**REVISION 1** 1 2 Vanadio-oxy-dravite, NaV<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, a new mineral 3 4 species of the tourmaline supergroup 5 FERDINANDO BOSI<sup>1</sup>, HENRIK SKOGBY<sup>2</sup>, LEONID REZNITSKII<sup>3</sup>, AND ULF HÅLENIUS<sup>2</sup> 6 7 8 <sup>1</sup>Dipartimento di Scienze della Terra, Sapienza Università di Roma, P.le A. Moro, 5, I-00185 Rome, 9 Italy 10 <sup>2</sup>Department of Geosciences, Swedish Museum of Natural History, Box 50007, SE-10405 Stockholm, 11 Sweden 12 <sup>3</sup>Russian Academy of Science. Siberian Branch, Institute of the Earth's crust, Lermontova str., 128, 13 Irkutsk, Russia 14 15 16 17 18 ABSTRACT 19 Vanadio-oxy-dravite, NaV<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, is a new mineral of the 20 tourmaline supergroup. It is found in metaguartzites of the Pereval marble guarry (Sludyanka, 21 Lake Baikal, Russia) in association with quartz, Cr-V-bearing tremolite and mica, diopside – 22 kosmochlor - natalyite, Cr-bearing goldmanite, escolaite - karelianite, dravite - oxy-23 vanadium-dravite, V-bearing titanite and rutile, ilmenite, oxyvanite - berdesinskiite, 24 shreyerite, plagioclase, scapolite, zircon, pyrite and an unnamed oxide of V, Cr, Ti, U and Nb. 25 Crystals are green, transparent with a vitreous luster, pale green streak and conchoidal fracture. Vanadio-oxy-dravite has a Mohs hardness of approximately 71/2, and a calculated 26 density of 3.14 g/cm<sup>3</sup>. In plane polarized light, vanadio-oxy-dravite is pleochroic (O = yellow 27 green and E = pale olive green) and uniaxial negative:  $\omega = 1.693(5)$ ,  $\varepsilon = 1.673(5)$  Vanadio-28 29 oxy-dravite is rhombohedral, space group R3m, with the unit-cell parameters a = 16.0273(3), c= 7.2833(1) Å, V = 1620.24(5) Å<sup>3</sup>, Z = 3. Crystal-chemical analysis resulted in the empirical 30 structural formula: 31

 ${}^{X}(Na_{0.70}Ca_{0.23}\square_{0.05}K_{0.02})_{\Sigma1.00} \ {}^{Y}(V^{3+}_{1.39}Mg_{1.16}Al_{0.35}Fe^{3+}_{\phantom{30}0.04}Ti^{4+}_{\phantom{30}0.04}Fe^{2+}_{\phantom{20}0.02})_{\Sigma3.00}$ 

 $^{Z}$ (Al<sub>3.74</sub>Mg<sub>1.28</sub>V<sup>3+</sup><sub>0.78</sub>Cr<sup>3+</sup><sub>0.20</sub>)<sub>Σ6.00</sub>  $^{T}$ (Si<sub>6.00</sub>O<sub>18</sub>)  $^{B}$ (BO<sub>3</sub>)<sub>3</sub>  $^{V}$ (OH)<sub>3</sub>  $^{W}$ [O<sub>0.74</sub>(OH)<sub>0.26</sub>]<sub>Σ1.00</sub>. The crystal structure of vanadio-oxy-dravite was refined to an R1 index of 1.70% using 1800 unique reflections collected with Mo $K\alpha$  X-radiation. Ideally, vanadio-oxy-dravite is related to oxy-dravite and oxy-vanadium-dravite by the homovalent substitution  $V^{3+} \leftrightarrow Al^{3+}$ . Tourmaline with chemical compositions classified as vanadio-oxy-dravite can be either Aldominant or V-dominant as a result of the compositional boundaries along the solid solution between Al and  $V^{3+}$  that are determined at  $^{Y+Z}$ (V<sub>1.5</sub>Al<sub>5.5</sub>), corresponding to Na $^{Y}$ (V<sub>1.5</sub>Al<sub>1.5</sub>) $^{Z}$ (Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, and  $^{Y+Z}$ (V<sub>5</sub>Al<sub>2</sub>), corresponding to Na $^{Y}$ (V<sub>2</sub>Al<sub>2</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O.

42 43

44

45

46 47

48

49

50

5152

53

54

55

56

57

58

59

60

61

6263

64

32

33

34

35

36

3738

39

40

41

#### Introduction

The tourmaline supergroup minerals are widespread, occurring in sedimentary, igneous and metamorphic rocks (Dutrow and Henry 2011). They are important indicator minerals that can provide information on the compositional evolution of their host rocks, chiefly due to their ability to incorporate a large number of elements (e.g., Novák et al. 2004; Agrosì et al. 2006; Lussier et al. 2011a; Novák et al. 2011; van Hinsberg et al. 2011; Bačík et al. 2012). However, the chemical composition of tourmalines is also controlled by short-range and long-range constraints (e.g., Hawthorne 1996, 2002a; Bosi and Lucchesi 2007; Bosi 2010, 2011; Henry and Dutrow 2011; Skogby et al. 2012; Bosi 2013) as well as by temperature (van Hinsberg and Schumacher 2011). Tourmaline supergroup minerals are complex borosilicates and their crystal structure and crystal chemistry have been extensively studied (e.g., Foit 1989; Hawthorne and Henry 1999; Bosi and Lucchesi 2007; Lussier et al. 2008; Bosi 2008; Bosi et al. 2010; Lussier et al. 2011b; Filip et al. 2012). In accordance with Henry et al. (2011), the general formula of tourmaline may be written as:  $XY_3Z_6T_6O_{18}(BO_3)_3V_3W$ , where  $X (\equiv {}^{[9]}X) =$  $Na^{+}, K^{+}, Ca^{2+}, \Box$  (=vacancy);  $Y (\equiv {}^{[6]}Y) = Al^{3+}, Fe^{3+}, Cr^{3+}, V^{3+}, Mg^{2+}, Fe^{2+}, Mn^{2+}, Li^{+}; Z (\equiv {}^{[6]}Y) = {}^{[6]}Y = {}$  $^{[6]}Z) = \text{Al}^{3+}, \text{ Fe}^{3+}, \text{ Cr}^{3+}, \text{ V}^{3+}, \text{ Mg}^{2+}, \text{ Fe}^{2+}; \text{ T} \ (\equiv {}^{[4]}T) = \text{Si}^{4+}, \text{ Al}^{3+}, \text{ B}^{3+}; \text{ B} \ (\equiv {}^{[3]}B) = \text{B}^{3+}; \text{ W} \ (\equiv {}^{[4]}T) = \text{Si}^{4+}, \text{ Al}^{4+}, \text{ Al$  $^{[3]}O1) = OH^{1-}, F^{1-}, O^{2-}; V (\equiv ^{[3]}O3) = OH^{1-}, O^{2-}$  and where, for example, T represents a group of cations  $(Si^{4+}, Al^{3+}, B^{3+})$  accommodated at the [4]-coordinated T sites. The dominance of these ions at one or more sites of the structure gives rise to a range of distinct mineral species. Recently, several new minerals of the tourmaline supergroup were approved by the Commission on New Minerals, Nomenclature and Classification (CNMNC) of the

International Mineralogical Association (IMA). Among these are a number of oxy-tourmalines related by complete solid solution in the Al<sup>3+</sup>-Cr<sup>3+</sup>-V<sup>3+</sup> subsystem: oxy-dravite, end-member formula NaAl<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2012-004a; Bosi and Skogby 2013), oxy-chromium-dravite, NaCr<sub>3</sub>(Cr<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2011-097; Bosi et al. 2012a); oxy-vanadium-dravite, NaV<sub>3</sub>(V<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 11-E; Bosi et al. 2013a).

A new species of oxy-tourmaline, vanadio-oxy-dravite, has been approved by the IMA-CNMNC (proposal no. 2012-074). The holotype specimen (sample PR73) is deposited in the collections of the Museum of Mineralogy, Earth Sciences Department, Sapienza University of Rome, Italy, catalogue number 33068. A formal description of the new species vanadio-oxy-dravite is presented here, including a full characterization of its physical, chemical and structural properties.

### OCCURRENCE, APPEARANCE AND PHYSICAL AND OPTICAL PROPERTIES

The crystals of vanadio-oxy-dravite are green and occur in metaquartzites in the Pereval marble quarry, Sludyanka crystalline complex, Southern Baikal region, Russia (51°37'N 103°38'E). The Pereval quarry is the type locality (see Bosi et al. 2012a for a more detailed description) for natalyite, florensovite, kalininite, magnesiocoulsonite, oxy-vanadium-dravite, oxy-chromium-dravite, vanadio-oxy-chromium-dravite, chromo-alumino-povondraite, batisivite, oxyvanite and cuprokalininite. Minerals associated with the holotype specimen are: quartz, Cr-V-bearing tremolite and muscovite-celadonite, diopside – kosmochlor – natalyite, Cr-bearing goldmanite, escolaite – karelianite, dravite, V-bearing titanite and rutile, ilmenite, oxyvanite – berdesinskiite, shreyerite, plagioclase, scapolite, zircon, pyrite and an unnamed oxide of V, Cr, Ti, U and Nb, The host rocks (quartz-diopside) are Cr-V-bearing carbonate-siliceous sediments, metamorphosed to granulite facies and partly diaphtorized (Salnikova et al. 1998) to amphibolite facies (retrograde stage). Vanadio-oxy-chromium-dravite was formed in the prograde stage (granulite facies). The crystals are euhedral, reaching up to 0.3 mm in length, and may be chemically zoned (for details, see Figure 3 of Bosi et al. 2013a), but homogeneous crystals also occur.

The morphology of vanadio-oxy-dravite consists of elongated  $\{10\overline{1}0\}$  and  $\{11\overline{2}0\}$  prisms terminated by a prominent  $\{0001\}$  pedion and small, minor  $\{10\overline{1}1\}$  pyramidal faces.

Crystals are green, with pale green streak, transparent and display vitreous luster. They are brittle and show conchoidal fracture. The Mohs hardness is approximately  $7\frac{1}{2}$  (Reznitsky et al. 2001). The calculated density is 3.14 g/cm<sup>3</sup>. In transmitted light, vanadio-oxy-dravite is pleochroic with O = yellow green and E = pale olive green. Vanadio-oxy-dravite is uniaxial negative with refractive indices, measured by the immersion method using white light from a tungsten source, of  $\omega = 1.693(5)$ ,  $\varepsilon = 1.673(5)$ . The mean index of refraction, density and chemical composition lead to excellent compatibility indices  $(1 - K_p/K_c = 0.026)$  (Mandarino 1976, 1981)

107 Methods

# Single-crystal structural refinement

A representative fragment of the type specimen was selected for X-ray diffraction measurements on a Bruker KAPPA APEX-II single-crystal diffractometer, at Sapienza University of Rome (Earth Sciences Department), equipped with a CCD area detector (6.2 × 6.2 cm<sup>2</sup> active detection area,  $512 \times 512$  pixels) and a graphite crystal monochromator, using Mo $K\alpha$  radiation from a fine-focus sealed X-ray tube. The sample-to-detector distance was 4 cm. A total of ca. 3265 exposures (step = 0.2°, time/step = 20 s) covering a full reciprocal sphere with a redundancy of about 8 was used. Final unit-cell parameters were refined by means of the Bruker AXS SAINT program using reflections with  $I > 10 \, \sigma(I)$  in the range  $5^{\circ} < 20 < 73^{\circ}$ . The intensity data were processed and corrected for Lorentz, polarization, and background effects with the APEX2 software program of Bruker AXS. The data were corrected for absorption using the multi-scan method (SADABS). The absorption correction led to a significant improvement in  $R_{\rm int}$ . No violations of R3m symmetry were noted.

Structural refinement was done with the SHELXL-97 program (Sheldrick 2008). Starting coordinates were taken from Bosi et al. (2004). Variable parameters were: scale factor, extinction coefficient, atomic coordinates, site scattering values and atomic-displacement factors. To obtain the best values of statistical indexes (R1, wR2), a fully ionized scattering curve for O was used, whereas neutral scattering curves were used for the other atoms. In detail, the occupancy of the X site was modeled by using the sodium scattering factor, the Y site vanadium and magnesium scattering factors, and the Z site using vanadium and aluminum factors. The T and B sites were modeled, respectively, with Si and B scattering factors and with a fixed occupancy of 1, because refinement with unconstrained occupancies

showed no significant deviations from this value. Three full-matrix refinement cycles with isotropic-displacement parameters for all atoms were followed by anisotropic cycles until convergence was attained. No significant correlations over a value of 0.7 between the parameters were observed at the end of refinement. Table 1 lists crystal data, data-collection information and refinement details; Table 2 gives the fractional atomic coordinates and site occupancies; Table 3 gives the displacement parameters; Table 4 gives selected bond distances.

## X-ray powder diffraction

Powder X-ray data were was derived from the single-crystal structural refinement since the available sample material was not sufficient for X-ray powder diffraction measurements. Data are listed in Table 5.

## **Electron-Microprobe analysis**

Electron-microprobe analyses of the crystal used for X-ray diffraction refinement were obtained by wavelength-dispersive spectrometer (WDS mode) with a Cameca SX50 instrument at the "Istituto di Geologia Ambientale e Geoingegneria (Rome, Italy), CNR", operating at an accelerating potential of 15 kV and a sample current of 15 nA, 10 μm beam diameter. Minerals and synthetic compounds were used as standards: wollastonite (Si, Ca), magnetite (Fe), rutile (Ti), corundum (Al), vanadinite (V) fluorphlogopite (F), periclase (Mg), jadeite (Na), K-feldspar (K), sphalerite (Zn), metallic Cr, Mn and Cu. The overlap corrections and the PAP routine were applied (Pouchou and Pichoir 1991). The results (Table 6) represent mean values of 10 spot analyses. In accord with the very low concentration of Li in dravitic samples (e.g., Henry et al. 2011), the Li<sub>2</sub>O content was assumed to be insignificant. Manganese, Zn, Cu and F were below their respective detection limits (0.03 wt%) in the studied sample.

### **Infrared spectroscopy**

A homogeneous vanadio-oxy-dravite crystal was measured by Fourier transform infrared (FTIR) absorption spectroscopy to characterize (OH) absorption bands in the wavenumber range 2000-5000 cm<sup>-1</sup> using a Bruker Equinox 55 spectrometer equipped with a NIR source, a CaF<sub>2</sub> beam-splitter, a wire-grid polarizer and an InSb detector. Polarized spectra with a resolution of 4 cm<sup>-1</sup> were acquired parallel and perpendicular to the crystallographic **c**-

axis direction using a circular measurement area of 100 µm diameter on a 36 µm thick doubly-polished crystal plate that had been oriented parallel the **c**-axis by morphology and optical microscopy. As often observed, fundamental (OH) absorption bands polarized parallel to the **c**-axis direction of tourmalines are exceptionally intense, and it was not possible to thin the sample sufficiently to avoid off-scale absorption intensity for the strongest band (Fig. 1).

# **Optical absorption spectroscopy**

Polarized, room-temperature optical-absorption spectra were recorded on the same 36 μm thick crystal platelet used for the FTIR measurements. The spectra were measured in the range 270-1100 nm (37037-9091 cm<sup>-1</sup>) at a resolution of 1 nm using an AVASPEC-ULS2048X16 spectrometer attached via a 400 μm UV optical fiber to a Zeiss Axiotron UV-microscope. A 75 W Xenon arc lamp served as illuminating source and Zeiss Ultrafluar 10x lenses served as objective and condenser. A UV-quality Glan-Thompson prism with a working range from 250 to 2700 nm (40000 to 3704 cm<sup>-1</sup>) was used as polarizer. The size of the circular measure aperture was 64 μm in diameter. The wavelength scale of the spectrometer was calibrated against Ho<sub>2</sub>O<sub>3</sub> doped and Pr<sub>2</sub>O<sub>3</sub>/Nd<sub>2</sub>O<sub>3</sub> doped standards (Hellma glass filters 666F1 and 666F7) with an accuracy better than 15 cm<sup>-1</sup> in the wavelength range 300-1100 nm. Recorded spectra were fitted using the Jandel PeakFit 4.12 software assuming Gaussian peak shapes.

### RESULTS AND DISCUSSION

### **General comment**

The present crystal-structure refinement and electron-microprobe analyses were obtained from the same single crystal. However, complementary optical and spectroscopic data were recorded from coexisting crystals. Small differences in composition are likely to occur between these crystals.

# **Determination of atomic proportions**

In agreement with the structural-refinement results, the boron content was assumed to be stoichiometric in the vanadio-oxy-dravite sample ( $B^{3+} = 3.00$  apfu). Both the site-scattering results and the bond lengths of *B* and *T* are consistent with the *B* site fully occupied by boron and no amount of  $B^{3+}$  at the *T* site (e.g., Hawthorne 1996; Bosi and Lucchesi 2007). In line

with the optical absorption results, which show the occurrence of both  $Fe^{2+}$  and  $Fe^{3+}$  (see below), the Fe contents were calculated on the basis of  $Fe^{3+}/\Sigma Fe = 0.68$ , measured by Mössbauer spectroscopy (Bosi et al. 2013b) for a Fe-bearing chromo-alumino-povondraite coming from the same locality as the present sample. Note that due to the relatively low concentrations of Fe, uncertainty on its oxidation state has little influence on the overall charge calculations. The (OH) content can then be calculated by charge balance with the assumption T + Y + Z = 15.00. The atomic proportions were calculated on this assumption (Table 6). The excellent match between the number of electrons per formula unit (epfu) derived from chemical and structural analysis supports this procedure, respectively: 251.2 epfu and 251.3 epfu.

## **Site populations**

The anion site populations in the studied sample follow the general preference suggested for tourmaline (e.g., Grice and Ercit 1993; Henry et al. 2011): the O3 site (V position in the general formula) is occupied by (OH), and the O1 site (W position in the general formula) is occupied by O2- and (OH). The cation distribution at the *T*, *Y* and *Z* sites was optimized by using a quadratic program to minimize the residuals between calculated and observed data (based on the chemical and structural analysis). Site-scattering values, octahedral and tetrahedral mean bond-distances (i.e., <Y-O>, <Z-O> and <T-O>) were calculated as the linear contribution of each cation multiplied by its ideal bond-distance (Table 7). More details about the ideal distances as well as about the optimization procedure may be found in Bosi et al. (2004) and Bosi and Lucchesi (2004; 2007). The robustness of this approach was confirmed by another optimization procedure (Wright et al. 2000), which led to very similar cation distributions (Table 7). This result represents another example of convergence of these two procedures to similar solutions for tourmaline (e.g., Bosi and Lucchesi 2007; Filip et al. 2012; Bosi et al. 2012a, 2013a).

The final structural formulae are as follows:

```
222 \qquad {}^{X}(Na_{0.70}Ca_{0.23}\square_{0.05}K_{0.02})_{\Sigma1.00} \ {}^{Y}(V^{3+}_{1.39}Mg_{1.16}Al_{0.35}Fe^{3+}_{\phantom{3}0.04}Ti^{4+}_{\phantom{3}0.04}Fe^{2+}_{\phantom{2}0.02})_{\Sigma3.00}
```

- $Z(Al_{3.74}Mg_{1.28}V^{3+}_{0.78}Cr^{3+}_{0.20})_{\Sigma 6.00} {}^{T}(Si_{6.00}O_{18}) {}^{B}(BO_{3})_{3} {}^{V}(OH)_{3} {}^{W}[O_{0.74}(OH)_{0.26}]_{\Sigma 1.00}.$
- The bond-valence analysis is consistent with the optimized structural formulae. Bond-valence
- 225 calculations, using the formula and bond-valence parameters from Brown and Altermatt
- 226 (1985), are reported in Table 8.

## Name and crystal chemistry

The chemical composition of sample PR73 is consistent with tourmalines belonging to the alkali group, oxy-subgroup 3 (Henry et al. 2011). They are Na-dominant at the X site, oxygen-dominant at W with  $O^{2-} > (OH+F)^{1-}$ . As  $V^{3+}$  is the dominant cation at Y and  $Al^{3+}$  is the dominant cation at Z, its end-member composition may be represented as NaV<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O. This is in accord with the dominant-valency rule (Hatert and Burke 2008) for which the dominant ion of the dominant valency at one site becomes the basis for naming the species. As no tourmalines have yet been reported with  $V^{3+}$ - and  $Al^{3+}$ -dominant at Y and Z, respectively, this tourmaline can be classified as a new species. The closest end-member composition of a valid tourmaline species is that of oxy-dravite. The name vanadio-oxy-dravite may hence be assigned for the chemical composition, following Henry et al. (2011). The prefix *vanadio* represents the substitution of  $3V^{3+}$  for 3Al in the root composition of oxy-dravite.

Although there exists a significant degree of  $V^{3+}$ , Al and Mg disorder over the Y and Z sites, the structural formula of sample PR73 indicates a clear preference of  $V^{3+}$  for the Y site and Al for the Z site, while Mg shows only a slight preference for the Z site. Chromium, on the other hand, seems completely ordered at the Z site. The O1 site is dominated by  $O^{2-}$  with a relatively minor concentration of (OH). The presence of only minor concentrations of (OH) at O1 (~ 0.26 apfu) are consistent with both the observation of weak absorption bands at 3761 and 3732 cm<sup>-1</sup> in the infrared spectrum (Fig. 1), i.e. the area typically ascribed to the O1 site (see below), and the equation proposed by Bosi (2013) to estimate the (OH) contents in tourmaline:  ${}^{W}(OH) = [2-1.01 \cdot BVS(O1) - 0.21 - F] = 0.27$  apfu.

# **End-member formula**

Although the amount of Mg at Z is larger than that at Y (1.28 and 1.16 apfu, respectively), the formula of vanadio-oxy-dravite may also be approximated as  $Na^{Y}(V^{3+}{}_{2}Mg)^{Z}(Al_{4}MgV^{3+})(Si_{6}O_{18})(BO_{3})_{3}(OH)_{3}O$ . Such a formula could represent the structural formula for vanadio-oxy-dravite, but is not an end-member formula. In fact, this composition has multiple cations at more than one site (i.e.,  $V^{3+}$  and Mg over Y and Z), and hence is inconsistent with the characteristics of an ordered end-member formula as defined by Hawthorne (2002b). The most ordered site allocation of  $V^{3+}$ , Al and Mg consistent with the empirical formula and the end-member definition is  $Na^{Y}(V_{3})^{Z}(Al_{4}Mg_{2})(Si_{6}O_{18})(BO_{3})_{3}(OH)_{3}O$ .

Note that the difference between these two formulae is solely in  $V^{3+}$ -Mg order-disorder; no difference occurs in chemical composition.

# **Infrared spectroscopy**

Spectra recorded in polarized mode perpendicular and parallel to the crystallographic **c** axis show an intense broad band around 3550 cm<sup>-1</sup> and two weaker bands at 3732 and 3761 cm<sup>-1</sup>, all strongly polarized in the **c** direction (Fig. 1). As it was not possible to thin the sample sufficiently to avoid off-scale absorption for the main band, any possible fine structure cannot be discerned. However, in line with previous studies (Bosi et al. 2012b; Bosi et al. 2013b), the band can be related to the local arrangement ( ${}^{Y}V^{3+}$   ${}^{Z}R$   ${}^{Z}R$ )-O3, i.e., to the occurrence of (OH) at the V position of the tourmaline general formula (O3 site in the structure). The two weaker bands at ca. 3732 and 3761 cm<sup>-1</sup> are consistent with the minor concentrations of (OH) (ca. 0.3 apfu) assigned to the W position (O1 site in the structure), and may be related to the local arrangements  ${}^{Y}(Mg Mg Mg)$  and  ${}^{Y}(Mg Mg R)$  (cf. Gonzalez-Carreño et al. 1988; Bosi et al. 2012b; Bosi et al. 2013b).

# **Optical-absorption spectroscopy**

The optical spectra of vanadio-oxy-dravite show two broad absorption bands at ca. 432 and 609 nm, superimposed on an intense UV absorption edge (Fig. 3). These absorption bands are ascribed to spin-allowed d-d transitions in octahedrally coordinated  $V^{3+}$ , in accord with similar bands observed for vanadio-oxy-chromium-dravite (Bosi et al. 2012c) at ca. 440 and 610 nm. In addition to the dominant  $V^{3+}$ -bands, two broad but weak absorption bands at ca. 765 and 1100 nm are observed in the  $\mathbf{E}||\mathbf{a}$ -spectrum. They are in accord with previous studies on the optical spectra of tourmaline (e.g., Smith 1978; Taran et al. 1993; Mattson and Rossman 1987) and are ascribed to electronic transitions in  $Fe^{2+}$ - $Fe^{3+}$  pairs. Consequently, the appearance of these absorption bands suggests that a small fraction of the total iron in the present sample occurs in the divalent state.

### COMPOSITIONAL BOUNDARIES OF VANADIO-OXY-DRAVITE

The plot of the Z- and Y-site cations in the ternary diagram for the Al-Cr- $V^{3+}$  subsystem show, of course, that vanadio-oxy-dravite is  ${}^{Z}Al$ -dominant and  ${}^{Y}V^{3+}$ -dominant.

More interesting, however, is the triangular plot in terms of Al-Cr-V<sup>3+</sup> at *Y* and *Z*, showing that vanadio-oxy-dravite can be either Al-dominant or V<sup>3+</sup>-dominant (Fig. 3). This latter plot type displays the occurrence of three end-members along the full solid solution between the Al<sup>3+</sup> and V<sup>3+</sup> apices: oxy-dravite, vanadio-oxy-dravite and oxy-vanadium-dravite. These end-members are related by the substitution V<sup>3+</sup>  $\leftrightarrow$  Al<sup>3+</sup> at the Y position (vanadio-oxy-dravite  $\leftrightarrow$  oxy-dravite) and V<sup>3+</sup>  $\leftrightarrow$  Al<sup>3+</sup> at the Z position (oxy-vanadium-dravite  $\leftrightarrow$  vanadio-oxy-dravite), while their compositional boundaries are at: (1) Y+Z(V<sub>5</sub>Al<sub>2</sub>), corresponding to Na<sup>Y</sup>(V<sub>3</sub>)<sup>Z</sup>(V<sub>2</sub>Al<sub>2</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O; (2) Y+Z(V<sub>1.5</sub>Al<sub>5.5</sub>), corresponding to Na<sup>Y</sup>(V<sub>1.5</sub>Al<sub>1.5</sub>)<sup>Z</sup>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O. Consequently, oxy-dravite is characterized by V<sup>3+</sup> contents less than 1.5 apfu, vanadio-oxy-dravite is characterized by V<sup>3+</sup> contents between 5 and 1.5 apfu, and oxy-vanadium-dravite is characterized by V<sup>3+</sup> contents larger than 5 apfu.

The discovery of the new mineral vanadio-oxy-dravite provides new information on the crystal chemistry of the tourmaline supergroup. The current chemical data supports complete exchange of  $V^{3+}$ ,  $Cr^{3+}$  and Al in species of the tourmaline supergroup (Reznitsky et al. 2001; Bosi et al. 2004, 2013a,b), and it shows that vanadio-oxy-dravite can be either  $V^{3+}$ -dominant or Al-dominant.

309310

311

312

313

314

315

316

293

294

295

296

297

298

299 300

301

302

303

304305

306

307308

### ACKNOWLEDGMENTS

Chemical analyses were done with the kind assistance of M. Serracino to whom the authors express their gratitude. L. Reznitskii was supported by a grant from the Russian Foundation for Basic Research (project 13-05-00258). We thank the reviewers Darrel Henry and Frank C. Hawthorne for useful suggestions that improved the manuscript. The manuscript handling by Daniel Harlov is acknowledged.

317318

319

### REFERENCES CITED

Agrosì, G., Bosi, F., Lucchesi, S., Melchiorre, G., and Scandale, E. (2006) Mn-tourmaline crystals from island of Elba (Italy): Growth history and growth marks. American Mineralogist, 91, 944-952.

Bačík P., Méres Š., Uher P. (2011) Vanadium-bearing tourmaline in metacherts from Chvojnica, Slovak Republic: crystal chemistry and multistage evolution. Canadian Mineralogist, 49, 195-206.

- 326 Bačík, P., Uher, P., Ertl, A., Jonsson, E., Nysten, P., Kanický, V., and Vaculovič, T. (2012)
- Zoned REE Enriched Dravite from a Granitic Pegmatite in Forshammar Bergslagen
- Province, Sweden an EMPA, XRD and LA-ICP-MS study. Canadian Mineralogist. 50,
- 329 825-841.
- Bosi, F. (2008) Disordering of Fe<sup>2+</sup> over octahedrally coordinated sites of tourmaline.
- 331 American Mineralogist, 93, 1647-1653.
- Bosi, F. (2010) Octahedrally coordinated vacancies in tourmaline: a theoretical approach.
- 333 Mineralogical Magazine, 74, 1037-1044.
- Bosi, F. (2011) Stereochemical constraints in tourmaline: from a short-range to a long-range
- structure. Canadian Mineralogist, 49, 17-27.
- Bosi, F. (2013) Bond-valence constraints around the O1 site of tourmaline. Mineralogical
- 337 Magazine, 77, 343-351.
- 338 Bosi, F. and Lucchesi, S. (2004) Crystal chemistry of the schorl-dravite series. European
- 339 Journal of Mineralogy, 16, 335-344.
- Bosi, F. and Lucchesi, S. (2007) Crystal chemical relationships in the tourmaline group:
- structural constraints on chemical variability. American Mineralogist, 92, 1054-1063.
- 342 Bosi, F. and Skogby, H. (2013) Oxy-dravite, Na(Al<sub>2</sub>Mg)(Al<sub>5</sub>Mg)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, a new
- mineral species of the tourmaline supergroup. American Mineralogist, 98,
- 344 dx.doi.org/10.2138/am.2013.4441.
- 345 Bosi F., Lucchesi, S., and Reznitskii, L. (2004) Crystal chemistry of the dravite-chromdravite
- series. European Journal of Mineralogy, 16, 345-352.
- Bosi, F., Balić-Žunić, T., and Surour, A.A., (2010) Crystal structure analysis of four
- tourmalines from the Cleopatra's Mines (Egypt) and Jabal Zalm (Saudi Arabia), and the
- role of Al in the tourmaline group. American Mineralogist, 95, 510-518.
- 350 Bosi, F., Reznitskii, L., and Skogby, H. (2012a) Oxy-chromium-dravite,
- NaCr<sub>3</sub>(Cr<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, a new mineral species of the tourmaline
- supergroup. American Mineralogist, 97, 2024-2030.
- 353 Bosi, F., Skogby, H., Agrosì, G., and Scandale, E. (2012b) Tsilaisite,
- NaMn<sub>3</sub>Al<sub>6</sub>(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>OH, a new mineral species of the tourmaline supergroup
- from Grotta d'Oggi, San Pietro in Campo, island of Elba, Italy. American Mineralogist,
- 356 97, 989**-**994.

- 357 Bosi,, F., Reznitskii, L., Skogby, H., and Hålenius, U. (2012c) Vanadio-oxy-chromium-dravite
- 358 IMA 2012-034. CNMNC Newsletter No. 14, October 2012, page 1286; Mineralogical
- 359 Magazine, 76, 1281-1288.
- 360 Bosi, F., Reznitskii, L., and Sklyarov, E.V. (2013a) Oxy-vanadium-dravite,
- Na $V_3(V_4Mg_2)(Si_6O_{18})(BO_3)_3(OH)_3O$ : crystal structure and redefinition of the
- 362 "vanadium-dravite" tourmaline. American Mineralogist. 98, 501-505.
- Bosi, F., Skogby, H., Hålenius, U., and Reznitskii, L. (2013b) Crystallographic and
- 364 spectroscopic characterization of Fe-bearing chromo-alumino-povondraite and its
- relations with oxy-chromium-dravite and oxy-dravite. American Mineralogist, 98,
- 366 dx.doi.org/10.2138/am.2013.4447.
- Brown, I.D. and Altermatt, D. (1985) Bond-valence parameters obtained from a systematic
- analysis of the Inorganic Crystal Structure Database. Acta Crystallographica, B41, 244–
- 369 247.
- Dutrow, B.L. and Henry, D.J. (2011) Tourmaline: A geologic DVD. Elements, 7(5), 301-306.
- Filip, J., Bosi, F., Novák, M., Skogby, H., Tuček, J., Čuda, J., and Wildner, M. (2012) Redox
- processes of iron in the tourmaline structure: example of the high-temperature treatment
- of Fe<sup>3+</sup>-rich schorl. Geochimica et Cosmochimica Acta, 86, 239-256.
- Foit, F.F. Jr. (1989) Crystal chemistry of alkali-deficient schorl and tourmaline structural
- 375 relationships. American Mineralogist, 74, 422-431.
- 376 Gonzalez-Carreño, T., Fernandez, M. and Sanz, J. (1988) Infrared and electron micropobe
- analysis in tourmalines. Physics and Chemistry of Minerals, 15, 452-460.
- 378 Grice, J.D. and Ercit, T.S. (1993) Ordering of Fe and Mg in the tourmaline crystal structure:
- the correct formula. Neues Jahrbuch für Mineralogie, Abhandlungen, 165, 245-266.
- Hatert, F. and Burke, E.A.J. (2008) The IMA-CNMNC dominant-constituent rule revisited and
- extended. Canadian Mineralogist, 46, 717-728.
- 382 Hawthorne, F.C. (1996) Structural mechanisms for light-element variations in tourmaline.
- Canadian Mineralogist, 34, 123-132.
- 384 Hawthorne, F. C. (2002a) Bond-valence constraints on the chemical composition of
- tourmaline. Canadian Mineralogist, 40, 789-797.
- 386 Hawthorne, F.C. (2002b) The use of end-member charge-arrangements in defining new
- 387 mineral species and heterovalent substitutions in complex minerals. Canadian
- 388 Mineralogist, 40, 699-710

- Hawthorne, F.C. and Henry, D. (1999) Classification of the minerals of the tourmaline group.
- European Journal of Mineralogy, 11, 201-215.
- 391 Henry, D.J. and Dutrow, B.L. (2011) The incorporation of fluorine in tourmaline: Internal
- 392 crystallographic controls or external environmental influences? Canadian Mineralogist,
- 393 49, 41-56.
- Henry, D.J., Novák, M., Hawthorne, F.C., Ertl, A., Dutrow, B., Uher, P., and Pezzotta, F.
- 395 (2011) Nomenclature of the tourmaline supergroup minerals. American Mineralogist, 96,
- 396 895-913.
- 397 Lussier, A.J., Aguiar, P.M., Michaelis, V.K., Kroeker, S., Herwig, S., Abdu, Y., and
- Hawthorne, F.C. (2008) Mushroom elbaite from the Kat Chay mine, Momeik, near
- Mogok, Myanmar: I. Crystal chemistry by SREF, EMPA, MAS NMR and Mössbauer
- spectroscopy. Mineralogical Magazine, 72, 747-761.
- 401 Lussier, A.J., Hawthorne, F.C., Aguiar, P.M., Michaelis, V.K., and Kroeker, S. (2011a)
- Elbaite-liddicoatite from Black Rapids glacier, Alaska. Periodico di Mineralogia, 80, 57-
- 403 73.
- 404 Lussier, A.J., Abdu, Y. Hawthorne, F.C., Michaelis, V.K., Aguiar, P.M., and Kroeker, S.
- 405 (2011b) Oscillatory zoned liddicoatite from Anjanabonoina, central Madagascar. I.
- 406 Crystal chemistry and structure by SREF and <sup>11</sup>B and <sup>27</sup>Al MAS NMR spectroscopy.
- 407 Canadian Mineralogist, 49, 63-88.
- 408 Mandarino, J.A. (1976) The Gladstone-Dale relationship. Part I: derivation of new constants.
- 409 Canadian Mineralogist, 14, 498-502.
- 410 Mandarino, J.A. (1981) The Gladstone-Dale relationship. Part IV: the compatibility concept
- and its application. Canadian Mineralogist, 19, 441-450.
- 412 Mattson, S.M. and Rossman, G.R. (1987) Fe<sup>2+</sup>-Fe<sup>3+</sup> interactions in tourmaline. Physics and
- 413 Chemistry of Minerals, 14, 163-171.
- 414 Novák, M., Povondra, P., and Selway, J.B. (2004) Schorl-oxy-schorl to dravite-oxy-dravite
- 415 tourmaline from granitic pegmatites; examples from the Moldanubicum, Czech
- 416 Republic. European Journal of Mineralogy, 16, 323-333.
- 417 Novák M., Škoda P., Filip J., Macek I., and Vaculovič T. (2011) Compositional trends in
- 418 tourmaline from intragranitic NYF pegmatites of the Třebíč Pluton, Czech Republic;
- 419 electron microprobe, Mössbauer and LA-ICP-MS study. Canadian Mineralogist, 49,
- 420 359-380.

- 421 Pouchou, J.L. and Pichoir, F. (1991) Quantitative analysis of homogeneous or stratified
- microvolumes applying the model "PAP." In K.F.J. Heinrich and D.E. Newbury, Eds.,
- Electron Probe Quantitation, p. 31–75. Plenum, New York.
- 424 Reznitsky, L.Z., Sklyarov, E.V., Ushchapovskaya, Z.V., Nartova, N.V., Kashaev, A.A.,
- Karmanov, N.S., Kanakin, S.V., Smolin, A.S., and Nekrosova, E.A. (2001)
- Vanadiumdravite, NaMg<sub>3</sub>V<sub>6</sub>[Si<sub>6</sub>O<sub>18</sub>][BO<sub>3</sub>]<sub>3</sub>(OH)<sub>4</sub>, a new mineral of the tourmaline
- 427 group. Zapiski Vsesoyuznogo Mineralogicheskogo Obshchestva, 130, 59-72 (in
- 428 Russian).
- 429 Salnikova, E.B., Sergeev, S.A., Kotov, A.B., Yakovleva, S.Z., Steiger, R.H., Reznitskiy, L.Z.,
- and Vasil'ev, E.P. (1998) U-Pb zircon dating of granulite metamorphism in the
- 431 Slyudyanskiy complex, Eastern Siberia. Gondwana Research, 1, 195–205.
- Sheldrick, G.M. (2008) A short history of SHELX. Acta Crystallographica, A64, 112-122.
- Smith, G. (1978) A reassessment of the role of iron in the 5,000-30,000 cm<sup>-1</sup> region of the
- 434 electronic absorption spectra of tourmaline. Physics and Chemistry of Minerals, 3, 343-
- 435 373.

447

- 436 Skogby, H., Bosi, F., and Lazor, P. (2012) Short-range order in tourmaline: a vibrational
- spectroscopic approach to elbaite. Physics and Chemistry of Minerals, 39, 811-816.
- 438 Taran, M.N., Lebedev, A.S., and Platonov, A.N. (1993) Optical absorption spectroscopy of
- 439 synthetic tourmalines. Physics and Chemistry of Minerals, 20, 209-220.
- 440 van Hinsberg, V.J. and Schumacher, J.C. (2011) Tourmaline as a petrogenetic indicator
- 441 mineral in the Haut-Allier metamorphic suite, Massif Central, France. Canadian
- 442 Mineralogist, 49, 177-194.
- van Hinsberg, V.J., Henry, D.J., and Marschall, H.R. (2011) Tourmaline: an ideal indicator of
- its host environment. Canadian Mineralogist, 49, 1-16.
- Wright, S.E., Foley, J.A., and, Hughes, J.M. (2000) Optimization of site occupancies in
- minerals using quadratic programming. American Mineralogist, 85, 524-531.

448	
449	LIST OF TABLES
450	TABLE 1. Single-crystal X-ray diffraction data details for vanadio-oxy-dravite.
451	TABLE 2. Fractional atom coordinates and site occupancy for vanadio-oxy-dravite.
452	<b>TABLE 3.</b> Displacement parameters $(\mathring{A}^2)$ for vanadio-oxy-dravite.
453	TABLE 4. Selected bond distances (Å) for vanadio-oxy-dravite.
454	TABLE 5. Simulated X-ray powder diffraction data for vanadio-oxy-dravite.
455	TABLE 6. Chemical composition of vanadio-oxy-dravite.
456	TABLE 7. Cation site populations (apfu), mean atomic number and mean bond distances (Å)
457	for vanadio-oxy-dravite.
458	TABLE 8. Bond valence calculations (valence unit) for vanadio-oxy-dravite.
459	TABLE 9. Comparative data for oxy-dravite, vanadio-oxy-dravite and oxy-vanadium-dravite.
460	
461	
462	LIST OF FIGURES AND FIGURE CAPTIONS
463	FIGURE 1. Polarized FTIR absorption spectra in the (OH)-stretching region of vanadio-oxy-
464	dravite, vertically offset for clarity. Sample thickness 36 µm. The main band
465	around 3550 cm $^{-1}$ is truncated in the <b>c</b> direction due to excessive absorption.
466	<b>FIGURE 2.</b> Polarized electronic absorption spectra for vanadio-oxy-dravite (sample PR73).
467	<b>FIGURE 3.</b> Ternary diagram in terms of Al-V-Cr at the $Y + Z$ sites for oxy-tourmalines. Black
468	circle: present sample. Open circles: V <sup>3+</sup> -bearing oxy-dravite samples with V <sup>3+</sup> >
469	0.20 apfu (Bačík et al. 2011); and sample N825, oxy-vanadium-dravite (Bosi et al.
470	2013a); black crosses = vanadio-oxy-dravite samples from Sludyanka (Reznitsky et
471	al. 2001).
472 473	

Table 1. Single-crystal X-ray diffraction data: details for vanadio-oxy-dravite

Table 1. Single-crystal X-ray diffraction data: details for vanadio-oxy-dravite				
Sample	PR73			
Crystal size (mm)	$0.10\times0.16\times0.24$			
a (Å)	16.0273(3)			
c (Å)	7.2833(1)			
$V(\mathring{A}^3)$	1620.24(5)			
Range for data collection, 2θ (°)	5 - 73			
Reciprocal space range hkl	–21 ≤ <i>h</i> ≤ 25			
	$-23 \le k \le 26$			
	<b>–</b> 12 ≤ <i>I</i> ≤ 11			
Total number of frames	3265			
Set of measured reflections	7988			
Unique reflections, R <sub>int</sub> (%)	1800, 2.10			
Redundancy	8			
Absorption correction method	SADABS			
Refinement method	Full-matrix last-squares on F <sup>2</sup>			
Structural refinement program	SHELXL-97			
Extinction coefficient	0.00007(1)			
Flack parameter	0.06(2)			
wR2 (%)	3.81			
R1 (%) all data	1.70			
R1 (%) for $I > 2\sigma(I)$	1.65			
GooF	1.060			
Largest diff. peak and hole (±e <sup>-</sup> /Å <sup>3</sup> )	0.40 and -0.30			

*Notes*:  $R_{\rm int}$  = merging residual value; R1 = discrepancy index, calculated from F-data; wR2 = weighted discrepancy index, calculated from  $F^2$ -data; GooF = goodness of fit; Diff. Peaks = maximum and minimum residual electron density. Radiation, Mo $K\alpha$  = 0.71073 Å. Data collection temperature = 293 K. Space group R3m; Z = 3.

TABLE 2. Fractional atom coordinates and site occupancy for vanadio-oxy-dravite

Site	X	у	Z	Site occupancy
X	0	0	0.22772(15)	Na <sub>1.137(6)</sub>
Υ	0.12347(2)	0.061737(12)	0.63625(6)	$V_{0.513(4)}\ Mg_{0.487(4)}$
Ζ	0.29820(2)	0.26190(2)	0.61148(6)	$AI_{0.852(3)}V_{0.148(3)}$
В	0.11004(6)	0.22007(12)	0.4539(2)	B <sub>1.00</sub>
T	0.191020(18)	0.189344(19)	0.000	Si <sub>1.00</sub>
O1 (≡W)	0	0	0.7691(3)	O <sub>1.00</sub>
O2	0.06078(4)	0.12157(8)	0.48340(16)	O <sub>1.00</sub>
O3 (≡V)	0.26119(9)	0.13059(4)	0.51120(17)	O <sub>1.00</sub>
O4	0.09258(4)	0.18516(8)	0.07085(16)	O <sub>1.00</sub>
O5	0.18236(8)	0.09118(4)	0.09011(16)	O <sub>1.00</sub>
O6	0.19403(5)	0.18472(5)	0.77971(12)	O <sub>1.00</sub>
07	0.28367(5)	0.28332(5)	0.07786(11)	O <sub>1.00</sub>
O8	0.20860(5)	0.26946(6)	0.44063(12)	O <sub>1.00</sub>
H3	0.2565(16)	0.1283(8)	0.393(3)	H <sub>1.00</sub>

TABLE 3. Displacement parameters (Å<sup>2</sup>) for vanadio-oxy-dravite

Site	U <sup>11</sup>	U <sup>22</sup>	$U^{33}$	U <sup>23</sup>	U <sup>13</sup>	$U^{12}$	$U_{\rm eq}/U_{\rm iso}{}^*$
Χ	0.0149(4)	0.0149(4)	0.0158(5)	0.000	0	0.00743(18)	0.0152(3)
Y	0.00704(16)	0.00604(13)	0.00951(17)	-0.00095(5)	-0.00189(10)	0.00352(8)	0.00742(10)
Z	0.00619(13)	0.00663(13)	0.00729(12)	0.00037(9)	-0.00002(9)	0.00314(10)	0.00673(8)
В	0.0075(4)	0.0083(6)	0.0088(6)	0.0008(5)	0.0004(2)	0.0041(3)	0.0081(3)
T	0.00521(12)	0.00494(11)	0.00708(12)	-0.00055(9)	-0.00045(9)	0.00252(9)	0.00575(6)
01	0.0104(5)	0.0104(5)	0.0077(8)	0.000	0	0.0052(2)	0.0095(3)
02	0.0088(3)	0.0052(4)	0.0120(5)	0.0014(3)	0.00071(17)	0.0026(2)	0.0091(2)
O3	0.0179(5)	0.0139(4)	0.0061(5)	0.00053(19)	0.0011(4)	0.0089(3)	0.0122(2)
04	0.0084(3)	0.0162(5)	0.0094(5)	-0.0009(4)	-0.00045(19)	0.0081(3)	0.0104(2)
O5	0.0158(5)	0.0078(3)	0.0097(5)	0.00071(19)	0.0014(4)	0.0079(3)	0.01019(19)
06	0.0102(3)	0.0085(3)	0.0068(3)	-0.0014(2)	-0.0008(2)	0.0053(3)	0.00825(13)
07	0.0079(3)	0.0069(3)	0.0101(3)	-0.0008(2)	-0.0028(3)	0.0014(3)	0.00929(14)
08	0.0045(3)	0.0097(3)	0.0178(4)	0.0025(3)	0.0009(3)	0.0029(3)	0.01094(15)
H3							0.018*

Notes: Equivalent ( $U_{eq}$ ) and isotropic ( $U_{iso}$ ) displacement parameters; H-atom was constrained to have a  $U_{iso}$  1.5 times the  $U_{eq}$  value of the O3 oxygen.

TABLE 4. Selected bond distances (Å) for vanadio-oxv-dravite

TABLE 4. OCICCI	TABLE 4. Gelected Bond distances (A) for Variatio-oxy-dravite					
<i>B</i> -O2	1.384(2)	Y-O1	1.9682(10)			
<i>B</i> -O8 <sup>A</sup> (× 2)	1.3715(11)	Y-O2 <sup>B</sup> (× 2)	2.0327(8)			
<b-o></b-o>	1.376	Y-O3	2.1174(12)			
		Y-O6 <sup>C</sup> (× 2)	2.0066(8)			
T-O4	1.6292(5)	<y-o></y-o>	2.027			
T-O5	1.6453(5)					
T-07	1.5995(7)	Z-O3	2.0164(6)			
<i>T</i> *-O6	1.6080(9)	<i>Z</i> -O6	1.9374(8)			
< <i>T</i> -O>	1.620	<i>Z</i> -O8 <sup>E</sup>	1.9178(8)			
		<i>Z</i> -O7 <sup>E</sup>	1.9349(8)			
X-O2 <sup>B,F</sup> (× 3)	2.5129(13)	<i>Z</i> -O7 <sup>D</sup>	1.9836(8)			
X-O4 <sup>B,F</sup> (× 3)	2.8126(13)	Z-08	1.9492(8)			
X-O5 <sup>B,F</sup> (× 3)	2.7223(12)	<z-o></z-o>	1.957			
<x-o></x-o>	2.683	О3-Н	0.86(2)			

*Notes*: Standard uncertainty in parentheses. Superscript letters: A = (y - x, y, z); B = (y - x, -x, z); C = (x, x - y, z); D = (y - x + 1/3, -x + 2/3, z + 2/3); E = (-y + 2/3, x - y + 1/3, z + 1/3); F = (-y, x - y, z). Transformations relate coordinates to those of Table 2.

<sup>\*</sup> Positioned in adjacent unit cell.

**TABLE 5**. Simulated X-ray powder diffraction data for vanadio-oxy-dravite ( $\lambda = 1.5418 \text{ Å}$ )

tor vanadio-oxy-di	for vanadio-oxy-dravite ( $\Lambda = 1.5418 \text{ A}$ )					
$I_{\rm calc} > 9 \ (\%)$	d <sub>calc</sub> (Å)	h	k	1		
37	6.4467	1	0	1		
19	5.0153	0	2	1		
19	4.6185	3	0	0		
52	4.2606	2	1	1		
66	4.0041	2	2	0		
47	3.5221	0	1	2		
10	3.0317	4	1	0		
67	2.9928	1	2	2		
100	2.5958	0	5	1		
15	2.3970	2	3	2		
15	2.3594	5	1	1		
13	2.2091	5	0	2		
12	2.1500	0	3	3		
11	2.0763	2	2	3		
43	2.0573	1	5	2		
28	1.9340	3	4	2		
20	1.6756	0	6	3		
13	1.6520	2	7	1		
17	1.6025	5	5	0		
15	1.5226	0	5	4		
15	1.4704	5	1	4		
10	1.4232	4	3	4		

*Notes*: I = calculated intensity, d = calculated interplanar spacing; hkl = reflection indices.

TABLE 6. Chemical composition of vanadio-oxy-dravite

	·	,	
	Weight %		apfu
SiO <sub>2</sub>	35.34(17)	Si	6.00(3)
TiO <sub>2</sub>	0.29(4)	Ti <sup>4+</sup>	0.037(5)
$B_2O_3^*$	10.23	В	3.00
$Al_2O_3$	20.36(37)	Al	4.08(6)
$Cr_2O_3$	1.48(28)	Cr <sup>3+</sup>	0.20(4)
$V_2O_3$	15.97(39)	V <sup>3+</sup>	2.18(5)
$Fe_2O_3^{\dagger}$	0.34	Fe <sup>3+</sup>	0.044(4)
FeO <sup>†</sup>	0.15	Fe <sup>2+</sup>	0.021(3)
MgO	9.65(10)	Mg	2.44(3)
CaO	1.24(7)	Ca	0.23(1)
Na <sub>2</sub> O	2.11(6)	Na	0.70(2)
$K_2O$	0.09(1)	K	0.019(2)
H <sub>2</sub> O*	2.86	ОН	3.26
Total	100.12		

<sup>\*</sup> Calculated by stoichiometry.

*Notes*: Errors for oxides are standard deviations (in brackets) of 10 spot analyses. Standard errors for the atomic proportions (in brackets) were calculated by error-propagation theory; apfu = atoms per formula unit.

<sup>†</sup> Calculated as  $Fe^{3+}/\Sigma Fe$  ratio = 0.68 (see text); measured FeO = 0.45(4).

TABLE 7. Cation site populations (apfu), mean atomic numbers and mean bond lengths (Å) for vanadio-oxy-dravite

Site	Site population	Mean ator	Mean atomic number		Mean bond length	
Site	Site population	refined	calculated	refined	calculated*	
X	0.70 Na + 0.23 Ca + 0.05 □ + 0.02 K	12.51(8)	12.52			
Υ	$1.39 \text{ V}^{3+} + 1.16 \text{ Mg} + 0.35 \text{ Al} + 0.04 \text{ Fe}^{3+} + 0.04 \text{ Ti}^{4+} + 0.02 \text{ Fe}^{2+}$	17.64(7)	17.65	2.027	2.032	
	$(1.39 \text{ V}^{3+} + 1.20 \text{ Mg} + 0.30 \text{ Al} + 0.04 \text{ Fe}^{3+} + 0.04 \text{ Ti}^{4+} + 0.02 \text{ Fe}^{2+})^{\dagger}$					
Z	3.73 AI + 1.28 Mg + 0.78 V <sup>3+</sup> + 0.20 Cr <sup>3+</sup>	14.48(4)	14.46	1.957	1.955	
	$(3.77 \text{ Al} + 1.24 \text{ Mg} + 0.79 \text{ V}^{3+} + 0.20 \text{ Cr}^{3+})^{\dagger}$					
T	6.00 Si	14 <sup>‡</sup>	14.00	1.620	1.620	
В	3 B	$5^{\ddagger}$	5			

Notes: apfu = atoms per formula unit;

\* Calculated using the ionic radii of Bosi and Lucchesi (2007).

† Site populations optimized by the procedure of Wright et al. (2000).

<sup>&</sup>lt;sup>‡</sup> Fixed in the final stages of refinement

**TABLE 8.** Bond-valence calculations (valence unit) for vanadio-oxy-dravite.

TABLE 6. Bond valence calculations (valence unit) for variatio oxy dravite.						
Site	Χ	Υ	Ζ	Τ	В	Σ
O1		0.51 <sup>×3</sup> →				1.51
O2	0.17 <sup>×3</sup> ↓	$0.42^{\times 2} \downarrow \rightarrow$			0.96	1.98
O3		0.34	$0.40^{\times 2} \rightarrow$			1.13
04	0.07 <sup>×3</sup> ↓			$0.99^{\times 2} \rightarrow$		2.05
O5	0.10 <sup>×3</sup> ↓			$0.94^{\times 2} \rightarrow$		1.98
O6		0.45 <sup>×2</sup> ↓	0.49	1.04		1.99
07			0.49	1.07		1.99
			0.43			
O8			0.48		1.00 <sup>×2</sup> ↓	2.00
			0.52			
$\Sigma$	1.02	2.59	2.81	4.04	2.96	
MFV	1.17	2.62	2.79	4.00	3.00	

\*MFV = mean formal valence from site populations.

**TABLE 9.** Selected properties of oxy-dravite, vanadio-oxy-dravite and oxy-vanadium-dravite.

	Oxy-dravite	Vanadio-oxy-dravite	Oxy-vanadium-dravite		
a (Å)	15.9273(2)	16.0273(3)	16.1908(4)		
c (Å)	7.2001(1)	7.2833(1)	7.4143(2)		
$V(Å^3)$	1581.81(4)	1620.24(5)	1683.21(7)		
Space group	R3m	R3m	R3m		
Optic sign	Uniaxial (-)	Uniaxial (-)	Uniaxial (–)		
ω	1.650(5)	1.693(5)	1.786(5)		
ε	1.620(5)	1.673(5)	1.729(4)		
Colour	Dark red	Green	Black		
Pleochroism	O = orange E = pink	O = yellow green E = pale olive green	O = deep brownish green E = yellow green		
Reference	(1)	(2)	(3) and (4)		
*(1) Bosi and Skogby (2013); (2) this study; (3) Reznitskii et al. (2001); (4) Bosi et al. (2013a).					

FIGURE 1

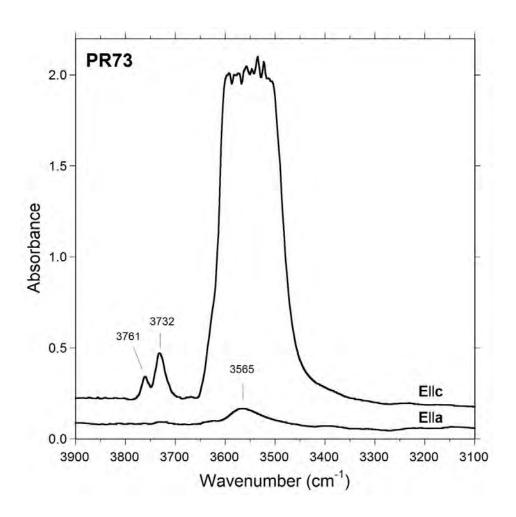


FIGURE 2

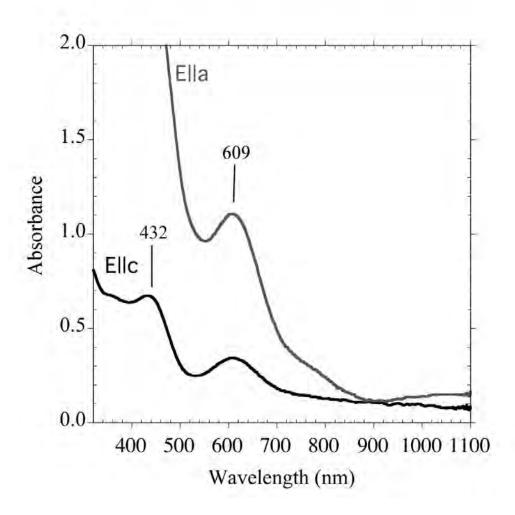


FIGURE 3

