

1 **Revision 1**

2

3

Bubble formation during decompression of andesitic melts

4 **Adrian Fiege**¹, François Holtz¹, and Sarah B. Cichy²

5

6 ¹Institut für Mineralogie, Leibniz Universität Hannover, Callinstraße 3, 30167 Hannover, Germany

7 ²Laboratoire Magmas et Volcans, UMR 65, Université Blaise Pascal, 5 rue Kessler, 63038

8 Clermont-Ferrand Cedex, France

9

10 *E-mail:* afiege@umich.edu, f.holtz@mineralogie.uni-hannover.de, [11 \[bpclermont.fr\]\(mailto:bpclermont.fr\)](mailto:s.cichy@opgc.univ-</p></div><div data-bbox=)

12 *Fax:* +49 511 762 3045

13 *Phone:* +49 511 762 5281

14

15

ABSTRACT

16 Bubble formation during continuous decompression from ~400 to ~70 MPa was

17 investigated experimentally in hydrous andesitic melts at $T = 1030^\circ\text{C}$ and at an oxygen fugacity

18 (fO_2) of about $\log(fO_2/\text{bar}) = \text{QFM}+1$ (QFM: quartz-fayalite-magnetite buffer). Experiments were

19 carried out at variable decompression rates (r), ranging from 0.0005 to 0.1 MPa/s. The samples

20 were directly quenched after decompression, allowing the investigation of the influence of r on

21 the bubble formation. The effect of variable annealing times (t_A) after decompression was also

22 investigated for experiments performed at a decompression rate of 0.1 MPa/s. These samples

23 were annealed for $t_A = 0$ to 72 h at final pressure (70 MPa) in order to study changes in

24 vesiculation during magma storage at shallow depths after fast ascent. Back scattered electron
25 (BSE) images of the samples were analyzed to determine bubble number densities (*BND*).

26 The *BND* values increase strongly with increasing r and vary from about $10^{2.2} \text{ mm}^{-3}$ at
27 0.0005 MPa/s to about $10^{4.6} \text{ mm}^{-3}$ at 0.1 MPa/s . After fast decompression ($r \sim 0.1 \text{ MPa/s}$), the
28 *BND* decrease significantly with t_A , i.e., from $\sim 10^{4.6} \text{ mm}^{-3}$ at $t_A = 0 \text{ h}$ to $\sim 10^{2.9} \text{ mm}^{-3}$ at $t_A = 72 \text{ h}$.
29 A comparison of the derived *BND* values with recently published experimental data demonstrates
30 the essential role of the decompression path on bubble formation. The *BND* are higher in
31 experiments with multi- or single-step decompression when compared to continuous
32 decompression. The new data show that H_2O -undersaturated andesitic melts are characterized by
33 1 to 2 log units higher *BND* values than H_2O -saturated rhyolitic melts after decompression with
34 the same rate, indicating a strong influence of melt composition on bubble nucleation. This
35 compositional effect is not predicted accurately by existing models and the interpretation of the
36 vesicularity of dacitic to andesitic melts may lead to overestimations of magma ascent rates by
37 about an order of magnitude.

38

39 **Keywords:** Bubble formation, bubble number density, andesite, continuous decompression

40

41

INTRODUCTION

42 Decompression induced degassing of a magma is a key factor controlling the style of a
43 volcanic eruption (e.g., review of Sparks et al., 1994). Thus, a good knowledge of bubble forming
44 processes in magmatic systems is a pre-requisite to improve our understanding of hazardous,
45 explosive eruptions and pyroclastic flows (see review of Sparks et al. 1994). However,
46 experimental data at geologically relevant conditions in magmatic systems are rare, especially at
47 low, constant decompression rates (r). Continuous decompression experiments, aiming to

2

48 understand the formation of bubbles in silicate melts, are typically performed at $r > 0.02$ MPa/s
49 (e.g., Mourtada-Bonnefoi and Laporte 2004; Cichy et al. 2011; Gondé et al. 2011), while in
50 natural systems r is often < 0.02 MPa/s, depending on the eruptive style (e.g., Cashman 2004).
51 Additionally, the experimental data available in literature to model bubble formation and growth
52 are limited to rhyolitic melt compositions (e.g., Mourtada-Bonnefoi and Laporte 1999; 2002;
53 2004; Mangan and Sisson 2000; Mangan et al. 2004a; Hamada et al. 2010; Cichy et al. 2011;
54 Gondé et al. 2011), while no information on the behavior of more mafic or depolymerized melts
55 exists.

56 In this study, continuous decompression experiments were carried out to explore the
57 bubble nucleation and growth in an andesitic melt and the results are compared to rhyolitic melts.
58 The experimental approach chosen in this study simulates the formation of bubbles in an initially
59 bubble- and crystal-free melt, thus enabling the interpretation of homogeneous bubble nucleation.
60 The compiled experiments cover a wide range of r (0.0005 to 0.1 MPa/s). We demonstrate that
61 the decompression path (continuous decompression vs. multi-step decompression) and the initial
62 H₂O contents (H₂O-saturated vs. H₂O-undersaturated) are crucial parameters for the
63 interpretation of bubble number densities.

64

65 **EXPERIMENTAL PROCEDURE AND ANALYTICAL METHOD**

66 **Experimental procedure**

67 In order to investigate the effect of r and t_A on bubble formation and on bubble number
68 densities (BND ; i.e., the number of bubbles per unit volume) in andesitic melts, the distribution of
69 bubbles in selected experimental products from a previous study by on sulfur distribution
70 between melts and fluids (Fiege et al. 2014) has been investigated.

71 The experimental strategy applied by Fiege et al. (2014) was as follows: In the *first step*, a
72 synthetic anhydrous glass with a composition corresponding to the Krakatau andesite
73 (Mandeville et al. 1996) was prepared by melting a mixture of oxide and carbonate powders for 2
74 h in a Pt₉₀Rh₁₀ crucible at 1600°C and 1 atm. The melt was quenched to a glass in a water bath,
75 ground and melted again to improve the homogeneity. The composition of the anhydrous
76 andesitic glass is listed in Table 1. In the *second step*, hydrous sulfur-bearing glasses were
77 synthesized at high pressure (~500 MPa; uncertainty of the pressure transducer: ±5 MPa) and
78 high temperature (~1030°C) under fluid absent conditions in internally heated pressure vessels
79 (IHPV). The syntheses were carried out in Au capsules. Glass powder and variable amounts of
80 gypsum powder (Ca[SO₄]×2H₂O) were mixed thoroughly in an agate mortar. The powder
81 mixture and required amounts of deionized H₂O (H₂O source) were filled stepwise into the Au
82 capsule and were compacted using a piston and a hammer to minimize entrapped air. In some
83 experiments, a 10 wt% HCl aqueous solution was also added to prepare Cl-bearing glasses. The
84 *third step* included the isothermal decompression experiments, conducted in IHPV at a
85 temperature (*T*) of 1030°C and QFM+0.8 to QFM+1.8.. Ground powder of the synthesized H₂O-
86 S±Cl-bearing starting glass was loaded into gold capsule, compressed to minimized entrapped
87 air, crimped flat and welded shut. This technique produces almost cylindrical melt pools, which
88 minimizes wall effects and allows the expansion of the loaded material during decompression.
89 Prior to decompression, the samples were annealed for ~5 min at a constant pressure (*P*) of ~400
90 MPa. Subsequently, pressure was released continuously from 400 MPa to 70 MPa at different
91 rates (*r*), varying from 0.0005 to 0.1 MPa/s. After the decompression, samples were either rapidly
92 quenched (following the method described by Berndt et al. 2002) to preserve strongly non-
93 equilibrium conditions or were annealed for various times (*t_A*) at final *P* and *T* before quenching
94 in order to approach near-equilibrium conditions for the distribution of volatiles (H₂O, S and Cl)

95 between melt and fluid. As reported in previous studies on andesitic melts (e.g., King and
96 Holloway 2002; Botcharnikov et al. 2006), the estimated cooling rate of about 150°C/s obtained
97 by applying the rapid quench technique (Berndt et al. 2002) is sufficient for our research
98 purposes.

99 The experiments selected for the investigation of *BND* are characterized by starting
100 glasses with similar initial water (H₂O) contents of about 6.5 wt%, while their initial sulfur
101 contents are ranging from ~140 to ~1050 ppm and chlorine is varying from 0 to ~1000 ppm. The
102 initial concentrations of H₂O, S and Cl in the starting glasses are given in Table 2 and details on
103 the experimental series GYCIA, GYM CIA, QFMA and QFM CIA as well as on the methods used
104 for the characterization of the volatile concentrations in the starting glasses [electron microprobe
105 (EMP), IR spectroscopy, Karl Fischer titration] are given in Fiege et al. (2014). Solubility
106 experiments of Botcharnikov et al. (2006) indicate that the H₂O solubility in andesitic melts at
107 our starting conditions of ~400 MPa and 1030°C is higher than 6.6 wt%. This is also confirmed
108 by the model of Burnham (1979), which predicts a H₂O solubility of 7.2 wt% for the investigated
109 andesitic melt composition at 400 MPa and 1030°C. The effect of small amounts of S and Cl in
110 the melt on H₂O solubility is presumably negligible (Webster et al. 2009). Moreover, BSE images
111 of the starting glasses as well as of test runs which were immediately quenched after short
112 annealing (~5 to ~10 min) at 1030°C and 400 MPa without decompression show no bubbles.
113 Hence, we conclude that no fluid phase was present prior to decompression as well as within the
114 first tenths of MPa of *P* release. The experimental series QFMA and QFM CIA (0.0005 to 0.1
115 MPa/s; *t_A* = 0 h; Table 2) were chosen to explore the influence of *r* on *BND*, while series GYCIA
116 and GYM CIA (0.1 MPa/s; *t_A* = 0 to ~72 h) were selected to evaluate possible changes in
117 vesiculation processes during annealing after a fast decompression (0.1 MPa/s). The *fO₂* of these

118 experiments was adjusted to \sim QFM+0.8 (QFMA and QFMCIA) or \sim QFM+1.8 (GYCIA and
119 GYMCI A).

120

121 **Estimation of bubble number densities (*BND*)**

122 Three BSE images (collected using EMP) of the experimental products were analyzed to
123 characterize bubble size and distribution, using the public computer program ImageJ
124 (<http://rsb.info.nih.gov/ij/>). The geometric information derived from ImageJ analyses was used to
125 estimate the *BND* of the samples, following the method of Noguchi et al. (2008) with a procedure
126 already applied by Cichy et al. (2011) and Nowak et al. (2011).

127 The ImageJ software allows us to determine, e.g., total area, average size, area fraction as
128 well as width and length of individual bubbles. Moreover, the (average) minor and major axes as
129 well as the angle of the ellipsoid fitted to the bubbles are automatically calculated and used for
130 subsequent *BND* estimations. As ImageJ analyses provide only 2D bubble size information, the
131 estimation of the *BND* involves as a final step a 3D correction using the CSD-Corrections 1.4
132 software (<http://depcom.uqac.ca/~mhiggins/csdcorrections.html>). The methods used in CSD-
133 Corrections 1.4 are described by Higgins (2000, 2002) and Higgins and Chandrasekharam
134 (2007). For the 3D correction, the roundness factor calculated by ImageJ, ranging from 0.6 to 0.9,
135 was applied while the aspect ratio was set to 1.0 - 1.1 - 1.2 (short axis - intermediate axis - long
136 axis).

137 Noteworthy, the BSE images were taken at different magnifications and from different
138 locations of the samples, which minimizes possible truncation effects (for details see section
139 “Accuracy of the *BND* estimation“ and, e.g., Armienti 2008). The image sizes are ranging
140 between \sim 300 \times \sim 220 μ m and \sim 2500 \times \sim 1850 μ m (resolution: \sim 0.5 to \sim 2.5 μ m/pixel,
141 respectively). We have chosen to process BSE images rather than reflected-light microscopic

142 images for our *BND* analyses because BSE images show a higher contrast between bubbles and
143 melt.

144

145 Table 1

146

147

RESULTS

148 Figure 1 shows a typical microscopic image of a partially degassed andesitic glass
149 (GYCIA-3). The size of the bubbles ranges from a few microns up to $\sim 80 \mu\text{m}$; a typical feature of
150 fast decompressed ($\sim 0.1 \text{ MPa/s}$) samples with 0 to $\sim 2 \text{ h}$ of annealing time at final *P-T* conditions.
151 The analyses of BSE images indicate that the amount of bubbles decreases while the size of
152 bubbles increases with decreasing *r* from 0.1 to 0.0005 MPa/s (e.g., QFMCI series; Fig. 2). A
153 comparable trend has been observed with increasing t_A after fast decompression ($\sim 0.1 \text{ MPa/s}$;
154 e.g., GYMCI series; Fig. 3). The bubbles appear often elongated and show a preferred orientation
155 (flow structures) – especially in experiments with low *r* ($< 0.02 \text{ MPa/s}$) and/or long term
156 annealing after decompression ($t_A > 5 \text{ h}$) – indicating that convection may have occurred during
157 the experimental runs. No significant influence of small variations in the S and/or Cl content in
158 the system on the *BND* was detected (see Fig. 4).

159

160 Figure 1, 2, 3 and 4

161

Accuracy of the *BND* estimation

163 As mentioned previously, the method applied for the estimation of the *BND* may result in
164 truncation effects which can affect *BND* values significantly (e.g., Armienti 2008). Such
165 truncation effects either arise from the resolution of the BSE image (left-hand truncation) or from

166 the image size (right-hand truncation). In case of left-hand truncation, the number of bubbles with
167 a size close to the resolution limit is reduced. However, the resolution of the BSE images used in
168 this study was found to be sufficient to detect bubbles with a diameter of $\geq 1 \mu\text{m}$, while
169 microscopic investigations indicate that the amount of bubbles with a diameter $< 1 \mu\text{m}$ is
170 negligible. Thus, a reduction of the number of the smallest bubbles is unlikely in our case,
171 implying that left-hand truncation can be ruled out. Right-hand truncation effects occur when the
172 largest bubbles in the sample are excluded or are underrepresented in the analyzed area of the
173 sample. In bubble size distribution (BSD) plots, this would result in a flattening of the size
174 distribution to nearly horizontal trends at large sizes. However, such trends are not observed in
175 the *BSD* plots, implying that right-hand truncation is unlikely (see Fig. A.1 in the appendix,
176 showing *BSD* plots of selected samples).

177 Another source of error is the treatment of bubbles which are located partly outside the
178 area covered by the image. To account for this problem, *BND* values were estimated using two
179 approaches. In the first approach, bubbles which were cut at the image edge were excluded from
180 the calculation. In the second approach, the area of the bubbles visible in the image was included
181 in the calculation. The *BND* values derived from these two procedures differ by less than 0.12
182 log. The *BND* values obtained from the three different images of one sample, taken at different
183 magnifications and from different locations, differ typically by < 0.8 log units, indicating a quite
184 homogeneous distribution of the bubbles within the samples. The average *BND* values of all six
185 analyses (i.e., 3 images at different magnifications and locations on the sample; bubbles at the
186 edges included or excluded) are listed in Table 2.

187

188 Table 2

189

190 **Influence of r and t_A on BND**

191 The BND data of experiments conducted at different r and with immediate quenching at
192 final P ($t_A = 0$ h) are shown in Figures 4a and 5. A roughly linear increase of $\log(BND \times \text{mm}^3)$
193 with $\log(r \times \text{s/MPa})$ is indicated by the data obtained in this study for hydrous, but initially not
194 volatile-saturated andesitic melts. Figure 5 shows that the BND values determined in previous
195 experimental studies (e.g., Mangan and Sisson 2000; Mourtada-Bonnefoi and Laporte 2004;
196 Hamada et al. 2010) on homogeneous nucleation in CO_2 -free, H_2O -saturated rhyolitic systems at
197 700 to 900°C are typically more than 1 log unit lower than those obtained in this study. BND
198 values determined by Gondé et al. (2011) for homogeneous nucleation in H_2O -undersaturated
199 rhyolitic melts at 730 to 1000°C and for r ranging from 0.06 to 2.3 MPa/s show no clear trend
200 and have to be interpreted with caution due to the H_2O loss during their experiments as already
201 mentioned by the authors. However, the data obtained by Gondé et al. (2011) for a r of 40 to 50
202 MPa/s seem to be in good agreement with other experimental results on H_2O -saturated rhyolitic
203 systems (e.g., Hamada et al. 2010; see Fig. 5). Changes in bulk Cl and S contents as well as slight
204 variations in fO_2 , varying from $\sim\text{QFM}+0.8$ to $\sim\text{QFM}+1.8$, have no significant effect on the
205 decompression-induced bubble formation in the studied andesitic melts.

206 The influence of t_A on BND has been studied at $\text{QFM}+1.8$ (experimental series GYCIA
207 and GYMCI A; BND values listed in Table 2) and at a constant r of ~ 0.1 MPa/s. Figure 4b
208 indicates that $\log(BND \times \text{mm}^3)$ decreases by 1 log unit or more within the first ~ 5 h of annealing
209 after decompression. BND values obtained from one experimental series with $t_A = 5$ h and $t_A = 72$
210 h are identical within error, and it can be concluded that annealing of more than 5 h after
211 decompression has a minor effect on BND .

212

213

214

DISCUSSION

215 **Role of microlites and decompression path on bubble nucleation**

216 In Figure 4a the *BND* data of this study for r ranging from 0.0005 to 0.1 MPa/s and
217 $t_A = 0$ h are compared to the experimental results of Cichy et al. (2011). Cichy et al. (2011)
218 investigated the vesiculation in a rhyodacitic system induced by a decompression from 300 to 50
219 MPa at 850°C (r ranging from 0.0002 to 20 MPa/s). It is emphasized that the differences in T
220 should have a negligible effect on *BND* at given r (see Fig. 5 and Hamada et al. 2010). As
221 mentioned above, a roughly linear increase of $\log(BND \times \text{mm}^3)$ with $\log(r \times \text{s/MPa})$ is indicated
222 by our experimental series QFMA and QFMClA (Fig. 4a; Table 2). In contrast, Cichy et al.
223 (2011) observed a minimum of $\log(BND \times \text{mm}^3)$ at $r \approx 0.02$ MPa/s. Noteworthy, the experiments
224 of Cichy et al. (2011) were conducted to investigate the combination of simultaneous
225 decompression-induced degassing and crystallization. The samples before decompression
226 contained significant amounts of crystalline phases (including Fe-Ti microlites). Thus,
227 heterogeneous bubble nucleation processes were studied by Cichy et al. (2011), while
228 homogeneous nucleation is assumed for our experiments. However, the results of Mangan and
229 Sisson (2000) indicate that *BND* in experiments with heterogeneous nucleation are comparable to
230 the values of homogenous nucleation at given r . Mangan and Sisson (2005) noted that the
231 presence or absence of heterogeneity in the melt dictates the supersaturation pressure ΔP , while
232 the nucleation rate (homogenous or heterogeneous) is mainly influenced by r . The microlite
233 number densities (*MND*) determined by Cichy et al. (2011) do not show any clear dependence on
234 r ($MND \approx 10^{5.5} \text{ mm}^{-3}$ in the range of 0.0002 MPa/s to 20 MPa/s), indicating that changes in
235 nucleation style (homogenous vs. heterogeneous), cannot explain the *BND* minimum at $r \approx 0.02$
236 MPa/s observed by Cichy et al. (2011). Cichy et al. (2011) calculated *MND* values either for
237 plagioclase only or for all other phases together, such as amphibole, pyroxene and oxides.

10

238 However, considering that Fe-Ti-oxides are very important bubble nucleation sites (Hurwitz and
239 Navon 1994), calculating *MND* for Fe-Ti-oxides only may provide useful information to interpret
240 the *BND* data of Cichy et al. (2011).

241 Cichy et al. (2011) applied a multi-step decompression technique to conduct experiments
242 with $r \leq 0.01$ MPa/s, implying that pressure was released stepwise over the whole run duration
243 (typical decompression steps ranged between 8.4 to 40 MPa). Nowak et al. (2011) revealed a
244 significant influence of the decompression style on homogeneous bubble nucleation in the same
245 rhyodacitic melt at T above the liquidus (1050°C). The authors show that the multi-step
246 decompression technique leads to *BND* values which are more than one log unit higher than those
247 obtained for continuous decompression at given r . *BND* values determined by Nowak et al.
248 (2011) for single-step decompression experiment are even higher. Thus, the differences between
249 our *BND* data and the multi-step decompression data of Cichy et al. (2011; see Fig. 4a) are
250 consistent with the observations made by Nowak et al. (2011), and the non-linear trend found by
251 Cichy et al. (2011) can probably be assigned to variations in the decompression style. A
252 comparison of our data with *BND* values determined by Gardner and Denis (2004) and Gardner
253 (2007) for heterogeneous bubble nucleation in CO₂-free, H₂O-saturated rhyolitic melts at 550 to
254 800°C, using single-step decompression technique, confirms this hypothesis. For instance,
255 Gardner and Denis (2004) calculated *BND* values of $10^{4.7}$ to $10^{5.0}$ mm⁻³ at $r \approx 0.007$ MPa/s, while
256 we observed about one log unit lower *BND* values ranging from $10^{3.4}$ to $10^{4.0}$ mm⁻³ for a slightly
257 higher r of about 0.01 MPa/s. However, at $r \approx 0.1$ MPa/s Gardner and Denis (2004) and Gardner
258 (2007) estimated *BND* values of $10^{4.6}$ to $10^{5.2}$ mm⁻³, while we determined *BND* values ranging
259 from $10^{4.2}$ to $10^{4.9}$ mm⁻³ for similar r , indicating that the influence of decompression style on *BND*
260 may decrease with increasing r .

261 The differences in bubble nucleation observed between continuous decompression and
262 multi-step decompression may arise from a high ΔP induced by instantaneous P drops of 50 MPa
263 during multi-step decompression experiment, leading to a predominance of bubble nucleation
264 (Nowak et al. 2011). With respect to continuous decompression, as stated by Nowak et al. (2011),
265 nucleation may require a certain ΔP and, thus, bubble growth will dominate over nucleation
266 processes upon further continuous decompression caused by favorable energetics (e.g., review of
267 Sparks et al. 1994; Nowak et al. 2011). Nevertheless, the increase of $\log(BND \times \text{mm}^3)$ with
268 decreasing multi-step decompression rate at $r < 0.01$ MPa/s observed by Cichy et al. (2011)
269 needs to be clarified by further investigations.

270

271 Figure 5

272

273 **Bubble formation during magma storage after a fast ascent**

274 The observed strong decrease of $\log(BND \times \text{mm}^3)$ with increasing t_A (series GYCIA and
275 GYMCI; see Fig. 4b) is probably linked to the very fast decompression ($r \sim 0.1$ MPa/s). It is
276 assumed that a fast decompression of the andesitic melt leads to a high supersaturation pressure
277 (ΔP). Thus, the gas ($\text{H}_2\text{O-S}\pm\text{Cl}$) pressure in the melt is significantly larger than the ambient P in
278 the vessel (Sparks et al. 1994). Furthermore, as a consequence of the high ΔP , the critical radius
279 size of the bubble embryo is lowered, leading to the domination of bubble nucleation over bubble
280 growth and/or coalescence upon decompression. The resulting significant decrease of BND by
281 about 1 log unit within the first few hours of annealing indicates that coalescence and/or bubble
282 growth become more important than nucleation.

283 Evidence of bubbles coalescence is preserved in our quenched experimental end-products
284 and can be seen in the BSE images of our series GYMCI in Figure 3. Common bubble

285 coalescence structures such as interpenetration shapes, dimples or flattened inter-bubble melt
286 films (Castro et al., 2012) can be observed especially in experiments with short annealing times
287 (Fig. 3a-d with t_A of 0 and 1h). In contrast, at longer t_A ($t_A=5h$; see Fig. 3e and 3f), such bubble
288 shapes were not observed, and only few large bubbles (diameters typically $> 200\mu\text{m}$) are
289 preserved in the quenched glass. This indicates that most previously small-sized vesicles merged
290 together within t_A of 1 and 5 h to form larger individuals, which may even have migrated to the
291 capsule wall.

292 The determination of the volatile concentrations in the melt after decompression is also
293 useful to interpret bubble growth and coalescence upon further annealing. The volatile contents
294 (H_2O , S, Cl) were determined by Fiege et al. (2014) using IR spectroscopy (H_2O content) as well
295 as EMP analyses (S, Cl) and the data are given in their Table 2. Figure 6 shows that the H_2O
296 concentrations of the partially decompressed andesitic melts remain nearly constant during
297 annealing at final P - T conditions. The small variations in H_2O (3.56 to 3.84 wt%; except
298 GYMClA-4: 1.98 wt%) concentrations in the melt are independent of t_A (and r) and within
299 uncertainty of the determination (see Fig. 6b, and 6f). The low H_2O concentration measured by
300 Fiege et al. (2014) in the glass of sample GYMClA-4 can possibly be explained by the loss of
301 small amounts of volatiles from the capsule at the end of the experiments, though no clear
302 evidence for such a leakage were observed. By contrast, minor amounts of S and Cl are still
303 released from the melt to the fluid phase within the first hour of annealing after decompression
304 (e.g., GYClA series: S_{melt} decreases from ~ 180 ppm to 90 ppm upon annealing; see Fig. 6e).
305 Nevertheless, Figure 6 clearly shows that most of the volatiles have already been exsolved upon
306 decompression. Thus, in agreement with observations made on BSE images (see previous
307 paragraph), it is proposed that coalescence dominates over bubble growth upon further annealing
308 at final P - T conditions. However, Lautze et al. (2011) noted that diffusive coarsening, also known

309 as Ostwald ripening, should be the dominant mechanism leading to decreasing BND with
310 increasing t_A . Hence, in addition to volatile diffusion, other processes (e.g., convection fluxes)
311 may have influenced the bubble formation in our decompression experiments.

312

313 Figure 6

314

315 **Comparison of experimental data with the model of Toramaru (2006) and with previous** 316 **results**

317 Toramaru (2006) developed a model to estimate r based on BND data of natural pumices.
318 Our experimental data are used to test this decompression-rate-meter for andesitic systems. The
319 following equation, based on the numerical simulations of Toramaru (1995; 2006) and provided
320 by Hamada et al. (2010), was applied to estimate the BND for diffusion-controlled bubble
321 nucleation in andesitic melts:

322

$$323 \quad BND \approx 34 \cdot C_{SAT} \cdot \left(\frac{16 \cdot \pi \cdot \sigma_{LB}^3}{3 \cdot k \cdot T \cdot P_{SAT}^2} \right)^{-2} \cdot \left(\frac{\Omega_L \cdot P_{SAT}}{k \cdot T} \right)^{\frac{1}{4}} \cdot \left(\frac{P_{SAT}^2 \cdot k \cdot T \cdot C_{SAT} \cdot D_{H2O}}{4 \cdot \sigma_{LB}^2 \cdot r} \right)^{\frac{3}{2}} \quad (1)$$

324

325 where C_{SAT} is the H₂O content at saturation pressure expressed as the number of H₂O-molecules
326 per unit volume of liquid, σ_{LB} is the surface tension of the bubble-liquid interface in N/m, k is the
327 Boltzmann constant (1.38×10^{-23} J/K), T is the temperature in K, P_{SAT} is the H₂O saturation
328 pressure in Pa, Ω_L is the molecular volume of H₂O in the liquid (fixed to 3×10^{-29} m³; Toramaru
329 1989; Hamada et al. 2010), D_{H2O} is the H₂O diffusivity in the melt in m²/s, and r is the
330 decompression rate in Pa/s.

331 In a first approximation, P_{SAT} is assumed to be the P at which a H_2O activity [$a(H_2O)$] of
332 ~ 1 is reached in the andesitic system upon decompression. Thus, P_{SAT} strongly depends on the
333 initial H_2O content in the melt (here: 6.3 to 6.6 wt%). Assuming that the S and Cl contents in the
334 andesitic system do not affect significantly the water solubility, P_{SAT} is estimated to range
335 between 3.04×10^8 Pa for ~ 6.3 wt% H_2O initial and 3.31×10^8 Pa for ~ 6.6 wt% H_2O initial at
336 1030°C (see Table A.1 in the appendix) using the model of Burnham (1979). This assumption is
337 confirmed by experiments of Webster et al. (2009) indicating that S (~ 0.02 to ~ 0.13 wt%) and Cl
338 (~ 0.16 to ~ 0.84 wt%) have a minor effect on the H_2O solubility at 200 MPa and ~ 900 to $\sim 1000^\circ\text{C}$
339 in phonolitic systems, as well as by Fig. 6a and 6b, which clearly show that the final H_2O
340 contents in the decompressed glasses are largely independent of the bulk S and Cl concentrations.
341 Subsequently, C_{SAT} at P_{SAT} is calculated to range between 4.67×10^{27} and 4.90×10^{27} m^{-3} for ~ 6.3 to
342 ~ 6.6 wt% H_2O using the model of Ochs and Lange (1999) to estimate the density of the melt
343 (ρ_{melt}) at P_{SAT} and 1030°C (see Table A.1 in the appendix). Furthermore, we have used the
344 relationship provided by Behrens et al. (2004) for the estimation of the H_2O diffusivity in dacitic
345 melts as a function of T and H_2O content in the melt in order to calculate D_{H_2O} . Assuming an
346 initial H_2O content in the melt of 6.5 wt% (prior to decompression) and a final H_2O content of 3.6
347 wt% (after decompression; see also Fiege et al., 2014), D_{H_2O} is estimated to decrease from
348 2.03×10^{-11} m^2/s to 1.12×10^{-11} m^2/s during the decompression process. An average D_{H_2O} value of
349 1.58×10^{-11} m^2/s has been used for our *BND* calculations.

350 We have chosen the equation of Behrens et al. (2004) proposed for dacitic melts rather
351 than that proposed for an andesitic melt because the melt composition used in this study
352 (Krakatau andesite) is closer to the dacite than to the andesite. D_{H_2O} estimated using the
353 relationship for andesitic systems provided by Behrens et al. (2004) or using the general model
354 developed by Zhang and Ni (2010) would be ~ 2.5 to ~ 4.8 times higher than the D_{H_2O} of a dacitic

355 melt, which would result in a decrease of the BND values by ~ 0.6 to ~ 0.8 log units lower at given
356 r .

357 The value of σ_{LB} is estimated using the following equation (e.g., Hirth et al. 1970; Blander
358 and Katz 1975; Hurwitz and Navon 1994; Mangan and Sisson 2005; Gardner and Ketcham
359 2011):
360

$$361 \quad J = \frac{2 \cdot n_0^2 \cdot \Omega_L \cdot D_{H_2O} \cdot [\sigma_{LB} / (k \cdot T)]^{1/2}}{a_0} \cdot \exp\left[\frac{-16 \cdot \pi \cdot \sigma_{LB}^3}{3 \cdot k \cdot T \cdot \Delta P^2}\right] \quad (2)$$

362

363 where J is the nucleation rate calculated by dividing the determined BND by the time
364 allowed for nucleation (t_N ; see, e.g., Gardner and Ketcham 2011). Here, t_N is the sum of the
365 duration of the decompression and t_A . The number density of dissolved water molecules n_0 is
366 estimated using the concentration of H_2O molecules measured by IR spectroscopy in the starting
367 glasses and the density of the starting glasses ρ_{glass} . The values for ρ_{glass} are estimated from the
368 glass composition using the Gladstone-Dale rule (Gladstone and Dale 1863; see also Mandeville
369 et al. 2002). The IR measurements are described in detail in Fiege et al. (2014) and the
370 determined concentrations for molecular H_2O [$c(H_2O)$] and OH groups [$c(OH)$] are given in
371 Table A.1 in the appendix. The mean distance between dissolved water molecules a_0 is calculated
372 assuming the stereological definition $a_0 = 0.667 (d/V_{frac})$ from Russ (1986) in which d is the
373 diameter of a water molecule (3.7×10^{-10} m) and V_{frac} is the volume fraction of water molecules
374 in the melt. ΔP is the critical supersaturation pressure required to overcome the kinetic barrier to
375 nucleation at a specific T and H_2O content (Mangan and Sisson 2005). ΔP cannot be determined
376 directly from our experiments. However, the lower viscosity of the studied andesitic melts when

377 compared, for instance, to rhyodacitic melts should lead to a lower surface tension and a lower
378 ΔP . This is in agreement with the results of Mangan and Sisson (2005), which show a remarkable
379 decrease of σ_{LB} with the addition of mafic components (e.g., CaO, FeO_{total}, MgO). Thus, we
380 suggest that the ΔP of 60 MPa determined by Mangan and Sisson (2005) for rhyodacitic melts
381 (~69 wt% SiO₂; ~5 wt% initial H₂O) at 1000°C should represent an upper limit for the ΔP of the
382 studied composition. Preliminary experimental data of Masotta et al. (2013) indicate that ΔP may
383 be < 20 MPa in andesitic melts and, thus, σ_{LB} was estimated to range from 0.028 to 0.063 N/m
384 (see Table 2) in our experiments on evolved andesitic compositions, assuming a ΔP of 20 and 60
385 MPa, respectively. These upper and lower σ_{LB} values are applied to predict *BND*-trends for the
386 studied andesitic melt at 1030°C, assuming an average initial H₂O content of 6.5 wt% and an
387 average ρ_{melt} of 2.225 g/cm³ (see Fig. 5). Figure 5 shows clearly that $\log(BND \times \text{mm}^3)$
388 obtained from experiments of this study, which were quenched directly after continuous
389 decompression, increases linearly with $\log(r \times s/\text{MPa})$, being in good agreement with the
390 calculated trend using the Toramaru (2006) model. The experimental data for the initially H₂O-
391 undersaturated andesitic melts of this study indicate a similar slope as predicted by the Toramaru
392 (2006) model; however, the *BND* values of the experiments are about 0.5 to 1.5 log unit higher
393 than those of the modeled trends which were calculated assuming a σ_{LB} of 0.028 or 0.063 N/m.
394 Experimental data on homogeneous nucleation in CO₂-free rhyolitic systems obtained by Mangan
395 and Sisson (2000), Mourtada-Bonnefoi and Laporte (2004), Hamada et al. (2010) as well as
396 Gondé et al. (2011) at 700 to 1000°C are plotted in Figure 5 for comparison. CO₂-bearing
397 experiments are not considered as CO₂ presumably affects bubble formation to a significant
398 extent (see, e.g., Cichy et al. 2011). The data for homogeneous nucleation during decompression
399 of initially H₂O-saturated rhyolitic systems (Mangan and Sisson 2000; Mourtada-Bonnefoi and

400 Laporte 2004; Hamada et al. 2010) show *BND* values which are about 1 to 2 log units lower at
401 given *r* than values determined in this study for initially H₂O-undersaturated andesitic systems.
402 These previously published *BND* data for H₂O-saturated rhyolitic systems indicate even slightly
403 lower *BND* values than predicted by the Toramaru (2006) model for rhyolitic compositions at
404 800°C ($\sigma_{LB} = 0.08$ N/m; see Fig. 5). Noteworthy, the large scatter of most results published by
405 Gondé et al. (2011) does not allow a clear interpretation of their data. Considering that *T* has most
406 likely a minor effect on decompression-induced bubble nucleation (see Fig. 5 and Hamada et al.
407 2010), the experimental results of andesitic and rhyolitic systems indicate that the compositional
408 effect on bubble formation is more pronounced than predicted by the Toramaru (2006) model
409 using equation 2 for the estimation of σ_{LB} . For instance, Figure 5 shows that the Toramaru (2006)
410 model may be able to predict our data if a low σ_{LB} of ~0.015 N/m is assumed. On the other hand,
411 Pichavant et al. (2013) determined a σ_{LB} of 0.18 N/m for H₂O-CO₂-bearing basaltic melts at 1140
412 to 1180°C indicating that σ_{LB} can also be significantly higher in mafic melts than suggested by
413 the data of Mangan and Sisson (2005). However, the high σ_{LB} values found by Pichavant et al.
414 (2013) are probably in part related to the positive correlation between *T* and σ_{LB} (e.g.,
415 Bagdassarov et al 2000; Mangan and Sisson 2005) and/or to the influence of CO₂ on bubble
416 formation (e.g., Mourtada-Bonnefoi and Laporte 2002).

417 Strictly, the dependence of *BND* on *r* reflects the influence of the degree of volatile
418 supersaturation on bubble formation and is a non-linear function of changing total *P* (see
419 equation 1, and Mangan et al. 2004b; Mangan and Sisson 2005). Considering that all variables
420 except ΔP required for the calculation of σ_{LB} using equation 2 are well constrained and/or have a
421 negligible influence on σ_{LB} (e.g., P_{SAT}), the ΔP of the investigated andesitic melt is possibly lower
422 than 20 MPa. However, very low ΔP (< 8 MPa) are required in order to obtain a σ_{LB} of about

423 0.015 N/m using equation 2. Such low values for ΔP are quite unlikely considering that we
424 studied an evolved andesitic melt which does not show strong differences in major element
425 composition when compared to the rhyodacite of Mangan and Sisson (2005). Here, it is important
426 to emphasize that equation 1 and 2 do not account for the influence of volatiles other than H₂O on
427 bubble formation. Although the $\log(BND \times \text{mm}^3)$ values determined for experiments directly
428 quenched after fast decompression at $r = 0.1$ MPa/s (GYCIA-1, GYMClA-1, QFMA-1;
429 QFMClA-1) are similar (see Table 2), indicating that the variation of the bulk S (~140 to ~1000
430 ppm) and Cl (0 to ~1000 ppm) content may have a minor effect on bubble formation, the
431 presence of these two volatile may play an important role. In addition, the fluid phase released by
432 the andesitic melt upon decompression contains also significant amounts of charge balancing
433 cations like, for instance, Na⁺ and/or Ca²⁺ (see e.g., Zajacz et al. 2012). Hence, the diffusivity of
434 S, Cl, and alkalis in the melt (D_S , D_{Cl} , $D_{alkalis}$) should be included in an updated approach for the
435 calculation of BND as a function of r (equation 1) and for the calculation of σ_{LB} as a function of J
436 and ΔP (equation 2). This is justified by the analysis of Figure 6, showing that minor amounts of
437 S and Cl are still released from the melt to the fluid phase during the first hour of annealing after
438 fast decompression (0.1 MPa/s).

439 The H₂O-activity [a(H₂O)] prevailing in the melt prior to decompression may
440 provide another explanation for the strong deviation observed between our data for andesitic
441 systems and the modeled trend. Noteworthy, our data as well as the problematic dataset of Gondé
442 et al. (2011) are obtained for melts which are not saturated in H₂O before decompression, which
443 is in contrast to all other literature data shown in Figure 5. Gondé et al. (2011) proposed that
444 H₂O-understaturated melts could explain BND values which are several log units higher when
445 compared to H₂O-saturated melts. The presence of a free-fluid phase prior to decompression, e.g.,
446 in the form of microbubble nuclei, can possibly explain the lower BND values derived in initially

447 H₂O-saturated experiments due to dominating bubble growth over bubble nucleation upon
448 decompression. However, this would imply that the H₂O-saturated starting materials of Mangan
449 and Sisson (2000), Mourtada-Bonnefoi and Laporte (2004) and Hamada et al. (2010) contain
450 microbubbles prior to decompression without being mentioned by the authors, which is quite
451 unlikely. Presuming that H₂O-activity prior to decompression plays an important role, and
452 considering that magmatic systems are often not H₂O-saturated before an eruption (e.g., Hervig et
453 al. 1989; Gardner et al. 1995), the commonly much higher *BND* values determined for natural
454 pumices when compared to most experimental data (Hamada et al. 2010) can, at least in part, be
455 explained by the experimental strategy applied in most previous works; i.e., initially H₂O-
456 saturated melts.

457 In conclusion, the most plausible explanation for the discrepancies observed between our
458 dataset for a H₂O+S±Cl-bearing andesitic system and the trend(s) predicted by the Toramaru
459 (2006) model is that equation 1 (Toramaru (2006) model) and 2 (estimation of σ_{LB}) are
460 inadequate to predict decompression-induced bubble nucleation within initially H₂O+S±Cl-
461 undersaturated systems because, e.g., D_S , D_{Cl} , $D_{alkalis}$, and $a(H_2O)$ are not considered.

462

463 **IMPLICATIONS FOR NATURAL SYSTEMS**

464 This study presents for the first time experimental data on bubble formation in andesitic
465 systems during continuous decompression, while previous experimental investigations (using
466 comparable techniques) exclusively focused on rhyolitic systems. Two main results are of
467 importance to the *BND* interpretation of pumices formed, for instance, during explosive Plinian
468 eruptions and can help to evaluate existing bubble nucleation models.

469 First, the strong influence of t_A on *BND* mentioned above (Fig. 4b) indicates that a good
470 knowledge of possible storage times during magma ascent is required to interpret magma ascent
471 rates on the basis of natural pumice samples.

472 Second, our data show that anhydrous melt composition is crucial to model bubble
473 formation. Considering that the composition of the studied dacitic andesite (~65 wt% bulk SiO₂
474 in the anhydrous starting glass) is similar to the composition of the material erupted by some
475 hazardous volcanic eruptions such as the 1883 eruption of Krakatau volcano (~64 to 69 wt% SiO₂
476 in the almost dry bulk samples, ~60 to ~72 wt% in the matrix glasses and glass inclusions;
477 Mandeville et al. 1996) our data are important for our understanding of the vesiculation in the
478 volcanic ejecta of this type of volcanism. Ascent rates estimated on the basis of *BND* in dacitic to
479 andesitic volcanic ejecta may be overestimated by about an order of magnitude if the model of
480 Toramaru (2006) is used to account for compositional variations. Thus, the interpretation of the
481 vesiculation of tephritic samples, which cover commonly a wide range of compositions (SiO₂
482 content of glass shards of the Mariana tephra ranges from ~53 to ~67 wt%; Straub 1995) needs
483 experimental calibrations for appropriate composition. A larger dataset, accounting for variations
484 in major elements as well as in initial volatile contents (H₂O ± CO₂ ± S ± Cl) would be helpful to
485 calibrate the Toramaru (2006) model. This improved “decompression-meter” would allow us to
486 obtain detailed insights on how (fast) magma travels from a deep reservoir to the surface based
487 on vesiculated samples.

488

489 **ACKNOWLEDGEMENTS**

490 This project was supported by the German Science Foundation (BE1720/25-1). We would
491 like to thank Harald Behrens (scientific discussions), Eric Wolff (electron microprobe) and Otto
492 Dietrich (sample preparation). The valuable comments of the journal editor F. Castro as well as

493 of the reviewers M. Mangan, C. Martel, and A. Toramaru have significantly improved the quality
494 of this manuscript.

495

496

REFERENCES

- 497 Armienti, P. (2008) Decryption of igneous rock textures: crystal size distribution tools. In K.D.
498 Putirka and F.J. Tepley III, Eds., Minerals, Inclusions and Volcanic Processes, 69, p. 623–649.
499 Reviews in Mineralogy and Geochemistry, Mineralogical Society of America, Chantilly,
500 Virginia.
- 501 Bagdassarov, N., Dorfman, A., and Dingwell, D.B. (2000) Effect of alkalis, phosphorus, and
502 water on the surface tension of haplogranite melt. *American Mineralogist*, 85, 33-40.
- 503 Behrens, H., Zhang, Y.X., and Xu, Z.G. (2004) H₂O diffusion in dacitic and andesitic melts.
504 *Geochimica et Cosmochimica Acta*, 68, 5139-5150.
- 505 Berndt, J., Liebske, C., Holtz, F., Freise, M., Nowak, M., Ziegenbein, D., Hurkuck, W., and
506 Koepke, J. (2002) A combined rapid-quench and H₂-membrane setup for internally heated
507 pressure vessels: Description and application for water solubility in basaltic melts. *American*
508 *Mineralogist*, 87, 1717-1726.
- 509 Blander, M., and Katz, J.L. (1975) Bubble nucleation in liquids. *American Institute of Chemical*
510 *Engineers Journal*, 21, 833-848.
- 511 Botcharnikov, R.E., Behrens, H., and Holtz, F. (2006) Solubility and speciation of C-O-H fluids
512 in andesitic melt at T=1100-1300 degrees C and P=200 and 500MPa. *Chemical Geology*, 229,
513 125-143.
- 514 Burnham, C.W. (1979) Magmas and hydrothermal fluids. In H. L. Barnes, Ed., *Geochemistry of*
515 *hydrothermal ore deposits*, p. 71-136. John Willey and Sons, New York.

- 516 Cashman, K.V. (2004) Volatile controls on magma ascent and eruption. In R.S.J. Sparks and C.J.
517 Hawkesworth, Eds., *The State of the Planet: Frontiers and Challenges in Geophysics*, 150, p.
518 109-124. Geophysical Monograph series, American Geophysical Union, Washington.
- 519 Castro, J.M., Burgisser, A., Schipper, C.I., and Mancini, S. (2012) Mechanisms of bubble
520 coalescence in silicic magmas. *Bulletin of Volcanology*, 74, 2339-2352.
- 521 Cichy, S.B., Botcharnikov, R.E., Holtz, F., and Behrens, H. (2011) Vesiculation and Microlite
522 Crystallization Induced by Decompression: a Case Study of the 1991–1995 Mt Unzen
523 Eruption (Japan). *Journal of Petrology*, 52, 1469-1492.
- 524 Fiege, A., Behrens, H., Holtz, F., and Adams, F. (2014) Kinetic vs. thermodynamic control of
525 degassing of H₂O-S±Cl bearing andesitic melts. *Geochimica et Cosmochimica Acta*, 125, 241-
526 254. (DOI: 10.1016/j.gca.2013.10.012)
- 527 Gardner, J.E. (2007) Heterogeneous bubble nucleation in highly viscous silicate melts during
528 instantaneous decompression from high pressure. *Chemical Geology*, 236, 1-12.
- 529 Gardner, J.E., and Denis, M.-H. (2004) Heterogeneous bubble nucleation on Fe-Ti oxide crystals
530 in high-silica rhyolitic melts. *Geochimica et Cosmochimica Acta*, 68, 3587-3597.
- 531 Gardner, J.E., and Ketcham, R.A. (2011) Bubble nucleation in rhyolite and dacite melts:
532 temperature dependence of surface tension. *Contributions to Mineralogy and Petrology*, 162,
533 929-943.
- 534 Gardner, J.E., Rutherford, M., Carey, S., and Sigurdsson, H. (1995) Experimental constraints on
535 pre-eruptive water contents and changing magma storage prior to explosive eruptions of
536 Mount St Helens volcano. *Bulletin of Volcanology*, 57, 1-17.
- 537 Gladstone, and Dale (1863) Researches on the Refraction, Dispersion, and Sensitiveness of
538 Liquids. *Philosophical Transactions of the Royal Society of London*, 153, 317-343.

- 539 Gondé, C., Martel, C., Pichavant, M., and Bureau, H. (2011) In situ bubble vesiculation in silicic
540 magmas. *American Mineralogist*, 96, 111-124.
- 541 Hamada, M., Laporte, D., Cluzel, N., Koga, K.T., and Kawamoto, T. (2010) Simulating bubble
542 number density of rhyolitic pumices from Plinian eruptions: constraints from fast
543 decompression experiments. *Bulletin of Volcanology*, 72, 735-746.
- 544 Hervig, R.L., Dunbar, N., Westrich, H.R., and Kyle, P.R. (1989) Pre-eruptive water content of
545 rhyolitic magmas as determined by ion microprobe analyses of melt inclusions in phenocrysts.
546 *Journal of Volcanology and Geothermal Research*, 36, 293-302.
- 547 Higgins, M.D. (2000) Measurement of crystal size distributions. *American Mineralogist*, 85,
548 1105-1116.
- 549 Higgins, M.D. (2002) Closure in crystal size distributions (CSD), verification of CSD
550 calculations, and the significance of CSD fans. *American Mineralogist*, 87, 171-175.
- 551 Higgins, M.D., and Chandrasekharam, D. (2007) Nature of Sub-volcanic Magma Chambers,
552 Deccan Province, India: Evidence from Quantitative Textural Analysis of Plagioclase
553 Megacrysts in the Giant Plagioclase Basalts. *Journal of Petrology*, 48, 885-900.
- 554 Hirth, J.P., Pound, G.M., and Pierre, G.R.St. (1970) Bubble nucleation. *Metallurgical*
555 *Transactions*, 1, 939-945.
- 556 Hurwitz, S., and Navon, O. (1994) Bubble nucleation in rhyolitic melts: Experiments at high
557 pressure, temperature, and water content. *Earth and Planetary Science Letters*, 122, 267-280.
- 558 King, P.L., and Jolloway, J.R. (2002) CO₂ solubility and speciation in intermediate (andesitic)
559 melts: The role of H₂O and composition. *Geochimica et Cosmochimica Acta*, 66, 1627-1640.
- 560 Lautze, N., Sisson, T., Mangan, M., and Grove, T. (2011) Segregating gas from melt: an
561 experimental study of the Ostwald ripening of vapor bubbles in magmas. *Contributions to*
562 *Mineralogy and Petrology*, 161, 331-347.

- 563 Mandeville, C.W., Carey, S., and Sigurdsson, H. (1996) Magma mixing, fractional crystallization
564 and volatile degassing during the 1883 eruption of Krakatau volcano, Indonesia. *Journal of*
565 *Volcanology and Geothermal Research*, 74, 243-274.
- 566 Mandeville, C.W., Webster, J.D., Rutherford, M.J., Taylor, B.E., Timbal, A., and Faure, K.
567 (2002) Determination of molar absorptivities for infrared absorption bands of H₂O in andesitic
568 glasses. *American Mineralogist*, 87, 813-821.
- 569 Mangan, M., and Sisson, T. (2000) Delayed, disequilibrium degassing in rhyolite magma:
570 decompression experiments and implications for explosive volcanism. *Earth and Planetary*
571 *Science Letters*, 183, 441-455.
- 572 Mangan, M., and Sisson, T. (2005) Evolution of melt-vapor surface tension in silicic volcanic
573 systems: Experiments with hydrous melts. *Journal of Geophysical Research*, 110, B01202.
- 574 Mangan, M., Mastin, L., and Sisson, T. (2004a) Gas evolution in eruptive conduits: combining
575 insights from high temperature and pressure decompression experiments with steady-state
576 flow modeling. *Journal of Volcanology and Geothermal Research*, 129, 23-36.
- 577 Mangan, M.T., Sisson, T.W., and Hankins, W.B. (2004b) Decompression experiments identify
578 kinetic controls on explosive silicic eruptions. *Geophysical Research Letters*, 31,
579 L08605. Masotta, M., Ni, H., and Keppler, H. (2013) In-situ observations of bubble growth in
580 basaltic, andesitic and rhyodacitic melts. Abstract V13H-04 presented at 2013 Fall Meeting,
581 San Fransisco, Calif., 9-13 Dec.
- 582 Mourtada-Bonnefoi, C.C., and Laporte, D. (1999) Experimental study of homogeneous bubble
583 nucleation in rhyolitic magmas. *Geophysical Research Letters*, 26, 3505-3508.
- 584 Mourtada-Bonnefoi, C.C., and Laporte, D. (2002) Homogeneous bubble nucleation in rhyolitic
585 magmas: An experimental study of the effect of H₂O and CO₂. *Journal of Geophysical*
586 *Research*, 107, ECV 2-1-EVC 2-19.

- 587 Mourtada-Bonnefoi, C.C., and Laporte, D. (2004) Kinetics of bubble nucleation in a rhyolitic
588 melt: an experimental study of the effect of ascent rate. *Earth and Planetary Science Letters*,
589 218, 521-537.
- 590 Noguchi, S., Toramaru, A., and Nakada, S. (2008) Relation between microlite textures and
591 discharge rate during the 1991-1995 eruptions at Unzen, Japan. *Journal of Volcanology and*
592 *Geothermal Research*, 175, 141-155.
- 593 Nowak, M., Cichy, S.B., Botcharnikov, R.E., Walker, N., and Hurkuck, W. (2011) A new type of
594 high-pressure low-flow metering valve for continuous decompression: First experimental
595 results on degassing of rhyodacitic melts. *American Mineralogist*, 96, 1373-1380.
- 596 Ochs, F.A., and Lange, R.A. (1999) The Density of Hydrous Magmatic Liquids. *Science*, 283,
597 1314-1317.
- 598 Pichavant, M., Carlo, I., Rotolo, S.G., Scaillet, B., Burgisser, A., Gall, N., and Martel, C.(2013)
599 Generation of CO₂-rich melts during basalt magma ascent and degassing. *Contributions to*
600 *Mineralogy and Petrology*, 166, 545-561.
- 601 Russ, J.C. (1986) *Practical Stereology*, Springer, New York.
- 602 Sparks, R.S.J., Barclay, J., Jaupart, C., Mader, H.M., and Phillips, J.C. (1994) Physical Aspects
603 of Magma Degassing – I. Experimental and theoretical constraints on vesiculation. In M.R.
604 Carroll and J.R. Holloway, Eds., *Volatiles in Magmas*, 30, p. 413–445. Reviews in
605 *Mineralogy*, Chantilly, Virginia.
- 606 Straub, S.M. (1995) Contrasting compositions of Mariana Trough fallout tephra and Mariana
607 Island arc volcanic: a fractional crystallization link. *Bulletin of Volcanology*, 57, 403-421.
- 608 Toramaru, A. (1989) Vesiculation process and bubble size distributions in ascending magmas
609 with constant velocities. *Journal of Geophysical Research*, 94, 17523-17542.

- 610 Toramaru, A. (1995) Numerical study of nucleation and growth of bubbles in viscous magmas.
611 Journal of Geophysical Research, 100, 1913-1931.
- 612 Toramaru, A. (2006). BND (bubble number density) decompression rate meter for explosive
613 volcanic eruptions. Journal of Volcanology and Geothermal Research, 154, 303-316.
- 614 Webster, J.D., Sintoni, M.F., and De Vivo, B. (2009) The partitioning behavior of Cl, S, and H₂O
615 in aqueous vapor - +/- saline-liquid saturated phonolitic and trachytic melts at 200 MPa.
616 Chemical Geology, 263, 19-36.
- 617 Zajacz, Z., Candela, P.A., Piccoli, P.M., and Sanchez-Valle, C. (2012) The partitioning of sulfur
618 and chlorine between andesite melts and magmatic volatiles and the exchange coefficients of
619 major cations. Geochimica et Cosmochimica Acta, 89, 81-101.
- 620 Zhang, Y., and Ni, H. (2010) Diffusion of H, C, and O Components in Silicate Melts. In Y.
621 Zhang and D.J. Cherniak, Eds., Diffusion in Minerals and Melts, 72, p. 171–225. Reviews in
622 Mineralogy, Chantilly, Virginia.

623

624

LIST OF FIGURE CAPTIONS

625 **Figure 1.** Microscopic image of sample GYCIA-3 ($r \sim 0.1$ MPa/s; $t_A = 5$ h). The bubble size
626 ranges from about ~ 5 to ~ 80 μm in diameter.

627

628 **Figure 2a-c.** BSE images of products of selected experiments with various decompression rates
629 r . a) QFM CIA-1, b) QFM CIA-3 and c) QFM CIA-5. Values for r are displayed in the upper right
630 corner of each image and $t_A = 0$ h for the three experiments. Decreasing r leads to a decreasing
631 number of bubbles and an increasing size of bubbles. Some bubbles are elongated and show a
632 preferred orientation (flow structures).

633

634 **Figure 3a-f.** Selected BSE images from experimental series GYMClA. a) and b) GYMClA-1; c)
635 and d) GYMClA-5; e) and f) GYMClA-3. The decompression rate r of all three experiments was
636 0.1 MPa/s.

637

638 **Figure 4a-b.** BND plotted as a function of decompression rate r and annealing time t_A . a)
639 $\log(BND \times \text{mm}^3)$ vs. r (*logarithmic scaling*). Samples of the experimental series QFMA,
640 QFMClA, GYClA and GYMClA which have been directly quenched after continuous
641 decompression ($t_A = 0$ h) are plotted together with the results of Cichy et al. (2011; Ci11)
642 obtained for a rhyodacitic melt compositions. Note that decompression experiments of Cichy et
643 al. (2011) with $r \geq 0.1$ MPa/s were conducted applying a continuous decompression technique
644 similar to that in this study, while those with $r \leq 0.01$ MPa/s are carried out using a multi-step
645 approach. b) $\log(BND \times \text{mm}^3)$ vs. t_A at constant $r = 0.1$ MPa/s (experimental series: GYClA and
646 GYMClA).

647 *Horizontal error bars* are smaller than symbol size.

648 a: QFMA experiments: ~ 140 ppm initial S, Cl-free, $\sim \text{QFM}+0.8$

649 b: QFMClA: ~ 240 ppm initial S, ~ 1000 ppm initial Cl, $\sim \text{QFM}+0.8$

650 c: GYClA: ~ 1000 ppm initial S, ~ 1000 ppm initial Cl, $\sim \text{QFM}+1.8$

651 d: GYMClA: ~ 1000 ppm initial S, 500 ppm initial Cl, $\sim \text{QFM}+1.8$.

652

653 **Figure 5.** Results of this study and previously published data based on continuous decompression
654 experiments and comparison to the trends predicted by the model of Toramaru (2006). The black
655 arrow indicates that two experiments of Mangan and Sisson (2000) conducted at $r = 0.025$ MPa/s
656 show BND of $10^{-0.59} \text{ mm}^{-3}$ and $10^{-1.32} \text{ mm}^{-3}$, respectively.

657 *Horizontal error bars* are smaller than symbol size.

658 Ci11: Experiments of Cichy et al. (2011); only CO₂-free, continuous decompression experiments
659 are considered; rhyodacite; 850°C

660 Go11: Gondé et al. (2011); haplogranite; 730 to 1000°C

661 Ha10: Hamada et al. (2010); rhyolitic obsidian; 700 to 800°C

662 MS00: Mangan and Sisson (2000); rhyolitic obsidian; 900°C

663 MBL04: Mourtada-Bonnefoi and Laporte (2004); rhyolite; 800°C

664 To06: The trends were calculated for rhyolitic and andesitic systems using the model of
665 Toramaru (2006), see text for further details.

666 *homogen.*: Homogeneous bubble nucleation

667 *heterogen.*: Heterogeneous bubble nucleation

668 *saturated*: The melt was H₂O saturated prior to decompression.

669 *a*: The andesite-trends were calculated for 1030°C and $\sigma_{LB} = 0.015, 0.028$ or 0.063 N/m,
670 assuming average values for the initial the H₂O content in the melt (6.5 wt%) and for ρ_{melt} (2.225
671 g/cm³); see text for details.

672 *b*: The rhyolite-trend was calculated for 800°C using the values given by Hamada et al. (2010)
673 for C_{SAT} , P_{SAT} , Ω_L and D_{H2O} and assuming a σ_{LB} of 0.08 N/m.

674

675 **Figure 6a-f.** H₂O, S, and Cl contents of the starting glasses and the partially degassed glasses. a)
676 H₂O content vs. r ; b) H₂O content vs. t_A ; c) S content vs. r ; d) S content vs. t_A ; e) Cl content vs. r ;
677 f) Cl content vs. t_A . The grey shaded area marks the decompression step. The H₂O, S, and Cl
678 contents of all glasses are given in Table 2 of Fiege et al. (2014). The H₂O concentrations within
679 the glasses were measured using IR spectroscopy, while the S and Cl contents were determined
680 using EMP (for details see Fiege et al 2014).

Table 1. Chemical composition (wt%) of the anhydrous Krakatau andesite used as a starting material.

SiO₂	64.79	(59)
TiO₂	1.28	(7)
Al₂O₃	15.62	(22)
FeO_{tot} (a)	4.94	(33)
MnO	0.23	(4)
MgO	1.4	(6)
CaO	4.92	(2)
Na₂O	3.92	(25)
K₂O	1.8	(7)
Total	98.91	(78)

Note: Errors are provided in parentheses: standard deviation (1 sigma) based on EMP analysis; number of EMP analyses: 20.

(a) FeO_{tot}: total Fe concentration in the glass given as FeO

682

683

Table 2. Bubble number densities (*BND*) and surface tension (σ_{LB}).

Sample ID	\log (<i>BND</i> × mm ³) (a)	r [MPa/s]	t_A [h]	σ_{LB} [N/m] (b)	Sample ID	\log (<i>BND</i> × mm ³) (a)	r [MPa/s]	t_A [h]	σ_{LB} [N/m] (b)
GYCIA; initial volatile content: 6.56 (24) wt% H ₂ O; 1018 (188) ppm S; 1033 (21) ppm Cl					QFMA; initial volatile content: 6.47 (13) wt% H ₂ O; 137 (15) ppm S				
GYCIA-1	4.92 (28)	0.10	-	0.028 / 0.059	QFMA-1	4.55 (31)	0.10	-	0.029 / 0.060
GYCIA-2	4.77 (30)	0.10	1.10	0.029 / 0.060	QFMA-2	4.67 (32)	0.05	-	0.029 / 0.060
GYCIA-3	4.00 (60)	0.10	5.00	0.029 / 0.061	QFMA-3	4.37 (11)	0.02	-	0.029 / 0.060
GYCIA-4	3.03 (40)	0.10	72.06	0.030 / 0.062	QFMA-4	3.98 (37)	0.01	-	0.029 / 0.061
GYCIA-5	4.88 (34)	0.10	1.05	0.029 / 0.060	QFMA-5	2.14 (67)	0.0005	-	0.030 / 0.063
GYMCIA; initial volatile content: 6.29 (24) wt% H ₂ O; 1043 (143) ppm S; 510 (10) ppm Cl					QFMCIA; initial volatile content: 6.58 (13) wt% H ₂ O; 260 (19) ppm S; 1042 (24) ppm Cl				
GYMCIA-1	4.68 (64)	0.10	-	0.029 / 0.060	QFMCIA-1	4.24 (35)	0.10	-	0.029 / 0.060
GYMCIA-3	2.62 (32)	0.10	5.00	0.030 / 0.062	QFMCIA-2	4.43 (25)	0.05	-	0.029 / 0.060
GYMCIA-4	2.78 (35)	0.10	72.06	0.030 / 0.063	QFMCIA-3	3.90 (76)	0.02	-	0.029 / 0.061
GYMCIA-5	3.61 (17)	0.10	1.05	0.029 / 0.061	QFMCIA-4	3.43 (26)	0.01	-	0.029 / 0.061
					QFMCIA-5	2.19 (56)	0.0005	-	0.030 / 0.063

Note: Errors are provided in parentheses: standard deviation (1 sigma) of the six analyses performed (3 BSE images; bubbles at the edges included/exclude; see text for details).

(a): logarithmized *BND* values [mm⁻³].

(b): calculated for $\Delta P = 20$ MPa (first value) or 60 MPa (second value); see text for details.

- : $t_A = 0$ h; relative error of r : ~7 %; error of σ_{LB} : ± 0.003 N/m

684

685

FIGURES

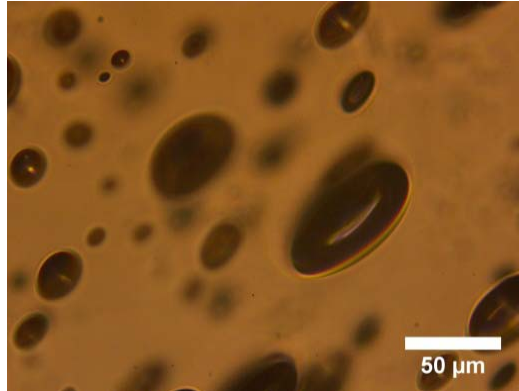


Figure 1. Microscopic image of sample GYCIA-3 ($r \sim 0.1$ MPa/s; $t_A = 5$ h). The bubble size ranges from about ~ 5 to ~ 80 μm in diameter.

686

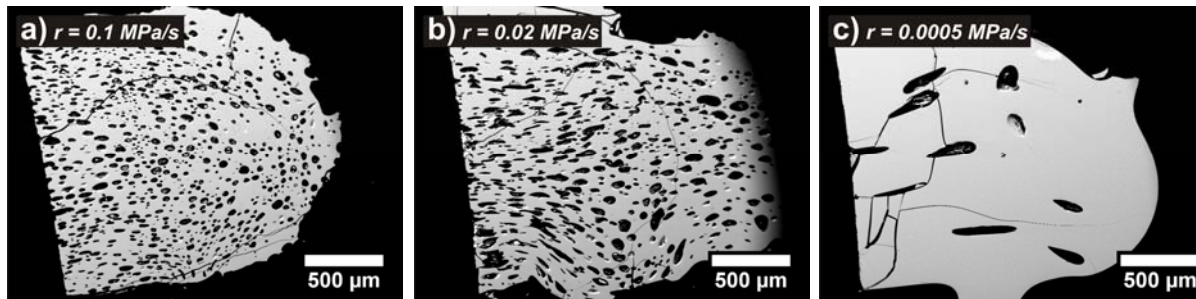


Figure 2a-c. BSE images of products of selected experiments with various decompression rates r . a) QFMCIA-1, b) QFMCIA-3 and c) QFMCIA-5. Values for r are displayed in the upper right corner of each image and $t_A = 0$ h for the three experiments. Decreasing r leads to a decreasing number of bubbles and an increasing size of bubbles. Some bubbles are elongated and show a preferred orientation (flow structures).

687

688

689

690

691

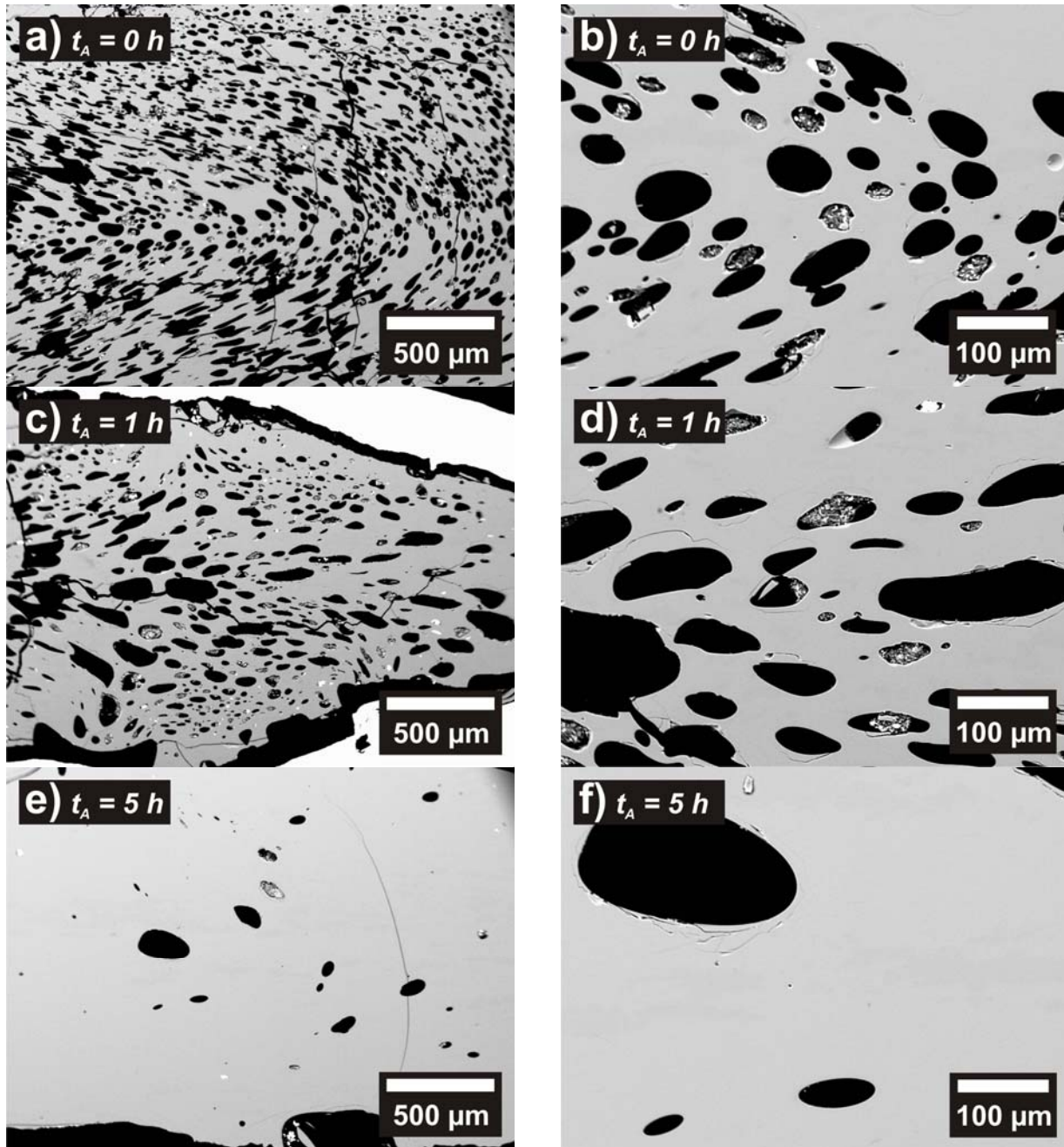


Figure 3a-f. Selected BSE images from experimental series GYMCIA. a) and b) GYMCIA-1; c) and d) GYMCIA-5; e) and f) GYMCIA-3. The decompression rate r of all three experiments was 0.1 MPa/s.

692

693

694

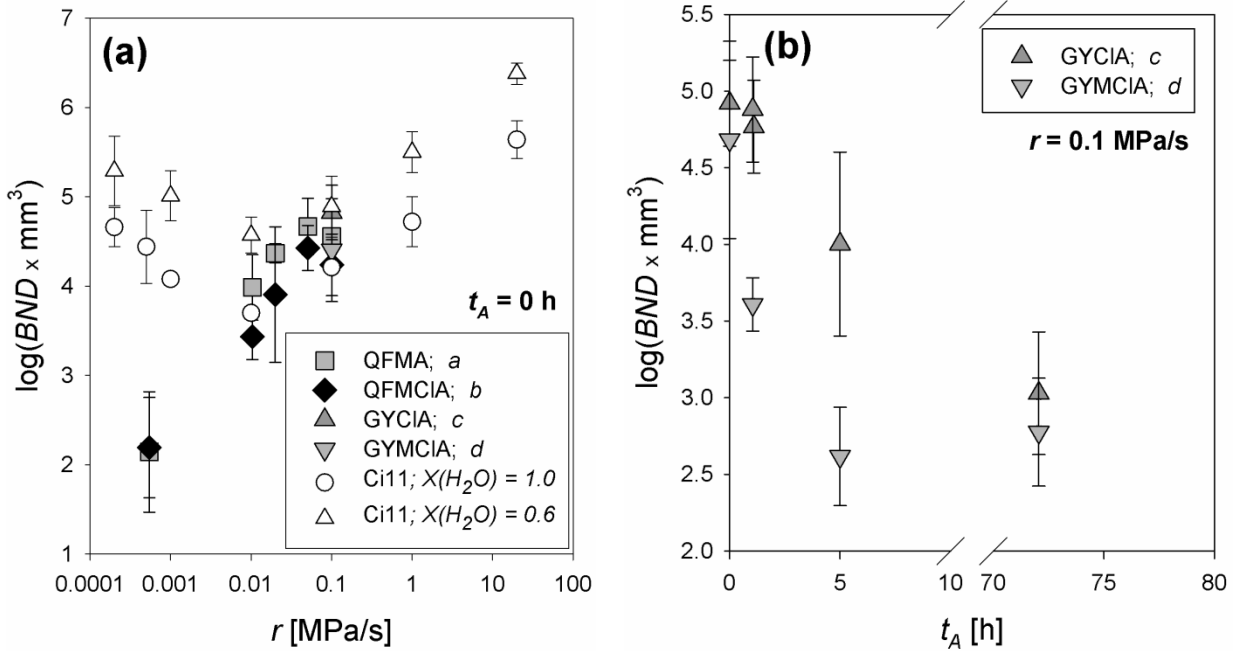


Figure 4a-b. BND plotted as a function of decompression rate r and annealing time t_A . a) $\log(BND \times \text{mm}^3)$ vs. r (logarithmic scaling). Samples of the experimental series QFMA, QFMClA, GYCIA and GYMClA which have been directly quenched after continuous decompression ($t_A = 0$ h) are plotted together with the results of Cichy et al. (2011; Ci11) on rhyodacitic melt compositions. Here, decompression experiments of Cichy et al. (2011) with $r \geq 0.1$ MPa/s were conducted applying the same continuous decompression technique, while those with $r \leq 0.01$ MPa/s are carried out using a multi-step approach. b) $\log(BND \times \text{mm}^3)$ vs. t_A at constant $r = 0.1$ MPa/s (experimental series: GYCIA and GYMClA).

Horizontal error bars are smaller than symbol size.

a: QFMA experiments: ~ 140 ppm initial S, Cl-free, $\sim \text{QFM} + 0.8$

b: QFMClA: ~ 240 ppm initial S, ~ 1000 ppm initial Cl, $\sim \text{QFM} + 0.8$

c: GYCIA: ~ 1000 ppm initial S, ~ 1000 ppm initial Cl, $\sim \text{QFM} + 1.8$

d: GYMClA: ~ 1000 ppm initial S, 500 ppm initial Cl, $\sim \text{QFM} + 1.8$.

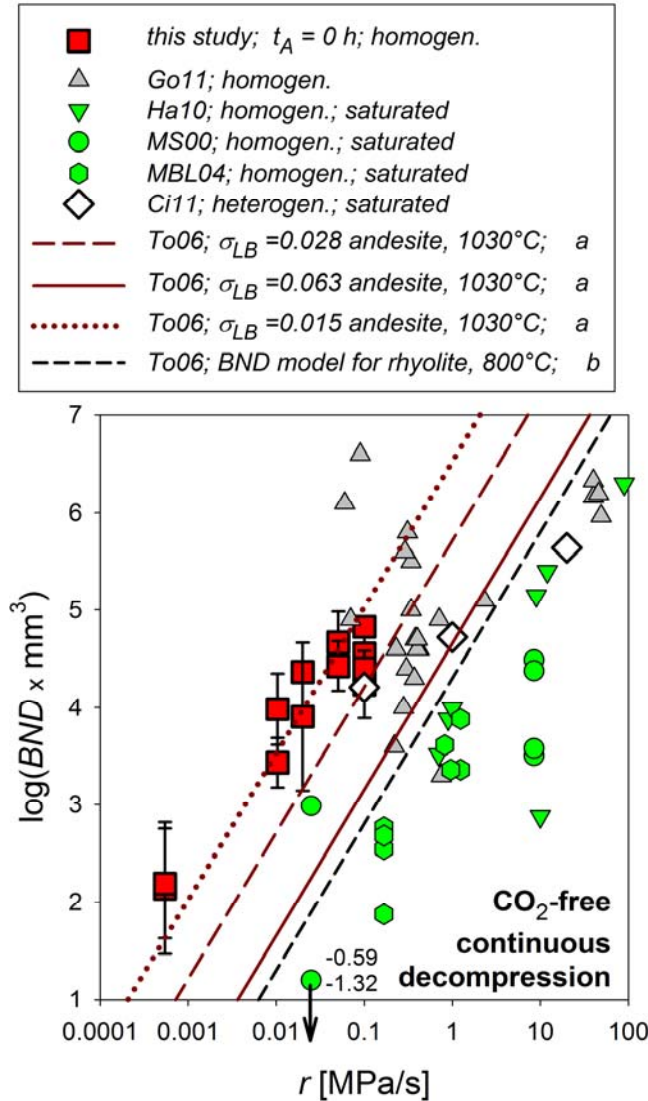


Figure 5. Results of this study and previously published data based on continuous decompression experiments and comparison to the trends predicted by the model of Toramaru (2006). The black arrow indicates that two experiments of Mangan and Sisson (2000) conducted at $r = 0.025$ MPa/s show BND of $10^{-0.59}$ mm⁻³ and $10^{-1.32}$ mm⁻³, respectively.

Horizontal error bars are smaller than symbol size.

Ci11: Experiments of Cichy et al. (2011); only CO₂-free, continuous decompression experiments are considered; rhyodacite; 850°C

Go11: Gondé et al. (2011); haplogranite; 730 to 1000°C

Ha10: Hamada et al. (2010); rhyolitic obsidian; 700 to 800°C

MS00: Mangan and Sisson (2000); rhyolitic obsidian; 900°C

MBL04: Mourtada-Bonnefoi and Laporte (2004); rhyolite; 800°C

To06: The trends were calculated for rhyolitic and andesitic systems using the model of Toramaru (2006), see text for further details.

homogen.: Homogeneous bubble nucleation

heterogen.: Heterogeneous bubble nucleation

saturated: The melt was H₂O saturated prior to decompression.

a: The andesite-trends were calculated for 1030°C and $\sigma_{LB} = 0.015, 0.028$ or 0.063 N/m, assuming average values for the initial H₂O content in the melt (6.5 wt%) and for ρ_{melt} (2.225 g/cm³); see text for details.

b: The rhyolite-trend was calculated for 800°C using the values given by Hamada et al. (2010) for $C_{SAT}, P_{SAT}, \Omega_L$ and D_{H2O} and assuming a σ_{LB} of 0.08 N/m.

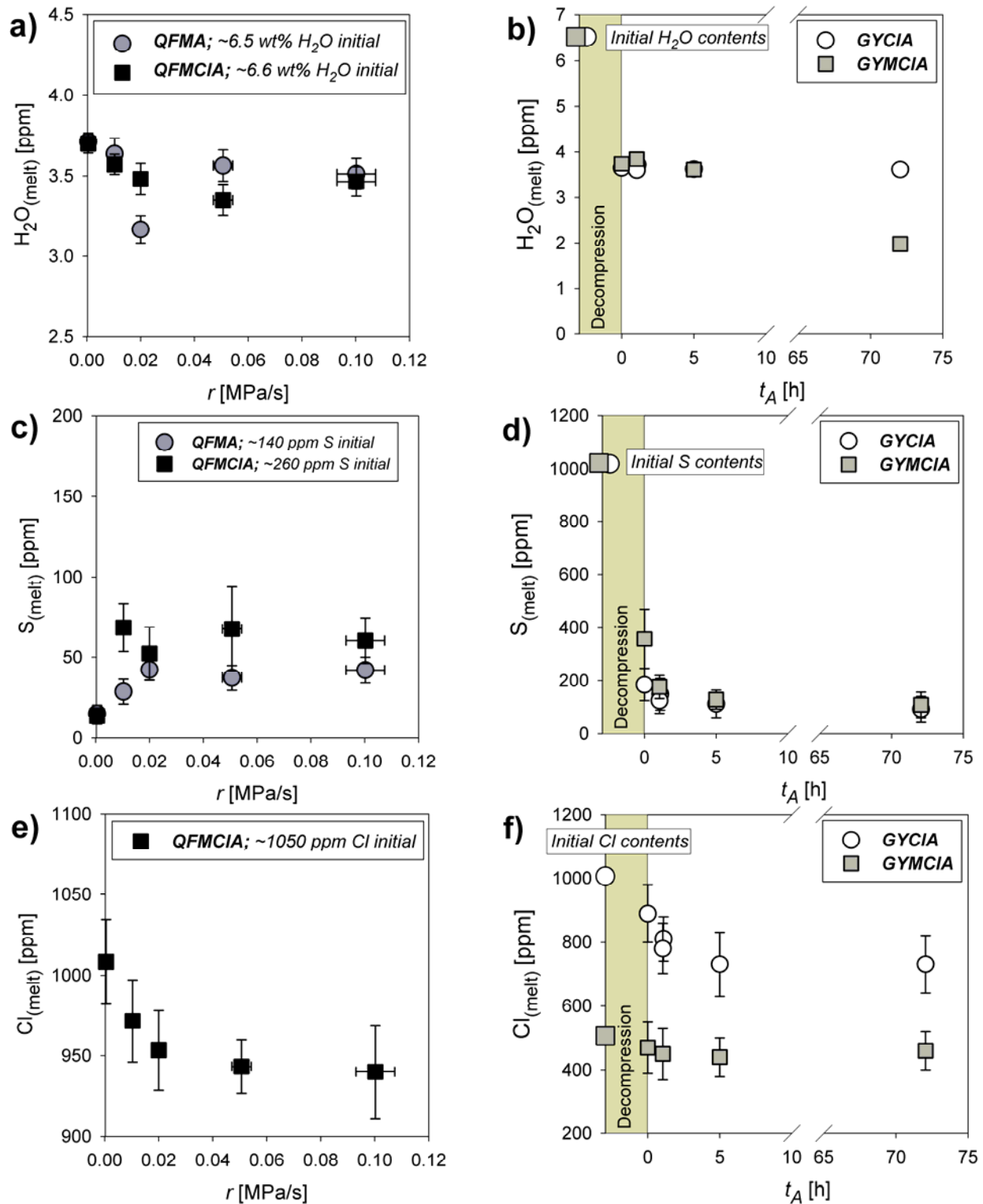
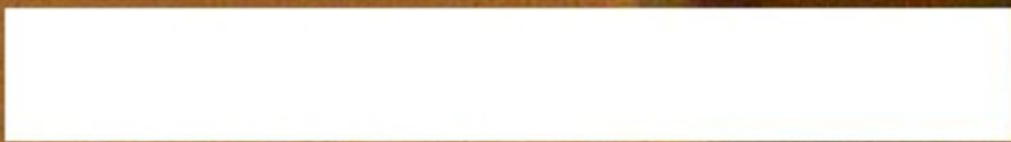
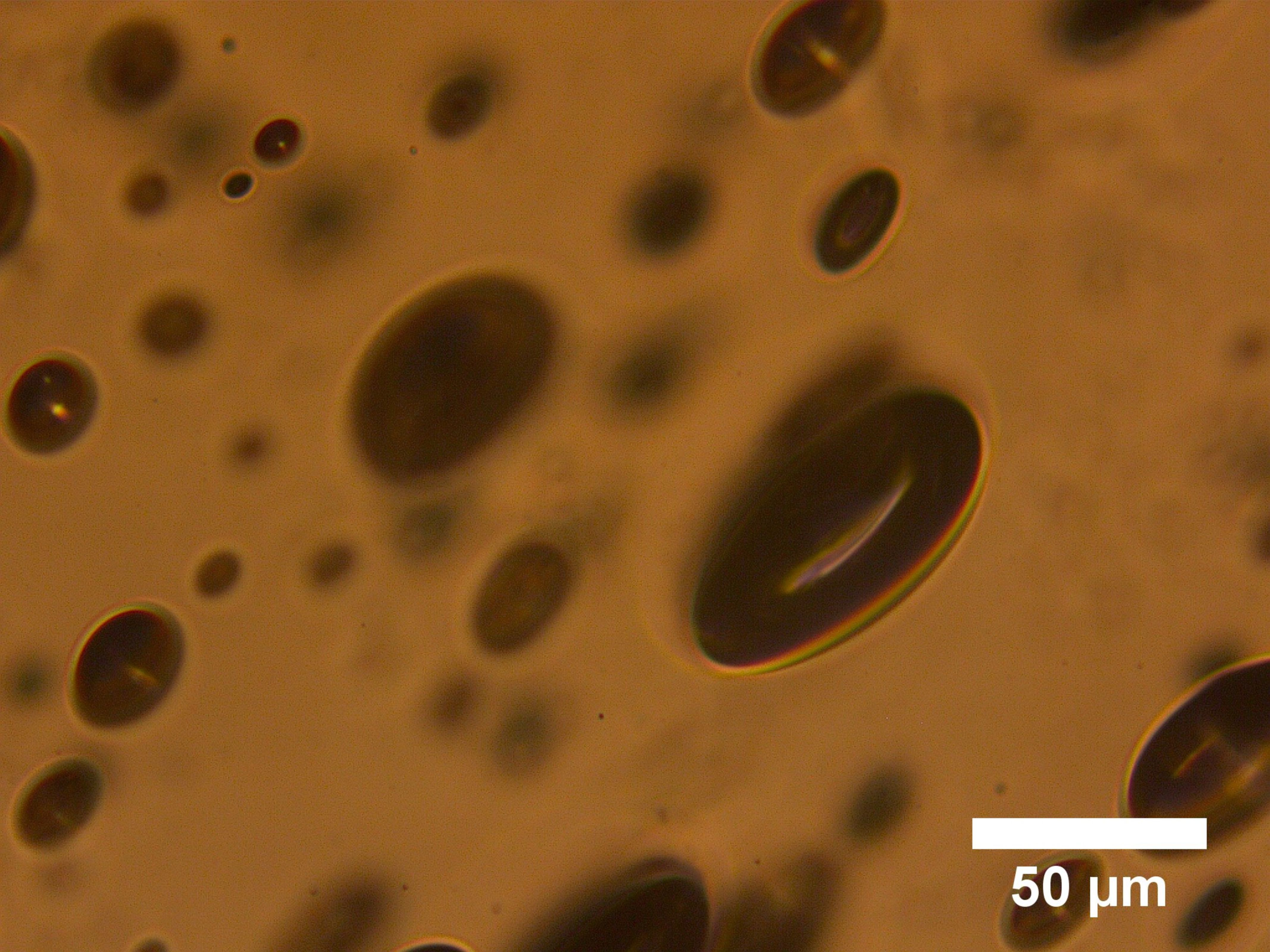


Figure 6a-f. H_2O , S, and Cl contents of the starting glasses and the partially degassed glasses. a) H_2O content vs. r ; b) H_2O content vs. t_A ; c) S content vs. r ; d) S content vs. t_A ; e) Cl content vs. r ; f) Cl content vs. t_A . The grey shaded area marks the decompression step. The H_2O , S, and Cl contents of all glasses are given in Table 2 of Fiege et al. (2014). The H_2O concentrations within the glasses were measured using IR spectroscopy, while the S and Cl contents were determined using EMP (for details see Fiege et al 2014).



50 μm

a) $r = 0.1 \text{ MPa/s}$



500 μm

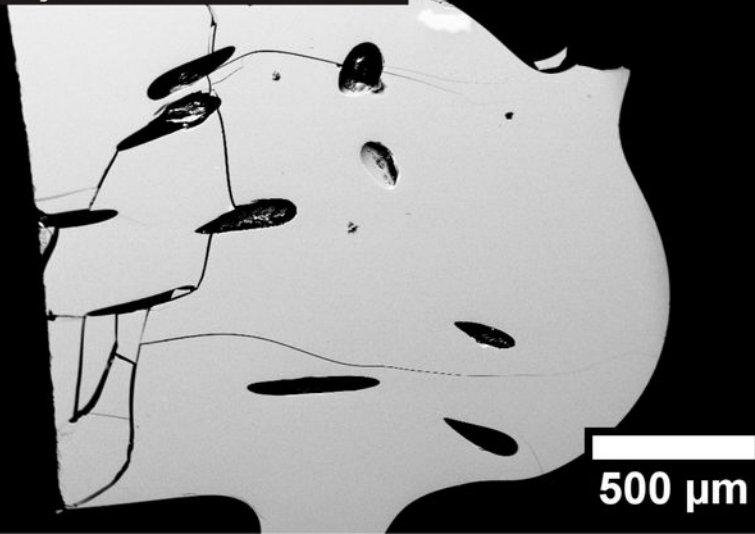
b) $r = 0.02 \text{ MPa/s}$



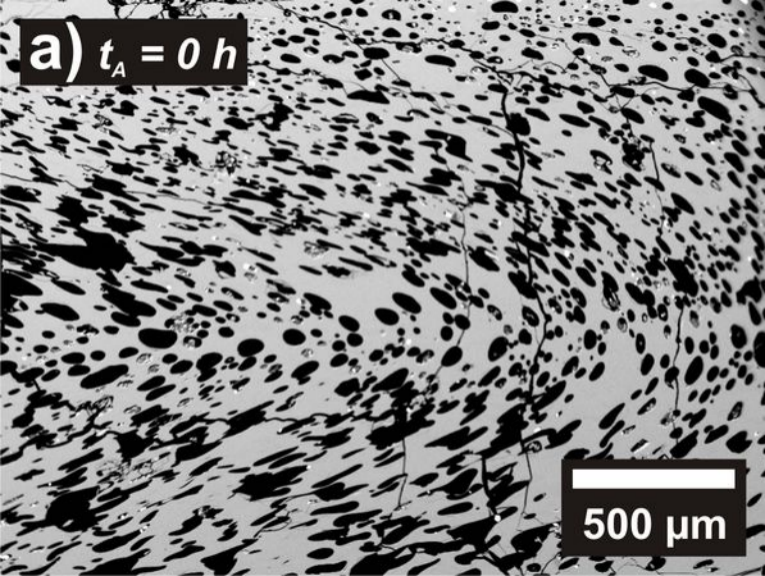
500 μm

A grayscale micrograph showing a porous, interconnected network of dark, irregularly shaped structures against a lighter background. The structures appear to be interconnected, forming a complex, porous lattice. The overall appearance is that of a porous material, possibly a foam or a porous polymer. The structures are distributed throughout the field of view, with some larger, more prominent features and many smaller, more numerous ones. The background is a light gray, providing contrast for the dark structures. The image is framed by a black border.

C) $r = 0.0005 \text{ MPa/s}$

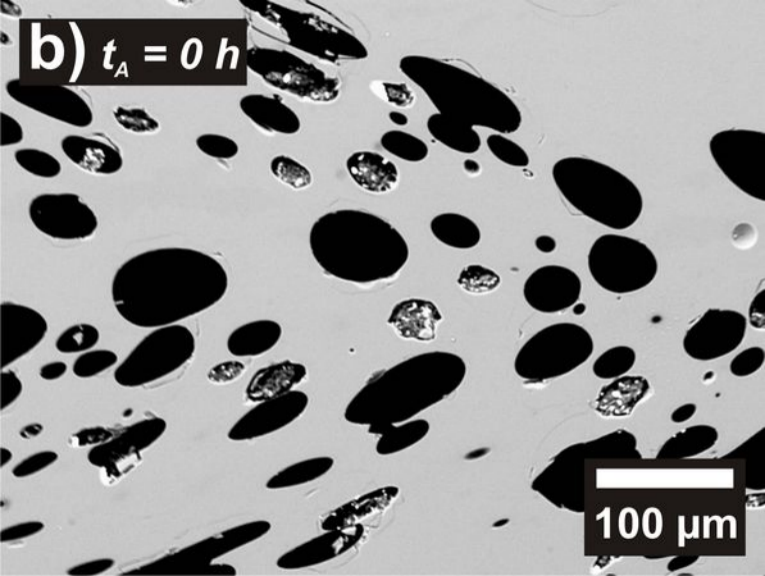


a) $t_A = 0 \text{ h}$



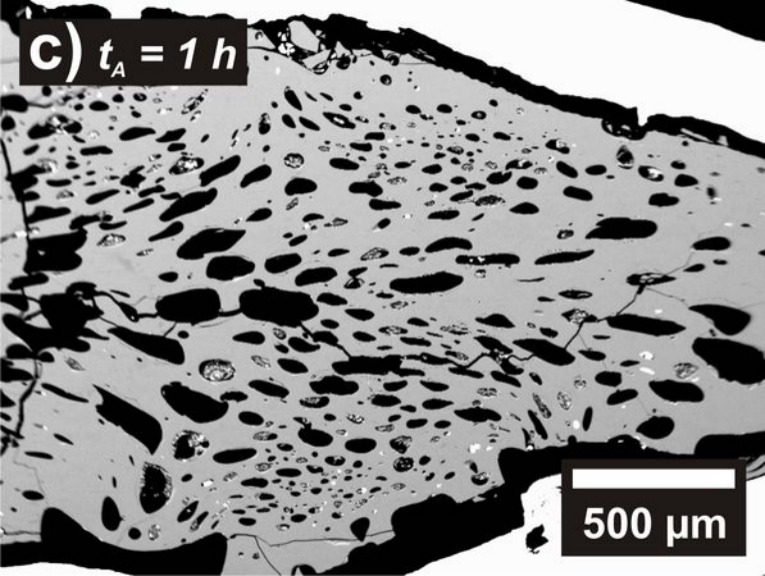
500 μm

b) $t_A = 0 \text{ h}$



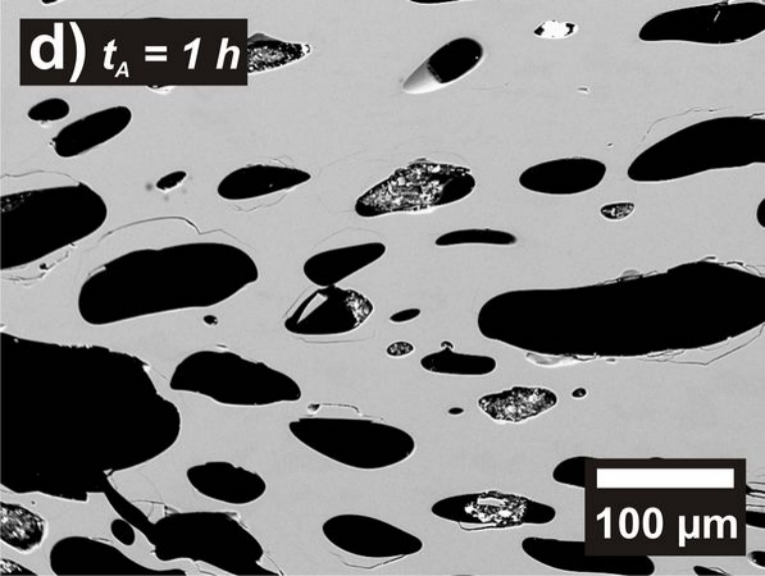
100 μm

C) $t_A = 1 h$



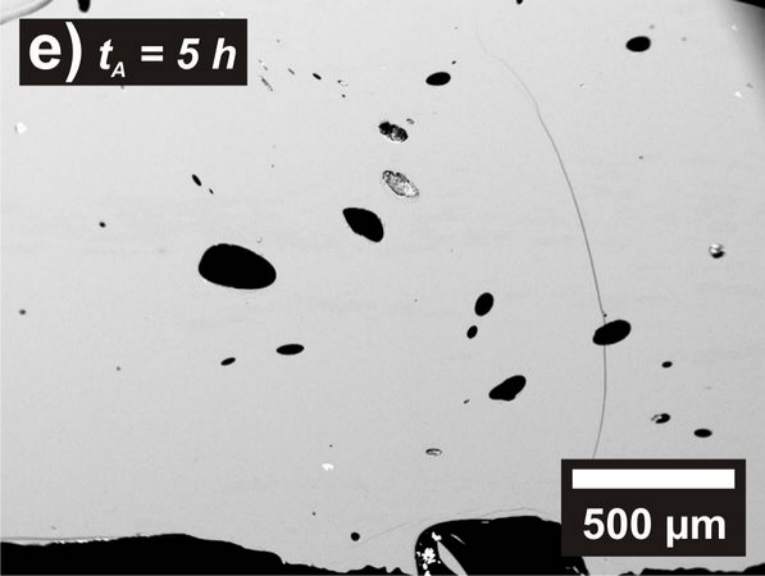
500 μm

d) $t_A = 1 h$



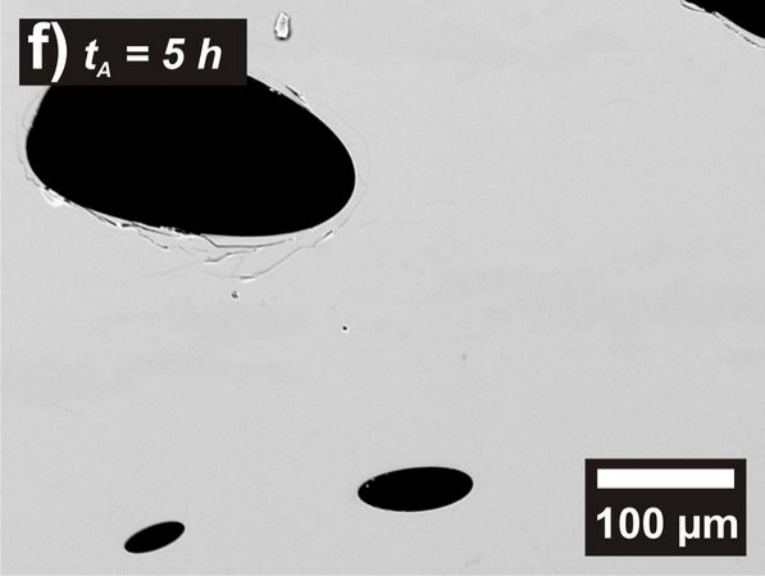
100 μm

e) $t_A = 5 h$

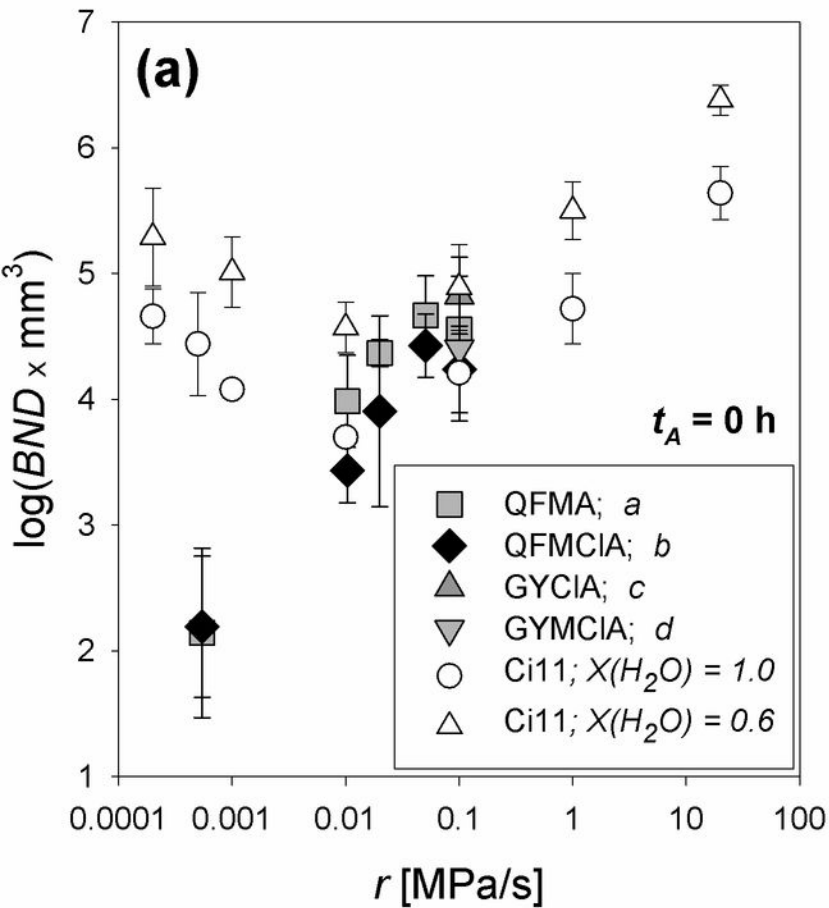


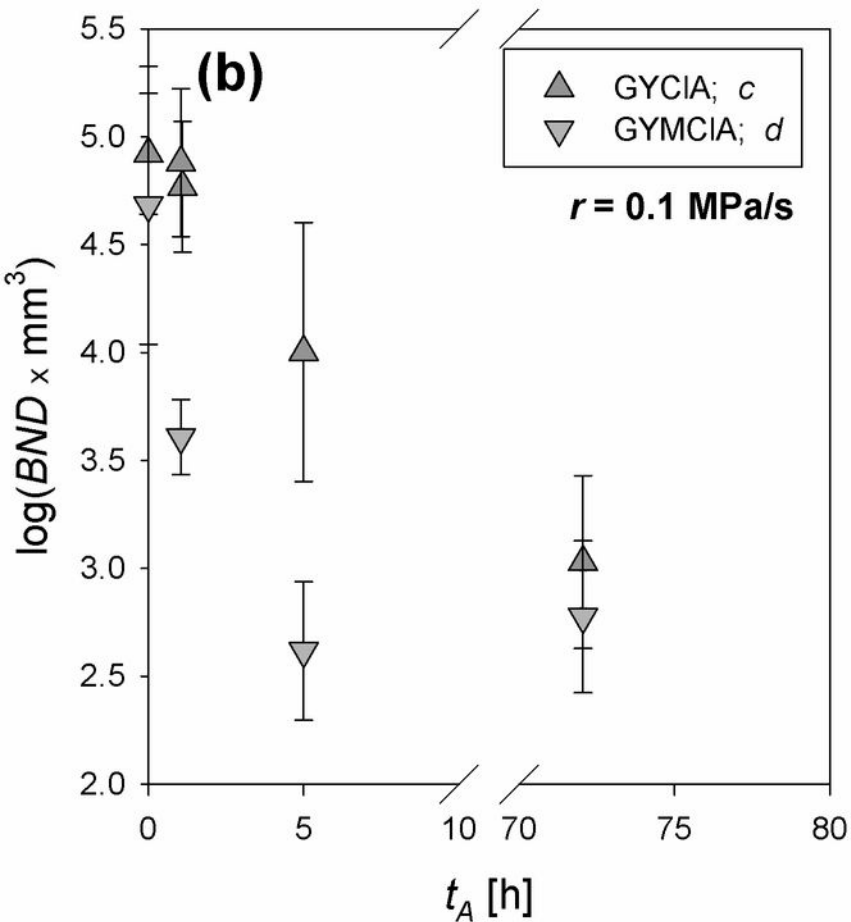
500 μm

f) $t_A = 5 h$



100 μm





- *this study; $t_A = 0$ h; homogen.*
- △ *Go11; homogen.*
- ▼ *Ha10; homogen.; saturated*
- *MS00; homogen.; saturated*
- *MBL04; homogen.; saturated*
- ◇ *Ci11; heterogen.; saturated*
- *To06; $\sigma_{LB} = 0.028$ andesite, 1030°C; a*
- *To06; $\sigma_{LB} = 0.063$ andesite, 1030°C; a*
- ⋯ *To06; $\sigma_{LB} = 0.015$ andesite, 1030°C; a*
- - - *To06; BND model for rhyolite, 800°C; b*

