

REVISION 1

Identifying the spin transition in Fe²⁺-rich MgSiO₃ perovskite from X-ray diffraction and vibrational spectroscopy

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Abstract

Based on numerical results from density-functional perturbation theory calculations, we show that the magnetic spin transition in Fe²⁺-rich MgSiO₃ perovskite can be identified as changes in the powder X-ray diffraction (XRD) pattern and the vibrational spectra. In particular theory predicts how the symmetry breaking and the volume reduction associated with the spin transition affect both structural and vibrational properties. The XRD measurements of (Mg_{0.5}Fe_{0.5})SiO₃ perovskite indeed demonstrated that the new diffraction peaks and the peak broadening formed during the spin transition can be explained by the associated symmetry breaking. We also show computationally that certain vibrational peaks exhibit a shift at the transition;

31 the Grüneisen parameters of certain modes are affected by the transition, thus bearing
32 on the thermodynamical properties. Raman and/or infrared measurements before and
33 after the spin transition could identify these changes.

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Introduction

36 Fe-bearing perovskite, the major mineral of the Earth's lower mantle, exhibits
37 a magnetic spin transition under compression. The transition corresponds roughly to
38 the reduction of the number of unpaired d electrons of Fe^{2+} iron: from 4 unpaired
39 electrons for high-spin state to 0 unpaired electrons for the low-spin state. For Fe^{3+}
40 iron these numbers decrease from 5 to 3 to 1 unpaired electrons into respectively the
41 high-spin, intermediate-spin and low-spin states. These numbers represent ideal
42 nominal molecular-like values; in a solid there are electronic bands and the magnetic
43 spin is given per unit cell as the difference between the total number of electrons with
44 the two different spins occupying the bands.

45 The transition depends on a series of factors, like pressure and temperature,
46 iron amount, iron crystallographic site and iron redox state. The ferrous iron on the
47 Mg site stays in high-spin up to large pressures (Bengtson et al. 2009; Hsu et al.
48 2010a, 2011) when in relatively low, mantle-like concentrations (Bengtson et al.
49 2008). The ferric iron can be incorporated in both the Mg and the Si sites, with the
50 one on the Si site undergoing a spin transition at very low pressures and the one on
51 the Mg site remaining in high-spin, until the transformation to post-perovskite (e.g.,
52 Stackhouse et al., 2007; Irifune et al., 2010; Caracas 2010; Hsu et al. 2011).

53 The high-spin to low-spin transition of ferrous iron has been documented
54 experimentally with several in situ measurements using nuclear or electronic
55 spectroscopies, like Mössbauer spectroscopy (Jackson et al. 2005; McCammon et al.
56 2008; Grocholski et al. 2009; Lin et al. 2012) or X-ray emission spectroscopy (XES)
57 (Badro et al. 2004; Li et al. 2004, 2006): a reduction in the local magnetization of the
58 iron atoms is seen either as a change in the Mössbauer pattern or as the disappearance
59 of the Fe $K\beta$ satellite peak in the XES spectra. Several reviews give some very good
60 overviews of the recent literature on this subject (Hsu et al. 2010b; Lin et al. 2013).

61 In parallel with the experimental effort, first principles calculations based on
62 density functional theory showed that the transition is associated with a structural
63 distortion corresponding to the reduction in size of the Fe atoms (Umemoto et al.

64 2010). Such distortion breaks the site symmetry of the iron atoms, allowing a lateral
65 displacement that accommodates the reduction in volume. The electronic changes
66 connected to the magnetic transition have been extensively studied for various iron
67 concentration and configuration in a series of papers (Bengtson et al. 2009; Hsu et al.
68 2010a; Umemoto et al. 2010). In recent phonon calculations we identified dynamical
69 instabilities in the perovskite structure with orthorhombic symmetry and low-spin
70 configuration. Following the unstable vibrational modes we were able to determine on
71 a strong physical basis the structural distortions and the symmetry breaking associated
72 to the spin transition (Caracas et al. 2011).

73 The geophysical implications of the spin transition in perovskite are yet a
74 matter of investigation. It was suggested that its effect on the bulk seismic wave
75 velocities is weak, such as at realistic iron concentrations it would not be visible in
76 seismic data (Li et al. 2005). However, the changes in seismic anisotropy, like the
77 ratio of the seismic compressional (V_p) and shear (V_s) wave velocities or the shear
78 wave splitting $|V_{s1}-V_{s2}|$, are sensitive to the spin state of iron. The spin transition can
79 induce differences of anisotropy between the top and the bottom of the lower mantle
80 that are large enough to be visible in seismic data and tomographic studies (Caracas et
81 al. 2010).

82 The investigations of the effects of the spin transition on the ternary phase
83 diagram of perovskite - post-perovskite - magnesiowüstite and on the iron
84 partitioning between the lower-mantle minerals are still in their incipient phase with
85 no final answer reached so far (Wolf et al. 2008; Metsue and Tsuchiya 2012;
86 Grocholski et al. 2012), but might be important, and one expects important effort to
87 be allocated towards this topics in the near future, from both experiments and
88 calculations. Consequently a rapid and inexpensive method of identification of the
89 spin transition in perovskite is highly desirable.

90 Because both the electronic and the nuclear spectroscopic methods are rather
91 tedious to perform and require a considerable experimental setup, here we investigate
92 alternate possibilities for identifying the spin transition. We concentrate and discuss
93 only the case of ferrous iron on the A site. We base on the structural distortions and
94 the symmetry breaking that are associated with this spin transition to quantify the
95 potential signature of the transition on X-ray diffraction (XRD) patterns and on
96 vibrational spectra. We show that changes in these spectra would make them viable

97 methods for identifying the spin transition if careful accurate measurements are
98 recorded.

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Methodology

101 We perform first-principles calculations based on the density functional theory
102 (DFT) in the ABINIT implementation (Gonze et al. 2002, 2009; Torrent et al. 2008).

103 First we employ the Planar Augmented Wavefunction (PAW) approach to
104 obtain the crystal structures and to investigate the symmetry breaking due to the spin
105 transition. In the PAW method an accurate, all-electron, representation of the
106 electronic system is placed in one-to-one correspondence with a pseudopotential-like
107 treatment of the valence electrons (Blochl 1994). We use the local density
108 approximation (LDA) for the exchange correlation energy. It was shown that the use
109 of more sophisticated exchange correlation formalisms, like the generalized gradient
110 approximation, does not change the physical description of the system, except for an
111 upward shift of the spin transition pressure (Hsu et al. 2010). We adopt $6\times 6\times 6$ grids
112 of special high-symmetry \mathbf{k} points (Monkhorst and Pack 1976) to sample the
113 electronic density in the reciprocal space. We employ a 16Ha (1 Ha=27.2116eV)
114 kinetic energy cut-off for the wave functions on the coarse mesh and a 36 Ha cut-off
115 for the wave functions on the finer grid inside the PAW spheres. This set of
116 parameters ensures an accuracy of the calculation better than 1GPa in pressure and
117 1mHa/unit cell in energy. We fully relax the structures (i.e. we minimize the
118 interatomic force constants and the non-hydrostatic stresses) at several pressures and
119 iron compositions, for both the high-spin and low-spin magnetic configurations. We
120 use the resulting theoretical structures to build the X-ray diffraction pattern, which we
121 compare with experimental patterns.

122 In a second set of calculations we analyze the lattice dynamical properties of
123 both high-spin and low-spin perovskite within the density-functional perturbation
124 theory (DFPT) (Baroni et al. 2002; Gonze et al. 2005). We use the theoretical
125 structures as determined above in PAW, to determine the phonons in the center of the
126 Brillouin zone. However, because of current implementation restrictions, we do not
127 use PAW for the dynamical calculations, but instead we need to employ Troullier-
128 Martins norm-conserving pseudopotentials. We relax the structures in the standard

129 norm-conserving formalism, using the same $6\times 6\times 6$ grids of special high-symmetry \mathbf{k}
130 points and 40 Ha kinetic energy cutoff.

131 We compute the low-spin structures of the end-member FeSiO_3 term, and then
132 dilute Fe by Mg. Eventually we look at the theoretical X-ray diffraction patterns for
133 $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$, and at the vibrational pattern of FeSiO_3 and $\text{Mg}_{0.75}\text{Fe}_{0.25}\text{SiO}_3$.

134 Complementary to calculations we obtained in situ X-ray diffraction patterns
135 on powder $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ perovskite in a diamond-anvil cell at the BL10XU of
136 SPring-8. Starting material was pyroxene with composition of $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ (Tateno
137 et al. 2007). The $\text{Fe}^{3+}/\Sigma\text{Fe}$ ratio of the starting material was not directly measured but
138 could be ~ 0.07 according to Sinmyo et al. (2011), in which pyroxene was synthesized
139 in a similar procedure. We compacted the sample powder into pellets and then coated
140 the surface of the sample with gold on both sides, which served as a laser absorber.
141 This coating technique reduces temperature gradients, and thus minimizes chemical
142 heterogeneity caused by Soret diffusion (Sinmyo and Hirose 2010). It was loaded into
143 a hole in a pre-indented rhenium gasket, together with thermal insulation layers of
144 dried NaCl. The assembly was compressed by beveled diamond anvils with 150 μm
145 culet. Angle-dispersive XRD patterns were collected on an imaging plate. A
146 monochromatic incident X-ray beam with a wavelength of 0.41351 \AA was collimated
147 to about 6- μm area (full-width of half maximum) on the sample. Two-dimensional
148 XRD images were integrated as a function of two-theta angle in order to have a
149 conventional one-dimensional diffraction profile using the IPAnalyzer (Seto et al.
150 2010). The sample was heated by a couple of fiber lasers using the double-sided
151 heating technique. Temperature was measured by fitting the thermal radiation
152 spectrum to the Planck radiation function using the spectro-radiometric method
153 (Ohishi et al. 2008). Pressure was determined from the unit-cell volume of gold, using
154 the equation of state proposed by Tsuchiya (2003). Pressure uncertainty arising from
155 the uncertainty in unit-cell volume was ± 0.6 GPa.

156 We also recorded several X-ray emission spectra on samples undergoing the
157 spin transition. XES is an established technique widely used to study high-spin to
158 low-spin crossover in transition-metal elements and their compounds. For samples
159 containing iron in a high-spin state, the emission spectra consist of $K\beta_{1,3}$ main peak at
160 7058 eV and satellite $K\beta'$ peak located at lower energy appearing as a result of the
161 exchange interaction between $3p$ core-hole and $3d$ in a final state of the emission
162 process. The iron spin crossover results in a loss of $3d$ magnetic moment, leading to a

163 disappearance of such satellite peak. High-resolution X-ray emission spectra of the Fe
164 $K\beta$ line were obtained for $Mg_{0.5}Fe_{0.5}SiO_3$ perovskite at the BL12XU beamline of
165 SPring-8 (Ozawa et al. 2011). Such measurements were carried out at room
166 temperature in a DAC at pressures between ambient pressure and 100 GPa (run #2),
167 and 93 and 98 GPa (run #3). We used beveled anvils with 200 μm culet. Starting
168 material was prepared by heating $Mg_{0.5}Fe_{0.5}SiO_3$ gel to 1273 K under the oxygen
169 fugacity slightly above the iron-wüstite buffer in run #2. In run #3, the pyroxene
170 starting material same as that in run #1 was used. The gel starting material may have
171 higher $Fe^{3+}/\Sigma Fe$ ratio of ~ 0.2 according to Ohta et al. (2008). The sample coated with
172 gold was loaded together with thermal insulation layers of dried NaCl, into an X-ray
173 translucent gasket that was composed of cubic boron nitride and beryllium. Such
174 composite gasket can increase the thickness of the sample chamber and reduce the
175 contribution from iron impurity in beryllium. The sample was compressed to high
176 pressure at 300 K and then heated by laser from both sides with scanning the laser
177 beam. The high-pressure phase was subsequently identified by XRD measurements at
178 BL10XU. Pressure was determined at room temperature based on the Raman shift of
179 the T_{2g} mode of the culet of diamond anvil (Akahama and Kawamura 2004). We used
180 monochromatic X-rays at 11 keV as an incident beam for XES measurements. It was
181 collimated to $\sim 15 \mu m$ in diameter at the sample. The pressure variations within the
182 area from which XES spectra were collected were less than 3 GPa. A 1m Rowland
183 circle type spectrometer, equipped with 1m spherical bent Ge (620) analyzer, was
184 used to acquire the emission spectra. These spectra were collected through the gasket
185 in order to avoid a loss of intensity due to absorption by the diamond. The spectra
186 covered the energy range from 7020 to 7080 eV with a resolution of 0.8 eV.

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189 Results

190 Structural implications of the spin transition

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192 In recent lattice dynamical calculations we showed that $FeSiO_3$ perovskite,
193 with orthorhombic $Pbnm$ symmetry, in both the intermediate spin and the low spin
194 state exhibits unstable phonon modes (Caracas et al. 2011). These instabilities
195 correspond to displacements of the Fe sub-lattice accompanied by a rearrangement of
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197 parts of the O sub-lattice. The spin transition is also accompanied by a density
198 difference of 4.9% at 20 GPa going down quasi-linearly to 3.8% at 120 GPa. Based
199 on the two MgSiO₃ and FeSiO₃ end-members we are now able to construct the correct
200 guess structures for the intermediate compositions along the Mg-Fe join in both high-
201 spin and in low-spin magnetic configurations. Then using these initial geometries the
202 appropriate structural relaxation eventually yields the correct final structures. For the
203 high-spin case, the orthorhombic symmetry is preserved along the whole Mg-Fe join.
204 The local replacement pattern of the two cations breaks the local symmetry in specific
205 configurations, but the large-scale average structure remains orthorhombic. The
206 starting geometry can be obtained by replacing Mg with Fe in the MgSiO₃ structure or
207 Fe with Mg in the FeSiO₃ structure, the final result being the same. For the low-spin
208 case the structure exhibits distortions that lower its symmetry down to monoclinic or
209 triclinic even for the pure FeSiO₃ end-member term. Actually, in order to be able to
210 preserve these distortions, and prevent the code from finding a local minimum, along
211 the solid solution, we need to start with the distorted low-spin structure of FeSiO₃ and
212 replace Fe with Mg in the desired amount. Indeed, starting the simulations from
213 orthorhombic FeSiO₃ fails to find the appropriate global minimum and the correct
214 low-spin crystal structures (Bengtson et al. 2009; Hsu et al. 2010a; Umemoto et al.
215 2010, Caracas et al., 2011).

216 The spin transition has thus two different structural facets: symmetry breaking
217 and volume reduction. The correlation between the spin transition and the distortions
218 of the orthorhombic structure should be reflected in changes of the symmetry of
219 certain physical properties, like the Raman and infrared spectra or the XRD patterns.
220 These changes could be observed as possible appearances of new peaks and/or
221 broadening of the existent ones. The volume and interatomic bond distance reductions
222 associated with the spin transition should be reflected in shifts of the positions of the
223 peaks in vibrational spectra and in the XRD patterns. Accurate in situ Raman/infrared
224 and XRD measurements could then capture at least some of these features. Here we
225 analyze these different changes and illustrate them for several particular terms along
226 the solid solution.

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X-ray powder diffraction pattern

230 Figures 1a and 1b show calculated X-ray diffraction patterns of the theoretical
231 structure of the middle term of the solid solution, $\text{Mg}_{0.50}\text{Fe}_{0.50}\text{SiO}_3$, in respectively
232 high-spin and low-spin configurations, computed at 90 GPa. We use one unit cell with
233 4 formula units, i.e. 4 A sites, where we replace half the A atoms with Mg and half
234 with Fe. Basic math tells us there are six possibilities, i.e. combinations of four by
235 two. In the high-spin phases, there are no distortions associated to the spin, the
236 orthorhombic symmetry is preserved (even though some forbidden reflections
237 conditions might be lifted), which makes the four crystallographic sites equivalent.
238 When we proceed with the Fe/Mg replacements, based on symmetry-determined
239 multiplicities, only three different Fe/Mg atomic configurations are needed to describe
240 the high-spin state. In the low-spin phase, there are distortions associated to the spin,
241 the orthorhombic symmetry is not preserved, and the four A sites are not equivalent
242 anymore. Eventually six configurations are needed to describe all possible Mg/Fe
243 distribution in one unit cell for the low-spin state.

244 For the high-spin magnetic cases we choose the ferromagnetic configuration,
245 as the structural differences between the ferromagnetic and the antiferromagnetic
246 orthorhombic FeSiO_3 are small enough not to be detectable in diffraction. Arrows in
247 panel *b* show the changes occurring due to the spin transition, like broadening of the
248 peaks and appearance of new weaker peaks and shoulders.

249 In a powder experiment of a solid solution crystal, each configuration would
250 contribute with its own diffraction pattern, to both peak position and intensities. For
251 our case, though within a limited configurational space, we may distinguish between
252 several such different atomic arrangements. The corresponding diffraction of each one
253 of them is represented with light red lines in Figures 1a and 1b. They are then
254 summed up to obtain the full powder diffraction spectrum, represented by the bold
255 blue lines.

256 All the major diffraction peaks preserve their position and intensity throughout
257 the spin transition. But the spectra before and after the transition are different in the
258 detail, as shown in Figures 1a and b due to symmetry breaking. Several new peaks
259 appear in the low-spin phase as a result of degeneracy lifting, and new shoulders
260 develop and grow to result in somewhat broader peaks than in the case of the high-
261 spin spectrum. This is due to the degeneracy lifting for the structural configurations.
262 In this way more individual contributions are recorded, and almost each one is but
263 slightly different from the others. This symmetry breaking lifts degeneracies between

264 diffraction peaks like the orthorhombic 110 and -110. They are the same in an
265 orthorhombic group, but different in a monoclinic or triclinic group occurring after
266 the spin transition. In the latter symmetry setup, this would result in two peaks. If the
267 non-orthorhombic distortions are small then they can overlap or appear as shoulders
268 and/or broadening of the same larger peak. Their intensity can increase too, as is the
269 case for the -110 peak at 7.4° 2θ angle and the 102 and 012 peaks above 9° 2θ angles.
270 The latter sets of peaks are invisible in MgSiO₃ perovskite, become very weakly
271 visible in the high-spin configuration because of the Mg/Fe substitution that lifts the
272 ban of this phase extinction, and have their intensity enhanced in the low-spin phase
273 because of the symmetry breaking and associated structural distortions.

274 Similarly, a set of new weak peaks appears due to lifting and changing the
275 selection rules as a result the symmetry breaking. For example the orthorhombic
276 space group bans a series of reflections, like $h00$, $0k0$, or $00l$ with h , k , or l even
277 numbers or the $0kl$ with l odd, namely peaks like 012, 102, etc. (Hahn and Arroyo
278 2006). Any distortion would then break at least some of these rules. Again, if the
279 distortion is small, the resulting new peaks are weak, but visible, and would occur
280 after the spin transition. Such peaks are the 02-1 around 11.5° 2θ angles, the 02-2
281 around 13° 2θ angles, and the 123 around 15.9° 2θ angles. They are all represented
282 with arrows in Figure 1b, and they all have their correspondence in both experiment
283 and theory.

284 The experimental XRD pattern, recorded at 89 GPa and 300 K after heating at
285 around 2000 K for 1 hour is represented in Figure 1c. Both theoretical and
286 experimental patterns use 0.414 \AA wavelength radiation. Refined experimental lattice
287 parameters are $a=4.4395\pm 0.0038 \text{ \AA}$; $b=4.6361\pm 0.0027 \text{ \AA}$; $c=6.4144\pm 0.0069 \text{ \AA}$;
288 $V=132.02\pm 0.20 \text{ \AA}^3$. These values are within 1% of the theoretical, which are
289 respectively 4.42 \AA , 4.64 \AA , and 6.39 \AA at 90 GPa. Consequently, the exact position
290 and intensity of the diffraction peaks are only slightly different between the
291 theoretical and the experimental structures because of these differences in
292 pressure/volume, and because of the precise iron amount and iron distribution. In
293 order to further decrease these discrepancies for our comparison, we use the
294 experimental lattice parameters to model the theoretical spectra in Figure 1a and 1b.

295 Apart from intensity rescaling, there is an excellent one-to-one correlation
296 between all the major peaks of the theoretical and the experimental patterns. All these

297 major peaks can be indexed using the orthorhombic setting. Stars in panel c denote
298 several weak peaks that cannot be indexed using the orthorhombic setting. This may
299 suggest that something happened to the structure, which affected the details of the
300 atomic arrangements, but not the general topology. At least two, possibly four, of
301 these peaks have a good match to the new theoretical peaks predicted to occur during
302 the spin transition. It is possible that the remaining unindexed peaks correspond to
303 particular Mg/Fe configurations that were not explicitly covered in our theoretical
304 analysis.

305 Finally we perform additional XES measurements (runs #2 and 3) to examine
306 the spin state of iron in $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ perovskite. In the first set of experiments, the
307 sample was squeezed to 93 GPa at room temperature. The sample was then heated at
308 1700-2100K for 43 minutes and the diffraction peaks from the perovskite phase
309 appeared upon heating. Subsequently, the spin state was examined by XES at 88 GPa
310 and 300 K. Next we increased the pressure to 100 GPa and obtained the X-ray
311 emission spectrum. Afterwards, the spectra were collected with decreasing pressure
312 down to ambient pressure. In the second run, we first compressed the sample to 93
313 GPa. During subsequent heating at 2000-2300 K, the peaks of perovskite appeared.
314 After heating for 110 mins, the X-ray emission spectrum was obtained at 93 GPa and
315 room temperature. Next we increased the pressure and collected the spectrum at 98
316 GPa.

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Vibrational spectra

320 A second set of alternative methods proposed here for the identification of the
321 spin transition is represented by the vibrational spectroscopies, like Raman or
322 infrared. The volume collapse at the transition induces a hardening of the phonons
323 visible as an upward shift of all modes at the transition. Furthermore, the symmetry
324 breaking participates to changing the Grüneisen parameters and also may induce
325 changes of the mode characters, which in turn can be reflected in their activation for
326 either Raman or infrared spectroscopy.

327 Figure 2 shows the correspondence of the theoretical vibration modes in the
328 zone-center between the high-spin and the low-spin phases. The calculations were
329 done for the pure FeSiO_3 end-member term and for one configuration of the
330 intermediate term $(\text{Mg}_{0.75}\text{Fe}_{0.25})\text{SiO}_3$. The frequencies of the modes in the low-spin

331 phases are represented on the horizontal axes and in the high-spin phase on the
332 vertical axes. If there were no changes at the spin transition, all the points would fall
333 on the black diagonal line. The spread and the negative shift of modes in the high-spin
334 with respect to the low-spin, suggest changes occur at the transition. These changes
335 are strong for the pure end-member term (panel a), but changes in specific peaks
336 could be visible even for large dilutions (panel b), with only 25% Fe on the Mg site.

337 For the end-member FeSiO_3 term, represented in Figure 2a there is a clear
338 shift of the majority of the phonon frequencies at the spin transition, e.g. a mode at
339 400 cm^{-1} in the high-spin phase is found in the low-spin not at 400 cm^{-1} (this would
340 fall on the black line), but above. This shift can actually be as large as 150 cm^{-1} or
341 more for some specific modes, notably the low-frequency modes dominated by the
342 displacements of the iron atoms. The shift in the higher-frequency modes, dominated
343 by O displacements is due to the modifications in the geometry and bond distances
344 within the Fe-O coordination polyhedral. The last highest-frequency modes are only
345 weakly affected by the spin transition as these modes are dominated by the breathing
346 modes of the octahedra, which do not suffer much from the magnetic collapse.

347 The same pattern is observed even for diluted systems, along the MgSiO_3 -
348 FeSiO_3 join. Figure 2b shows the shifts for a composition of 75%Mg and 25%Fe, not
349 far from realistic iron concentrations. Even here the frequencies of the phonon modes
350 exhibit a raise in wavevector number. The most significant changes are for the modes
351 whose eigenvectors are dominated by the displacements of the iron atoms; in these
352 cases the upward shifts can reach almost 100 cm^{-1} .

353 The inset in Figure 2(b) shows the pressure dependence of the frequencies for
354 several modes, whose derivatives yield the Grüneisen parameters, clearly different
355 between the high-spin and the low-spin phases.

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Discussion

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360 The occurrence of the spin transition for an intermediate term in a solid
361 solution, like MgSiO_3 - FeSiO_3 , is a continuous process over a certain pressure (and
362 temperature) range (Tsuchiya et al. 2006). Consequently, the peak shifts of both
363 diffraction and vibrational patterns associated to the volume collapse due to the spin
364 transition are smoothed out; then one would not expect to be able to detect them.

365 Similarly, other changes of the diffraction, like appearance of new peaks, broadening
366 of others, etc., are not sudden but continuous and would be hard to follow
367 experimentally. However differences should occur in the patterns before and after the
368 transition, i.e. between the high-spin and the low-spin phases, in terms of general
369 topology and/or pressure dependence.

370 The smoothness of such a transition is clearly observed in the experimentally
371 obtained X-ray emission spectra. The relatively low-pressure range of spin transition
372 of ferrous iron in Fe-rich perovskite is consistent with theoretical predictions
373 (Bengtson et al. 2008; Umemoto et al. 2008). Figure 3 shows such a spectrum
374 recorded in situ for pressures up to 100 GPa. The thermodynamic path was mostly
375 decompression. A satellite $K\beta'$ peak is clearly observed occurring with decreasing
376 pressure, a mark for the presence of a high-spin state at low pressures. At higher
377 pressure the peak does not vanish completely, when compared to the FeO low-spin
378 reference. This suggests either a non-complete spin transition, or a possibility of some
379 residual Fe^{3+} contributing to the electronic density.

380 Finally the smoothness of the transition would make any shift in Raman or
381 infrared frequencies hardly visible in an experimental measurement. However, as
382 shown in the inset of Figure 2b, the pressure derivatives, i.e. the Grüneisen
383 parameters, of some particular modes are different in the high-spin and the low-spin
384 phases. This would make them preferred tracers of the spin transition: the Grüneisen
385 parameters of these modes would have two distinct values at low and high pressures
386 with an average intermediate trend as the transition advances.

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389 Implications

390 We have shown that the spin transition of iron in Fe^{2+} -bearing MgSiO_3
391 perovskite bears specific signatures in both XRD patterns and Raman/infrared spectra.
392 Our assertions are based on results from first-principles calculations performed within
393 the density-functional theory. The volume collapse and the symmetry breaking
394 associated to this transition make that new peaks may appear in diffraction.
395 Experimental recordings of powder diffraction data on iron-rich perovskite samples
396 after the spin transition support our numerical results.

397 However, caution must be paid, because the relative low intensity of these
398 extra peaks makes them hard to observe in more Fe-diluted systems. Already for the
399 large Fe concentrations of our samples, the strong NaCl and Au reflections together
400 with a high background, make the observation and identification of such peaks a
401 difficult task (hence both solid and dashed arrows in Figure 1c). For example recent
402 single-crystal experimental data (Boffa Ballaran et al. 2012) on diluted samples (4
403 mol % Fe) did not observe the distortions seen here. Moreover some of the theoretical
404 peaks overlap with the reflections of Au, and thus cannot be observed experimentally.

405 Certain Raman and infrared modes exhibit an upward frequency shift during
406 the transition, and different Grüneisen parameters in the high-spin and the low-spin
407 phases. While this still waits for experimental confirmation, the predicted changes in
408 Grüneisen parameters could signify differences in the thermodynamical properties of
409 perovskite before and after the spin transition, differences not only due to volume, but
410 intrinsic to the structure.

411 This can have further implications for modeling the Earth's mantle, if such a
412 transition takes place at large depths. But even if the transition in perovskite occurs
413 beyond the thermodynamical limits of the Earth's mantle, our study suggests that
414 more attention should be paid to the effect of the spin transition on physical properties
415 of mantle minerals other than seismic.

416

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428 **Bibliography**

429 Akahama, Y., Kawamura, H. (2004), High-pressure Raman spectroscopy of
430 diamond anvils to 250 GPa: Method for pressure determination in the multimegabar
431 pressure range, *Journal of Applied Physics*, 96, 3748-3751.

- 432 Badro, J., Rueff, J. P., Vankó, G., Monaco, G., Fiquet, G., and Guyot, F.
433 (2004), Electronic transitions in perovskite: Possible nonconvecting layers in the
434 lower mantle, *Science*, 305, 383–386.
- 435 Baroni, S., de Gironcoli, S., Dal Corso, A., and Giannozzi, P. (2001) Phonons
436 and related crystal properties from density-functional perturbation theory. *Reviews in*
437 *Modern Physics*, 73, 515–562.
- 438 Bengtson, A., Persson, K., and Morgan, D. (2008), Ab initio study of the
439 composition dependence of the pressure induced spin cross-over in perovskite
440 ($\text{Mg}_{1-x}\text{Fe}_x$) SiO_3 , *Earth and Planetary Science Letters*, 265, 535–545.
- 441 Bengtson, A., Li, J., and Morgan, D. (2009), Mössbauer modeling to interpret
442 the spin state of iron in (Mg,Fe) SiO_3 perovskite, *Geophysical Research Letters*, 36,
443 L15301, doi:10.1029/2009GL038340.
- 444 Blochl, P.E. (1994), Projector augmented-wave method. *Physical Review B*,
445 50, 17953-17979.
- 446 Boffa Ballaran, T., Kurnosov, A., Glazyrin, K., Frost, D. J., Merlini, M.,
447 Hanfland, M., and Caracas, R. (2012) Effect of chemistry on the compressibility of
448 silicate perovskite in the lower mantle. *Earth and Planetary Science Letters* 333–334,
449 181–190.
- 450 Caracas, R. (2010), Spin and structural transitions in AlFeO_3 and
451 FeAlO_3 perovskite and post-perovskite, *Physics of the Earth and Planetary Interiors*,
452 182, 10-17.
- 453 Caracas, R. and Cohen, R.E. (2005), Effect of the chemistry of the stability
454 and elasticity of the perovskite and post-perovskite phases in the MgSiO_3 - FeSiO_3 -
455 Al_2O_3 system and implications for the lowermost mantle, *Geophysical Research*
456 *Letters*, 32, L16310, doi:10.1029/2005GL023164, 2005a.
- 457 Caracas, R., and Cohen, R.E. (2007), Effect of chemistry on the physical
458 properties of perovskite and post-perovskite, in *Post-Perovskite: The Last Mantle*
459 *Phase Transition*, *Geophysical Monography. Series 174*, doi:10.1029/174GM10.
- 460 Caracas, R., and Cohen, R.E. (2008), Ferrous iron in post-perovskite from
461 first-principles calculations, *Physics of the Earth and Planetary Interiors*, 168, 147-
462 152.
- 463 Caracas, R., Mainprice, D., and Thomas, T. (2010) Is the spin transition in
464 Fe^{2+} -bearing perovskite visible in seismology? *Geophysical Research Letters*, 37,
465 L13309, doi:10.1029/2010GL043320

- 466 Caracas, R., Ozawa, H., and Hirose, K. (2011) Identification of the Spin
467 Transition in Fe-bearing Perovskite and Implications for the Lower Mantle, Abstract
468 MR34A-08, presented at 2011 Fall Meeting AGU.
- 469 Gonze, X., Beuken, J.-M., Caracas, R., Detraux, F., Fuchs, M., Rignanese, G.-
470 M., Sindic, L., Verstraete, M., Zerah, G., Jollet, F., Torrent, M., Roy, A., Mikami, M.,
471 Ghosez, Ph., Raty, J.-Y., and Allan, D. C. (2002) First-principle computation of
472 material properties the ABINIT software project. Computational Materials Science,
473 25, 478–492. [<http://www.abinit.org>]
- 474 Gonze, X., Rignanese, G.-M., and Caracas, R. (2005) First-principles studies
475 of the lattice dynamics of crystals, and related properties. Zeitschrift für
476 Kristallographie, 220, 458–472.
- 477 Gonze, X., Amadon, B., Anglade, P.-M., Beuken, J.-M., Bottin, F., Boulanger,
478 P., Bruneval, F., Caliste, D., Caracas, R., Côté, M., Deutsch, T., Genovesi, L.,
479 Ghosez, Ph., Giantomassi, M., Goedecker, S., Hamann, D.R., Hermet, P., Jollet, F.,
480 Jomard, G., Leroux, S., Mancini, M., Mazevet, S., Oliveira, M.J.T., Onida, G.,
481 Pouillon, Y., Rangel, T., Rignanese, G.-M., Sangalli, D., Shaltaf, R., Torrent, M.,
482 Verstraete, M.J., Zerah G., and Zwanziger, J.W. (2009) ABINIT: First-principles
483 approach to material and nanosystem properties. Computer Physics Communications,
484 180, 2582-2615.
- 485 Grocholski, B., Catalli, K., Shim, S. H., and Prakapenka, V. (2012)
486 Mineralogical effects on the detectability of the post-perovskite boundary.
487 Proceedings of the National Academy of Science U.S.A., 109, 2275-2279.
- 488 Grocholski, B., Shim, S. H., Sturhahn, W., Zhao, J., Xiao, Y., and Chow, C.
489 (2009), Spin and valence states of iron in $(\text{Mg}_{0.8}\text{Fe}_{0.2})\text{SiO}_3$ perovskite, Geophysical
490 Research Letters, 36, L24303, doi:10.1029/2009GL041262.
- 491 Hahn, Th., and Arroyo, M. I. (2006) *International Tables for*
492 *Crystallography* (2006). Volume A, *Space-group symmetry*, International Union of
493 *Crystallography*.
- 494 Hirose, K., Takafuji, N., Fujino, K., Shieh, S. R., and Duffy, T. S. (2008), Iron
495 partitioning between perovskite and post- perovskite: A transmission electron
496 microscope study, American Mineralogist, 93, 1678–1681.
- 497 Hsu H., Umemoto, K., Blaha, P., and Wentzcovitch, R. (2010a), Spin states
498 and hyperfine interactions of iron in $(\text{Mg,Fe})\text{SiO}_3$ perovskite under pressure. Earth
499 and Planetary Science Letters, 294,19-26.

- 500 Hsu, H., Umemoto K., Wu, Z., and Wentzcovitch, R. M. (2010b), Spin-State
501 Crossover of Iron in Lower-Mantle Minerals: Results of DFT+U Investigations,
502 Reviews in Mineralogy and Geochemistry, 71, 169-199.
- 503 Hsu, H., Blaha, P., Cococcioni, M., and Wentzcovitch, R. M. (2011), Spin-
504 state crossover and hyperfine interactions of ferric iron in MgSiO₃ perovskite,
505 Physical Review Letters, 106, 118501.
- 506 Irifune, T., Shinmei T., McCammon, C., Miyajima, N., Rubie, D. C., and
507 Frost, D. J. (2010) Iron Partitioning and Density Changes of Pyrolite in Earth's Lower
508 Mantle, Science, 327, 193-195.
- 509 Jackson, J. M., Sturhahn, W., Shen, G., Zhao, J., Hu, M.Y., Errandonea, D.,
510 Bass, J. D., and Fei, Y. (2005), A synchrotron Mössbauer spectroscopy study of
511 (Mg,Fe)SiO₃ perovskite up to 120 GPa, American Mineralogist, 90, 199–205.
- 512 Li, J., Struzhkin, V. V., Mao, H. K., Shu, J., Hemley, R. J., Fei, Y., Mysen, B.,
513 Dera, P., Prakapenka, V., and Shen, G. (2004), Electronic spin state of iron in lower
514 mantle perovskite, Proceedings of the National Academy of Science U.S.A., 101,
515 14027–14030.
- 516 Li, J., Sturhahn, W., Jackson, J. M., Struzhkin, V. V., Lin, J. F., Zhao, J., Mao,
517 H. K., and Shen, G. (2006), Pressure effect on the electronic structure of iron in
518 (Mg,Fe)(Si,Al)O₃ perovskite: A combined synchrotron Mössbauer and X-ray
519 emission spectroscopy study up to 100 GPa, Physics and Chemistry of Minerals, 33,
520 575–585.
- 521 Li, L., J. P. Brodholt, S. Stackhouse, D. J. Weidner, M. Alfredsson, and G. D.
522 Price (2005b), Elasticity of (Mg,Fe)(Si,Al)O₃-perovskite at high pressure, Earth and
523 Planetary Science Letters, 240, 529–536.
- 524 Lin, J. F., Alp, E. E. Mao, Z. Inoue, T. McCammon, C. Xiao, Y. Chow, P. and
525 Zhao J. (2012), Electronic spin and valence states of iron in the lower-mantle silicate
526 perovskite by synchrotron Mössbauer spectroscopy, American Mineralogist, 97, 592–
527 597.
- 528 Lin, J. F., Speziale, S., Mao, Z., and Marquardt, H. (2013) Effects of the
529 electronic spin transitions of iron in lower mantle minerals: Implications for deep
530 mantle geophysics and geochemistry, Reviews in Geophysics, 51,
531 doi:10.1002/rog.20010.
- 532 Mao, W.L., Meng, Y., Shen, G., Prakapenka, V.B., Campbell, A. J., Heinz, D.
533 L., Shu, J., Caracas, R., Cohen, R.E., Fei, Y., Hemley R. J., and Mao, H.-k. (2005),

- 534 Iron-rich silicates in the Earth's D" layer. Proceedings of the National Academy of
535 Science, 102, 9751-9753.
- 536 McCammon, C., Kantor, I., Narygina, O., Rouquette, J., Ponkratz, U.,
537 Sergueev, I., Mezouar, M., Prakapenka, V. B., and Dubrovinsky, L. (2008), Stable
538 intermediate-spin ferrous iron in lower mantle perovskite, Nature Geoscience, 1, 684–
539 687.
- 540 McCammon, C., Dubrovinsky, L., Narygina, O., Kantor, I., Wu, X., Glazyrin,
541 K., Sergueev, I., and Chumakov, A. I. (2010), Low-spin Fe²⁺ in silicate perovskite
542 and a possible layer at the base of the lower mantle, Physics of the Earth and
543 Planetary Interiors, 180, 215–221.
- 544 Metsue, A., and Tsuchiya, T. (2012), Thermodynamic properties of
545 (Mg,Fe²⁺)SiO₃ perovskite at the lower-mantle pressures and temperatures: An
546 internally consistent LSDA + U study, Geophysical Journal International, 190, 310–
547 322.
- 548 Monkhorst, H. J., and Pack, J. D. (1976), Special points for Brillouin-zone
549 integrations. Physical Review B, 13, 5188-5192.
- 550 Ohishi, Y., Hirao, N., Sata, N., Hirose, K., and Takata, M. (2008), Highly
551 intense monochromatic X-ray diffraction facility for high-pressure research at SPring-
552 8, High Pressure Research, 28, 163-173.
- 553 Ozawa, H., Hirose, K., Ohta, K., Ishii, H., Hiraoka, N., Ohishi, Y., Seto, Y.
554 (2011), Spin crossover, structural change, and metallization in NiAs-type FeO at high
555 pressure, Physical Review B, 84, 134417.
- 556 Seto, Y., Nishio-Hamane, D., Nagai, T., and Sata, N. (2010), Development of
557 a software suite on X-ray diffraction experiments, The Review of High Pressure
558 Science and Technology, 20, 269-276.
- 559 Sinmyo, R., and Hirose, K. (2010), The Soret diffusion in laser-heated
560 diamond-anvil cell, Physics of the Earth and Planetary Interiors, 180, 172-178.
- 561 Stackhouse, S., Brodholt, J. P., and Price, D. G. (2007), Electronic spin
562 transitions in iron-bearing MgSiO₃ perovskite, Earth and Planetary Science Letters,
563 253, 282-290.
- 564 Tateno, S., Hirose, K., Sata, N., and Ohishi, Y. (2007), Solubility of FeO in
565 (Mg,Fe)SiO₃ perovskite and the post-perovskite phase transition, Physics of the Earth
566 and Planetary Interiors, 160, 319-325.

567 Torrent, M., Jollet, F., Bottin, F., Zerah G., and Gonze, X. (2008),
568 Implementation of the projector augmented-wave method in the ABINIT code:
569 Application to the study of iron under pressure, Computational Material Science, 42,
570 337-351.

571 Tsuchiya, T. (2003), First-principles prediction of P - V - T equation of state of
572 gold and the 660-km discontinuity in Earth's mantle, Journal of Geophysical
573 Research, 108, 2462, DOI:10.1029/2003JB002446.

574 Tsuchiya, T., Wentzcovitch, R. M., da Silva, C. R. S., de Gironcoli, S. (2006)
575 Spin transition in magnesiowuestite in Earth's lower mantle. Physical Review Letters
576 96, 198501.

577 Umemoto, K., Wentzcovitch, R. M., Yu, Y. G., and Requist, R. (2008), Spin
578 transition in (Mg,Fe)SiO₃ perovskite under pressure, Earth and Planetary Science
579 Letters, 276, 198–206.

580 Umemoto, K., Hsu, H. and Wentzcovitch, R. M. (2010), Effect of site
581 degeneracies on the spin crossovers in (Mg, Fe)SiO₃ perovskite, Physics of the Earth
582 and Planetary Interiors 180, 209–214.

583 Wolf, A. S., Caracas, R., and Asimow, P.D. (2008) Thermodynamic Phase
584 Relations in the MgO-FeO-SiO₂ System in the Lower Mantle, Eos Trans. AGU,
585 89(53), Fall Meet. Suppl., Abstract MR42a-03.

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589 Figure 1. Theoretical X-ray diffraction patterns for the $(\text{Mg}_{0.5}\text{Fe}_{0.5})\text{SiO}_3$
590 perovskite in high- (a) and low- (b) spin configurations. The symmetry breaking that
591 is associated to the spin transition induces changes in the diffraction pattern, like
592 appearance of new peaks, shoulders and peak broadening. Red arrows in (b) point to
593 such new peaks. (c) Room-temperature XRD patterns of $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ collected
594 after heating at 89 GPa. The diffraction peaks from $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ perovskite were
595 indexed on an orthorhombic $Pbnm$ cell. Pv, perovskite; Au, gold. In situ experimental
596 diffraction patterns show clearly the occurrence of such new peaks after the spin
597 transition. Red arrows in (c) point to the correspondence between the new diffraction
598 peaks predicted by calculations and the experimental peaks, which cannot be indexed
599 using a standard orthorhombic structure for perovskite: solid and dashed lines
600 correspond to certain and less certain correspondence. Theoretical diffraction patterns
601 are obtained using the experimental lattice parameters, and theoretical reduced atomic
602 positions, to account for the symmetry breaking. This minimizes the peak position
603 mismatch due to the standard discrepancies between DFT and experimental volumes
604 and pressures.

605

606 Figure 2. Correspondence between the theoretical frequencies of the vibrational mode
607 in zone center for pure FeSiO_3 perovskite (a) and for $(\text{Mg}_{0.75}\text{Fe}_{0.25})\text{SiO}_3$ perovskite
608 (b), computed at several pressures. The frequencies in the low-spin are on the
609 horizontal axis and in the high-spin are on the vertical axis. If there were no changes
610 due to the spin transition all the points would fall on the black lines, which is clearly
611 not the case. The inset in panel (b) shows the pressure (horizontal axis in GPa)
612 dependence of the frequencies (vertical axis in wavenumbers in cm^{-1}) of several
613 modes in both the high-spin (solid line) and the low-spin (dashed line) phases. Their
614 derivatives, i.e. the Grüneisen parameters, are clearly different between the two spin
615 states.

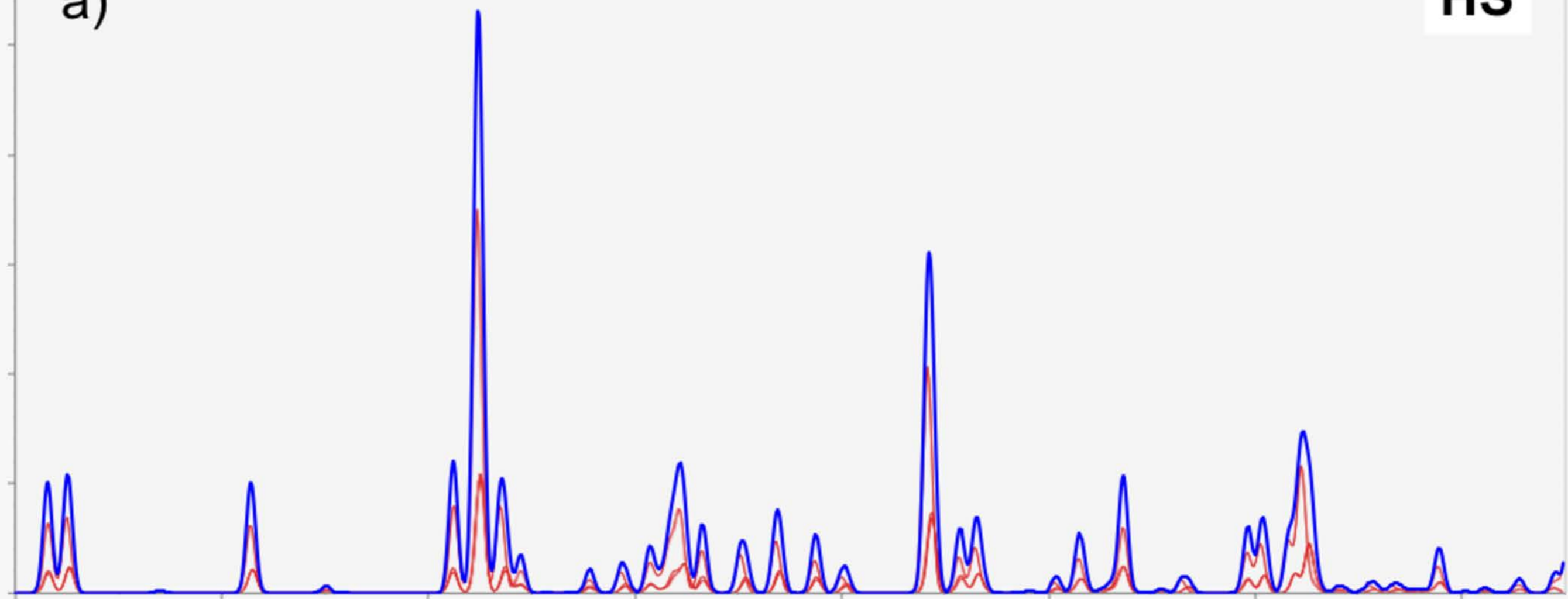
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617 Figure 3. (a) Evolution of X-ray emission spectra of $\text{Mg}_{0.5}\text{Fe}_{0.5}\text{SiO}_3$ perovskite (solid
618 squares, run #2; open squares, run #3). Measurements were conducted at 300 K. All
619 spectra are normalized to transmitted intensity, and shifted so that the weighted
620 average of main ($K\beta$) plus satellite ($K\beta'$) emission lines is set to 7058 eV. X-ray
621 emission spectrum of FeO at 146 GPa is also shown as a reference of Fe^{2+} in low-spin
622 state (Ozawa et al. 2011). (b) Normalized intensity of the $K\beta'$ peak as a function of

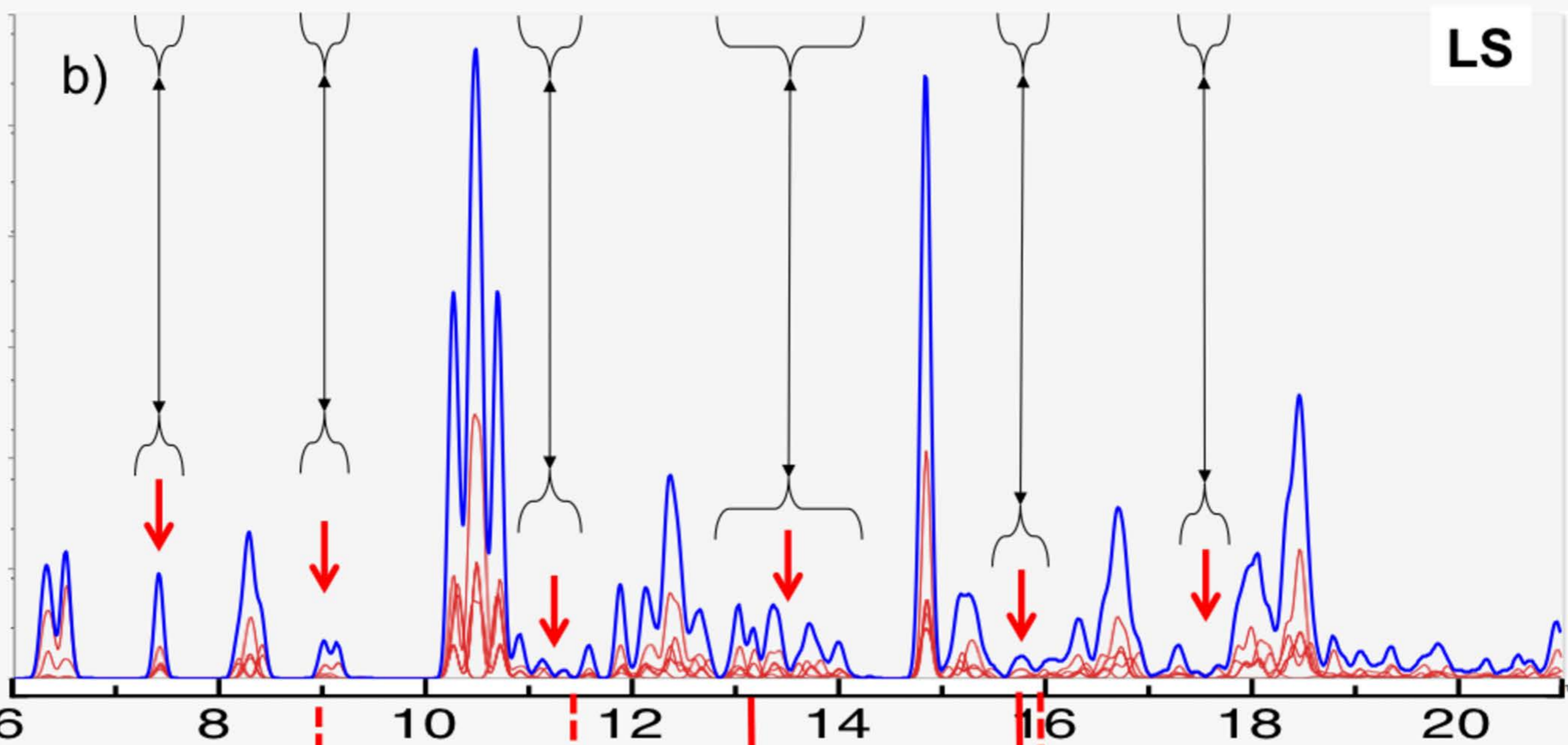
623 pressure; same symbols as (a). The intensity of the satellite peak was obtained by
624 subtracting each spectrum from the reference spectrum shown in (a). The intensity of
625 the $K\beta'$ peak is proportional to the relative abundance of the high-spin and low-spin
626 species (Badro et al. 2004). The intensity of the satellite Fe $K\beta$ peak clearly
627 diminishes with pressure, but fails to disappear completely, suggesting that some
628 residual iron did not yet undergo the spin transition.

HS

a)

**LS**

b)

**Exp.**