

1 **Revision 1**

2 **Crystal structures of laihunite and intermediate phases between laihunite-1M**  
3 **and fayalite: Z-contrast imaging and *ab initio* study**

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## ABSTRACT

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Crystals of laihunite from Xiaolaihe of Liaoning Province, NE China, were studied using

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selected-area electron diffraction (SAED), high-resolution transmission electron microscopy

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(HRTEM), and Z-contrast imaging. Z-contrast images directly reveal ordered vacancies in M1

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sites. The results confirm early structural models for 1-layer laihunite (or laihunite-1M) with

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ideal stoichiometry of  $[\square]_{0.5}\text{Fe}^{2+}_{0.5}\text{Fe}^{3+}\text{SiO}_4$ . 2-layer laihunite and 3-layer laihunite are found to

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be chemically different from laihunite-1M. The 2-layer laihunite can be viewed as a periodic

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intergrowth of laihunite and fayalite in the 1:1 ratio. The 3-layer laihunite can be considered to

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be a periodic intergrowth of laihunite and fayalite in the 1:0.5 ratio along the *c*-axis. Ideal

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stoichiometries for the 2-layer structure and the 3-layer structure are  $[\square]_{0.5}\text{Fe}^{2+}_{2.5}\text{Fe}^{3+}[\text{SiO}_4]_2$ , and

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$[\square]_{1.0}\text{Fe}^{2+}_{3.0}\text{Fe}^{3+}_{2.0}[\text{SiO}_4]_3$ , respectively. The structural intergrowth of the 3-layer laihunite and

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the 1-layer laihunite results in chemical compositions falling within the range between the two

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aforementioned structures, such as the chemical formula of  $[\square]_{0.4}\text{Fe}^{2+}_{0.8}\text{Fe}^{3+}_{0.8}\text{SiO}_4$ , reported

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earlier in literatures.

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The crystal structures of the 1-layer laihunite (1M), the 2-layer laihunite (2M), and the 3-

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layer laihunite (3Or) determined from Z-contrast images and *ab initio* calculations using the

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density functional theory (DFT) have space groups of  $P2_1/b$ ,  $P2_1/b$ , and  $Pbnm$ , respectively. The

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previously reported monoclinic symmetry for the 3-layer laihunite may be an artifact due to

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overlapping diffraction spots from both, the laihunite-3Or and the laihunite-1M. Our study

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demonstrates that the method of combining Z-contrast imaging and *ab initio* calculation can be

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effectively used for identifying structures of nano-phases in host crystals.

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## INTRODUCTION

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Laihunite that has a distorted olivine-type structure with ferric and ferrous irons and

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ordered distribution of vacancies was first discovered in a high-grade metamorphosed banded

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iron formation (BIF) (Laihunite Research Group, 1976; Fu et al., 1982). The occurrence of

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laihunite is not limited to metamorphosed BIFs. Laihunite was also found in certain black

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fayalite (Schaefer, 1983; 1985), some ferric-fayalite from granite, and Fe-bearing olivine from

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volcanic and intrusive rocks (Sueno et al., 1985; Banfield et al., 1990, 1992; Janney and

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Banfield, 1998; Putnis, 1979; Xu et al., 1992; Konishi and Xu, 2012). The laihunite, which

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coexists with fayalite, magnetite, quartz, ferrosilite, almandine and hedenbergite, was formed in

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the process of oxidation of fayalite during metamorphism (Fu et al., 1982; Wang, 1982;

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Kitamura et al., 1984). The structure refinement of the 1-layer laihunite shows  $P2_1/b$  symmetry

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and ordered arrangement of vacancies in half of the  $M_1$  sites of a fayalite structure (Fu et al.,

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1982). A proposed ideal structural formula for the 1-layer laihunite is  $[\ ]_{0.5}Fe^{2+}_{0.5}Fe^{3+}SiO_4$  (Fu

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et al., 1982). However, structural details regarding symmetry and the number of vacancies in the

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structures are still a matter of controversy (Fu et al., 1982; Tamada et al., 1983; Ferrifayalite

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Research Group, 1976). To address these questions, it was proposed that the 1-layer structure is

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an average of the 3-layer laihunite that has a tripled periodicity along the  $c$ -axis with respect to

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fayalite structure (Shen et al., 1984; Tamada et al., 1983). A recent *ab initio* calculation result

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even suggested that 1-layer laihunite structure may have a triclinic symmetry (Chatterjee and

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Saha-Dasgupta, 2010).

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Transmission electron microscopy (TEM) and some X-ray studies have shown that there

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are also domains of fayalite, the 3-layer, and the 2-layer structures that are intergrown with

69 laihunite (Li et al., 1981; Fu et al., 1982; Kitamura et al., 1984; Sueno et al., 1985; Chou, 1985;  
70 Kondoh et al., 1985). Single-crystal structure refinement of a 3-layer laihunite was carried out  
71 based on  $P2_1/b$  symmetry (Shen et al., 1982; 1984). It was found that a 3-layer laihunite  
72 synthesized by oxidizing fayalite at high temperature has a higher Fe content than natural 2-layer  
73 and 3-layer laihunite structures do (Kondoh et al., 1985). At the same time, the synthetic 3-layer  
74 structure is different from the natural 3-layer laihunite in composition (Kondoh et al., 1985). All  
75 the reported TEM data suggest intergrowth of the laihunite structures at nanometer and sub-  
76 micrometer scales (1-layer, 2-layer, and 3-layer domains). Proposed models for the 3-layer  
77 laihunite superstructure are also controversial regarding the number of vacancies and the  
78 distributions of vacancies (Li et al., 1981; Shen et al., 1984; Chou, 1985). A 3-layer laihunite  
79 structure with monoclinic symmetry determined using X-ray single crystal diffraction method  
80 was proposed based on observation of overlapping diffraction peaks from multiple phases with  
81 the average composition of  $[\text{Fe}]_{0.4}\text{Fe}^{2+}_{0.8}\text{Fe}^{3+}_{0.8}\text{SiO}_4$  (Shen et al., 1984). In this paper, we use  
82 state-of-the-art methods of high-resolution Z-contrast imaging, in-situ X-ray EDS analyses, and  
83 *ab initio* calculation based on the density function theory (DFT) method to provide both  
84 structures and compositions of the laihunite (laihunite-1M) and intermediate phases between  
85 laihunite-1M and fayalite.

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## SAMPLE AND EXPERIMENT

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Samples for TEM and associate X-ray energy-dispersive spectroscopy (EDS)

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investigation of laihunite crystals were obtained from Xiaolaihe, Liaoning Province, NE China,

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where in fact the laihunite was first discovered. The host rock of the above minerals is Archean

92 metamorphosed BIF. Other coexisting minerals are fayalite, laihunite, magnetite, quartz,  
93 ferrosilite and hedenbergite (Laihunite Research Group, 1976; Ferrifayalite Research Group,  
94 1976).

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96 Our first TEM experiments were carried out with a Philips 420ST electron microscope  
97 equipped with an EDAX energy-dispersive X-ray spectrometer and a Princeton Gamma-Tech  
98 System-4000 analyzer as that described by Livi and Veblen (1987). Transmission electron  
99 microscopy and scanning transmission electron microscopy (STEM) studies were carried out  
100 using a FEI Titan 80-200 aberration corrected scanning/transmission electron microscope  
101 operated at 200 kV coupled with an EDAX high resolution EDS detector and Gatan image  
102 filtering system. This instrument is capable of imaging single atoms with  $\sim 0.1$  nm spatial  
103 resolution in STEM mode. Probe current was set at 24.5 pA. Collection angle of HAADF  
104 detector for acquiring all the Z-contrast images ranges from 54 to 270 mrad (corresponding to  
105  $7.5$  ( $1/\text{\AA}$ ) to  $38.2$  ( $1/\text{\AA}$ ) in reciprocal space).

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107 The scanning transmission electron microscopy (STEM) method uses the high-angle  
108 annular dark-field (HAADF) detector to give the most highly localized 1s Bloch state imaging,  
109 which eliminates most of the obvious effects of dynamical diffraction. Z-contrast images are  
110 HAADF images with atomic resolution (Kirkland, 1998). The intensity of Z-contrast images is  
111 dependent on numbers of atoms ( $n$ ) in an atomic column (or occupancy) and atomic number of  
112 atoms in the sites through  $\sim Z^2$  (Pennycook, 2002). Local composition and occupancy may be  
113 obtained from measured intensities. The relationship between intensity and atomic number for

114 the experimental condition based on calculated intensities for atoms columns in a Co-Mn-Si  
115 alloy is  $\sim Z^{2.27}$  (Shi, 2013).

116

117 Specimens for TEM/STEM investigation were prepared by crushing the selected crystals  
118 in alcohol and then depositing a drop of crystal suspension on holy-carbon Cu-grids. Because of  
119 fairly well developed {100} and {010} cleavages of fayalite and laihunite, it is not difficult to get  
120 crystal grains containing 00 $l$  diffraction in their diffraction patterns. Fayalite (Fe<sub>2</sub>SiO<sub>4</sub>) and  
121 forsterite (Mg<sub>2</sub>SiO<sub>4</sub>) were used as a standard for determination of the K-factor of Fe and Mg.  
122 All settings for the X-ray energy dispersion spectra (EDS) collections were the same as for the  
123 K-factor determination. Fe<sup>2+</sup> and Fe<sup>3+</sup> of the laihunite were calculated from the  
124 (Fe+Mg+Mn+Ca+Ni)/Si ratio deduced from the EDS results, by considering charge balance and  
125 assuming all tetrahedral sites are occupied by Si. X-ray EDS analysis from a small area with  
126 known structure can be obtained by combining selected area electron diffraction (SAED) and  
127 HRTEM images collected from the same area. This analysis is possible because SAED and  
128 HRTEM can provide structural information from very small areas, such as domains of 1-layer  
129 laihunite, 2-layer laihunite, and 3-layer laihunite.

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131 The DFT calculations were performed by using Vienna *Ab Initio* Simulation Package  
132 (VASP) (Kresse et al., 1996). The general gradient approximation (GGA) with the Perdew,  
133 Burke, and Ernzerhof (PBE) parameters was employed (Perdew et al., 1996). The projector-  
134 augmented wave (PAW) method with the energy cutoff of 600 eV was used. K-point meshes of  
135 4×4×2, 4×2×2, and 4×1×2 were found to be sufficient for 1-layer, 2- layer, and 3-layer laihunite  
136 structures, respectively. In order to take into account the on-site Coulomb repulsion of 3d

137 electron in Fe atoms, we employed a simplified (rotationally invariant) approach known as  
138 GGA+U (Dudarev et al., 1998). In Dudarev's method, an effective U parameter,  $U_{\text{eff}} = U - J$ , is  
139 used. According to previous DFT calculations of fayalite (Cococcioni et al., 2003; and  
140 Stackhouse et al., 2010),  $U_{\text{eff}}=4.8\text{eV}$  is a reasonable value and thus it was used in this study. All  
141 initial structures with P1 symmetry proposed in this study were relaxed using the static energy  
142 minimization scheme, where both the shape and volume of the cell were allowed to relax. The  
143 final structures with lowest energy states are chosen as possible structures for the 3 phases. The  
144 electron diffraction patterns of calculated structures were generated by SingleCrystal<sup>TM</sup> software.

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## 146 **RESULTS AND DISCUSSIONS**

147

### 148 **TEM results**

149 TEM images show that the crystals are heterogeneous in structure with nano-cracks  
150 between neighboring domains (Figure 1). Selected-area electron diffraction (SAED) patterns  
151 show that the crystal contains laihunite (or 1-layer laihunite), 3-layer laihunite domains, 2-layer  
152 laihunite domains, and fayalite domains (Figure 2). We found many areas with intergrown 3-  
153 layer and 1-layer laihunite structures (Figure 2D, 2E). The SAED pattern (E) shows diffractions  
154 from 1-layer laihunite and 3-layer laihunite. Positions of 001 and 003 are not from the 3-layer  
155 laihunite domain, but from 1-layer laihunite domain, because they are off the center between 00  
156  $2/3$  and 00  $4/3$ , and between 00  $8/3$  and 00  $10/3$  reflections (fractional indices are based on  
157 fayalite unit cell and setting). Periodicity along the *c*-axis for the 3-layer laihunite is  $1.5c$  of the  
158 1-layer laihunite. Diffraction spots of 001 and 003 from a 1-layer laihunite should not be  
159 considered as 003 and 009 diffractions of the 3-layer laihunite structure. Unit cell parameters can

160 be measured by comparing the diffraction difference and Fast Fourier transform (FFT) patterns  
161 between known structure (fayalite) and unknown laihunite phases with different periodicities  
162 along their *c*-axes (Table 1).

163

164 X-ray EDS spectra of laihunite with different structures (different periodicities along the  
165 *c*-axis) were collected from relatively ordered domains and confirmed by SAED patterns and  
166 HRTEM images. Using a general formula for laihunite  $[\ ]_x\text{Fe}^{2+}_{2-3x}\text{Fe}^{3+}_{2x}\text{SiO}_4$ , we found the  
167 number of vacancies for the 1-layer laihunite range from 0.47-0.52; the amount of vacancies for  
168 the 3-layer laihunite range from 0.30-0.39; and the amount of vacancies for the 2-layer laihunite  
169 range from 0.21-0.25. The observed laihunite domains with difference periodicities are  
170 chemically distinct (Table 2), and they are not in simple substructure and superstructure  
171 relationships.

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### 173 **Z-contrast imaging results**

174 Z-contrast image of a fayalite area shows the positions of Fe (bright spots) and Si (much  
175 less bright spots among 3 Fe atoms). A structure model of fayalite is also overlaid on the image.  
176 Positions of oxygen are omitted for clarity because scattering from oxygen atom is much weaker  
177 than those from Fe and Si atoms. The diffraction spots of *00l* (*l*=odd) appear in SAED pattern  
178 (Figure 2C) and FFT pattern (Figure 4C) from fayalite due to multiple diffraction of coherent  
179 scattered electrons at low-angle. Z-contrast images that use the high-angle annular dark-field  
180 (HAADF) detector to collect non-coherent scattered electrons at high angle eliminate most of the  
181 obvious effects of dynamical diffraction (Pennycook, 2002). FFT pattern from the Z-contrast



182 image of Figure 3 does not show 001, 003 spots due to 2-fold screw axis along the *c*-axis in  
183 fayalite structure (Figure 4D).

184

185 A [100]-zone-axis Z-contrast image at the interface between fayalite and laihunite shows  
186 Fe vacancies at M1 sites within some of the (001) planes. Periodic distribution of the vacancies  
187 results in locally ordered domains of 2-layer laihunite (labeled 2) and 3-layer laihunite (labeled  
188 1.5) (Fig. 5). A [110]-zone-axis Z-contrast image shows intergrowth of 1-layer laihunite, 2-layer  
189 laihunite, and 3-layer laihunite domains (Fig. 6). The Z-contrast images directly show positions  
190 of Fe vacancies in some of the M1 sites of fayalite-based structure. Structure models for 1-layer,  
191 2-layer and 3-layer laihunite can be obtained based on observed distribution of vacancies in  
192 fayalite-based structure. Figure 7 is a [100]-zone-axis Z-contrast image of 1-layer laihunite  
193 domain. Half of the M1 sites are vacancies in the 1-layer laihunite. Structure model for the 1-  
194 layer laihunite is also overlaid on the original image (Fig. 7A) and the noise-filtered image (Fig.  
195 7B). The structure of the 1-layer laihunite is found to have a monoclinic symmetry. Some areas  
196 show a twin relationship between neighboring 1-layer laihunite domains and interface between  
197 1-layer and 3-layer laihunite (Figure 8). Stoichiometry for the 1-layer laihunte domains is  
198  $[\text{0.5Fe}^{2+}\text{0.5Fe}^{3+}\text{SiO}_4]$  based on the number of vacancies in half of the M1 sites. Fe in M2 site  
199 neighboring the vacancies will be ferric Fe in order to maintain the local charge balance. This  
200 assertion is consistent with chemical compositions found by analyzing laihunite-1M domains  
201 (Table 2).

202

203 A [110]-zone-axis Z-contrast image from an area dominated by 3-layer laihunite shows  
204 that (001) vacancy layers occur every one and half fayalite units along the *c*-axis. Periodicity

205 along the *c*-axis is tripled with respect to the fayalite structure (Figure 9). FFT pattern from the  
206 image shows 002, and 004 spots, which indicates a 2-fold screw axis along the *c*-axis of the 3-  
207 layer laihunite structure. The 3-layer laihunite will keep orthorhombic symmetry (or, laihunite-  
208 3Or). Stoichiometry for the laihunite-3Or will be  $[ ]_{1.0} \text{Fe}^{2+}_{3.0} \text{Fe}^{3+}_{2.0} [\text{SiO}_4]$ . This is consistent  
209 with chemical compositions from analyzed laihunite-3Or domains (Table 2). The structural  
210 intergrowth of 3-layer laihunite and 1-layer laihunite will results in chemical compositions that  
211 lie between the ideal 1-layer laihunite and 3-layer laihunite, such as the reported chemical  
212 formula of  $[ ]_{0.4} \text{Fe}^{2+}_{0.8} \text{Fe}^{3+}_{0.8} \text{SiO}_4$  (Shen et al., 1984).

213

## 214 **DFT CALCULATIONS**

215 *Z*-contrast images provide structure models that are based on compositions and  
216 distributions of vacancies among the M1 sites of fayalite-based structure. In order to obtain  
217 detailed atomic positions for the proposed structure models, we carried out *ab initio* calculation  
218 using DFT. In these calculations, it is important to consider magnetic properties of Fe in a given  
219 structure. According to published data, there are two possible antiferromagnetic (AF)  
220 configurations of fayalite: AF interaction between edge-sharing octahedra and AF interaction  
221 between corner-sharing octahedra (Cococcioni et al., 2003). In DFT calculations of 1-layer  
222 laihunite (Chatterjee and Saha-Dasgupta, 2010), the latter configuration was found to have a  
223 lower energy. Our calculations confirm that the latter configuration is more stable for all three  
224 laihunite structures (see appendix). Therefore, the laihunite structures reported below were  
225 calculated using the second magnetic configuration.

226

227 The distributions of vacancies in the initial structure models were based on the STEM  
228 observation that all vacancies occupy the M1 sites of fayalite-based structure. The optimized  
229 configurations with lowest energy for 1-layer, 2-layer and 3-layer laihunite structures are shown  
230 in Figures 10 and 11. A monoclinic structure model (laihunite-3M) based on Shen et al., (1984)  
231 is also illustrated for comparison. Obviously, the “laihunite-3M” model does not fit observed Z-  
232 contrast images. The unit cell parameters calculated from DFT are reported in Table 1 and they  
233 show a good agreement with the experimentally measured values. Because of the vacancy  
234 positions, the symmetries of 1-layer and 2-layer laihunite structures are reduced to monoclinic  
235 (the alpha angle are 91.39° and 90.79°, respectively), whereas 3-layer laihunite structure retains  
236 the orthorhombic symmetry. The 1-layer and 2-layer laihunite structures have space group  $P2_1/b$   
237 (no. 14) (Table 3). The 3-layer laihunite structure resumes the same space group of  $Pbnm$  (no.  
238 62) as fayalite (Table 3). Here, we use laihunite-1M, laihunite-2M, and laihunite-3Or to  
239 represent the observed 1-layer, 2-layer, and 3-layer laihunite structures, although they are not  
240 polymorphs with the same composition.

241

242 In laihunite-1M, the M1-O<sub>6</sub> octahedra are larger than the M2-O<sub>6</sub> octahedra. In addition,  
243 the M2-O<sub>6</sub> octahedra are slightly more distorted (Table 4), which is consistent with conclusions  
244 from earlier calculations, although the earlier calculations suggest a triclinic symmetry for the 1-  
245 layer laihunite (Chatterjee and Saha-Dasgupta, 2010). Comparing all the structure models with  
246 fayalite structure, the positions for Si and Fe2 are shifted away from their ideal positions in  
247 fayalite (fractional coordinate  $z=0,25$  for Si and Fe2), because of the presence of Fe(III) in those  
248 Fe2 sites that neighbor the vacancies (Table 5). However, fractional coordinates for Si and Fe2 in  
249 all the reported structures are different (Table 5). Our calculated structure shows that both Fe2

250 and Si atoms are slightly shifted towards vacancy sites, instead of opposite shifts along the *c*-axis  
251 with respect the Si and Fe positions in fayalite structure.

252

253 For laihunite-2M we find that, while there are similar M1-O<sub>6</sub> and M2-O<sub>6</sub> (Fe<sup>3+</sup>-O<sub>6</sub>)  
254 octahedra adjacent to the vacancies as in laihunite-1M, the octahedra away from vacancies  
255 resemble those in fayalite. The laihunite-2M domains will also form twins, similarly to those  
256 observed in laihunite-1M.

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258

259 In laihunite-3Or, there is just one type of M1-O<sub>6</sub> octahedra due to its orthorhombic  
260 symmetry (Table 4). The calculated laihunite-3Or structure is different from a previously  
261 proposed structure with monoclinic symmetry (Shen et al., 1984). The laihunite crystals contain  
262 nanometer to sub-micrometer scales of domains of all 3 laihunite structures and even residual  
263 fayalite. The crystal analyzed by She et al. (1984) contains domains of laihunite-1M and  
264 laihunite-3Or. X-ray diffraction peaks from the crystal will overlap with peaks from laihunite-  
265 1M and laihunite-3Or. If odd *00l* reflections from laihunite-1M are considered as 003 and 009  
266 reflections (based on 3-layer laihunite setting), the obtained structure will have a monoclinic  
267 symmetry (*P2<sub>1</sub>/b*) instead of an orthorhombic symmetry (*Pbnn*). Structural intergrowth of 3-  
268 layer laihunite and 1-layer laihunite will result in average chemical compositions between the  
269 laihunite-1M and laihunite-3Or. The reported formula of  $[\ ]_{0.4}\text{Fe}^{2+}_{0.8}\text{Fe}^{3+}_{0.8}\text{SiO}_4$  (Shen et al.,  
270 1984) is an average composition of laihunite-1M and laihunite-3Or.

271

272 Chou (1985) proposed 2 types of 3-layer laihunite structures based on HRTEM images  
273 from [110]-zone-axis (for  $3C_1$  structure odd and even  $00l$  reflections) and [010]-zone axis (for  
274  $3C_2$  structure with even  $00l$  reflections). According to our Z-contrast images and calculated  
275 structure ( $Pbnm$ ) for the laihunite-3Or, odd  $00l$  reflections and tripled periodicity along the  $c$ -axis  
276 in [110]-zone-axis SAED pattern and high-resolution transmission electron microscopic  
277 (HRTEM) images may be artifacts from multiple diffractions. In [010]-zone-axis SAED pattern  
278 and HRTEM image, there will be no odd  $00l$  reflections due to  $n$ -glide plane perpendicular to  
279 (010). There is only one type of 3-layer structure, i.e., laihunite-3Or.

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281 The observed 2-layer and 3-layer structures are not simple superstructures of laihunite-  
282 1M, because their compositions are different from that of laihunite-1M. The originally proposed  
283 laihunite structure has 1-layer periodicity with ideal stoichiometry of  $[ ]_{0.5}Fe^{2+}_{0.5}Fe^{3+}SiO_4$   
284 (Laihunite Research Group, 1976; Fu et al., 1982). The observed intermediate phases are very  
285 similar to those in mixed-layer clay minerals, such as interstratified chlorite/serpentine minerals  
286 (Banfield and Bailey, 1996; Xu and Veblen, 1996). The observed intermediate structures in the  
287 laihunite – fayalite system and forsterite – Mg-laihunite system may be also described as  
288 interstratified olivine / laihunite minerals.

289

290 We propose that the observed laihunite structures are stable only within the fayalite host  
291 and we propose the following scenario for formation of these structures. Generally, the laihunite  
292 domains are formed during oxidative reaction of the fayalite in the solid state. Domains formed  
293 in early stage contain fewer vacancies and they serve as host phases to stabilize domains formed



317

## 318 IMPLICATIONS

319           It is very challenging to solve crystal structures of nano-precipitates in host minerals  
320 using X-ray diffraction and TEM. Our results provide a new way for solving the crystal  
321 structures of nano-crystals in or intergrown with host minerals, and locally ordered domains in  
322 interstratified clay minerals. Z-contrast imaging can provide information about topology of the  
323 crystal structure, and DFT calculation provides accurate coordinates of the atoms within a unit  
324 cell. The observed vacancy ordering in Fe-rich olivine may provide information about highly  
325 anisotropic diffusion of atoms in Fe<sup>3+</sup>-bearing olivine crystals. Vacancy layers in Fe<sup>3+</sup>-bearing  
326 olivine may behave as a “highway” for atom diffusion. Z-contrast imaging is a powerful tool for  
327 direct observation of vacancies and other defects. The method can be used for studying oxidation  
328 of iron and vacancy ordering in other minerals, such as Fe-bearing olivine minerals, and chain  
329 silicate minerals, and Fe-bearing oxide minerals. The combination of complementary methods  
330 described in this paper can be used for observing and quantifying local ordering of Fe atoms and  
331 vacancies in partially oxidized Mg-rich olivine crystals reported by Janney and Banfield (1998)  
332 and Konishi and Xu (2012).

333

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422  
423

424 **Figure captions**

425

426 Figure 1: (A) TEM image showing a structural heterogeneity of a laihunite crystal with nano-  
427 cracks or nano-tunnels (indicated by arrows) between neighboring domains. Nano-cracks formed  
428 during the transformation from fayalite to laihunite due to the difference in the  $c$ -dimensions  
429 between fayalite and laihunite. Inserted FFT pattern shows streaking along the  $c^*$  direction and  
430 intensity maxima at  $00\ \frac{1}{2}$  and  $00\ \frac{2}{3}$  positions (indicated by 2 short arrows). (B) An HRTEM  
431 image showing intergrowth of 2-layer and 3-layer laihunite domains and nano-tunnels.

432

433

434 Figure 2: SAED patterns from 3-layer laihunite (A), intergrown 2-layer and 3-layer laihunite (B),  
435 intergrown fayalite and 2-layer laihunite (C),  $[010]$ -zone axis of intergrown 1-layer and 2-layer  
436 laihunite (D), and tilted crystal with only  $00l$  reflections to avoid multiple diffraction from  $hkl$   
437 spots (E).

438 All the indices are based on fayalite unit cell setting. Arrows in (A) indicate extra spots from  
439 tripled periodicity along the  $c$ -axis. Long arrows and short arrows in (B) indicate diffraction  
440 spots from 3-layer laihunite and 2-layer laihunite, respectively. SAED (C) shows splitting in high  
441 order  $00l$  reflections ( $005$ ,  $006$ ) from fayalite and the 2-layer laihunite domain. All the indices  
442 are based on fayalite setting.

443

444 Figure 3: Z-contrast image of fayalite ( $[100]$ -zone axis) shows positions of Fe in M1 and M2  
445 sites. Fayalite model is also overlaid on the image. Si atoms at the center among 3 Fe atoms (2  
446 Fe at M1 site and one Fe in M2 site). Oxygen atoms are omitted because they are not visible in

447 Z-contrast images due to the weak intensity with respect to Fe and Si. The Fe atoms at M1, M2  
448 sites, and Si atoms are colored in yellow, orange, and blue, respectively.

449

450

451 Figure 4: a [100]-zone-axis HRTEM image showing an interface between fayalite (Fa) and 2-  
452 layer laihunite (2M). FFT patterns from the 2M domain (B) and fayalite (C) areas show 001 and  
453 003 spots that violate 2-fold screw axis due to the multiple diffraction effect in HRTEM imaging.  
454 The two arrows indicate spots from the 2-layer laihunite domain. It should be noted that the FFT  
455 pattern from the Z-contrast image of Figure 3 does not show 001 and 003 spots because Z-  
456 contrast image uses non-coherent scattered electrons at high angle and eliminates multiple  
457 diffraction effect (D).

458

459 Figure 5: A [100]-zone-axis Z-contrast image showing interface between fayalite and laihunite  
460 domains with doubled (2) and tripled (1.5) periodicities along the *c*-axis.

461

462 Figure 6: A [110]-zone-axis Z-contrast image showing intergrowth of 1-layer laihunite (1M), 2-  
463 layer laihunite (2M), and 3-layer laihunite (3Or). Inserted in the upper left and lower right  
464 comers are FFT patterns from 3-layer laihunite and 1-layer laihunite domains, respectively. An  
465 outline box area shows an interface between the 1-layer laihunite (1M) and the 3-layer laihunite  
466 (3Or).

467

468 Figure 7: Z-contrast images of 1-layer laihunite ([100]-zone-axis) showing positions of high  
469 density spots (corresponding to Fe, Si) and low-density sites of vacancies. Noise-filtered image

470 (lower one) better shows the positions of the atoms and vacancy sites. A unit cell of 1-layer  
471 laihunite model with the  $[\text{0.5Fe}^{2+}\text{0.5Fe}^{3+}\text{SiO}_4]$  stoichiometry is also overlaid on the image. The  
472 Fe atoms at M1, M2 sites, and Si atoms are colored in yellow, orange, and blue, respectively.

473

474 Figure 8: A [110]-zone-axis Z-contrast image of 1-layer laihunite (a raw image and a noise-  
475 filtered image) showing twinning relationship between the two neighboring 1-layer laihunite  
476 domains. M2 sites are brighter than M1 sites due to overlapping of Si and  $\text{Fe}^{3+}$  in the M2 sites  
477 along the [110] direction. It is also possible that a small number of Mg atoms migrated into M1  
478 sites. The twin boundary (TB) is a unit cell of fayalite that was proposed by Fu et al (1982). A  
479 unit cell of 1-layer laihunite model with the  $[\text{0.5Fe}^{2+}\text{0.5Fe}^{3+}\text{SiO}_4]$  stoichiometry is also overlaid  
480 on the image. An outlined area in the upper left corner shows an interface between 1-layer  
481 laihunite and 3-layer laihunite. Models for the interface (upper) and twin relationship (lower) are  
482 shown on the right hand side of the images. The Fe atoms at M1, M2 sites, and Si atoms are  
483 colored in yellow, orange, and blue, respectively.

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486 Figure 9: [110]-zone-axis Z-contrast image of an area dominated by 3-layer laihunite with  
487 domains of 1-layer laihunite (at top) and 2-layer laihunte. FFT pattern from the 3-layer laihunite  
488 shows even  $00l$  spots ( $l = 2n$ , indexing is based on 3-layer laihunite setting). There are two  
489 vacancy layers per unit cell along (001). Ideal stoichiometry for the 3-layer laihunite is  
490  $[\text{1.0Fe}^{2+}\text{3.0Fe}^{3+}\text{2}[\text{SiO}_4]\text{3}]$ , that is two 1-layer laihunite units+ one fayalite unit. Both the image  
491 and FFT pattern indicate that there is a 2-fold screw axis parallels to the  $c$ -axis, and two mirror

492 planes parallel to (001) within one unit cell at the position of the dashed lines. A projection of 3-  
493 layer laihunite structure (Fe and Si atoms only) is also overlaid on the image.

494

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496 Figure 10: Polyhedral models along the *a*-axis projection for laihunite-1M (A), laihunite-2M (B),  
497 laihunite-3Or (C), and a previously proposed “laihunite-3M” (D) structure for the 3-layer  
498 laihunite. The “laihunite-3M” (D) is not a correct model, because it does not fit the observed Z-  
499 contrast images. The polyhedral for Fe atoms at M1, M2 sites are colored in yellow and orange,  
500 respectively.

501

502 Figure 11. Projections of Fe and Si atoms of laihunite-1M, laihunite-2M, laihunite-3Or, and a  
503 previously proposed “laihunite-3M.” Panels (A) and (B) illustrate projections of the proposed  
504 structures along *a*-axis and [110]-zone axis, respectively. The Fe atoms at M1, M2 sites, Si are  
505 colored in yellow, orange, and blue, respectively. The O atoms are omitted due to the weak  
506 intensity. Fe(III) in M2B sites neighboring vacancy sites are slightly smaller than Fe(II) in Fe2A  
507 sites in order to show the difference between them.

508

509 Figure 12: (A) A low magnification Z-contrast image showing nano-tunnels or nano-cracks  
510 (indicated by arrows) in the laihunite. (B) A HRTEM image showing a magnetite nano-  
511 precipitate in a nano- tunnel.

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518 Table 1. Unit cell parameters of 1-layer, 2-layer and 3-layer laihunite structures. The  
519 experimental values are measured from electron diffraction and FFT patterns. Values for fayalite  
520 are from Fujino et al (1981). Uncertainty for the measured unit cell parameters is about  $\pm 0.02$  Å.  
521

	Experiment	DFT calculations
Laihunite-1M	4.80 Å, 10.20 Å, 1x5.80 Å	4.808 Å, 10.230Å, 1x5.808 Å
Laihunite-3Or	4.81 Å, 10.25 Å, 3x5.85 Å	4.824 Å, 10.291 Å, 3x5.882 Å
Laihunite-2M	4.82 Å, 10.30 Å, 2x5.93 Å	4.826 Å, 10.319 Å, 2x5.931 Å
Fayalite	4.8195 Å, 10.4788 Å, 6.0873 Å	

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527 Table 2 Structural formulae for fayalite and laihunite domains

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Structural formula

Analyzed domain

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532  $(\square_{0.00}\text{Fe}^{2+}_{1.93}\text{Mg}_{0.05}\text{Mn}_{0.01}\text{Ca}_{0.00}\text{Ni}_{0.01})\text{SiO}_4$  fayalite

533  $(\square_{0.21}\text{Fe}^{2+}_{1.28}\text{Fe}^{3+}_{0.42}\text{Mg}_{0.07}\text{Mn}_{0.00}\text{Ca}_{0.01}\text{Ni}_{0.01})\text{SiO}_4$  2-layer laihunite

534  $(\square_{0.21}\text{Fe}^{2+}_{1.28}\text{Fe}^{3+}_{0.43}\text{Mg}_{0.06}\text{Mn}_{0.01}\text{Ca}_{0.00}\text{Ni}_{0.01})\text{SiO}_4$  2-layer laihunite

535  $(\square_{0.22}\text{Fe}^{2+}_{1.28}\text{Fe}^{3+}_{0.44}\text{Mg}_{0.05}\text{Mn}_{0.00}\text{Ca}_{0.01}\text{Ni}_{0.00})\text{SiO}_4$  2-layer laihunite

536  $(\square_{0.225}\text{Fe}^{2+}_{1.11}\text{Fe}^{3+}_{0.45}\text{Mg}_{0.06}\text{Mn}_{0.01}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  2-layer laihunite

537  $(\square_{0.25}\text{Fe}^{2+}_{1.15}\text{Fe}^{3+}_{0.50}\text{Mg}_{0.09}\text{Mn}_{0.00}\text{Ca}_{0.0}\text{Ni}_{0.01})\text{SiO}_4$  2-layer laihunite

538  $(\square_{0.25}\text{Fe}^{2+}_{1.16}\text{Fe}^{3+}_{0.49}\text{Mg}_{0.09}\text{Mn}_{0.01}\text{Ca}_{0.00}\text{Ni}_{0.01})\text{SiO}_4$  2-layer laihunite

539  $(\square_{0.30}\text{Fe}^{2+}_{1.00}\text{Fe}^{3+}_{0.60}\text{Mg}_{0.08}\text{Mn}_{0.01}\text{Ca}_{0.01}\text{Ni}_{0.00})\text{SiO}_4$  3-layer laihunite

540  $(\square_{0.33}\text{Fe}^{2+}_{0.91}\text{Fe}^{3+}_{0.66}\text{Mg}_{0.08}\text{Mn}_{0.01}\text{Ca}_{0.01}\text{Ni}_{0.00})\text{SiO}_4$  3-layer laihunite

541  $(\square_{0.34}\text{Fe}^{2+}_{0.93}\text{Fe}^{3+}_{0.67}\text{Mg}_{0.06}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  3-layer laihunite

542  $(\square_{0.35}\text{Fe}^{2+}_{0.84}\text{Fe}^{3+}_{0.70}\text{Mg}_{0.10}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.01})\text{SiO}_4$  3-layer laihunite

543  $(\square_{0.35}\text{Fe}^{2+}_{0.88}\text{Fe}^{3+}_{0.69}\text{Mg}_{0.06}\text{Mn}_{0.01}\text{Ca}_{0.00}\text{Ni}_{0.01})\text{SiO}_4$  3-layer laihunite

544  $(\square_{0.37}\text{Fe}^{2+}_{0.82}\text{Fe}^{3+}_{0.74}\text{Mg}_{0.07}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  3-layer laihunite

545  $(\square_{0.39}\text{Fe}^{2+}_{0.75}\text{Fe}^{3+}_{0.77}\text{Mg}_{0.07}\text{Mn}_{0.01}\text{Ca}_{0.01}\text{Ni}_{0.01})\text{SiO}_4$  3-layer laihunite

546  $(\square_{0.47}\text{Fe}^{2+}_{0.50}\text{Fe}^{3+}_{0.95}\text{Mg}_{0.08}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihunite

547  $(\square_{0.48}\text{Fe}^{2+}_{0.48}\text{Fe}^{3+}_{0.96}\text{Mg}_{0.08}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihunite

548  $(\square_{0.49}\text{Fe}^{2+}_{0.49}\text{Fe}^{3+}_{0.98}\text{Mg}_{0.04}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihunite

549  $(\square_{0.51}\text{Fe}^{2+}_{0.38}\text{Fe}^{3+}_{1.02}\text{Mg}_{0.09}\text{Mn}_{0.00}\text{Ca}_{0.0}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihunite

550  $(\square_{0.52}\text{Fe}^{2+}_{0.36}\text{Fe}^{3+}_{1.04}\text{Mg}_{0.07}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihunite

551  $(\square_{0.52}\text{Fe}^{2+}_{0.36}\text{Fe}^{3+}_{1.04}\text{Mg}_{0.08}\text{Mn}_{0.00}\text{Ca}_{0.00}\text{Ni}_{0.00})\text{SiO}_4$  1-layer laihuite

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555 Table 3. Fractional coordinates of atoms and vacancies (Vac) in laihunite-1M, laihunite-2M, and  
 556 laihunite-3Or structures based on fayalite setting.

<b>Laihunite-1M</b>							
Space group $P2_1/b$				$\alpha = 91.39^\circ, \beta=90^\circ, \gamma=90^\circ$			
<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>	<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>
Fe1	0.00000	0.00000	0.00000	Fe2	0.50625	0.22833	0.25940
Si1	0.56010	0.09479	0.74789	Vac	0.00000	0.00000	0.50000
O1	0.22426	0.10578	0.76075	O2	0.32784	0.05564	0.24005
O3	0.70433	0.17868	0.53945	O4	0.72661	0.17100	0.96877

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<b>Laihunite-2M</b>							
Space group $P2_1/b$				$\alpha = 90.79^\circ, \beta=90^\circ, \gamma=90^\circ$			
<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>	<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>
Fe1A	0.00000	0.00000	0.50000	Fe1A	0.50331	0.49330	0.23409
Fe2A	0.01657	0.72036	0.37275	Fe2B	0.50465	0.77065	0.11864
Vac	0.00000	0.00000	0.00000				
Si1	0.57023	0.90246	0.37007	Si2	0.06434	0.59508	0.11892
O1	0.23210	0.90769	0.37049	O2	0.76982	0.10599	0.12715
O3	0.17407	0.44476	0.12487	O4	0.78700	0.54831	0.37052
O5	0.29257	0.17861	0.01647	O6	0.71221	0.83176	0.47971
O7	0.23020	0.66883	0.22619	O8	0.78605	0.32914	0.25917

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<b>Laihunite-3Or</b>							
Space group <i>Pbnm</i>				$\alpha = 90^\circ, \beta=90^\circ, \gamma=90^\circ$			
<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>	<b>Atom</b>	<b>x</b>	<b>y</b>	<b>z</b>
Fe1A	0.50432	0.22994	0.08029	Fe2A	0.01683	0.28260	0.25000
Fe2B	0.99655	0.00782	0.15910	Vac	0.00000	0.00000	0.00000
Si1	0.56779	0.10031	0.25000	Si2	0.43560	0.90602	0.08000
O1	0.22939	0.09520	0.25000	O2	0.73010	0.39511	0.08605
O3	0.17528	0.55647	0.08450	O4	0.20721	0.32165	0.01079
O5	0.22973	0.33279	0.15182	O6	0.71193	0.17355	0.17585
O7	0.28912	0.04515	0.75000				

562

563

564 Table 4. Structural data of Fe-O octahedra in fayalite, laihunite-1M, laihunite-2M, and laihunite-  
 565 3Or.

	Fayalite	1M	2M	3Or		
M1-O (Å)	2.16	2.15	2.16	2.16	2.16	
Volume of octahedra (Å <sup>3</sup> )	12.79	12.38	12.59	12.79	12.58	
Bond length variation	±0.02	±0.02	±0.03	±0.02	±0.02	
M2-O (Å)	2.17(Fe <sup>2+</sup> )	2.04 (Fe <sup>3+</sup> )	2.04 (Fe <sup>3+</sup> )	2.16 (Fe <sup>2+</sup> )	2.04 (Fe <sup>3+</sup> )	2.16 (Fe <sup>2+</sup> )
Volume of octahedra (Å <sup>3</sup> )	13.02	10.90	10.89	12.91	10.89	12.90
Bond length variation	±0.04	±0.03	±0.02	±0.04	±0.03	±0.04

566 Structure of fayalite is from Fujino et al. (1981).

567

568

569 Table 5: Difference between Si and Fe in their z coordinates among proposed models of 1-layer laihunite  
 570 and fayalite.

Authors	Si	Fe2	difference
Fayalite by Fujino et al.,1981	0.25	0.25	0
Ferrifayalite group, 1976	0.2220	0.2730	0.0510
Fu et al., 1982	0.2430	0.2740	0.0310
Tamada et al., 1983	0.2491	0.2579	0.0088
Chatterjee and Saha-Dasgupta, 2010	0.2501	0.2596	0.0095
This study	0.2521	0.2594	0.0073

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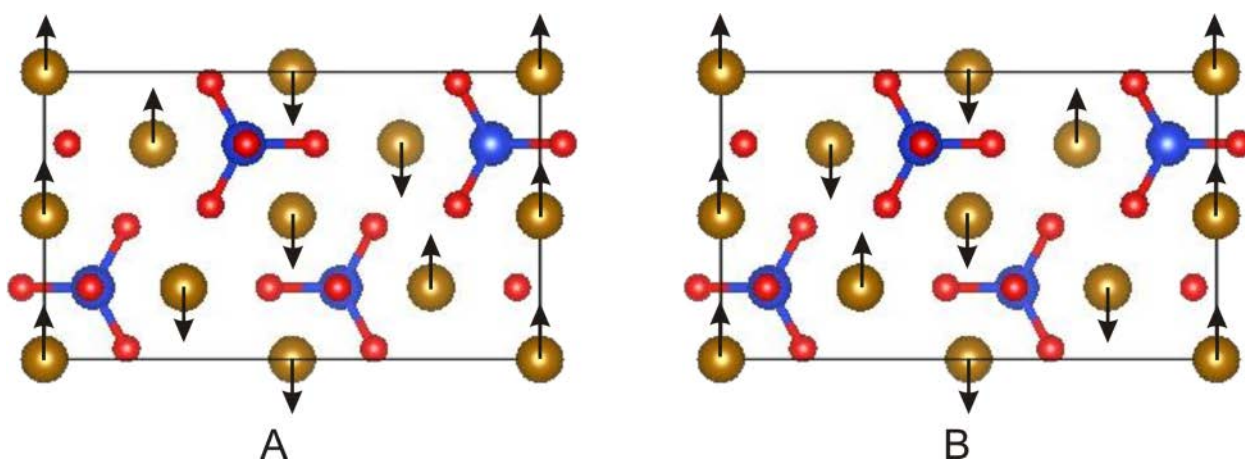
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575 **Appendix**

576 Figure A1. Two possible spin configurations in fayalite, modified from Chatterjee and Saha-

577 Dasgupta (2010).



580 Table A1. Total energy of laihunite-1M, laihunite-2M, and laihunite-3Or in the two possible

581 antiferromagnetic configurations together with fayalite.

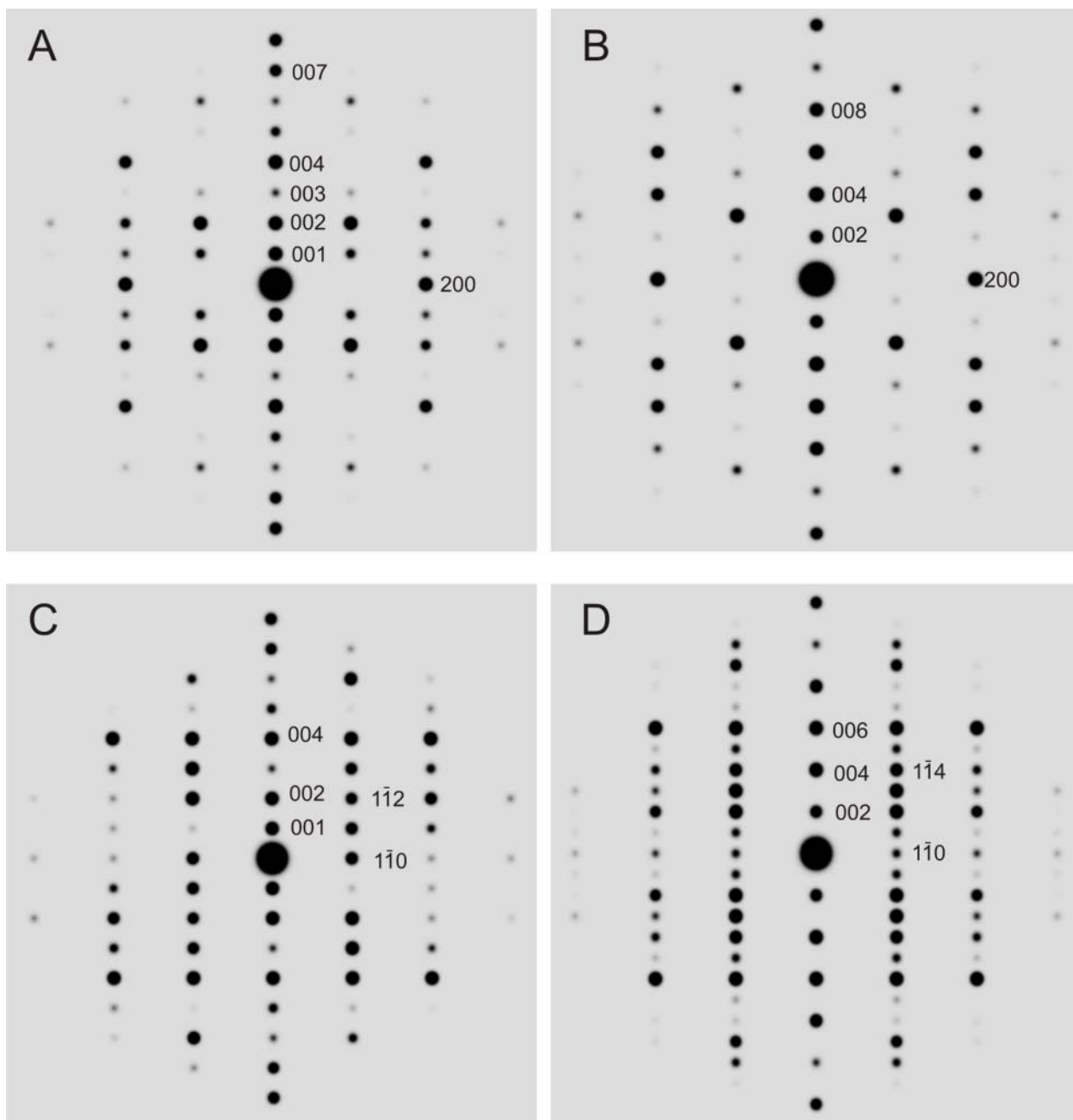
Structure	Configuration A (eV)	Configuration B (eV)
Laihunite-1M	-189.79	-189.86
Laihunite-2M	-395.08	-395.21
Laihunite-3Or	-585.00	-585.09
Fayalite	-205.11	-205.17

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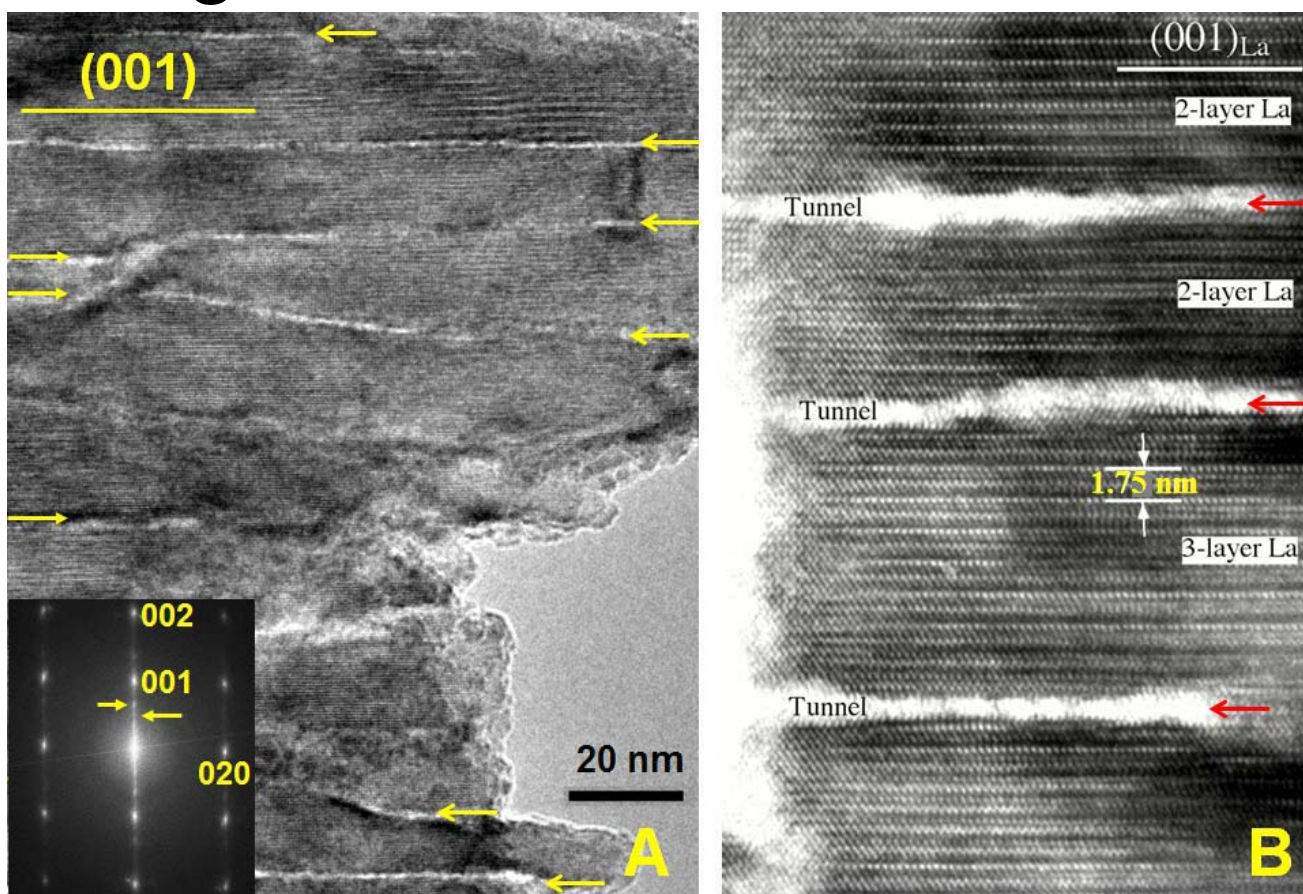
584 Figure A2. Simulated electron diffraction patterns of laihunite-2M (A, C) and laihunite-3Or (B,  
585 D). A and B are [001]-zone-axis patterns. C and D are [110]-zone-axis patterns.

586



587

# Fig. 1





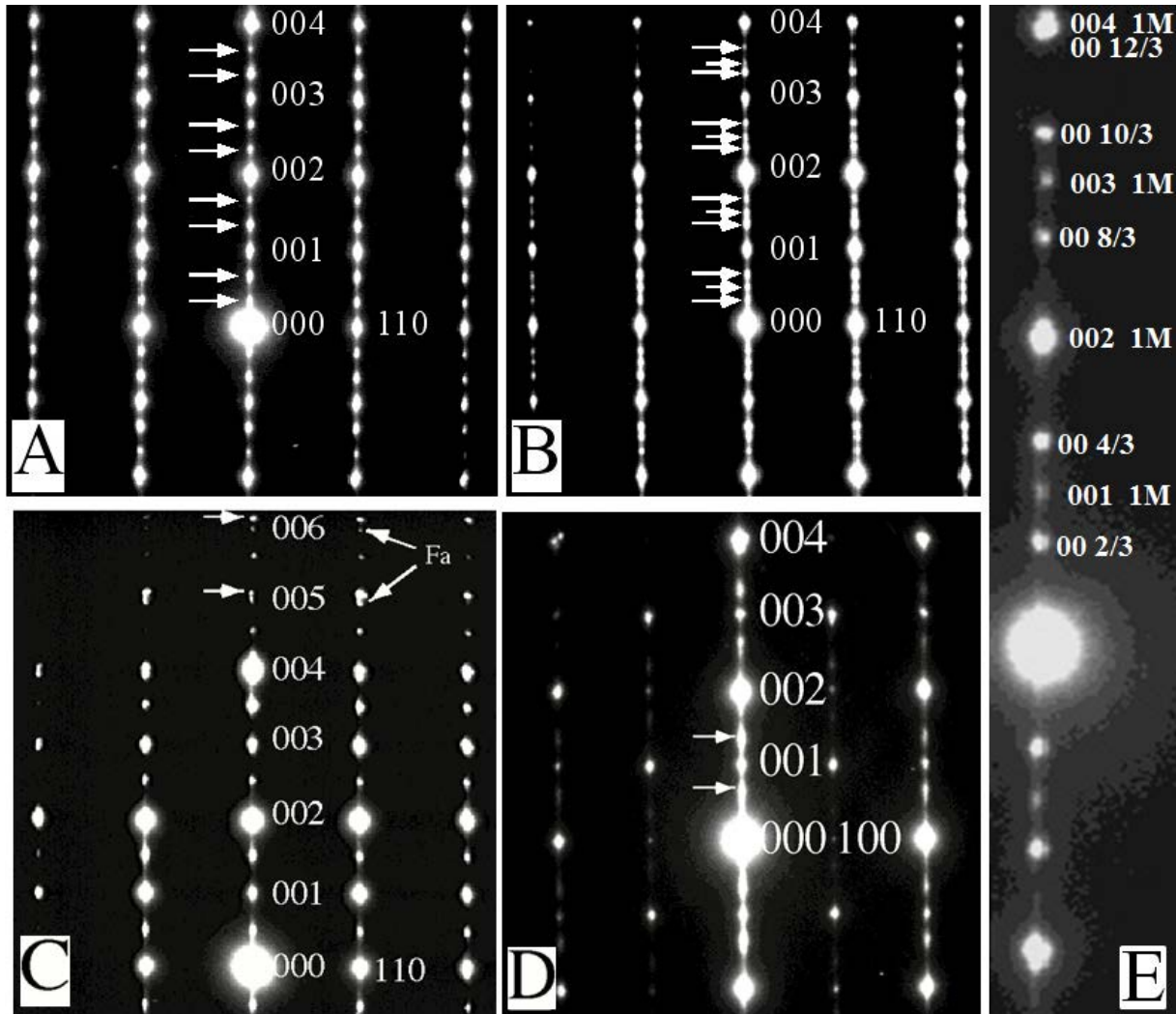
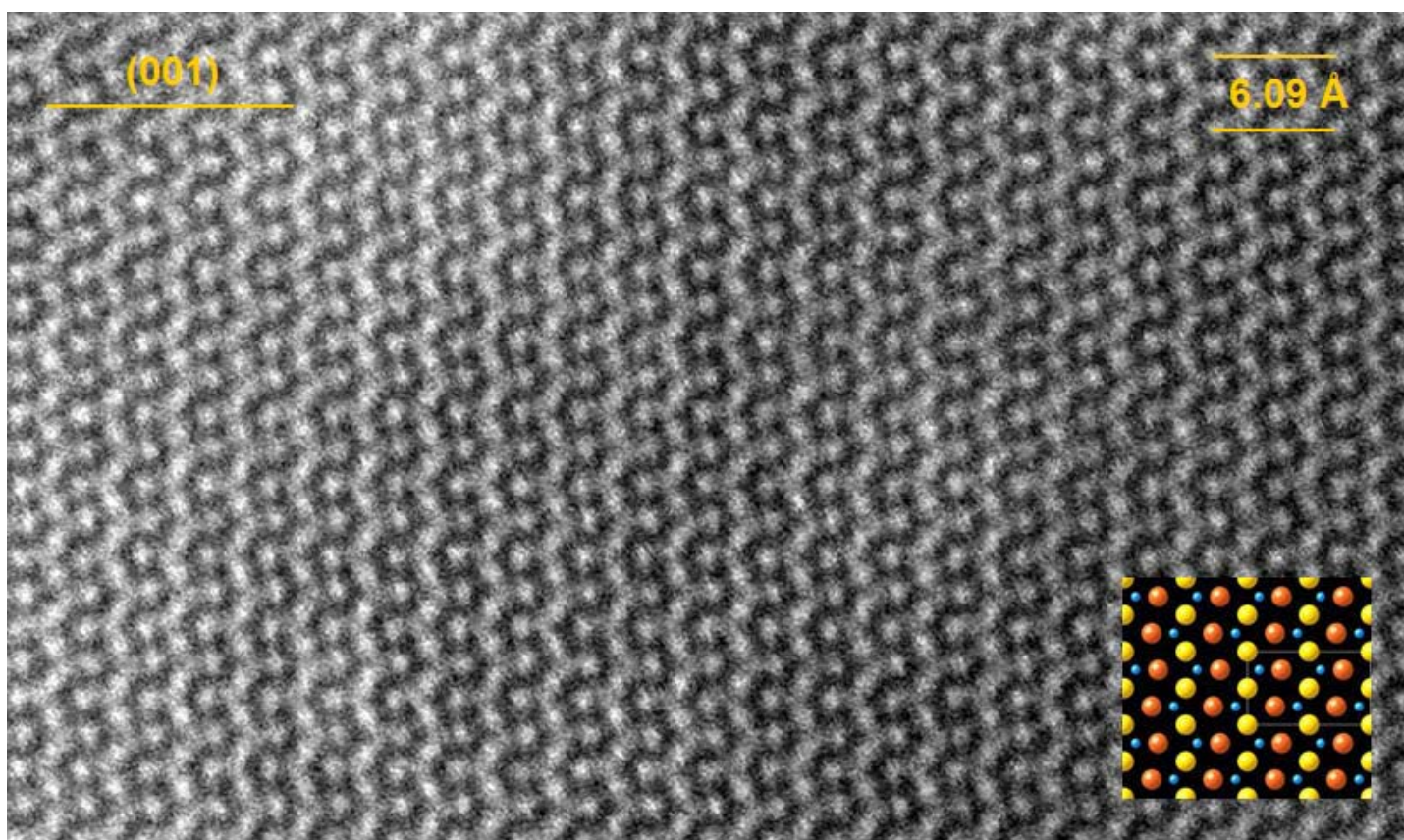


Fig. 2

# Fig. 3



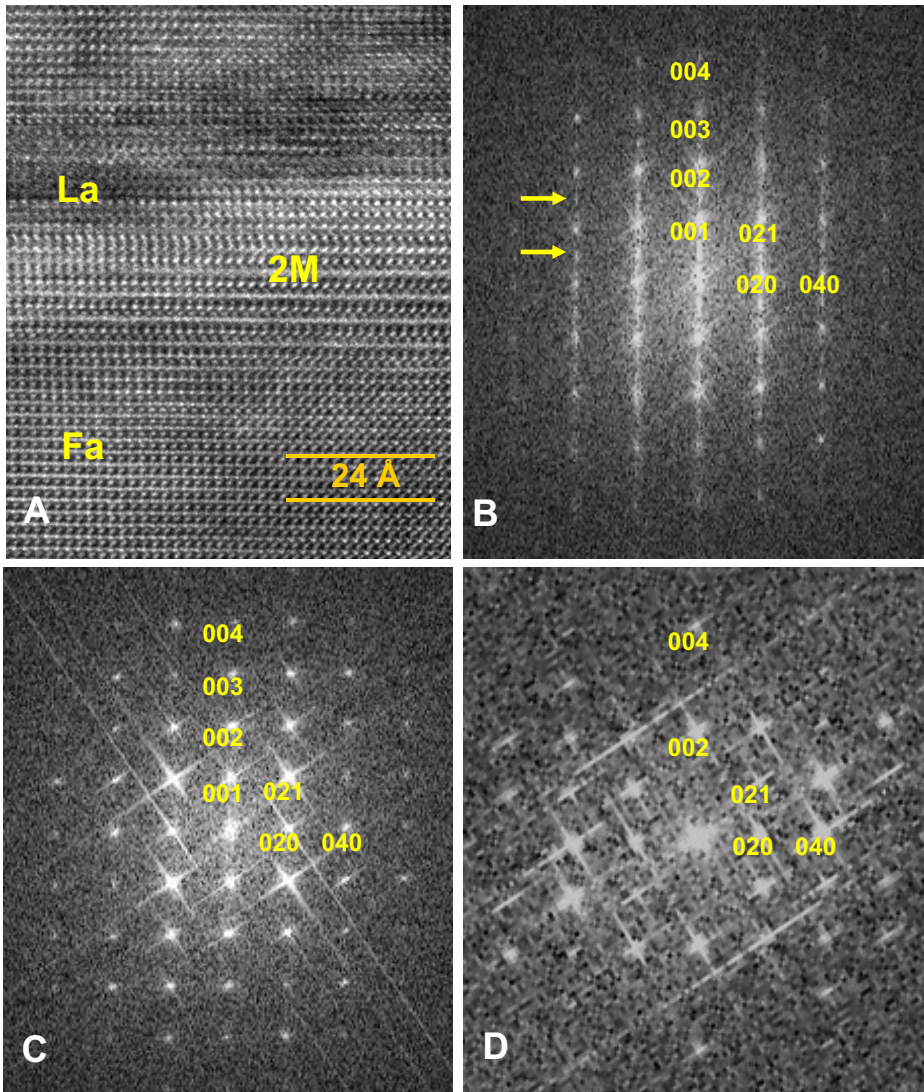
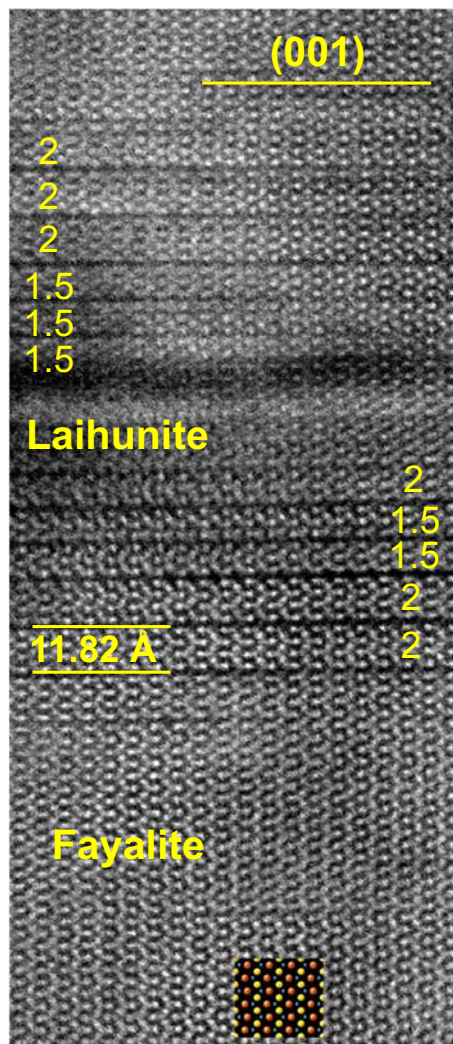
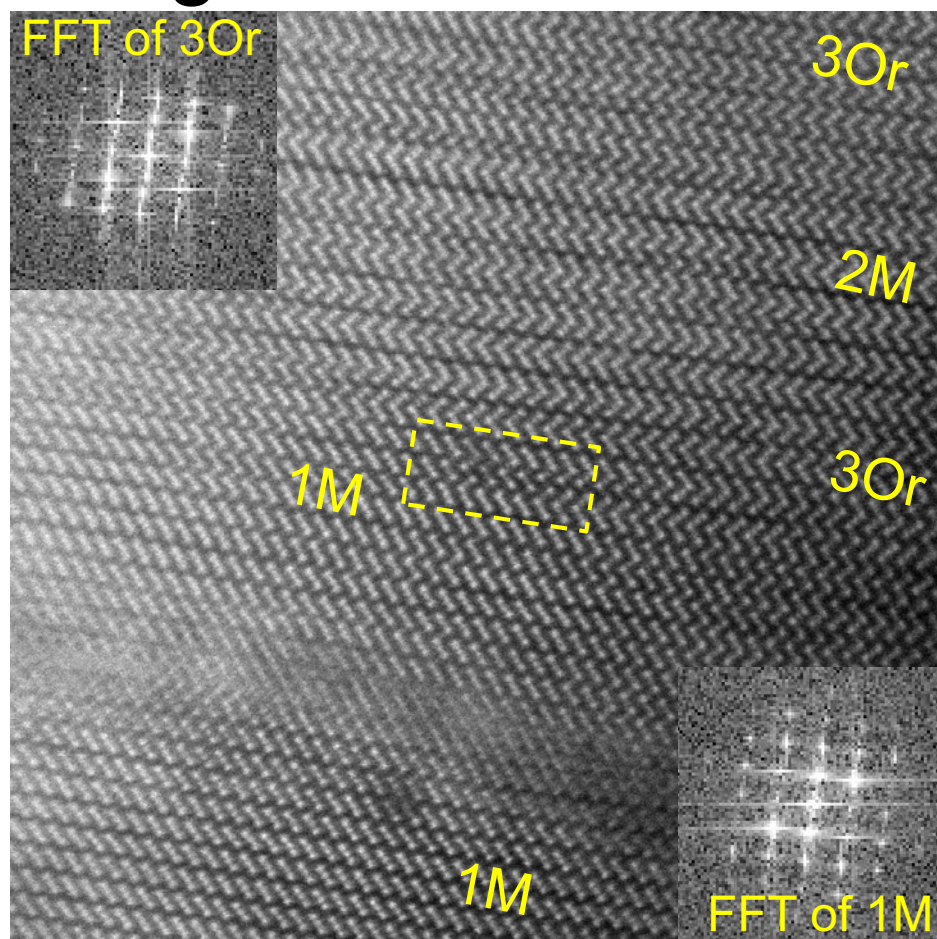


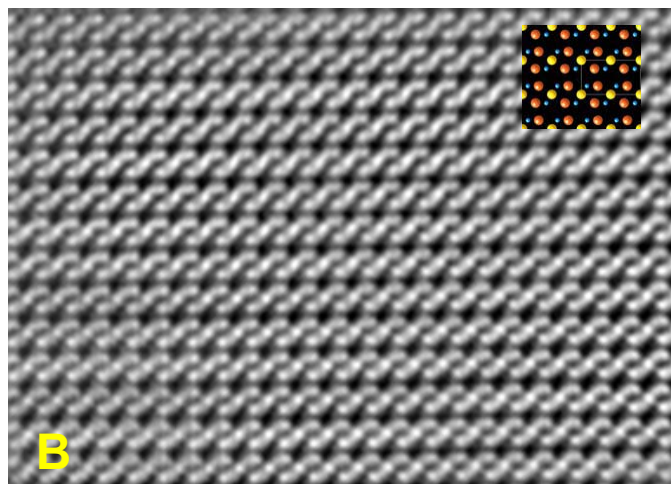
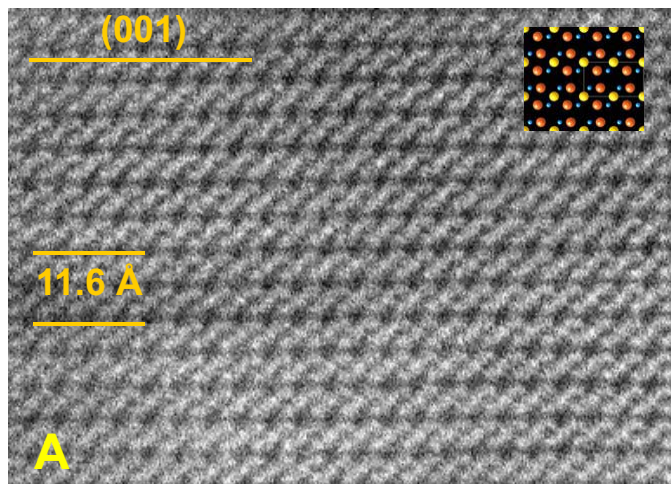
Figure 4

Figure 5



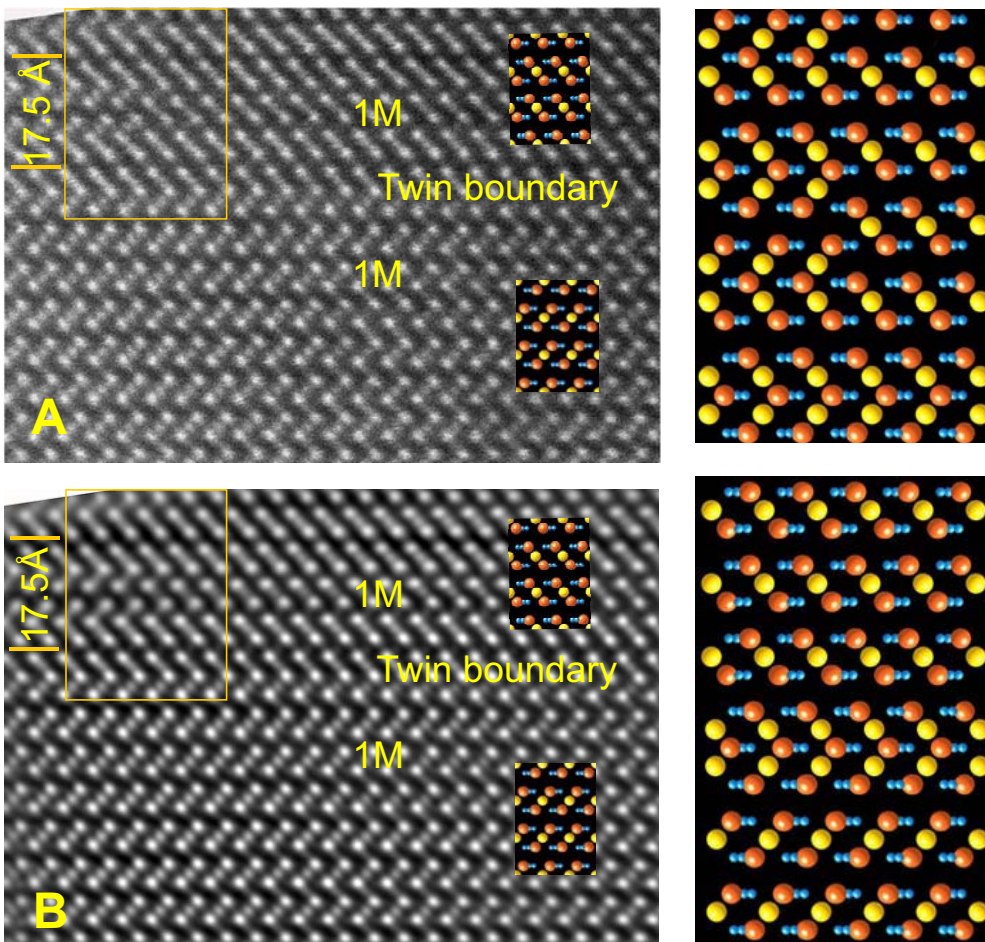
# Figure 6





# Figure 7

# Figure 8







# Figure 10

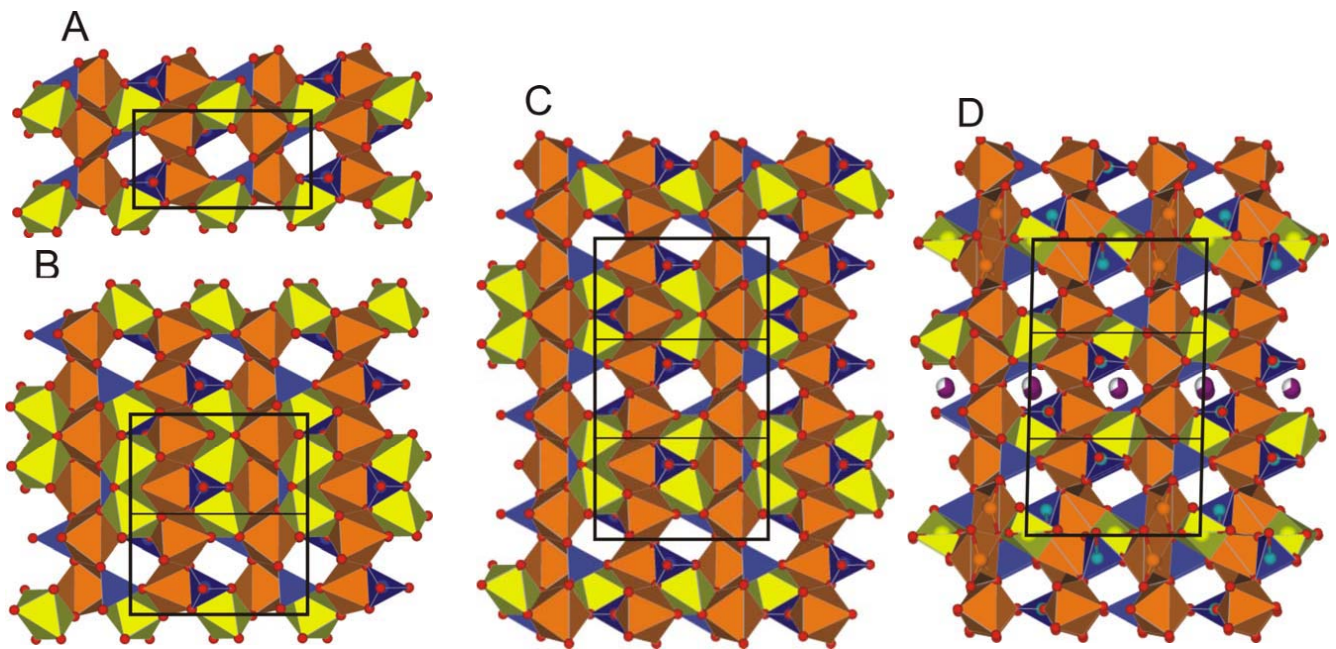
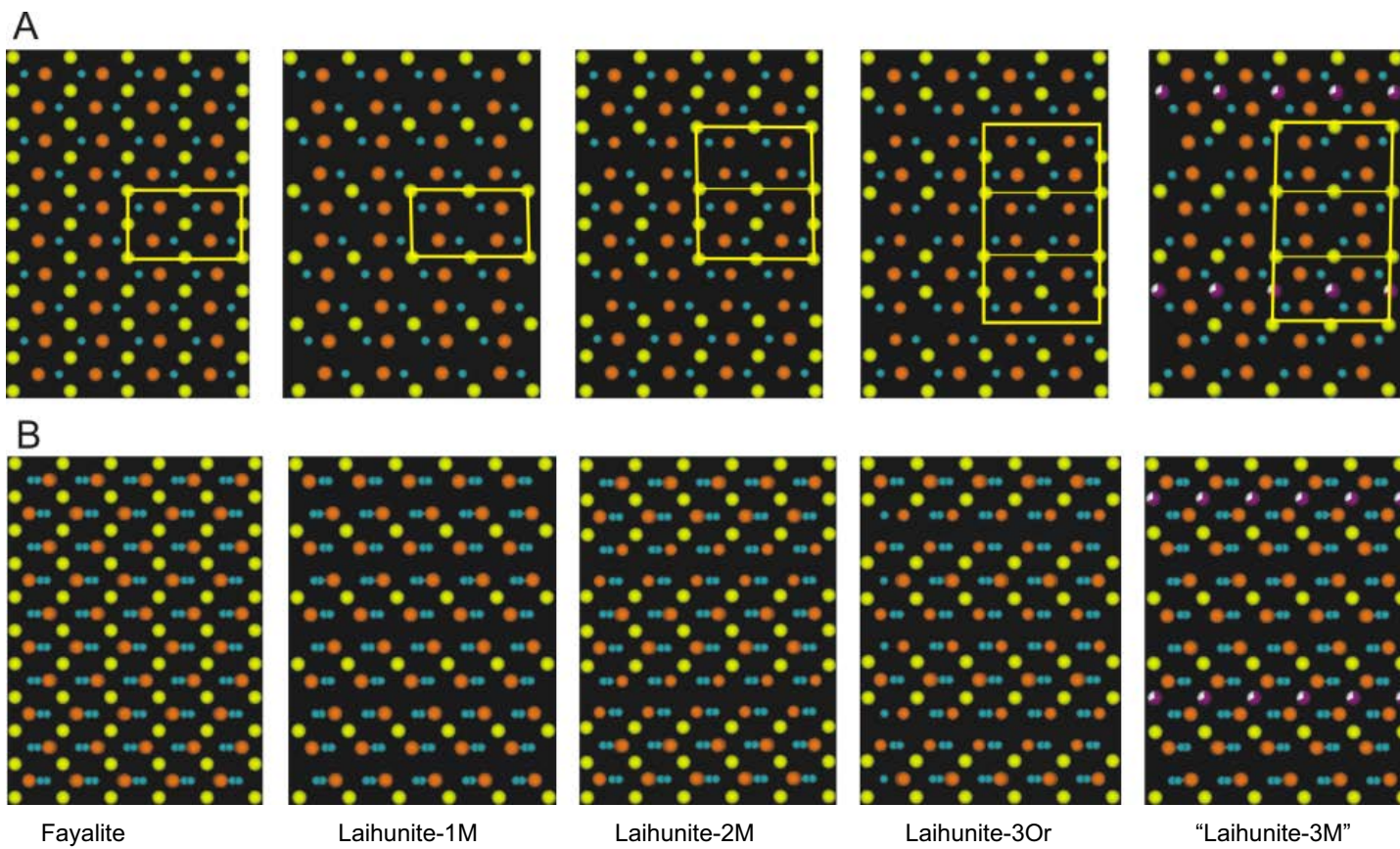


Figure 11



# Fig. 12

