

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26

Review 1

Sb⁵⁺ and Sb³⁺ substitution in segnitite: a new sink for As and Sb in the environment and implications for acid mine drainage

Stuart J. Mills¹, Barbara Etschmann², Anthony R. Kampf³, Glenn Poirier⁴ & Matthew Newville⁵

¹Geosciences, Museum Victoria, GPO Box 666, Melbourne 3001, Victoria, Australia

²Mineralogy, South Australian Museum, North Terrace, Adelaide 5000 Australia + School of Chemical Engineering, The University of Adelaide, North Terrace 5005, Australia

³Mineral Sciences Department, Natural History Museum of Los Angeles County, 900 Exposition Boulevard, Los Angeles, California 90007, U.S.A.

⁴Mineral Sciences Division, Canadian Museum of Nature, P.O. Box 3443, Station D, Ottawa, Ontario, Canada, K1P 6P4

⁵Center for Advanced Radiation Studies, University of Chicago, Building 434A, Argonne National Laboratory, Argonne. IL 60439, U.S.A.

*E-mail: smills@museum.vic.gov.au

27 **Abstract**

28 A sample of Sb-rich segnitite from the Black Pine mine, Montana, USA has been studied by
29 microprobe analyses, single crystal X-ray diffraction, μ -EXAFS and XANES spectroscopy.
30 Linear combination fitting of the spectroscopic data provided $\text{Sb}^{5+}:\text{Sb}^{3+} = 85(2):15(2)$, where
31 Sb^{5+} is in octahedral coordination substituting for Fe^{3+} and Sb^{3+} is in tetrahedral coordination
32 substituting for As^{5+} . Based upon this $\text{Sb}^{5+}:\text{Sb}^{3+}$ ratio, the microprobe analyses yielded the
33 empirical formula
34 $\text{Pb}_{1.02}\text{H}_{1.02}(\text{Fe}^{3+}_{2.36}\text{Sb}^{5+}_{0.41}\text{Cu}^{2+}_{0.27})_{\Sigma 3.04}(\text{As}^{5+}_{1.78}\text{Sb}^{3+}_{0.07}\text{S}^{6+}_{0.02})_{\Sigma 1.88}\text{O}_8(\text{OH})_{6.00}$. The crystal
35 structure refinement and bond valence analysis are consistent with these cation site
36 assignments. The formation of Sb-rich segnitite opens new possibilities for Sb sinks within
37 the supergene zone. Segnitite may, in fact, be an ideal host for the sequestering of a number
38 of toxic elements for $\text{pH} < 2$. At higher pH values, As is more likely to be incorporated into
39 schwertmannite and ferrihydrite.

40

41 Keywords: EXAFS, XANES, crystal structure, segnitite, antimony, valency, alunite
42 supergroup, oxidized zone.

43

44

45

46

47

48

49

50

51

52 **Introduction**

53 Segnitite, ideally $\text{PbFe}^{3+}_3(\text{AsO}_4)_2(\text{OH},\text{H}_2\text{O})_6$ (Birch et al. 1992), is the Pb□Fe-rich member of
54 the dussertite group, within the alunite supergroup (Mills et al. 2009a; Bayliss et al. 2010). It
55 forms solid solution series with both kintoreite and beudantite (e.g. Rattray et al. 1996;
56 Jambor 1999; Mills 2007), and it has been observed at more than 100 localities worldwide
57 (www.mindat.org).

58

59 Minerals in the alunite supergroup have the general formula $AB_3(\text{TO}_4)_2X_6$, where the *A*-site
60 can be occupied by monovalent (Na^+ , K^+ , Ag^+ , NH_4^+ or H_3O^+), divalent (Pb^{2+} , Ca^{2+} or Ba^{2+}),
61 or trivalent (Bi^{3+} or REE^{3+}) cations; the *B*-site can be occupied by either Fe^{3+} , Al^{3+} or Ga^{3+} ,
62 whilst the *T* cation site can be occupied by P^{5+} , S^{6+} or As^{5+} (e.g. Jambor, 1999). The
63 supergroup has always been of significant interest to mineral scientists especially because of
64 its relevance to acid mine drainage (e.g., Nordstrom et al. 2000; Welch et al. 2007, 2008,
65 2009) and the mobility of toxic elements (e.g., Kolitsch & Pring 2001) and because of its
66 structural variability (e.g., Grey et al. 2008; Mills et al. 2008). To date, however, there has
67 been only one reported example of Sb substitution within a member of the supergroup.
68 Kolitsch et al. (1999) reported the structure of “antimonian” dussertite from the Clara mine,
69 Germany, and attributed all of the Sb to Sb^{5+} substituting for Fe^{3+} within the octahedral site.
70 The find of an unusual Sb-rich segnitite from the Black Pine mine, Montana, USA, has
71 prompted further investigation by us into the valency and structural role of Sb within
72 members of the alunite supergroup.

73

74 **Experimental**

75 *Sample*

76 The Sb-rich segnitite occurs at the Black Pine mine, 14.5 km NW of Philipsburg, Granite Co.,
77 Montana, USA (46°26'52"N, 113°21'56"W), and was discovered by Mr. John Dagenais of
78 Vancouver, B.C., Canada. The mineral occurs as flattened, tabular or rhombohedral
79 yellowish crystals up to about 100 μm across. Crystals are commonly intergrown to form
80 botryoidal groups up to about 0.3 mm across and also form as coatings lining quartz vughs
81 and on quartz crystals. Associated minerals are hidalgoite and tetrahedrite. The Black Pine
82 mine is the type locality for philipsburgite (Peacor *et al.*, 1985), joëlbruggerite (Mills *et al.*,
83 2009a) and auriacusite (Mills *et al.*, 2010). The studied specimen has been deposited in the
84 collections of Mineral Sciences Department, Natural History Museum of Los Angeles
85 County, catalogue number 64096.

86

87 *X-ray Absorption Spectroscopy*

88 Sb K-edge (30491 eV) X-ray Absorption Near Edge Structure (XANES) and micro Extended
89 X-ray absorption Fine Structure (μ -EXAFS) spectra were measured at beam line 13-ID-C
90 (GSE-CARS) at the Advanced Photon Source (APS), at Argonne, USA. The APS is a 7 GeV
91 ring and operates in top-up mode with a current of 102 mA. 13-ID-C is an undulator beam
92 line with a Si(111) monochromator and an energy resolution ($\Delta E/E$) of 1.4×10^{-4} at 10 keV.
93 A focused beam size of $5 \mu\text{m}^2$ was used. A 16 element solid-state Ge detector was used for
94 detecting fluorescence data.

95

96 XANES and EXAFS data were analyzed with the HORAE package (Ravel and Newville
97 2005), calculations being performed using FEFF version 9 (Rehr *et al.* 2010). Self-absorption
98 effects were checked for the fluorescence data using the SABCOR (Booth and Bridges 2005)
99 correction routine that is incorporated into ATHENA (part of the HORAE package) and were
100 found to be negligible. The data used in the EXAFS fits ranged from $k = 2.0$ to 10.0 \AA^{-1} . The

101 fitting was done in R-space in the range $1.0 \square 4.0 \text{ \AA}$, with a Hanning window and multiple
102 $k^{1,2,3}$ weighting.

103

104 *Chemical analysis*

105 Quantitative wavelength-dispersive electron-microprobe analyses (6 points) were carried out
106 with a JEOL733 electron microprobe at the Mineral Sciences Division, Canadian Museum of
107 Nature. Operating conditions were 20 kV, with beam current of 20 nA and a 2 μm beam
108 diameter. Raw intensity data were corrected using a PAP matrix correction (Pouchou and
109 Pichoir 1984). No other elements were detected by energy dispersive spectroscopy. Sb^{3+} and
110 Sb^{5+} were apportioned based on the EXAFS and XANES reported below. H_2O was calculated
111 based on the site populations. The results, as well as the standards used, are shown in **Table**
112 **1**.

113

114 *X-ray diffraction*

115 The single-crystal X-ray diffraction data were obtained on a Rigaku R-Axis Rapid II curved
116 imaging plate microdiffractometer utilizing monochromatized $\text{MoK}\alpha$ radiation. The Rigaku
117 CrystalClear software package was used for processing of the diffraction data, including the
118 application of an empirical multi-scan absorption correction using ABSCOR (Higashi 2001).
119 The structure was solved by direct methods using SHELXS-97 software (Sheldrick 2008),
120 and SHELXL-97 (Sheldrick 2008) was used for the refinement of the structure. The
121 occupancies of the Fe and As sites were constrained to match the ratios obtained from the
122 chemical and spectroscopic analyses below. The final model, anisotropically refined,
123 converged to $R_1 = 0.0303$ for all 114 observed reflections [$F_o > 4\sigma F_o$] and 0.0333 for all 125
124 unique reflections. Details concerning data collection and structure refinement are provided

125 in **Table 2**. Fractional coordinates and atom displacement parameters are provided in **Table**
126 **3**, selected interatomic distances in **Table 4** and bond valence sums (BVS) in **Table 5**.

127

128 **Results**

129 *μ-EXAFS and XANES*

130 Linear combination fitting of the XANES spectra using standards Sb_2O_3 and $\text{KSb}(\text{OH})_6$
131 (Mitsunobu et al. 2008) resulted in a ratio of $\text{Sb}^{5+}:\text{Sb}^{3+}$ 85(2):15(2) ($\chi^2_{\text{red}} = 0.15$). The fraction
132 of $\text{Sb}^{5+}:\text{Sb}^{3+}$ was refined in the EXAFS data by fitting a linear combination of octahedrally
133 coordinated $\text{Sb}^{5+}\square\text{O}$ and tetrahedrally coordinated $\text{Sb}^{3+}\square\text{O}$ (**Table 6**). To reduce the number
134 of variables, the spectra for three grains were co-refined. The best fit resulted with 100%
135 $\text{Sb}^{5+}\square\text{O}$; however, it is possible to obtain a fit that is visually as good and with a χ^2_{red} that is
136 not statistically significantly different (Kelly et al. 2008) by fitting 85% octahedrally
137 coordinated $\text{Sb}^{5+}\square\text{O}$ and 15% tetrahedrally coordinated $\text{Sb}^{3+}\square\text{O}$ (**Figure 1**). Attempts to fit
138 more $\text{Sb}^{3+}\square\text{O}$ resulted in unreasonable fit parameters. It is important to note that the
139 microprobe analyses do not support having only As in the tetrahedral site, and tetrahedral
140 vacancies are not known to occur in the alunite supergroup. Thus, all lines of evidence
141 provide strong support for 15% $\text{Sb}^{3+}\square\text{O}$ being present in these grains.

142

143 *Microprobe analyses*

144 Based on the results of the μ -EXAFS and XANES above, we can recast the Sb_2O_5 in the
145 microprobe analyses as $\text{Sb}_2\text{O}_5 + \text{Sb}_2\text{O}_3$ in the proportion 85:15. In doing so, we obtain the
146 empirical formula
147 $\text{Pb}_{1.02}\text{H}_{1.02}(\text{Fe}^{3+}_{2.36}\text{Sb}^{5+}_{0.41}\text{Cu}^{2+}_{0.27})_{\Sigma 3.04}(\text{As}^{5+}_{1.78}\text{Sb}^{3+}_{0.07}\text{S}^{6+}_{0.02})_{\Sigma 1.88}\text{O}_8(\text{OH})_{6.00}$. It is interesting
148 to note that although $\text{Sb}^{5+} + \text{Cu}^{2+} > 0.5$ apfu, because of the difference in valence between
149 Sb^{5+} and Cu^{2+} , this phase does not qualify as a new mineral under the rule of *valency-imposed*

150 *double site occupancy* (Hatert et al. 2008), similar to beavertite-(Cu) or beavertite-(Zn)
151 (Bayliss et al. 2010; Sato et al. 2011). If Sb^{5+} were > 0.5 *apfu* (substituting for Fe^{3+}), this
152 mineral would qualify as a new mineral species.

153

154 *Crystal structure*

155 Segnitite has a rhombohedral alunite-type structure (e.g. Blount, 1974), consisting of layers
156 of corner-sharing $\text{FeO}_2(\text{OH})_4$ octahedra and AsO_4 tetrahedra parallel to (0001) and stacked
157 along *c*. Pb atoms are displaced from the origin, as in other Pb species of the supergroup (e.g.
158 philipsbornite, Cooper and Hawthorne 2012; plumbogummite, Mills et al. 2009c; kintoreite,
159 Kharisun et al. 1997; Grey et al. 2009), and are in 12-fold coordination. The $\text{Fe}(\text{O},\text{OH})_6$
160 octahedra each share corners with four neighbours to form a planar kagomé network with 3-
161 and 6-fold rings, which can also be described as hexagonal tungsten bronze (HTB)-type
162 layers. The O^{2-} anions of the $\text{Fe}(\text{O},\text{OH})_6$ octahedra in each 3-ring link to an arsenate, either
163 above or below the plane of the Fe layer. Successive layers are interconnected via hydrogen
164 bonds from H to arsenate O1 of an adjacent layer (Figure 2).

165

166 The average bond lengths for Sb-rich segnitite are $\langle \text{Fe}-\text{O} \rangle = 2.007 \text{ \AA}$, $\langle \text{As}-\text{O} \rangle = 1.685 \text{ \AA}$
167 and $\langle \text{Pb}-\text{O} \rangle = 2.84 \text{ \AA}$. The average bond lengths for $\langle \text{Fe}-\text{O} \rangle$ and $\langle \text{As}-\text{O} \rangle$ are
168 indistinguishable from those reported by Kolitsch et al. (1999) for “antimonian” dussertite,
169 2.009 and 1.684 \AA , respectively. It is not be surprising that they are the same, given that there
170 are many substitutions occurring in the Fe and As sites within the minerals. In the case of Sb-
171 rich segnitite, the incorporation of Sb^{5+} in the Fe site would be expected to slightly shorten
172 the average $\text{Fe}-\text{O}$ bond length; while, Cu^{2+} would be expected to elongate it. These
173 competing effects effectively cancel each other out. The BVS analysis provides a formal
174 valency for the site of 2.13 *v.u.*, with a site charge of 3.23+, which matches well with the

175 theoretical of 3.17+. The $\langle \text{As-O} \rangle$ is close to the general average for arsenate tetrahedra of
176 $\sim 1.683 \text{ \AA}$ (Shannon and Calvo, 1973), despite the fact that Sb^{3+} at the site should result in
177 longer bonds (the range observed for $\text{Sb}^{3+}\text{-O}$ bonds is $1.80\text{-}3.50 \text{ \AA}$; Mills et al. 2009b). It
178 may be that the small amount of S (the general average S-O in sulfates is 1.459; Hawthorne
179 et al. 2000) counteracts the effect of the Sb^{3+} . It is also noteworthy that the BVS for Pb and
180 O2 show overbonding, which is typical in the Pb-rich members of the alunite supergroup
181 (Mills et al. 2009d).

182

183 **Implications**

184 It is interesting to note that there are only two known minerals containing both Sb^{3+} and Sb^{5+} :
185 cervanite, $\alpha\text{-Sb}_2\text{O}_4$, a very common secondary mineral known from over two hundred
186 localities (www.mindat.org), and clinocervantite, $\beta\text{-Sb}_2\text{O}_4$, a very rare secondary mineral
187 only known from two localities in Italy (Basso et al. 1999). “Stibiconite” and
188 “bismutostibiconite” may be mixed valency; however, both are considered questionable
189 species (Atencio et al. 2010), and there is no spectroscopic data available to suggest that both
190 Sb^{3+} and Sb^{5+} are present in these phases.

191

192 Antimony is considered a toxic heavy metal, and because of this, many studies have been
193 aimed at understanding its behaviour in the supergene zone (e.g. Filella et al. 2009 and
194 references therein). In the supergene zone, Sb is often found as a weathering bi-product
195 associated with the mining of Sb sulphides and sulphosalts, such as stibnite, berthierite and
196 tetrahedrite. The presence of Sb is often also coupled with the presence of As, and this may
197 lead to severe environmental problems (e.g. Casiot et al. 2007; Majzlan et al. 2011; Mok and
198 Wai 1990). Leverett et al. (2012) recently showed that the minerals schafarzikite, FeSb_2O_4 ,
199 and tripuhyite, FeSbO_4 , act as important sinks for the element, while Majzlan et al. (2011)

200 showed that tripuyite was the main sink for Sb in the mine drainage tailings near Pezinok,
201 Slovakia, and that As was concentrated in amorphous oxide phases.

202

203 The formation of Sb-rich segnitite, therefore, opens new possibilities for Sb sinks within the
204 supergene zone. The Sb-rich segnitite contains on average about 28.93 wt % PbO, 25.88 wt
205 % As₂O₅, 6.33 wt % Sb₂O₅ and 2.44 wt % Sb₂O₃, showing that substantial amounts of toxic
206 elements can be locked away within the crystal structure of the mineral. In the case of the
207 neutral (pH range 6.5–8.0) tailings reported by Majzlan et al. (2011), it may be the case that
208 the mining tailings were at too high a pH for the formation of crystalline samples and/or
209 members of the alunite supergroup. Mills (2007) showed that segnitite could be synthesised
210 by hydrothermal methods in the pH range 1.0–1.5, indicating that under more acidic
211 conditions this could be a preferred sink for Pb, Fe, Sb and As. Acidic river waters with low
212 pH, such as those observed at Rio Tinto in Spain (pH range 1.5–2.7), exhibit a number of
213 jarosite group species (e.g. jarosite, natrojarosite and plumbojarosite) and show the potential
214 for sequestering a large number of toxic elements (Hudson-Edwards et al. 1999). Welch et al.
215 (2008) showed that jarosite dissolution is slow, making jarosite-group minerals good hosts
216 for sequestering toxic elements as long as the prevailing fluids remain at high pH. Insofar as
217 the precipitation of segnitite or other alunite-supergroup species, the main barrier appears to
218 be competition with As incorporation into schwertmannite and ferrihydrite (e.g. Carlson et al.
219 2002), which precipitate at higher pH values, 2.8–4.5 for schwertmannite and > 6.5 for
220 ferrihydrite (Bigam et al. 1996).

221

222 At the Black Pine mine, Mills et al. (2009c) deduced that joëlbruggerite,
223 Pb₃Zn₃(Sb⁵⁺, Te⁶⁺)As₂O₁₃(OH, O), formed in a highly oxidizing environment, in water in near
224 equilibrium with atmospheric oxygen [$\log a_{\text{O}_2(\text{aq})} = -3.58$] and a pH < 3. While cervantite

225 appears to form over a wide stability range (Roper et al. 2012), Sb-rich segnitite appears to
226 have formed in a much narrower range typified by highly oxidising conditions and $\text{pH} < 2$.

227

228 **Acknowledgements**

229 We thank Ritsuro Miyawaki, an anonymous reviewer and Associate Editor Diego Gatta who
230 provided helpful comments on the manuscript. The GeoSoilEnviroCARS (GSECARS)
231 beamline at the APS is supported by the National Science Foundation – Earth Sciences
232 (EAR-0622171) and Department of Energy – Geosciences (DE-FG02-94ER14466). Part of
233 this study was also funded by the John Jago Trelawney Endowment to the Mineral
234 Sciences Department of the Natural History Museum of Los Angeles County. Peter Williams
235 is also thanked for his insight on the role of Sb in the environment.

236

237 **References**

- 238 Atencio, D., Andrade, M.B., Christy, A.G., Gieré, R., and Kartashov, P.M. (2010) The
239 pyrochlore supergroup of minerals: nomenclature. *Canadian Mineralogist*, 48, 673–698.
240
- 241 Basso, R., Lucchetti, G., Zefiro, L., and Palenzona, A. (1999) Clinocervantite, $\beta\text{-Sb}_2\text{O}_4$, the
242 natural monoclinic polymorph of cervantite from the Cetine Mine, Siena, Italy. *European*
243 *Journal of Mineralogy*, 11, 95–100.
244
- 245 Bayliss, P., Kolitsch, U., Nickel, E.H., and Pring, A. (2010) Alunite supergroup:
246 recommended nomenclature. *Mineralogical Magazine*, 74, 919–927.
247
- 248 Bigham, J.M., Schwertmann, U., Traina, S.J., Winland, R.L., and Wolf, M. (1996)
249 Schwertmannite and the chemical modeling of iron in acid sulfate waters. *Geochimica et*
250 *Cosmochimica Acta*, 60, 2111–2121.
251
- 252 Birch, W.D., Pring, A., and Gatehouse, B.M. (1992) Segnitite, $\text{PbFe}_3\text{H}(\text{AsO}_4)_2(\text{OH})_6$, a new
253 mineral in the lusungite group from Broken Hill, New South Wales, Australia. *American*
254 *Mineralogist*, 77, 656–659.
255
- 256 Blount, A.M. (1974) The crystal structure of crandallite. *American Mineralogist*, 59, 41–47.
257
- 258 Booth, C.H., and Bridges, F. (2005) Improved self-absorption correction for fluorescence
259 measurements of extended X-ray absorption fine-structure. *Physica Scripta*, T115, 202–204.
260

- 261 Brown, I.D., and Altermatt, D. (1985) Bond-valence parameters from a systematic analysis of
262 the inorganic crystal structure database. *Acta Crystallographica*, B41, 244–247.
263
- 264 Carlson, L., Bigham, J.M., Schwertmann, U., Kyek, A., and Wagner, F. (2002) Scavenging of
265 As from acid mine drainage by schwertmannite and ferrihydrite: a comparison with synthetic
266 analogues. *Environmental Science & Technology*, 36, 1712–1719.
267
- 268 Casiot, C., Ujevic, M., Munoz, M., Seidel, J. L., and Elbaz-Poulichet, F. (2007) Antimony
269 and arsenic mobility in a creek draining an antimony mine abandoned 85 years ago (upper
270 Orb basin, France). *Applied Geochemistry*, 22, 788–798.
271
- 272 Cooper, M.A., and Hawthorne, F.C. (2012) Refinement of the crystal structure of zoned
273 philipsbornite–hidalgoite from the Tsumeb mine, Namibia, and hydrogen bonding in the
274 $D^{2+}(T^{5+}O_4)(TO_3OH)(OH)_6$ alunite structures. *Mineralogical Magazine*, 76, 839–849.
275
- 276 Filella, M., Williams, P. A., and Belzile, N. (2009) Antimony in the environment: knowns
277 and unknowns. *Environmental Chemistry*, 6, 95–105.
278
- 279 Grey, I.E., Mumme, W.G., Bordet, P., and Mills, S.J. (2008) A new crystal-chemical
280 variation of the alunite-type structure in monoclinic $PbZn_{0.5}Fe_3(AsO_4)_2(OH)_6$. *Canadian*
281 *Mineralogist*, 46, 1355–1364.
282
- 283 Grey, I.E., Mumme, W.G., Mills, S.J., Birch, W.D., and Wilson, N.C. (2009) The crystal
284 chemical role of Zn in alunite-type minerals: structure refinements for kintoreite and zincian
285 kintoreite. *American Mineralogist*, 94, 676–683.
286
- 287 Hatert, F., and Burke, E.A. (2008) The IMA–CNMNC dominant-constituent rule revisited
288 and extended. *Canadian Mineralogist*, 46, 717–728.
289
- 290 Hawthorne, F. C., Krivovichev, S. V., and Burns, P. C. (2000) The crystal chemistry of
291 sulfate minerals. *Reviews in Mineralogy and Geochemistry*, 40, 1–112.
292
- 293 Higashi, T. (2001) ABSCOR. Rigaku Corporation, Tokyo, Japan.
294
- 295 Hudson-Edwards, K.A., Schell, C., and Macklin, M.G. (1999) Mineralogy and geochemistry
296 of alluvium contaminated by metal mining in the Rio Tinto area, southwest Spain. *Applied*
297 *Geochemistry*, 14, 1015–1030.
298
- 299 Jambor, J.L. (1999) Nomenclature of the alunite supergroup. *Canadian Mineralogist*, 37,
300 1323–1341.
301
- 302 Kelly, S., Hesterberg, D., and Ravel, B. (2008) Analysis of soils and minerals using X-ray
303 absorption spectroscopy. In: *Methods of Soil Analysis, Part 5. Mineralogical Methods*. Soil
304 Sciences Society of America, Madison, USA, 387–463.
305
- 306 Kharisun, Taylor, M.R., Bevan, D.J.M., and Pring, A. (1997) The crystal structure of
307 kintoreite, $PbFe_3(PO_4)_2(OH,H_2O)_6$. *Mineralogical Magazine*, 61, 123–129.
308

- 309 Kolitsch, U., and Pring, A. (2001) Crystal chemistry of the crandallite, beudantite and alunite
310 groups: a review and evaluation of the suitability as storage materials for toxic metals.
311 Journal of Mineralogical and Petrological Sciences, 96, 67–78.
312
- 313 Kolitsch, U., Slade, P.G., Tiekink, E.R.T., and Pring, A. (1999) The structure of antimonian
314 dussertite and the role of antimony in oxysalt minerals. Mineralogical Magazine, 63, 17–26.
315
- 316 Krivovichev, S. V., and Brown, I. D. (2001) Are the compressive effects of encapsulation an
317 artifact of the bond valence parameters? Zeitschrift für Kristallographie, 216, 245–247.
318
- 319 Leverett, P., Reynolds, J.K., Roper, A.J., and Williams, P.A. (2012) Tripuhyite and
320 schafarzikite: two of the ultimate sinks for antimony in the natural environment.
321 Mineralogical Magazine, 76, 891–902.
322
- 323 Majzlan, J., Lalinská, B., Chovan, M., Bläß, U., Brecht, B., Göttlicher, J., Steininger, R.,
324 Hug, K., Ziegler, S., and Gescher, J. (2011) A mineralogical, geochemical, and microbiological
325 assessment of the antimony- and arsenic-rich neutral mine drainage tailings near Pezinok,
326 Slovakia. American Mineralogist, 96, 1–13.
327
- 328 Mills, S.J. (2007) *The crystal chemistry and geochronology of minerals from Broken Hill*.
329 Unpublished PhD Thesis, School of Earth Sciences, University of Melbourne.
330
- 331 Mills, S.J., Christy, A.G., Chen, E.C.C., and Raudsepp, M. (2009b) Revised values of the
332 bond valence parameters for $^{[6]}Sb(V)-O$ and $^{[3-11]}Sb(III)-O$. Zeitschrift für Kristallographie,
333 224, 423–431.
334
- 335 Mills, S.J., Grey, I.E., Mumme, W.G., Miyawaki, R., Matsubara, S., Bordet, P., Birch, W.D.,
336 and Raudsepp, M. (2008) Kolitschite, $Pb[Zn_{0.5}, \square_{0.5}]Fe_3(AsO_4)_2(OH)_6$, a new mineral from the
337 Kintore opencut, Broken Hill. Australian Journal of Mineralogy, 14, 63–67.
338
- 339 Mills, S.J., Hatert, F., Nickel, E.H., and Ferraris, G. (2009a) The standardisation of mineral
340 group hierarchies: application to recent nomenclature proposals. European Journal of
341 Mineralogy, 21, 1073–1080.
342
- 343 Mills, S.J., Kampf, A.R., Poirier, G., Raudsepp, M., and Steel, I.M. (2010) Auriacusite,
344 $Fe^{3+}Cu^{2+}(As,Sb)O_4O$, the first M^{3+} member of the olivenite group, from the Black Pine mine,
345 Montana, USA. Mineralogy and Petrology, 99, 113–120.
346
- 347 Mills, S.J., Kampf, A.R., Raudsepp, M., and Christy, A.G. (2009d) The crystal structure of
348 Ga-rich plumbogummite from Tsumeb, Namibia. Mineralogical Magazine, 73, 837–845.
349
- 350 Mills, S.J., Kolitsch, U., Miyawaki, R., Groat, L.A., and Poirier, G. (2009c) Joëlbruggerite,
351 $Pb_3Zn_3(Sb^{5+}, Te^{6+})As_2O_{13}(OH, O)$, the Sb^{5+} analogue of dugganite, from the Black Pine mine.
352 Montana. American Mineralogist, 94, 1012–1017.
353
- 354 Mitsunobu, S., Takahashi, Y., Sakai, Y., and Inumaru, K. (2008) Interaction of synthetic
355 sulfate green rust with antimony (V). Environmental Science & Technology, 43, 318–323.
356

- 357 Mok, W.M. and Wai, C.M. (1990) Distribution and mobilization of arsenic and antimony
358 species in the Coeur d'Alene River, Idaho. *Environmental Science & Technology*, 24,
359 102–108.
360
- 361 Nordstrom, D.K., Alpers, C.N., Ptacek, C.J., and Blowes D.W. (2000) Negative pH and
362 extremely acidic mine waters from Iron Mountain, California. *Environmental Science &*
363 *Technology*, 34, 254–258.
364
- 365 Peacor, D.R., Dunn, P.J., Ramik, R.A., Sturman, B.D., and Zeihan, L.G. (1985)
366 Philipsburgite, a new copper zinc arsenate hydrate related to kipushite, from Montana.
367 *Canadian Mineralogist*, 23, 255–258.
368
- 369 Pouchou, J. L., and Pichoir, F. (1984) A New Model for Quantitative X-Ray Microanalysis. I.
370 Application to the Analysis of Homogeneous Samples. *Rech Aerospaciale*, 3, 167–192.
371
- 372 Rattray, K.J., Taylor, M.R., Bevan, D.J.M., and Pring, A. (1996) Compositional segregation
373 and solid solution in the lead-dominant alunite-type minerals from Broken Hill, NSW.
374 *Mineralogical Magazine*, 60, 779–786.
375
- 376 Ravel, B., and Newville, M. (2005) ATHENA, ARTEMIS, HEPHAESTUS: data analysis for
377 X-ray absorption spectroscopy using IFEFFIT. *Journal of Synchrotron Radiation*, 12, 537–
378 541.
379
- 380 Rehr, J.J., Kas, J.J., Vila, F.D., Prange, M.P., and Jorissen, K. (2010) Parameter-free
381 calculations of X-ray spectra with FEFF9. *Physical Chemistry Chemical Physics*, 12, 5503–
382 5513.
383
- 384 Roper, A.J., Williams, P.A., and Filella, M. (2012) Secondary antimony minerals: phases that
385 control the dispersion of antimony in the supergene zone. *Chemie der Erde-Geochemistry*,
386 72, 9–14.
387
- 388 Sato, E., Nakai, I., Terada, Y., Tsutsumi, Y., Yokoyama, K., Miyawaki, R., and Matsubara, S.
389 (2011) Beaverite-(Zn), $\text{Pb}(\text{Fe}_2\text{Zn})(\text{SO}_4)_2(\text{OH})$, a new member of the alunite group, from
390 Mikawa Mine, Niigata Prefecture, Japan. *Mineralogical Magazine*, 75, 375–377.
391
- 392 Shannon, R.T., and Calvo, C. (1973) Refinement of the crystal structure of low temperature
393 Li_3VO_4 and analysis of mean bond lengths in phosphates, arsenates, and vanadates. *Journal of*
394 *Solid State Chemistry*, 6, 538–549.
395
- 396 Sheldrick, G.M. (2008) A short history of SHELX. *Acta Crystallographica*, A64, 112–122.
397
- 398 Welch, S.A., Christy, A.G., Kirste, D., Beavis, S.G., and Beavis, F. (2007) Jarosite
399 dissolution I – trace cation flux in acid sulfate soils. *Chemical Geology*, 245, 183–197.
400
- 401 Welch, S.A., Kirste, D., Christy, A.G., Beavis, F.R., and Beavis, S.G. (2008) Jarosite
402 dissolution II – reaction kinetics, stoichiometry and acid flux. *Chemical Geology*, 254,
403 73–86.
404
- 405 Welch, S.A., Christy, A.G., Isaacson, L., and Kirste, D. (2009) Mineralogical control of rare
406 earth elements in acid sulfate soils. *Geochimica et Cosmochimica Acta*, 73, 44–64.

407 Table 1. Chemical analytical data for the Sb-rich segnitite.

	average	min	max	standards
PbO	28.84	28.17	29.48	galena
CuO	2.73	2.59	2.84	cuprite
Fe ₂ O ₃	23.86	23.01	25.34	hematite
Sb ₂ O ₅ *	9.85	7.86	11.19	Sb ₂ Te
Sb ₂ O ₅ *	8.37	6.68	9.51	
Sb ₂ O ₃ *	1.33	1.06	1.51	
As ₂ O ₅	25.92	25.54	26.26	GaAs
SO ₃	0.25	0.16	0.30	pentlandite
H ₂ O _{calc}	8.01			
Total	99.31			

408 *recast into Sb₂O₅ and Sb₂O₃ based on 85:15 ratio.

409

410 Table 2. Data collection and structure refinement details for the Sb-rich segnitite.

411

Space group	<i>R-3m</i>
Unit cell dimensions	
<i>a</i> (Å)	7.3730(14)
<i>c</i> (Å)	17.228(3)
<i>V</i> (Å ³)	811.0(3)
<i>Z</i>	3
Absorption coefficient	26.452
<i>F</i> (000)	1073
° range	3.40□20.80
Index ranges	-7 ≤ <i>h</i> ≤ 7, -7 ≤ <i>k</i> ≤ 7, -16 ≤ <i>l</i> ≤ 17
Reflections collected/unique	1342 / 125
Reflections with <i>F</i> > 4σ(<i>F</i>)	114
Refinement method	Full-matrix least-squares on <i>F</i> ²
Parameters refined	31
GoF	1.22
Final <i>R</i> indices [<i>F</i> _o > 4σ(<i>F</i>)]	0.0303
<i>R</i> indices (all data)	0.0333
Extinction coefficient	0.0006(3)
Largest diff. peak / hole	+0.58, -0.81 e Å ⁻³

412

413

414

415

416 Table 3. Fractional coordinates and atomic displacement parameters for the Sb-rich segnitite.
 417

	<i>x</i>	<i>y</i>	<i>z</i>	U_{eq}	<i>occy.</i>	U_{11}	U_{22}	U_{33}	U_{23}	U_{13}	U_{12}
Pb	0.044(11)	0	0	0.032(17)	1	0.030(14)	0.03(5)	0.0357(17)	-0.006(8)	-0.003(4)	0.01(3)
Fe1	0.5	0	0.5	0.0097(11)	0.78	0.0088(15)	0.0070(18)	0.0129(18)	0.0023(13)	0.0012(6)	0.0035(9)
Sb1	0.5	0	0.5	0.0097(11)	0.13	0.0088(15)	0.0070(18)	0.0129(18)	0.0023(13)	0.0012(6)	0.0035(9)
Cu1	0.5	0	0.5	0.0097(11)	0.09	0.0088(15)	0.0070(18)	0.0129(18)	0.0023(13)	0.0012(6)	0.0035(9)
As2	0	0	0.31346(18)	0.0194(12)	0.95	0.0180(14)	0.0180(14)	0.022(2)	0	0	0.0090(7)
Sb2	0	0	0.31346(18)	0.0194(12)	0.04	0.0180(14)	0.0180(14)	0.022(2)	0	0	0.0090(7)
S2	0	0	0.31346(18)	0.0194(12)	0.02	0.0180(14)	0.0180(14)	0.022(2)	0	0	0.0090(7)
O1	0	0	0.5890(12)	0.030(6)	1	0.038(9)	0.038(9)	0.014(12)	0	0	0.019(4)
O2	0.2093(8)	-0.2093(8)	-0.0536(7)	0.019(3)	1	0.019(5)	0.019(5)	0.024(7)	0.001(3)	-0.001(3)	0.013(6)
O3	0.1261(8)	-0.1261(8)	0.1349(6)	0.017(3)	1	0.009(5)	0.009(5)	0.034(8)	-0.004(3)	0.004(3)	0.007(5)
H1	0.196(2)	-0.196(2)	0.129(9)	0.02	1						

418 Table 4. Selected bond lengths (Å) in the Sb-rich segnitite.
 419

Fe1	O3	1.999(4)	×4
	O2	2.022(11)	×2
	<Fe1□O>	2.007	
As2	O1	1.68(2)	
	O2	1.687(11)	×3
	<As2□O>	1.685	
Pb	O2	2.57(7)	×2
	O3	2.68(4)	×2
	O3	2.84(4)	×2
	O2	2.85(4)	×2
	O3	3.00(4)	×2
	O2	3.10(4)	×2
<Pb□O>	2.84		

420

421

422 Table 5. Bond valence sums for the Sb-rich segnitite. Values are expressed in valence units.
 423

	Pb		Fe		As		H	Σ
O1					1.29		0.3	1.59
O2	0.55	↓×2	0.52	↓×2	1.24	↓×3		2.31
O3	0.52	↓×2	0.55	↓×4			0.7	2.03
Σ	2.14		2.13		5.00			

424 Notes: Pb–O bond strengths from Krivovichev and Brown (2001); Sb–O from Mills et al. (2009b);
 425 Fe–O, Cu–O, As–O and H–O bond strengths from Brown and Altermatt (1985).

426

427

428

429

430

431

432 Table 6. EXAFS refinement parameters.

433

434

Sample	Fit	Site	Ligand	R (Å)	σ^2 (Å ²)	ΔE_0 (eV)	χ^2_{red}
Grain2	1. 100% Sb ⁵⁺	Oct	Sb ⁵⁺ □O ₆	1.965(7)	0.0027(5)	5.9(8)	103.5
		Oct	Multiple scattering paths	3.365(7) 3.930(7)	0.0027(5)		
		Oct	Sb ⁵⁺ □O ₆	1.969(8)	0.0027(5)		
Oct	Multiple scattering paths	3.370(8) 3.937(8)	0.0027(5)				
Grain3	1. 100% Sb ⁵⁺	Oct	Sb ⁵⁺ □O ₆	1.96(1)	0.0027(5)	5.9(8)	103.5
		Oct	Multiple scattering paths	3.36(1) 3.93(1)	0.0027(5)		
		Oct	Sb ⁵⁺ □O ₆	1.963(7)	0.0004(5)		
Oct	Multiple scattering paths	3.361(7) 3.925(7)	0.0004(5)				
Tet	Sb ³⁺ □O ₄	1.80 (fix)					
Grain2	2. 85% Sb ⁵⁺ + 15% Sb ³⁺	Oct	Sb ⁵⁺ □O ₆	1.965(8)	0.0004(5)	4.9(7)	114.5
		Oct	Multiple scattering paths	3.365(8) 3.931(8)	0.0004(5)		
		Tet	Sb ³⁺ □O ₄	1.80 (fix)			
Grain3	2. 85% Sb ⁵⁺ + 15% Sb ³⁺	Oct	Sb ⁵⁺ □O ₆	1.96(1)	0.0004(5)	4.9(7)	114.5
		Oct	Multiple scattering paths	3.36(1) 3.92(1)	0.0004(5)		
		Tet	Sb ³⁺ □O ₄	1.80 (fix)			
Grain1	2. 85% Sb ⁵⁺ + 15% Sb ³⁺	Oct	Sb ⁵⁺ □O ₆	1.96(1)	0.0004(5)	4.9(7)	114.5
		Oct	Multiple scattering paths	3.36(1) 3.92(1)	0.0004(5)		
		Tet	Sb ³⁺ □O ₄	1.80 (fix)			

435

436

437

438

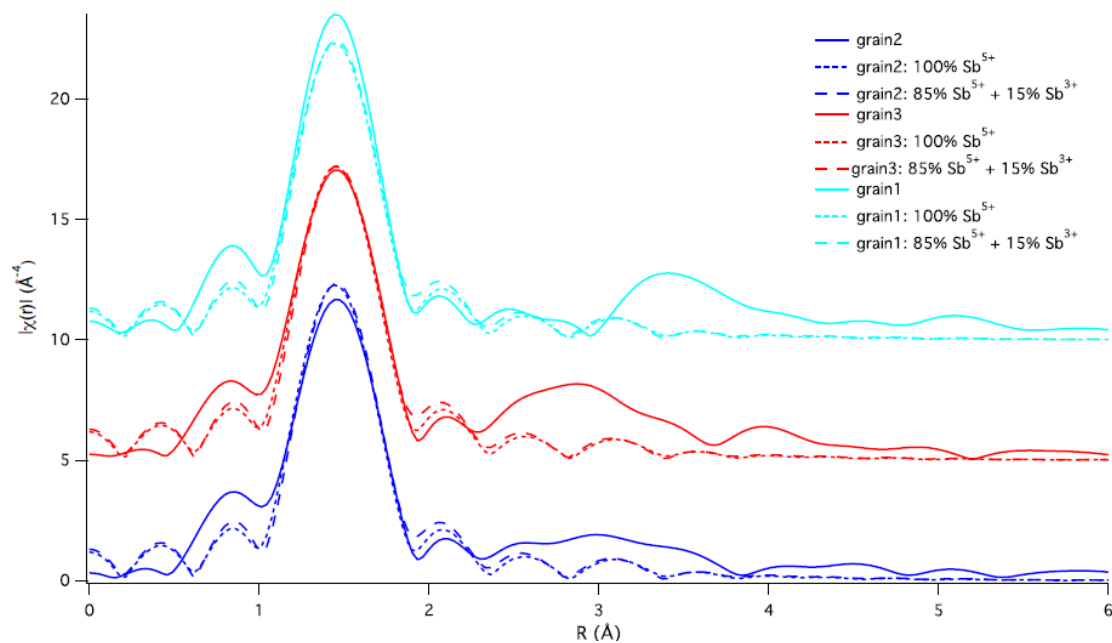
439

440

441

442 Figure 1. Sb K-edge EXAFS fits for three Sb-rich segnitite grains.

443



444

445

446

447

448

449

450

451

452

453

454

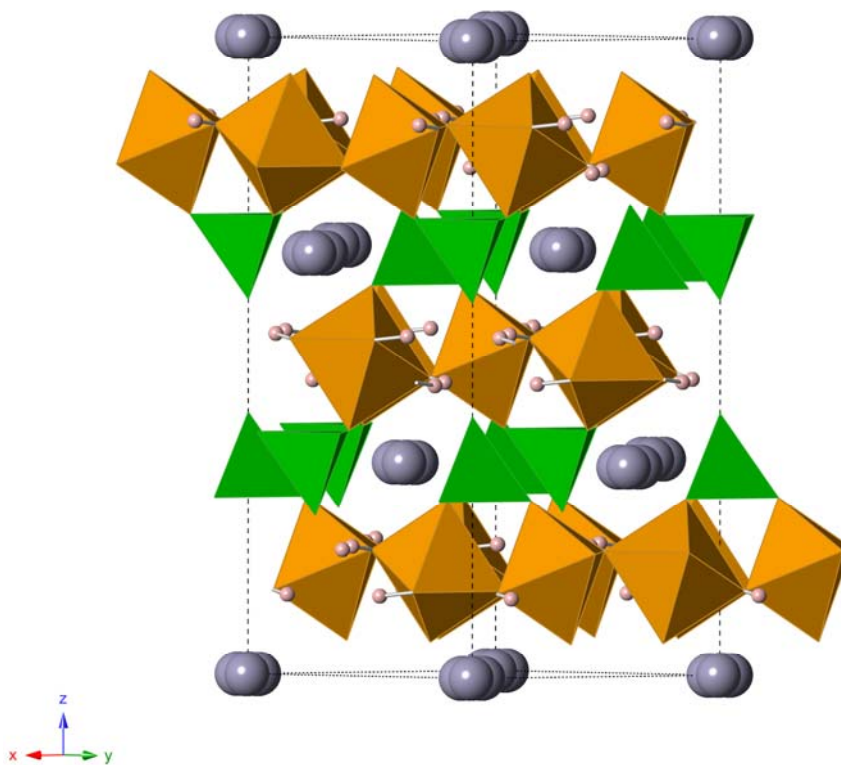
455

456

457

458

459 Figure 2. The crystal structure of Sb-rich segnitite viewed down [110]. FeO₆ octahedra are
460 orange, AsO₄ tetrahedra green, Pb atoms are grey and H atoms pink.



461

