Revision 4: Compositional variation and timing of aluminum phosphate-sulfate minerals in the basement rocks along the P2 fault and in association with the McArthur River uranium deposit, Athabasca Basin, Saskatchewan, Canada

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Abstract

The Athabasca Basin hosts world class uranium deposits, such as the McArthur River deposit. This paper presents the occurrence of aluminum phosphate-sulfate (APS) minerals in the metasedimentary rocks along the P2 fault, the main ore-hosting fault of the McArthur River deposit. It compares the APS minerals along the P2 fault with those outside the fault, examined in this study, and those from other deposits of the Athabasca Basin and from other Paleo- to Mesoproterozoic basins worldwide.

APS minerals are common along the P2 fault, rare outside of the P2 fault zone in the basement, and sparse along the unconformity between the Athabasca sandstones and the basement. The APS minerals along the P2 fault occur with sudoite (± illite, magnesiofoitite) and are zoned with Sr-, Ca- and S-rich cores (solid solution between svanbergite, crandallite and goyazite) and LREE- and P-rich rims (close to florencite composition). APS minerals in the Bleached Zone (altered rocks along the unconformity consisting predominantly of kaolin and illite) are Sr-, Ca- and S-rich (high svanbergite component) and occur with kaolin. APS minerals in the Red-Green Zone (mingled red hematitic and green chloritic basement rocks below the
Bleached Zone) occur with sudoite and clinohlore. They contain relict cores of LREE- and As-rich arsenoflorencite-(Ce), and rims of svanbergite-goyazite-crandallite solid solution.

The occurrence of svanbergite-crandallite-goyazite along the unconformity suggests their formation by relatively oxidizing fluids during diagenesis of the overlying sandstones. The relict cores of arsenoflorencite-(Ce) in the Red-Green Zone are interpreted to be the product of paleo-weathering before the deposition of the Athabasca sandstones. Florencitic APS minerals are found along the entire studied strike length (7 km) of the P2 fault, including the ore zone and non-mineralized areas, but absent outside the fault zone. The florencitic APS minerals contain low \( \text{SO}_4^{2-} \) in the ore zone, suggesting relatively reducing conditions during their crystallization.

Zoned APS minerals (with svanbergitic cores and florencitic rims) proximal to ore contain elevated \( \text{U} \) (up to 16 ppm). These features suggest that diagenetic, oxidizing, and uranium-bearing fluids travelled along the P2 fault and became relatively reduced, especially within the ore zone. It also suggests florencitic APS minerals are contemporaneous with uranium mineralization. The restricted occurrence of florencitic APS mineral along the P2 fault in the basement suggests their use in identifying fertile basement structures associated with uranium mineralization.

**Keywords:** hydrothermal alteration, APS, uranium mineralization, florencite, svanbergite, arsenoflorencite, diagenesis, paleo-weathering, unconformity-type uranium deposits

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**Introduction**

The Athabasca Basin hosts numerous large uranium deposits (Fig. 1), which are classified as unconformity-related uranium deposits. The prevalent model for the mineralization was first proposed by Hoeve and Sibbald (1978) and modified later by many researchers (e.g. Alexandre et
al. 2005; Derome et al. 2005; Richard et al. 2011; Mercadier et al. 2012). This model suggests that sea-water derived, uranium-bearing oxidizing brines (25-35 wt.% eq. NaCl) precipitated uranium ore at the unconformity when it reacted with a reducing fluid of currently unknown origin. Many deposits occur in the proximity of deformation zones in the basement and it is suggested that re-activated basement faults served as conduits for uranium-bearing fluids (Jefferson et al. 2007) or reducing basement fluids to reach the unconformity (McGill et al. 1993). The McArthur River deposit, the largest discovered high-grade (average grade of 16.46 %U₃O₈; Bronkhorst et al. 2012) uranium deposit on Earth, is situated along the P2 fault, a 13 km long reverse fault constrained to graphitic metapelite below the Athabasca Basin. Although geophysical and structural studies have been carried out on the P2 fault (e.g., Hajnal et al., 2010), the exact role of the fault in the mineralization remains uncertain.

In the Athabasca and Thelon basins of Canada, aluminum phosphate-sulfate (APS) minerals (also known as alumino-phosphate sulfate) spatially associated with uranium deposits have been reported (Quirt et al. 1991, Gaboreau et al. 2007, Riegler et al. 2014). APS minerals occur in a variety of environments (Dill, 2001) and form many solid solution series with over 40 end-members (Jambor 1999; Stoffregen et al. 2000; Dill 2001). APS minerals have the general formula: AB₃(XO₄)₂(OH)₆, where A = mono-, di-, tri- or, more rarely, tetravalent cations (K⁺, Na⁺, Rb⁺, Ca²⁺, Sr²⁺, REE³⁺, Th⁴⁺ etc.); B = trivalent cations (Al³⁺, Fe³⁺); and X = P⁵⁺, S⁶⁺, or As⁵⁺. Since incorporation of ions with different valences requires coupled substitution to accommodate charge balance, compositional zoning of APS minerals is commonly well preserved. This makes APS minerals excellent candidates for evaluating the chemical and physical characteristics of fluids from which they formed (Dill 2001; Beaufort et al. 2005; Gaboreau et al. 2005, 2007).
The focus of this paper is to present the occurrence and compositional variation of APS minerals in basement rocks along the P2 fault, discuss the role of the fault during mineralization and the timing of APS mineral formation with respect to uranium mineralization. In addition, the chemistry of APS minerals is compared with those along and outside of the fault zone in the basement, from other APS minerals studied in Athabasca Basin (within deposits or not), and from other Paleo- to Mesoproterozoic basins worldwide. The potential use of APS mineral in exploration for unconformity-related uranium deposits is also discussed.

**Geological Setting**

**Regional Geology**

The basement rocks in the eastern Athabasca Basin area are part of the Wollaston and Mudjatik Domains. The rocks of the Wollaston Domain consist of metapelite, metasemipelitie, meta-arkose, calc-silicate and quartzite (e.g., Annesley et al. 2005). They were metamorphosed under upper amphibolite to lower granulite conditions during the Trans-Hudson Orogeny (ca. 1.8-1.9 Ga; Lewry and Sibbald 1980, McKechnie et al. 2012). Abundant granitic pegmatite lenses intruded the metasedimentary rocks during peak metamorphism (Annesley et al. 2005; Jefferson et al. 2007).

Alteration immediately below the unconformity is laterally extensive and consists of a deeper Green Zone with sudoite (Al-Mg chlorite) and illite ± Fe-Mg chlorite, a transitional Red-Green Zone with illite, sudoite and hematite, an upper Red Zone with hematite and kaolin ± illite, and, locally, there is Bleached Zone of kaolin and illite, immediately beneath the unconformity (Macdonald 1985; Quirt 2001; Adlakha et al. 2014). In addition to the Red-Green transitional Zone, the Red and Green Zones are mingled in places. The origin of these Zones has been in
discussion. Macdonald (1985) considered them as the products of paleo-weathering before sandstone deposition, whereas Cuney (2003) proposed their formation after the deposition of sandstones. Quirt (2001) and Adlakha et al (2014) pointed out extensive overprinting of alteration minerals by subsequent hydrothermal fluids, which makes it difficult to determine the condition of their initial crystallization. Rocks of the Bleached Zone are altered to white due to loss of hematite and interpreted to be a reaction product of the Red Zone with late diagenetic fluids from the overlying basin (Macdonald 1985; Quirt 2001). The least-altered metapelite occurs > 50 m below the unconformity and away from faults.

The sedimentary rocks of the Athabasca Group are comprised of dominantly fluvial to shallow marine, quartz pebble conglomerate and quartz arenite, with a preserved maximum thickness of 1500 m in the centre of the basin (Ramaekers et al. 2007). Early diagenesis included silica cementation on hematite-dusted detrital quartz grains (Hoeve and Quirt 1984; Jefferson et al. 2007), alteration of kaolinite by dickite with minor illite and locally traces of sudoite (Quirt 2001; Laverret et al. 2006). Uranium deposits are localized near the intersection of the unconformity and basement faults or down-dip along these faults into the basement. These deposits are surrounded by alteration haloes of sudoite, illite and magnesio-foitite (alkali-deficient dravite). Sandstone-hosted unconformity-type deposits show the inner halo of sudoite with an outer alteration halo of illite (e.g., Hoeve et al. 1981; Hoeve and Quirt 1984; Alexandre et al. 2012). Basement-hosted unconformity-type deposits are accompanied by the inner alteration halo of illite and the outer halo of sudoite. The different proportions of sudoite and illite between two types of unconformity-type deposits are considered to be related to differences in fluid circulation (i.e. Ingress vs Egress; Hoeve and Quirt 1984).
Structure and alteration of the McArthur River deposit and P2 fault

The McArthur River deposit is comprised of six ore bodies, Zones 1-4 and Zones A and B, which collectively have a strike length of approximately 1.7 km near the intersection of the P2 fault (~050°/45-60°) and the unconformity, ~500 m below surface (Fig. 2c) (McGill et al. 1993; Alexandre et al. 2005; Bronkhorst et al. 2012). The total reserves of the McArthur River deposit as of August 31, 2012, were 385.5 Mlbs U₃O₈ (Bronkhorst et al. 2012). All ore bodies are hosted in the sandstone immediately above the unconformity except for the Zone 2 ore body which occurs immediately below the unconformity and is bounded by the vertical quartzite (VQ) fault and ~140° trending series of cross faults (Fig. 2b). Late minor remobilizations of uraninite formed veinlets and fractures with pyrite, chalcopyrite and minor nickel-cobalt sulfarsenides (McGill et al. 1993).

The P2 fault zone is a series of reverse faults with a strike length of approximately 13 km (McGill et al. 1993) and has been traced seismically at least 2 km below the unconformity (Hajnal et al. 2010). In the basement, the several fault planes of the P2 fault are mostly constrained to graphite-bearing metapelite and trend parallel to the basement foliation (McGill, 1993). Reactivation of P2 faulting formed differences in the thickness of the basal conglomerate at the unconformity and broad fracture and breccia zones in the sandstones and the reverse movement of the P2 fault raised a wedge of basement rocks above the unconformity with the vertical displacement of up to 80 m (McGill et al. 1993; Figs. 2a,b). The basement wedge is also referred to as the “middle block” by workers (Bronkhorst et al. 2012) as it is offset by differential movement of fault planes (Figs 2a,b). These offsets at the apex (or the “nose”) commonly hosts the uranium mineralization of McArthur River deposit and sporadic occurrences of low-grade (<1 %U₃O₈) mineralization along the P2 fault (Fig. 2a). The P2 fault and basement rocks show minor
lateral displacement by three sets of steeply dipping faults that strike 100 – 110° in the mine-site area (Fig. 2b). These faults and offsets likely provided pathways for mineralizing fluids, but also truncate high-grade zones, especially the Zone 2 ore body (Fig. 2b; McGill et al. 1993; Bronkhorst et al. 2012).

Basement rocks along the P2 fault were intensely altered to form sudoite, illite, magnesiofoitite, Fe-Mg chlorite, and sulfides (McGill et al. 1993; Adlakha et al. 2014). The hanging wall basement rocks, including the ore-hosting wedge, commonly show intense alteration forming sudoite and illite ± kaolin. Formation of illite and/or kaolin makes rocks creamy white in colour, and this alteration is called “bleaching”. Bleaching is common, especially in close proximity to the fault, mineralization and the unconformity (McGill et al. 1993). The most intense alteration of the basement rocks is found in close proximity to the Zone 2 ore body where sudoite and magnesiofoitite are the dominate alteration minerals (Adlakha et al. 2014). In comparison, illite with minor sudoite are the dominate alteration minerals in the moderately altered basement rocks in unmineralized areas (the south-west portion of the P2 fault) and near the weakly mineralized P2 Main deposit (Fig. 2b), that is “perched” higher in sandstones. APS minerals only occur in highly altered rocks where primary metamorphic minerals and textures are altered.

**Sampling**

A total of 192 samples of metasedimentary rocks, predominately metapelites with fewer pegmatite and quartzite, were collected from 27 different drill-holes along the P2 fault zone from the McArthur River mine site to ~7 km southwest (Fig. 3). Of these, 139 samples were selected for examination of APS minerals in thin sections. Out of the 139 samples, APS minerals were found in only 17 samples. Among these 17 samples, APS minerals in 4 samples were too small for quantitative analysis with electron microprobe analyzer (EPMA). Of these remaining 13
samples (Table 1), nearly all were from metapelites located along the P2 fault (11 of 13) that showed intense/pervasive alteration in which clay-sized minerals replaced most metamorphic minerals and metamorphic textures were obliterated by alteration, fractures and/or veinings. In this paper, samples collected within 50 m of the centre of the fault are considered to be associated with P2 fault, as the fault is accompanied by splay structures and alteration extending outside the fault zone.

The samples associated with the P2 fault were collected from varying distances to the McArthur River uranium deposit (Fig. 2c; Table 1). MAC425 was collected within the Zone 2 ore body (Fig. 3A), MAC152 < 15 m from the Zone 2 ore body, MAC67 (Fig. 3b) and MAC85 < 20 m from the Zone 2 ore body, MAC63 < 50 m from the Zone 2 ore body, MAC43 and MAC44 < 50 m from the Zone 3 ore body. MAC29 was within 1 m of low grade (< 1 wt% U₃O₈) uranium mineralization disseminated within the sandstone (Fig. 2a), MAC404 was approximately 50 m from low grade uranium mineralization within the sandstone. MAC117 and MAC35 are > 500 m away from low grade uranium mineralization in the sandstone, > 1 km from the low grade P2 Main deposit and > 2 km from the McArthur River deposit.

Although 5 samples from the Bleached Zone were examined in detail, APS minerals were found only in one sample (MAC17; Fig. 3c) of quartzite along the unconformity, > 250 m from low grade mineralization and > 1 km from the McArthur River deposit (Fig. 2a,c). Twenty-two samples of metasedimentary rocks of the Green, Red-Green and Red Zones outside the P2 fault were examined and APS minerals were found only in one sample (MAC412) of strongly hematized and chloritized metapelite from the Red-Green Zone in the footwall near the P2 Main deposit (Figs. 2c, 3d). Here we use field terminology “Bleached Zone” and the “Red-Green Zone” to reflect the samples position within the basement rocks as these terms were commonly
used by geologists in the area. The Bleached Zone occurs along the unconformity and it refers to a creamy-white colored zone where kaolin and illite predominate. The Red-Green Zone is a physical mixture of hematitic Red Zone and chloritic Green Zone.

**Methods**

Thin sections were examined using an optical petrographic microscope with reflected and transmitted light sources. Due to the very fine-grained nature of minerals in the samples, detailed textural analysis was also carried out on carbon-coated polished thin sections using a JEOL 6610LV scanning electron microscope (SEM) at the University of Ottawa. Minerals were identified using energy-dispersive spectroscopy with the spectrum acquisition time of 40 s and an accelerating voltage of 20 kV.

Quantitative analysis for the major and minor elemental compositions of the APS minerals was carried out using a JEOL 8230 EMPA at the University of Ottawa. APS minerals were analyzed at an accelerating voltage of 15 kV, a beam current of 20 nA using a focused beam of 5 μm (to account for alkali migration). Counting times were 20s on-peak and 10s off-peak for all elements except for S (10s on and 5 s off), F (50s on and 25s off), Pr (20s on and 5s off) Cl (25s on 12s off), P (10s on and 5s off), Si (10s on and 5s off) and As (20s on and 20 s off). The mineral standards used were sanidine (Si, Al, K), hematite (Fe), apatite (Ca, P, F), albite (Na), turgutite (Cl), celestine (Sr, S), synthetic NdPO₄ (Nd), PrPO₄ (Pr), CePO₄ (Ce), LaPO₄ (La), ThO₂ (Th) and GaAs (As). The elemental concentrations were calculated after matrix interference correction using ZAF. Elemental maps of selected APS grains were acquired at an accelerating voltage of 10 kV with a 10 nA current at a counting time of 0.5 s per pixel with a total of 10000 pixels per image.
APS minerals are small (mostly < 20 μm, rarely up to 40 μm) and surrounded by quartz and fine-grained alteration minerals (Fig. 4; Table 1-2), but it was necessary to use a broad electron beam of 5 μm to minimize the loss of alkalis from APS minerals during the EPMA analysis. This made it unavoidable for the surrounding minerals to contribute to the analytical results. To evaluate the possible contributions of surrounding minerals in the chemical composition obtained for the APS, their compositions were measured by EMPA. The contents of SiO₂ were used to monitor the contribution of surrounding minerals because APS minerals rarely contain Si. For the data showing Si, the amounts of K, Al, and Na of the surrounding phases were subtracted from the analytical results. For example, for Al₂O₃ using the equation, \( \text{Al}_2\text{O}_3^{\text{APS corrected}} = \text{Al}_2\text{O}_3^{\text{APS analyzed}} - (\text{SiO}_2^{\text{APS analyzed}} \times \text{Al}_2\text{O}_3^{\text{matrix}} / \text{SiO}_2^{\text{matrix}}) \). The remaining concentration values were normalized to the original analytical total. Table 2 shows the contents of SiO₂ before subtracting the contribution of surrounding minerals. The contribution of surrounding minerals never exceeded 5% of the total.

Trace element contents in the APS minerals were determined using polished thick (>60 μm) sections by laser ablation inductively coupled plasma mass spectrometer (LA-ICP-MS) at the Geological Survey of Canada, Ottawa, using a Photon-Machines Analyte193 ArF excimer laser (\( \lambda = 193 \mu m \)) with Helex ablation cell and Agilent 7700x quadrupole ICP-MS. Helium gas was used to transport ablated sample from the ablation cell and was mixed with Ar plasma (flow rate of 1.05 mL/min) via a T-connector before entering the ICP. The detailed operating conditions for the LA-ICP-MS instrument and the isotopes quantified are listed in Appendix Table 1. During analysis, ~40 s of background signal with the laser turned off was collected and then ablation was performed by a laser focused to 10 μm for a data acquisition time of ~60 s (total time of 100 s for one analysis) for a total time of 100 s for one analytical cycle. The isotope...
counts were monitored during the analysis, and the acquisition time for the data of APS minerals was adjusted (Appendix Fig. 1). Aluminum contents are similar among grains of APS minerals. Therefore, the counts of these isotopes and their count ratios to Al in the sample and the reference NIST610 (National Institute of Standards and Technology) were used to calculate the element abundances (Longerich et al. 1996). There is no international standard that has the composition similar to APS minerals. Therefore, the standard NIST610 was used as it contains high Al and it is the most used standard. The elemental concentrations were obtained based on the count ratios of elements to Al and the average Al contents in APS minerals of the sample determined by EMPA (Table 2). The Al content was used because Al does not vary between the core and rim of APS minerals (Table 2). The same grains could not be analyzed both by EPMA and LA-ICP-MS because electron beam decomposes APS minerals (Appendix Fig. 2). EMPA and LA-ICP-MS data is comparable (Appendix Fig. 3). One round of LA-ICPMS analysis consisted of seven analyses of grains, bracketed by two analyses of the NIST610 reference and one analysis of the BCR2G (Basalt, Columbia River) reference from USGS. The analysis of NIST610 shows that a counting precision of > 98% for all elements. Accuracy, determined by BCR2G, was > 90% for all elements except for Sc, Ti, Fe, Zn, Ce, Pr, Sm, Ho, Lu, Gd, Tb, Pb and Th which were > 85% and Cu, Y, Zr, Cr, Tm, Yb and Hf which were > 75%.

Results

Mineral occurrence and distribution

APS minerals were found in metapelites along the 7 km section of the P2 fault examined (Fig. 2). APS minerals were found only in one sample from the Bleached Zone (sample MAC17 in quartzite) (Figs. 2a, 4i; Table 1). APS minerals are rare in the basement rocks distal from the
unconformity and outside the P2 fault. They were found in only one sample (sample MAC 412) in metapelite that contains hematite and chlorite alteration (Fig. 3d) in the footwall 100 m from the P2 fault, 50 m below the unconformity and 100 m away from uranium mineralization (Fig. 4h; Table 1). APS minerals were not observed in samples along the VQ fault (Fig. 2b) or in samples of the least-altered basement rocks.

APS minerals along the P2 fault occur together with fine-grained sudoite and/or illite intergrown with magnesiofoitite (Adlakha et al. 2014; Figs. 4a-f). Less commonly, APS minerals occur with kaolin and illite, or a mixture of clinochlore and sudoite. In the Bleached Zone sample, APS minerals occur with kaolin and illite (Fig. 4i). APS minerals are rare in the transitional Red-Green Zone and occur predominately with clinochlore (Fig. 4h).

APS minerals are small (mostly < 20 µm, rarely up to 40 µm), transparent, pseudo-cubic (rarely hexagonal) crystals, disseminated and/or clustered in fine-grained clay minerals (Fig. 4). There is no evidence of dissolution between APS minerals and surrounding clay minerals. The texture suggest that they are in equilibrium (Figs. 4a-j,l). APS minerals are commonly zoned. Zoning may be observed using plane polarized light microscopy as some cores show high relief. SEM BSE-images commonly show darker cores with lighter rims (Fig. 4c), although grains with lighter cores and darker rims are also observed (Fig. 4h). Darker parts in BSE images are due to greater amounts of Sr and Ca, and brighter parts contain higher LREE.

In all samples, APS minerals commonly occur in contact and/or are spatially associated with zircon (Figs. 4f,g), primary apatite (Fig. 4l), or sulfide minerals, such as pyrite (Fig. 4k) and lesser chalcopyrite (Fig. 4d). Zircon grains associated with APS minerals are commonly fractured and/or partially dissolved and rimmed by xenotime (cf. Quirt et al. 1991) (Fig. 4g). Apatite forms
subhedral, rounded crystals and APS minerals also occur with small hydrothermal fluorapatite
crystals as disseminations or clusters. APS minerals commonly contain inclusions of sulfides or
Fe-oxide (Figs. 4h,k,l). APS minerals in the Zone 2 ore body are in close proximity (within the
same thin section) with uraninite.

APS mineral chemistry – major elements

The APS mineral formulae were calculated based on 6 cations (Table 2) as the ideal
formula is expressed as AB₃(XO₄)₂(OH)₆, where A = mono-, di-, tri- or, more rarely, tetravalent
cations (K⁺, Na⁺, Rb⁺, Ca²⁺, Sr²⁺, REE³⁺, Th⁴⁺ etc.); B = Al³⁺, Fe³⁺; and X = P⁵⁺, S⁶⁺, and As⁵⁺.

Cation charges were allocated using the following:

i. the X-site was filled with P⁵⁺, S⁶⁺, As⁵⁺,

ii. the B-site was filled by Al³⁺, then Fe³⁺ to have 3 cations,

iii. the A site was then filled by remaining cations.

High contents of trivalent cations in the A site create excess cation charges. To compensate for
this, PO₃(OH)²⁻ is assigned for (PO₄)³⁻, as has been used by previous researchers (e.g., Gaboreau
et al. 2007).

Along the P2 fault, APS minerals occur predominately with sudoite and lesser illite and
magnesiofoitite, in intensely altered metapelite and pegmatite. These APS minerals contain high
ΣLREE (0.35 – 0.58 apfu; LREE = La + Ce + Pr + Nd) relative to Sr (0.16 – 0.32 apfu) and Ca
(0.18 – 0.28), and high PO₄³⁻ (1.70 - 1.97 apfu) relative to SO₄²⁻ (0.06 – 0.35 apfu) and AsO₄³⁻
(<0.01 apfu) (Table 2). Ce is high (0.15 – 0.27 apfu) relative to La (0.09 – 0.17), Pr (<0.01 – 0.03
apfu), producing [Ce]N / [Ce]N* of ~0.8 – 1.6, where [Ce]N* = \sqrt{([La]N*+[Pr]N*)}. Cores of zoned
grains are enriched in Sr-, Ca-, and S, and rims contain high LREE and P (Fig. 5a; Table 2).
Within the cores, Ca and P contents are inversely correlated with Sr and S contents (Fig. 5a). Narrow width of zoning and small size of grains prevented acquisition of quantitative compositions of individual zones excluding coarse grains in MAC152 (Table 2). Therefore, most data presented in Table 2 represents mixtures (referred to as mixed data) of compositional zones identified in SEM-BSE.

APS minerals in the sample of Bleached Zone quartzite (sample MAC17) contain high Sr (0.68 apfu) relative to Ca (0.16 apfu) and ΣLREE (0.13 apfu) and high SO$_4^{2-}$ (0.54 apfu; Table 2).

APS minerals are abundant in the Red-Green Zone sample of the basement footwall (MAC412) near the P2 Main deposit. APS minerals occur with clinohlore and sudoite and contain high AsO$_4^{3-}$, (~0.63 apfu), Fe$^{3+}$ (0.16 apfu), and Ce (~0.32 apfu) showing positive Ce anomalies ([Ce]/[Ce]*) of up to 6. APS minerals from this sample are also zoned with high-Ce cores, and a Ca- and P-rich inner rim and Sr-rich outer rim (Fig. 5b). APS minerals in this sample were relatively coarse-grained; therefore, individual core and rims were analyzed in two grains (Table 2). The data shows that the cores contain high As (0.78 apfu) and low P (0.92 apfu) relative to the rim (0.26 and 1.46 apfu, respectively; Table 2).

**APS mineral chemistry – trace elements**

LA ICP-MS analyses were conducted on relatively coarse-grained (>10 μm) APS minerals proximal to the Zone 3 (sample MAC44) and Zone 2 ore deposits (samples MAC63 and MAC85). Counts of all elements were monitored during the analysis to detect mineral inclusions and contributions of surrounding minerals. Mineral inclusions should give sharp increases in the counts of elements, but no such spikes were observed during the analysis, confirming the lack of mineral inclusions. None of surrounding minerals contain any significant concentrations of
elements described below. Counts of elements, such as U, Y, Pb and Th were at background levels in surrounding minerals and only high when the laser beam hit APS grains. Unlike clays, APS minerals show high counts of LREE, Sr, Ca and P (Appendix Fig. 1). This confirms that these elements are in the APS crystal structure and were not incorporated from surrounding minerals. Furthermore, different grains from individual sections yielded comparable concentrations of minor and trace elements, suggesting that these trace elements originate from APS minerals and that the values are minima since the dilution effects by the surrounding minerals were not corrected. The dilution effects by surrounding minerals are confirmed by the detection of a significant Si (48000 – 99000 ppm) during the ablation of the APS minerals.

APS minerals in P2 fault proximal to Zone 3 (MAC44) ore deposit contain appreciable Se (130 ppm), U (0.39 – 0.88 ppm), Pb (160 - 180 ppm), and Th (1100 - 1600 ppm). Similarly, APS from sample MAC63 (~50 m) from Zone 2 ore and ~ 20 m into the hanging-wall of the P2 fault show appreciable Se (100 – 160 ppm), U (1.7 – 4.2 ppm), Pb (82 – 190 ppm), and Th (1000 – 1400 ppm). Sample MAC85, from within 20 m of Zone 2 ore and directly within the P2 fault, contains APS with relatively high U (16 ppm), Pb (260 ppm), and Th (2300 ppm); this sample also contains high LREE (0.54 apfu; Table 2). Similar to samples MAC44 and MAC63, sample MAC85 showed appreciable Se (130 ppm). Combined with the EPMA data of major elements (La, Ce, Pr, Nd), these APS minerals from the P2 fault area contain high LREE relative to HREE ([LREE]_N/[HREE]_N = 800 – 2500; Fig. 6) and show low S/Se weight ratios, ranging from 100 – 150.
Discussion

Major element compositions of APS minerals

The following elemental relationships are observed:

i. the P contents positively correlate with LREE and both inversely correlate with Sr and S (Figs. 5a, 7c,d); the LREE inversely correlate with Ca (Figs. 5a, 7d);

ii. the S contents positively correlate with Sr and inversely correlate with Ce (Figs. 5a, 7b).

iii. high Ca correlates with high P (Figs. 5a,b);

iv. the As contents inversely correlate with P;

v. high As APS minerals contain high Ce relative to other REE (Fig. 7a).

The above relationships and cations in A-site and X-site (Fig. 8) suggest the presence of the following end-members:

i. high LREE and P contents correspond to high proportions of florencite (REE)Al$_3$(PO$_4$)$_2$(OH)$_6$ (Figs. 8a-c);

ii. high S and Sr contents correspond to high proportions of svanbergite SrAl$_3$(PO$_4$)(SO$_4$)(OH)$_6$ with a minor goyazite component SrAl$_3$(PO$_4$)$_2$(OH)$_6$ (Figs. 8a-d);

iii. high Ca contents correspond to high proportions of crandallite CaAl$_3$(PO$_4$)$_2$(OH)$_6$ (Figs. 8a,b,d);

iv. high As and high Ce contents correspond to high proportions of arsenoflorencite (Ce, REE)Al$_3$(AsO$_4$)$_2$(OH)$_6$ (Figs. 8a,b).

APS minerals show wide compositional variations, but all can be expressed using the above five end members (crandallite, goyazite, svanbergite, florencite, arsenoflorencite).
Spatial, chemical and temporal variation of APS minerals and associated clay minerals

Along the entire studied range of the P2 fault, APS minerals occur with sudoite (± illite, magnesiofoitite) in metapelite and pegmatite. The APS minerals are florencitic, but contain significant Ca, Sr, and SO\textsubscript{4}\textsuperscript{2-} reflecting a minor component of svanbergite, mainly as core (Fig. 9).

Gaboreau et al. (2005, 2007) found that APS minerals proximal to uranium deposits contain high LREE and P relative to those in intermediate and distal areas (Fig. 9). The present study found a large compositional variation of APS minerals even within the ore compared to the range presented by Gaboreau et al. (2005, 2007) and compositional variation do not correlate with distance from uranium mineralization (Fig. 9). For example, the (LREE+P)/(Ca+Sr+S) ratios of APS minerals near (<1 m) the uranium ore range from 1.6 to 5.6, whereas those from barren samples far (>500 m) from mineralization vary from 2.8 - 4 (Fig. 9). We suggest that our data reflects the evolution of fluids in the ore zone where the earlier fluids contained high in Ca, Sr and SO\textsubscript{4}\textsuperscript{2-}, and evolved to have high LREE and P. This proposed interpretation is supported by compositional zoning of APS minerals with Ca, Sr and S -rich cores and LREE and P -rims (Fig. 5a).

APS minerals are abundant with kaolin and illite in one sample of the Bleached Zone (MAC17) along the unconformity. These grains are Sr- and S-rich and similar in composition within the sample but not with the APS minerals found along the P2 fault (Fig. 9). The average composition of the APS minerals is

\[
\text{Ca}_{0.16}\text{Sr}_{0.68}\text{Ce}_{0.05}\text{La}_{0.03}\text{Nd}_{0.04}\text{Pr}_{0.01}(\text{Al}_{2.96}^{\text{III}}\text{Fe}_{0.02})(\text{PO}_3\text{O}_{0.94}\text{OH}_{0.06})_{1.50}(\text{SO}_4)_{0.54}(\text{AsO}_4)_{0.01}(\text{OH})_{6},
\]
svanbergite$_{0.54}$-crandallite$_{0.16}$-goyazite$_{0.14}$-florencite$_{0.13}$. The data are comparable to APS mineral compositions in distal sandstones by Gaboreau et al. (2007) (Fig. 9).

One Red-Green Zone sample of intensely altered metapelite (MAC412) footwall to the P2 fault, contain APS minerals with As- and Ce-rich cores and Ca, Sr, and S-rich rims together with clinochlore and sudoite. The grains show similar Ca, Sr, and S contents with varying LREE and P due to the presence of As that substitutes for P (Figs. 7a and 9). The average composition of these grains is

$$(Ca_{0.31}Sr_{0.22}Ce_{0.32}La_{0.04}Nd_{0.02}Fe_{0.15})_2(Al_{2.92}Fe_{0.08})(PO_{3.83}OH_{0.17})_1(SO_4)_{0.18}(AsO_4)_{0.68}(OH)_6$$

arsenoflorencite(-Ce)$_{0.39}$-crandallite$_{0.31}$-svanbergite$_{0.18}$-goyazite$_{0.12}$. This is the first documented occurrence of arsenoflorencite(-Ce) in Athabasca Basin.

APS minerals from the McArthur River deposit and the P2 fault show cores of svanbergite-crandallite-goyazite rimmed by florencite (Figs. 4c-e,k,l, 5a). APS minerals from the Red-Green Zone contain arsenoflorencite(-Ce) cores rimmed by svanbergite-crandallite-goyazite (Figs. 4h, 5b, 8b). APS minerals of the Bleached Zone are svanbergite-crandallite-goyazite (Fig. 8). The zoning suggests the crystallization sequence of APS minerals in the study area. Early crystallization of arsenoflorencite (Ce) was followed by svanbergite-crandallite-goyazite, and finally florencite.

**Nature of fluids responsible for APS mineral formation**

APS minerals in the Bleached Zone sample contain high SO$_4^{2-}$ contents, up to 0.59 apfu (Figs. 7b,d, 8b-d), indicating relatively oxidizing conditions for their formation. These APS minerals have close to svanbergite end-member composition (Fig. 10), with a minor goyazite and crandallite component (Figs. 7b,d-f; 8a-d). The grains coexist with kaolin and illite (Fig. 4i), which is the diagenetic assemblage in the Athabasca Group sandstones (Hoeve and Quit 1984;
Quirt 2001). Furthermore, the composition of the APS minerals is similar to those of diagenetic origin reported in Athabasca and Kombolgie sandstones (Gaboreau et al. 2005, 2007), and other sedimentary basins worldwide (Spötl, 1990; Gaboreau et al. 2005, 2007; Pe-Piper and Dolansky 2005; Gall and Donaldson 2006; Fig. 10). Therefore, the APS minerals of the Bleached Zone along the unconformity most likely crystallized from an oxidizing, Ca-, Sr- and SO₄-rich, diagenetic fluid of the overlying sandstones.

The oxidizing conditions of diagenetic fluids from sandstones are recorded as relict cores of the APS minerals along the P2 fault in the basement (Figs. 4, 5a). These cores contain high Ca, Sr and S and are similar in composition to diagenetic APS minerals in sandstones described above. The S/P ratios of APS minerals as whole grains along the P2 fault are relatively low (<0.16) compared to those of the Bleached Zone (0.36; Fig. 11). The lowest S/P ratios (<0.06) are found in areas of Zone 2 ore in the basement (Fig. 11b). The data shows that the SO₄²⁻ contents in fluids decreased from the unconformity to the P2 fault and from non-mineralized areas along the P2 fault to ore zone. This indicates the change from a relatively oxidized to a relatively reduced environment. This modification of fluid chemistry may have taken place due to interaction of fluids with graphite-rich basement rocks, as has been suggested by Gaboreau et al (2007). Alternatively, it was caused by reduced fluids originated from the graphite-rich basement rocks and supplied through faults. In either case, this redox change closer to the uranium ore is consistent with the reduction of U⁶⁺ to U⁴⁺ necessary for the precipitation of uraninite. Therefore, S/P ratios of APS minerals record the reduction of fluids associated with the uranium mineralization.

The spatial association of florencitic APS minerals with the McArthur River uranium deposit (Fig. 2b) and their abundance within the deposit suggest the close relationship between
florencite and the uranium mineralization. This observation is consistent with that of Gaboreau (2005, 2007). The APS minerals in metapelites and pegmatites along the entire 7 km of the P2 fault studied are predominantly florencite with minor components of svanbergite-svanbergite-crandallite-goyazite (Figs. 4d-e, 9). The alteration assemblage of sudoite, magnesiofoitite and illite is also similar all along the P2 fault including the ore zone. The evidence suggests that similar fluids passed along the entire P2 fault. These fluids were likely uranium-bearing because APS minerals contain uranium (up to 16 ppm; Table 3).

The Red-Green Zone sample of metapelite (sample MAC412; Fig. 3d), 100 m laterally away from the P2 fault contains As- and LREE-rich APS minerals (Figs. 7a,d, 8b; Table 2). Arsenic in APS minerals occur as As$^{5+}$, replacing P$^{5+}$ in the X-site of the crystal structure. These As-bearing crystals show prominent zoning with cores of arsenoflorencite-(Ce) and rims of crandallite-goyazite and svanbergite solid-solution (Figs. 5b, 7a,c; Table 2) with sharp boundaries. The zoning suggests that arsenoflorencite-(Ce) formed earlier than the svanbergite-goyazite of diagenetic origin. This arsenoflorencite-(Ce) may have formed in the very early stage of diagenesis. If this is the case, As-rich APS minerals should be observed in sandstones. However, there is no record of As-rich APS minerals in the Athabasca sandstone. Furthermore, sharp boundaries between cores and with rims suggests that As-rich APS minerals formed before the sandstone diagenesis. We suggest that, arsenoflorencite(-Ce) cores may have formed during the paleo-weathering before the deposition of the Athabasca sandstones. This proposed interpretation is further supported by the presence of As$^{5+}$ in the APS minerals. Although As contents are high in organic-rich shale (e.g., Ketris and Yudovich 2009), As is mostly fixed in sulphide minerals (e.g., Ryan et al., 2013). For As to be incorporated in the APS minerals, it
must be oxidized to $\text{As}^{5+}$, which is stable in oxidizing, near-surface conditions (Smedley and Kinniburgh 2002).

**APS minerals related to uranium deposits and role of P2 fault**

It has been suggested that APS minerals with florencitic compositions are contemporaneous with uranium mineralization based on their occurrences within uranium deposits (Quirt 1991; Gaboreau et al. 2005, 2007; Mercadier et al. 2011). Here we have shown that APS minerals of florencitic composition are found along a 7 km strike length of the P2 fault, even in areas barren of mineralization (Table 2; Fig. 9); although, as previously discussed, APS minerals with the highest florencite component (reflecting relatively reducing conditions) are found in proximity to ore (Figs. 9, 11). The APS minerals also contain elevated uranium contents (Table 3; Fig. 6), suggesting that florencitic APS minerals crystallized from uranium-bearing fluids that became reduced. The evidence supports that florencite was contemporaneous with uranium mineralization.

Florencitic APS minerals are not found in the basement greater than ~50 m from the centre of the P2 fault (Table 1; Fig. 9). The distribution of the APS minerals constrains the extent of the fluid footprint, and suggests that the entire P2 fault was a conduit for an oxidizing uranium-bearing fluid during mineralization. The P2 fault was also a site of redox change where an initial oxidizing fluid became modified and relatively reducing.

**Implication**

APS minerals are very small in size, but ubiquitous in the altered metasedimentary rocks along the P2 fault, which hosts the McArthur River uranium deposit. Previous researchers noted the spatial association of florencitic APS minerals and uranium deposits and suggested their use
as an exploration vector (e.g., Gaboreau et al. 2007). Our study reveals that APS minerals record the timing and characteristics of geological events, from paleoweathering, diagenesis to uraniferous hydrothermal activity. We conclude that oxidizing diagenetic fluids travelled along P2 fault and formed svanbergite-crandallite-goyazite. The fluid evolved to become reducing through interaction with the basement rocks and formed florencite rims. Our study shows that APS minerals are an excellent mineralogical tool to fingerprint fluid pathways. APS minerals with similar chemistry (high florencite component) are found along a 7 km of strike length of the P2 fault, including areas barren of uranium mineralization. This suggests that uranium-bearing fluids likely passed along the entire P2 fault, but only produced uranium mineralization in localized areas. This study confirms that APS minerals are good indicators in identifying alteration related to uranium mineralization, thus they may be used, especially when having a florencitic composition, to identify fertile structures associated with uranium deposits.

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Morishita, T., Ishida, Y., Arai, S. (2005) Simultaneous determination of multiple trace element compositions in thin (<30 μm) layers of BCR-2G by 193 nm ArF excimer laser ablation-


**Figure Captions**

**Figure 1.** A map showing the location of the major uranium deposits (red squares), including McArthur River deposit, and basement structures (dashed lines), including the P2 fault, in the Athabasca Basin (modified from Jefferson et al., 2007). The south-eastern margin of the Athabasca Basin is underlain by the basement rocks of the Wollaston and Mudjatik Domains of the western Churchill Province. An insert shows the locations of the Athabasca basin relative to other two Mesoproterozoic sedimentary basins: Hornby and Thelon basins. Major shear zones: BB = Black Bay, BLSZ= Black Lake Shear Zone, CB = Cable Bay, GR = Grease River Shear Zone, RO = Robbilard, VRSZ = Virgin River Shear Zone.

**Figure 2.** Schematic cross-sections (a, b) showing the distribution of APS minerals, some sample locations (stars) and a geologic map (c) at the unconformity showing sample locations. (a) APS minerals are found in the P2 fault proximal to low-grade mineralization (sample MAC29) and in the Bleached Zone (BZ) in quartzite (sample MAC17); (b) APS minerals are found in the P2
fault proximal to the Zone 2 ore body (samples MAC63, 67, 85, 152, 425). APS minerals were not found in samples from the footwall and hanging-wall (indicated `No APS`) nor in samples collected from the VQ fault. (c) An interpreted geological strip map at the unconformity showing the P2 fault, the locations of sampled drill-holes and ore bodies, sample locations, and the sites of the cross sections AB. The P2 fault is essentially barren in the southwestern part. Low-grade mineralization occurs along the middle and southeast (P2 Main deposit) portions of the P2 fault. APS minerals were found in one sample (sample MAC412) of the Red-Green Zone in the footwall proximal to the P2 Main mineralization. Z1 - 4 = Zone numbers of ore body at the McArthur River deposit; uc = unconformity.

**Figure 3.** Photographs of selected drill-core samples; (a) uranium ore of the Zone 2 ore body and P2 fault (MAC425); (b) intensely sudoite-altered pegmatite from the P2 fault (MAC67); (c) quartzite from the Beached Zone along the unconformity (MAC17); and (d) hematitized and chloritized metapelite from the Red-Green Zone (MAC412). Sample locations are shows in Fig. 2.

**Figure 4.** Photomicrographs (a, b, f) and SEM BSE (c-e, g-l) images of APS minerals. (a) Disseminated APS grains in fine-grained illite (Ilt) and with hematite (Hem); (b) Clustered APS grains in contact with an aggregate of feather-textured, magnesiofoitite (Mgf) and a mixture of sudoite+illite+magnesiofoitite (Sud/Ilt/Mgf); (c) Zoned APS grains in magnesiofoitite; (d) Oscillatory zoned APS crystals disseminated in a mixture of clinochlore+illite+magnesiofoitite (Clc-Ilt-Mgf) with chalcopyrite (Ccp); (e) Zoned APS crystals clustered in magnesiofoitite+sudoite; (f) Zoned APS crystals, apatite, and magnesiofoitite in a mixture of sudoite+illite; (g) Partially corroded zircon (Zrn) mantled by xenotime (Xtm) and APS grains; (h) Aggregate of euhedral, zoned APS crystals in a clinochlore+sudoite matrix; (i) Disseminated,
sub-hedral APS grains in fine-grained kaolin (Kln); 

(j) A mixture of APS minerals and quartz rimming coarse quartz grain; 

(k) APS grain in contact with pyrite (Py); 

(l) APS crystal with inclusions of Fe-oxide in contact with apatite (Ap). Sample number and location are indicated at the bottom left corner of each image. Qz = Quartz

**Figure 5.** X-ray maps of APS minerals showing zoning with respect to Ca, Ce, Sr, S and P obtained by wave-length dispersive spectra of EPMA: 

(a) sample MAC425 from the P2 fault and Zone 2 ore body and 

(b) sample MAC412 from the Red-Green Zone (locations shown in Fig. 2a and 2b; Table 2 for quantitative chemical data). svan-goy-crand = svanbergite-goyazite-crandallite. Note that very dark areas in BSE images contain low mass elements and they are not holes or other mineral inclusions.

**Figure 6.** Chondrite-normalized REE, U, Th, Pb and W for APS grains from sample MAC44 (along the P2 fault proximal to the Zone 3 ore body of the McArthur River deposit) and from samples MAC63 and MAC85 (close to Zone 2 ore). The abundance of these elements was obtained using EPMA (La, Ce, Pr and Nd) and LA-ICP-MS (Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu, W, Pb, Th and U). Error bars are not shown in the diagram because the uncertainty of the analysis (< 10 %) is smaller than the size of symbols. Sample locations are in Figure 2c and Table 1. Each line represents one APS grain. Chondrite values from McDonough and Sun (1995).

**Figure 7.** Binary plots showing substitution and elemental relationships of APS minerals based on EPMA data: 

(a) As vs P (black) and As vs. Ce (gray); 

(b) S vs. Sr (black) and S vs. Ce (gray); 

(c) P vs Sr (black) and P vs. LREE (gray); 

(d) S vs. As (black) and S vs. P (gray); and 

(e) LREE vs. Sr (black) and LREE vs. Ca (gray). Each data point represents one grain. P2 fault = P2 fault samples (triangles). BZ = Bleached Zone sample (stars), RGZ = Red-Green Zone sample (crosses).
Figure 8. Compositions of APS minerals plotted on ternary diagrams. End member minerals are indicated at the apexes of the diagrams. (a) Ca-Sr-LREE; (b) P-S-As; (c) S-Sr-LREE; (d) S-Sr-Ca. Each data point represents one grain. P2 fault = P2 fault samples (triangles). BZ = Bleached Zone sample (stars), RGZ = Red-Green Zone sample (crosses).

Figure 9. Cross-plot showing coupled substitution of Ca-Sr-S for LREE-P for APS minerals along the P2 fault (triangles), and within the Bleached Zone (BZ) sample (stars) and Red-Green Zone (RGZ) (crosses). P2 fault samples are color coded with respect to distance in metre from the uranium mineralization (McArthur River deposit and low-grade): red filled triangles represent sample < 1 m from ore (MAC425 and MAC29), orange filled triangles were < 15 m from ore (MAC152), purple filled triangles were < 20 m from ore (MAC67 and MAC85), open green triangles were < 50 m from ore (MAC43, MAC44, MAC63 and MAC404) and open blue triangles were of unmineralized areas >500 m from the ore (MAC117 and MAC135). Each data point represents one grain. For comparison, the diagram shows the compositional ranges of APS minerals determined by Gaboreau et al (2005, 2007) from the Athabasca Basin (fields with various shades of gray) and Kombolgie Basin, Australia (field with various shades of blue).

Figure 10. An alunite-goyazite-crandallite [S-Sr-Ca] ternary plot showing compositions of APS minerals from the P2 fault (P2; triangles), the Bleached Zone (BZ; stars) sample, and the Red-Green Zone (RGZ; crosses) sample. Each data point represents one grain. For comparison, the compositional ranges of diagenetic APS minerals from various sedimentary units are shown: diagenetic APS minerals of the Athabasca and Kombolgie basins are represented by grey dotted field (Gaboreau et al. 2005, 2007) and diagenetic APS minerals of the Chaswood Formation (Piper and Dolansky, 2005), Hornby Basin and Thelon Basins (Gall and Donaldson, 2006) and Mitterberg Formation (Spötl, 1990) are collectively represented by the grey field. APS minerals
associated with uranium deposits of the Athabasca and Kombolgie Basin (Gaboreau et al. 2005, 2007) are represented by the black dashed field.

**Figure 11:** Spatial variation of S/P ratios for APS minerals. The average S/P ratio (mixed data, see Table 2) of APS minerals within a sample is expressed as a red, bold, italicized number next to its corresponding sample ((a) and (b)) or drill-core (c) location.

**Appendix**

**Figure 1:** (a) A graph of the LA-ICP-MS analysis of an APS grain from sample MAC85 showing element counts (major and trace) vs. time in seconds. High counts of Y, Th and U correlate with high counts of the major elements in APS minerals (eg., Ce, La, Al, and Sr). This graph clearly shows that trace elements (such as Y, Th and U) are incorporated in the APS crystal structure and were not present as inclusions or contributed by surrounding minerals of sudoite (Sud) or illite (Ilt). (b) The same graph as shown in (a) but counts/sec of each element is normalized to that of Sr.

**Figure 2:** An APS mineral (a) before EMPA analysis and (b) after analysis showing beam damage. (c) A laser pit after LA ICP-MS analysis. (a) and (c) are BSE images and (b) is a secondary electron image.

**Figure 3:** A binary plot comparing the average concentrations of Ca (green), Fe (red), Pr (cyan) and Nd (blue) in APS mineral grains from samples MAC44 (diamonds), MAC63 (x), and MAC85 (crosses) obtained with EMPA (Table 2) and LA ICP-MS (Table 3). Horizontal and vertical bars represent one standard deviation of the data (Tables 2 and 3). LA ICP-MS data of sample MAC85 is of one grain.
Table 1: Sample descriptions.

<table>
<thead>
<tr>
<th>zone</th>
<th>DDH</th>
<th>sample</th>
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<th>proximity to P2 fault</th>
<th>proximity to uc</th>
<th>proximity to U-min</th>
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<tr>
<td>P2</td>
<td>MC370</td>
<td>MAC29</td>
<td>560.2</td>
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<td>u/c (Fig. 2A)</td>
<td>&lt;1 m from lg U-min in ss</td>
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<td></td>
<td>H729</td>
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<td>&lt;50 m from Z3 ore</td>
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<td></td>
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<td>MAC44</td>
<td>64.5</td>
<td>~15 m laterally from P2 in fw</td>
<td>~15 m</td>
<td>&lt;50 m from Z3 ore</td>
</tr>
<tr>
<td></td>
<td>H201</td>
<td>MAC63</td>
<td>7.6</td>
<td>~30 m laterally from P2 in hw</td>
<td>~20 m</td>
<td>&lt;50 m from Z2 ore</td>
</tr>
<tr>
<td></td>
<td>H201</td>
<td>MAC67</td>
<td>30.8</td>
<td>the P2 fault zone</td>
<td>~20 m</td>
<td>&lt;20 m from Z2 ore</td>
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<td>MAC85</td>
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<td></td>
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<td>MAC404</td>
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<td>~15 m</td>
<td>~50m laterally from lg in ss</td>
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<td>MC361</td>
<td>MAC117</td>
<td>523.3</td>
<td>in the P2 fault zone</td>
<td>~25 m</td>
<td>&gt;500 m from U-min</td>
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<tr>
<td></td>
<td>MC362</td>
<td>MAC135</td>
<td>520.2</td>
<td>~40 m above P2 in hw</td>
<td>~10 m</td>
<td>&gt;500 m from U-min</td>
</tr>
<tr>
<td>GRZ</td>
<td>MC346</td>
<td>MAC412</td>
<td>669</td>
<td>~100 m laterally from P2</td>
<td>~50 m</td>
<td>~100 m below P2 Main</td>
</tr>
<tr>
<td>BZ</td>
<td>MC344</td>
<td>MAC17</td>
<td>499.4</td>
<td>~250 m laterally from P2 in hw</td>
<td>&lt;1 m</td>
<td>&gt;250 m from U-min</td>
</tr>
</tbody>
</table>

Notes:
- P2 = P2 fault sample; RGZ = Red-Green Zone sample; BZ = Bleached Zone sample
- DDH = diamond drill-hole number
- Drillhole with the prefix H are collared underground, therefore sample depths are not from the surface
- U-min = uranium mineralization in the McArthur River deposit; fw = footwall; hw = hangingwall; lg = low-grade (<1 wt% U3O8); ss
- Ap = apatite; Clc = clinochlore; Dck = dickite; Drv = dravite; Fe-Mg chl = Fe-Mg Chlorite; fg = fine-grained; Gr = graphite; Itt = illi
<table>
<thead>
<tr>
<th>degree of alteration and rock-type</th>
<th>APS occurrence</th>
<th>clay</th>
<th>mineralogy</th>
</tr>
</thead>
<tbody>
<tr>
<td>moderately altered mp</td>
<td>isolated, 3-10 µm</td>
<td>Sud</td>
<td>Mg, Xtm, Qz</td>
</tr>
<tr>
<td>green intensely altered graphitic mp</td>
<td>isolated, 3-10 µm</td>
<td>Sud/Ilt</td>
<td>Mg, Gr, Py, Ap, Zrn, Xtm</td>
</tr>
<tr>
<td>red and white intensely altered peg</td>
<td>zoned, clustered, 5-20 µm</td>
<td>Sud/Ilt</td>
<td>Hem, Mnz, Zrn, Qz, Gr, Py</td>
</tr>
<tr>
<td>green intensely altered mp</td>
<td>zoned, clustered, fragmented, 1-20 µm</td>
<td>Sud</td>
<td>Ilt, Rt, Ap, Fe-Mg Chl, Ilt, Py, Qz</td>
</tr>
<tr>
<td>green intensely altered peg</td>
<td>zoned, clustered, 5-20 µm</td>
<td>Sud</td>
<td>Mg, Ilt, Py, Ap, Qz</td>
</tr>
<tr>
<td>intensely altered graphitic mp w/ peg boudins</td>
<td>zoned, disseminated, 5-20 µm</td>
<td>Sud/Ilt</td>
<td>Drv, Py</td>
</tr>
<tr>
<td>green-white intensely altered mp</td>
<td>zoned, disseminated, 5-30 µm</td>
<td>Sud/Clc</td>
<td>Mg, Ill, Py, Qz, Ccp</td>
</tr>
<tr>
<td>red moderately altered mp w/ peg boudins</td>
<td>disseminated, &lt; 1 µm</td>
<td>Kln</td>
<td>Hem, Ap</td>
</tr>
<tr>
<td>green intensely altered mp with U-ore</td>
<td>some cores missing, isolated, 3-10 µm</td>
<td>Mgf/Sud</td>
<td>U, Hem</td>
</tr>
<tr>
<td>white-green intensely altered peg</td>
<td>isolated, 3-10 µm</td>
<td>Ilt</td>
<td>Fe-Mg Chl, fg Qz, Qtz</td>
</tr>
<tr>
<td>red-green intensely altered peg</td>
<td>disseminated, &lt; 1 µm</td>
<td>Ilt</td>
<td>Hem, Rt, fg Qz</td>
</tr>
<tr>
<td>red-green intensely altered mp</td>
<td>large clusters up to .5 mm, 10-40 µm</td>
<td>Sud/Clc</td>
<td>Hem</td>
</tr>
<tr>
<td>quartzite with bleached matrix</td>
<td>disseminated, &lt; 1 µm</td>
<td>Kln</td>
<td>Ilt, Zrn, Qz</td>
</tr>
</tbody>
</table>

= sandstone; uc = unconformity; peg = pegmatite; mp = metapelite
ite; Mgf = magnesiofotite, Py = pyrite; Qz = quartz; Rt = rutile sud = sudoite, U = uraninite; xtm = xenotime; zrn = zircon
Table 2: EMPA major element oxide data and atomic proportion formula units for APS minerals.

<table>
<thead>
<tr>
<th>Sample</th>
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Notes:

Due to small grain size of APS minerals and the relatively large electron beam size (5 µm) for the analysis, it was difficult to obtain precise composition of individual compositional zones in grains, as some rims are narrower than 5 µm. Most cores are also smaller than 5 µm (Figs. 4,5). Therefore, most of the compositional data presented in Table 2 reflect mixtures of different zones. Rim and core analysis was conducted on two samples containing relatively large APS (MAC152 and MAC412).

a: adjacent minerals, Ilt=illite, Sud=sudoite, Mgf = magnesiofoitite; Clc = clinochlore
b: Area of analysis in grains, ie. rim, core or mixed. Compositional zones are narrower than the beam size of 5 µm. The mixed area refers to the data from different compositional zones defined in SEM-BSE images
c: contribution of surrounding minerals in the analytical data was evaluating using SiO2 contents (see the Results for detailed description)
d: \([Ce]_N = \sqrt{([La]_N-[Pr]_N)}\)
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Table 3: LA-ICPMS data of APS (ppm).

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<td>&lt;0.161</td>
<td>0.6 0.007</td>
<td>1.56</td>
<td></td>
</tr>
<tr>
<td>W</td>
<td>0.79 0.36</td>
<td>0.9 0.029</td>
<td>1</td>
<td>0.5 0.03</td>
<td>0.61</td>
<td></td>
</tr>
<tr>
<td>Pb</td>
<td>160 10</td>
<td>120 39</td>
<td>260</td>
<td>10.3 0.005</td>
<td>0.92</td>
<td>12</td>
</tr>
<tr>
<td>Th</td>
<td>1300 200</td>
<td>1100 140</td>
<td>2300</td>
<td>5.4 0.035</td>
<td>0.91</td>
<td>13</td>
</tr>
<tr>
<td>U</td>
<td>0.62 0.19</td>
<td>2.4 1.1</td>
<td>16</td>
<td>1.7 0.065</td>
<td>1.46</td>
<td>1.9</td>
</tr>
</tbody>
</table>

*BCR2G from US Geological Survey*

*based on the data of BCR2G provided by USGS*

data is the average of analyzed grains (# grains) in the sample (± 1 standard deviation)
Figure 2

a

- No APS footwall
- P2 fault
- No APS

b

- Sample (MAC##)
- P2 fault
- MAC412
- MAC135
- MAC29
- MAC404
- Z4, Z3
- ZAZ1, ZA, ZB
- Cross-section A
- Cross-section B
- McArthur River deposit

Legend:
- Quartzite
- Metapelite
- Graphitic
- Pegmatite
- Mineralization
- UC
- P2 Fault
- Fault
- Drill-hole collar

Symbols:
- Hanging wall
- Fault

Grid:
- 0 50 meters
- 100 meters

Northwest to Southeast

Figure 2 represents geological cross-sections A and B of the McArthur River deposit, highlighting the distribution of mineralization, faulting, and other geological features.
Figure 4
Figure 5

a) MAC425 florencite rim; svanbergite-crandallite core

b) MAC412 svan-goy-crand rim; arsenoflorencite-(Ce) core
Figure 6
Figure 7
Figure 9

![Image of a graph with data points and labels indicating different categories of U-deposits and their corresponding distances from the P2 fault and RGZ. The graph shows variations in LREE + P apfu, Ca + Sr + S apfu, and data from Kombolgie Basin and Athabasca Basin.](image)

- **fault-hosted**
- **florencite**
- **svanbergite**
- Data from Athabasca Basin: U-proximal bsmt, U-proximal ss, intermediate, distal bsmt, distal ss
- Data from Kombolgie Basin: proximal to U-deposits, intermediate, distal
Svanbergite
SrAl(PO₃)(SO₄)(OH)₃·4H₂O

Goyazite
SrAl(PO₃)(OH)·(H₂O)₃·2H₂O

Crandallite
CaAl(PO₃)(SO₄)(OH)₃·4H₂O

Woodhouseite
CaAl(PO₃)(OH)·(H₂O)₃·2H₂O

Alunite
KAl(SO₄)(OH)₃·2H₂O

This study

Diagenetic of the Athabasca & Kombolgie basins

Diagenetic in other basins

Associated with U deposits

Figure 10
Figure 11

a.

b.

data:

sample (MAC##)

S/P ratio

P2 fault

C.

0.16-21

0.03

0.05-0.16

0.17

0.12

0.36

0.11

0.13

0.16

quartzite

metapelite

graphitic

pegmatite

mineralization

uc

P2 fault

fault

drill-hole collar

0 50 100 meters

0 500 1000 1500 2000 m

NW

SE

Figure 11

0.06

0.12

0.05

0.06

0.12

0.03

0.17

0.11

0.06

0.16

0.36

0.05-0.16

0.03

0.17

0.11

0.36