\text{Co}_2\text{Si}_2\text{O}_6$, and $\text{Fe}_2\text{Si}_2\text{O}_6$. *Zeitschrift für Kristallographie*, 158, 279-298.
582
- 583 Shannon, R.D. (1976) Revised effective ionic radii and systematic studies of interatomic distances
584 in halides and chalcogenides. *Acta Crystallographica*, A32, 751-767.
585
- 586 Sheldrick, G.M. (1996) SADABS, program for empirical absorption correction of area detector
587 data. University of Göttingen, Göttingen, Germany.
588

- 589 Sheldrick, G.M. (1997) SHELXL-97, program for Crystal Structure Analysis. University of
590 Göttingen, Göttingen, Germany.
- 591
- 592 Syono, Y., Akimoto, S. I., and Matsui, Y. (1971) High pressure transformations in zinc silicates.
593 Journal of Solid State Chemistry, 3, 369-380.
- 594
- 595 Thompson, R.M., Downs, R.T., and Redhammer, J.G. (2005) Model pyroxenes III: Volume of *C2/c*
596 pyroxenes at mantle *P*, *T*, and *x*. American Mineralogist, 90, 1840-1851.
- 597
- 598 Toby, B.H. (2001) EXPGUI, a graphical user interface for GSAS. Journal of Applied
599 Crystallography, 34, 210-213.
- 600
- 601 Tribaudino, M. (1996) High-temperature crystal chemistry of *C2/c* clinopyroxenes along the join
602 CaMgSi₂O₆-CaAl₂SiO₆. European Journal of Mineralogy, 8, 273-279.
- 603
- 604 Tribaudino, M., Benna, P., and Bruno, E. (1989) Average structure and M2 site configurations in
605 *C2/c* clinopyroxenes along the Di-En join. Contributions to Mineralogy and Petrology, 103, 452-
606 456.
- 607
- 608 Tribaudino, M., Nestola, F., Cámara, F., and Domeneghetti, M. C. (2002) The high-temperature
609 *P2₁/c*-*C2/c* phase transition in Fe-free pyroxene (Ca_{0.15}Mg_{1.85}Si₂O₆): Structural and thermodynamic
610 behavior. American Mineralogist, 87, 648-657.
- 611
- 612 Tribaudino, M., Nestola, F., Meneghini, C., and Bromiley, G.D. (2003) The high-temperature
613 *P2₁/c*-*C2/c* phase transition in Fe-free Ca-rich *P2₁/c* clinopyroxenes. Physics and Chemistry of
614 Minerals, 30, 527-535.

615
616
617
618
619
620
621
622
623
624
625
626
627
628
629
630
631
632
633
634
635
636
637
638
639
640

Tribaudino, M., Nestola, F., and Meneghini, C. (2005a) The structure behavior of intermediate pyroxenes along the diopside-enstatite join. *The Canadian Mineralogist*, 43, 1411-1421.

Tribaudino, M., Nestola, F., and Ohashi, H. (2005b) High temperature single crystal investigation in a clinopyroxene of composition $(\text{Na}_{0.5}\text{Ca}_{0.5})(\text{Cr}_{0.5}\text{Mg}_{0.5})\text{Si}_2\text{O}_6$. *European Journal of Mineralogy*, 17, 297-304.

Tribaudino, M., Bromiley, G., Ohashi, H., and Nestola, F. (2009) Synthesis, TEM characterization and thermal behavior of $\text{LiNiSi}_2\text{O}_6$ pyroxene. *Physics and Chemistry of Minerals*, 36, 527-536.

641 **Figure captions**

642 **Figure 1: (a)** Crystal structure of one of the *C2/c* clinopyroxenes synthesized ($\text{Ca}_{0.5}\text{Zn}_{1.5}\text{Si}_2\text{O}_6$),
643 viewed perpendicular to the *c* axis and electron density map at the M2 site (*z* section); the cation
644 occupancy of the site is reported; **(b)** CaM2 and ZnM2 site positions are always located,
645 respectively, at the M2 and M2' sites.

646

647 **Figure 2:** Unit cell parameters vs. composition in Ca-Zn and Ca-Co pyroxenes: the former are
648 represented by the green points, and the latter are represented by purple squares. Solid and dashed
649 lines show the trend for the *C2/c* and *P2₁/c* pyroxenes, respectively [Data for Ca-Co pyroxenes
650 from Mantovani et al. (2014)].

651

652 **Figure 3:** M2-O bond lengths vs. composition in Ca-Zn and in Ca-Co pyroxenes **(a)**; M2'-O bond
653 lengths with composition in Ca-Zn pyroxenes and in Ca-Co pyroxenes **(b)** [Data for Ca-Co
654 pyroxenes from Mantovani et al. (2013) and Ghose et al. (1987); for Ca-Zn pyroxenes from Ohashi
655 et al. (1996), Morimoto et al. (1975) and this work].

656

657 **Figure 4:** U_{eqO2} vs. composition [Data for Ca-Co pyroxenes from Mantovani et al. (2013) and
658 Ghose et al. (1987); for Ca-Zn pyroxenes from Ohashi et al. (1996), Morimoto et al. (1975) and this
659 work; for Ca-Mg from Bruno et al. (1982) and Tribaudino et al. (1989); for Ca-Fe from Ohashi et
660 al. (1975) and Nestola et al. (2007)].

661

662 **Figure 5: (a)** ΔV (\AA^3) of the cell, M2 eightfold and M1 polyhedra with composition (mol %) for
663 Ca-Zn and **(b)** Ca-Co pyroxenes; **(c)** volume of the M2 eightfold polyhedron vs. average M2 ionic
664 radius for Ca-Co, Ca-Mg, Ca-Fe, and Ca-Zn pyroxenes [For references see Fig. 4].

665

666 **Figure 6:** M1-O bond length variation with composition in Ca-Zn and Ca-Co pyroxenes [For
667 references see Fig. 3].

668

669 **Figure 7:** M1 polyhedral volume vs. cation radius predicted by an “ionic model” [Data for Ca-Co
670 pyroxenes from Ghose et al. (1987); for Ca-Zn pyroxenes from Ohashi et al. (1996); for Ca-Mg
671 from Tribaudino et al. (1989); for Ca-Fe from Nestola et al. (2007)].

672

673 **Figure 8:** Average M1-O bond lengths vs. composition (Ca content in the M2 site) for Ca-Zn,
674 Ca-Co, Ca-Fe and Ca-Mg pyroxenes. It is assumed that polyhedral bond lengths are not different in
675 the monoclinic and orthorhombic phases for $\text{Co}_2\text{Si}_2\text{O}_6$, $\text{Mg}_2\text{Si}_2\text{O}_6$ and $\text{Fe}_2\text{Si}_2\text{O}_6$ (Ohashi et al. 1996;
676 Sasaki and Takéuki, 1981; Hugh-Jones et al. 1994, Morimoto et al. 1960) [For references see Fig.
677 4].

678

679 **Figure 9:** T-O bond lengths vs. composition in Ca-Zn and in Ca-Co pyroxenes. [For references see
680 Fig. 3].

681

682 **Figure 10:** (a) Difference between average bridging and non-bridging T-O distances vs. Ca content
683 in the M2 site; (b) O3-O3-O3 kinking angle with composition; (c) tetrahedral chain shift vs. Ca
684 content in the M2 site: tetrahedral chain shift $S = 2a \cos(\beta)(x_{\text{O}3} - 0.5) + 2cz_{\text{O}3}$. [For references see Fig.
685 4].

686

687 **Figure 11:** Bond valence of the O3 atoms coming from M2-M2' cations [For references see Fig. 3].

688

689 **Figure 12:** The ionic partitioning curves for M1(a) and M2 sites (b) and divalent cations of
690 different ionic radius calculated using the parameters reported for a Ca poor lherzolitic pyroxene by

691 Adam and Green (2006). The distribution coefficients for the cations are taken from the same
692 paper.

693

694 **Tables:**

695 **Table 1:** Synthesis conditions, cell parameters and phase abundance from Rietveld refinement of
696 the X-ray diffraction powder pattern.

697 **Table 2:** Structure refinement data and unit cell parameters for Ca-Zn pyroxenes.

698 **Table 3:** Fractional atomic coordinates and isotropic displacement parameters for Ca-Zn pyroxenes.

699 Data for $\text{CaZnSi}_2\text{O}_6$ and $\text{Zn}_2\text{Si}_2\text{O}_6$ are taken from Ohashi et al. (1996) and Morimoto et al. (1975),
700 respectively, here and in Table 4.

701 **Table 4:** Bond lengths (\AA), polyhedral volumes (\AA^3) and angles for Ca-Zn pyroxenes.

702

703

704

705

706

707

708

709

710

Sample	Ca ₁ Zn ₁	Ca _{0.8} Zn _{1.2}	Ca _{0.5} Zn _{1.5}	Ca _{0.3} Zn _{1.7}	Ca _{0.2} Zn _{1.8}	Zn ₂
						711
P (GPa)	4	5	5	5	5	4
T (°C)	1000	1200	1200	1200	1200	1000
<i>a</i> (Å)	9.816(1)	9.801(1)	9.787(1)	9.783(1)	9.789(2)	9.807(2)
<i>b</i> (Å)	8.990(1)	9.004(1)	9.079(1)	9.125(1)	9.157(1)	9.197(1)
<i>c</i> (Å)	5.258(1)	5.251(1)	5.2652(1)	5.273(1)	5.288(2)	5.306(2)
β (°)	105.892(1)	106.494(8)	108.223(2)	109.777(5)	110.433(3)	111.454(4)
<i>V</i> (Å ³)	446.34(1)	444.27(3)	444.40(7)	443.01(9)	443.93(2)	445.48(2)
Diffractometer	Bruker D2	Thermo ARL X'tra	Thermo ARL X'tra	Bruker D2	Bruker D2	Bruker D2
Scan range (°)	7-80° 2 θ	10-100° 2 θ	10-100° 2 θ	7-80° 2 θ	7-80° 2 θ	7-80° 2 θ
pyroxene ^a	99.8%	97.1%	99.8%	96.2%	90.9%	97.7%
coesite ^a	0.2%	2.9%	0.2%	3.8%	9.1%	2.3%

718 ^awt% as determined by Rietveld refinement.

719 .

720 Table 1

	Sample	Ca_{0.5}Zn_{1.5}	Ca_{0.3}Zn_{1.7}	Ca_{0.2}Zn_{1.8}
721	Space group	<i>C2/c</i>	<i>C2/c</i>	<i>C2/c</i>
	<i>Z</i>	4	4	4
722	<i>a</i> (Å)	9.794(6)	9.773(6)	9.779(6)
723	<i>b</i> (Å)	9.071(5)	9.116(5)	9.146(5)
	<i>c</i> (Å)	5.264(3)	5.272(2)	5.284(3)
724	β (°)	108.26(8)	109.74(8)	110.47(8)
	<i>V</i> (Å ³)	444.1(5)	442.2(5)	442.8(4)
725	Diffractionmeter	Bruker AXS Smart	Bruker AXS Smart	Bruker AXS Smart
	Radiation	MoK α	MoK α	MoK α
726	No. of indep. refl. $I > 4\sigma$	712	736	724
	$2\theta_{\max}$ (°)	63.34	64.38	64.28
727	$R_{4\sigma}$ (F)%	3.3	3.6	4.4
	wR ₂ (%)	8.4	8.6	10.0

729 **Table 2**

730

731

732

		Ca₁Zn₁[*]	Ca_{0.5}Zn_{1.5}	Ca_{0.3}Zn_{1.7}	Ca_{0.2}Zn_{1.8}	Zn₂[§]
733	M2 (Ca)	<i>y/b</i> 0.30015(5)	0.2955(3)	0.2883(7)	0.2838(10)	
		U _{eq} 0.0079(1)	0.0119(6)	0.0122(14)	0.0085(3)	
734	M2' (Zn)	<i>y/b</i>	0.2439(2)	0.2394(2)	0.2377(1)	0.2361(1)
		U _{eq}	0.0118(4)	0.0092(3)	0.0085(3)	0.0067(4)
735	M1 (Zn)	<i>y/b</i> 0.90583(3)	0.90033(6)	0.89680(7)	0.8949(9)	0.8919(1)
		U _{eq} 0.00625(6)	0.0141(2)	0.0130(2)	0.0122(2)	0.0073(2)
736	T (Si)	<i>x/a</i> 0.28631(5)	0.2936(10)	0.29784(13)	0.2996(2)	0.3016(1)
		<i>y/b</i> 0.09265(4)	0.08905(9)	0.08692(11)	0.0858(1)	0.0849(1)
		<i>z/c</i> 0.22916(9)	0.2455(2)	0.2562(2)	0.2609(3)	0.2668(2)
737		U _{eq} 0.00424(9)	0.0114(2)	0.0098(3)	0.0083(3)	0.0042(2)
738	O1	<i>x/a</i> 0.1164(1)	0.1195(3)	0.1212(3)	0.1228(4)	0.1241(3)
		<i>y/b</i> 0.0889(1)	0.0880(2)	0.0872(3)	0.0866(3)	0.0868(3)
		<i>z/c</i> 0.1442(2)	0.1442(5)	0.1452(6)	0.1458(7)	0.1473(6)
739		U _{eq} 0.0061(3)	0.0126(5)	0.0113(6)	0.0101(7)	0.0077(5)
740	O2	<i>x/a</i> 0.3602(1)	0.3695(3)	0.3747(4)	0.3761(4)	0.3787(3)
		<i>y/b</i> 0.2479(1)	0.2440(3)	0.2409(3)	0.2403(4)	0.2393(3)
		<i>z/c</i> 0.3212(2)	0.3478(6)	0.3625(7)	0.3668(8)	0.3719(6)
741		U _{eq} 0.0082(3)	0.0235(6)	0.0193(7)	0.0139(7)	0.0090(5)
742	O3	<i>x/a</i> 0.3500(1)	0.3523(2)	0.3530(3)	0.3547(4)	0.3533(3)
		<i>y/b</i> 0.0190(1)	0.0219(3)	0.0239(3)	0.0241(4)	0.0238(3)
743		<i>z/c</i> 0.9932(2)	1.0077(5)	1.0173(6)	1.0226(7)	1.0273(5)
744		U _{eq} 0.0066(3)	0.0152(5)	0.0150(6)	0.0125(7)	0.0084(5)

745

746

Notes: *x/a* and *z/c* for the M2 and M1 special position are 0 and 0.25. * Data from Ohashi et al. (1996).
[§]Data from Morimoto et al. (1975).

747

748

749 **Table 3**

750

751

		Ca_1Zn_1^*	$\text{Ca}_{0.5}\text{Zn}_{1.5}$	$\text{Ca}_{0.3}\text{Zn}_{1.7}$	$\text{Ca}_{0.2}\text{Zn}_{1.8}$	Zn_2^{\S}
M2	Ca-O1A1, B1×2	2.356(1)	2.374(3)	2.347(6)	2.337(8)	
	Ca-O2C2, D2×2	2.336(1)	2.133(4)	2.014(4)	1.975(4)	
	Ca-O3C1, D1×2	2.601(1)	2.604(3)	2.644(6)	2.665(8)	
	Ca-O3C2, D2×2	2.729(1)	2.812(3)	2.889(5)	2.934(7)	
	<Ca-O>	2.505	2.480	2.474	2.478	
M2'	Zn-O1A1, B1×2		2.023(3)	2.018(3)	2.029(4)	2.031(3)
	Zn-O2C2, D2×2		2.105(4)	2.004(4)	1.973(4)	1.933(2)
	Zn-O3C1, D1×2		2.987(3)	3.018(4)	3.022(4)	3.027(2)
	Zn-O3C2, D2×2		3.110(3)	3.174(3)	3.205(4)	3.255(3)
	<Zn-O>		2.556	2.554	2.557	2.561
	$V_{\text{M2}}(\text{\AA}^3)$ 8-coord.	25.91(11)	24.91(11)	24.39(13)	24.29(12)	24.27(12)
M2 along b axis ^a	3.690	3.737	3.763	3.780	3.800	
M1	Zn-O1A2, B2×2	2.075(1)	2.041(3)	2.025(4)	2.022(4)	2.018(3)
	Zn-O1A1, B1×2	2.158(1)	2.233(2)	2.272(3)	2.299(4)	2.333(3)
	Zn-O2C1, D1×2	2.075(1)	2.079(3)	2.089(3)	2.090(4)	2.084(3)
	<Zn-O>	2.103	2.117	2.129	2.137	2.145
	$V_{\text{Zn}}(\text{\AA}^3)$	12.26(5)	12.51(5)	12.62(5)	12.76(5)	12.83(7)
	OAV ^b	18.5	30.7	36.5	42.2	51.5
M1 along b axis ^a	1.427	1.415	1.401	1.406	1.402	
T	Si-O1	1.604(1)	1.619(3)	1.625(4)	1.620(4)	1.617(3)
	Si-O2	1.585(1)	1.603(3)	1.601(3)	1.605(4)	1.604(3)
	Si-O3A1	1.668(1)	1.649(3)	1.633(3)	1.634(4)	1.625(3)
	Si-O3A2	1.687(1)	1.659(3)	1.645(3)	1.640(4)	1.627(3)
	<Si-O>	1.636	1.632	1.626	1.625	1.618
	$V_{\text{T}}(\text{\AA}^3)$	2.24(1)	2.22(2)	2.19(2)	2.19(2)	2.16(2)
	TAV ^b	27.9	16.2	13.1	12.0	11.6
	O3A1-Si-O3A2	104.2(1)	107.2(1)	109.2(1)	109.8 (1)	111.2(1)
	Si-O3-Si	135.7(1)	138.0(2)	139.6(2)	139.6(2)	141.6(2)
	O3-O3-O3	165.2(7)	162.8(2)	161.2(3)	161.0(3)	161.3(1)
	O3A1-O3A2	2.647(1)	2.662(2)	2.672(2)	2.679(2)	2.683(4)
T along b axis ^a	2.263	2.322	2.362	2.375	2.397	

752

753

754

755

Notes : ^aM2 contribution: $(0.5-y/bO1)b$; M1 contribution: $(y/bO2-y/bO1)b$; T contribution: $(0.5-y/bO2)b$.

^bOAV and TAV are angle variance for tetrahedra and octahedra (Robinson et al. 1971).

* Data from Ohashi et al. (1996). [§]Data from Morimoto et al. (1975).

756

757

758 Table 4

759

760























