

1 **Revision 1**

2

3 **Thermal diffusivities of MgSiO₃ and Al-bearing MgSiO₃ perovskites**

4

5 **KENJI OHTA,^{1,*} TAKASHI YAGI,² AND KEI HIROSE³**

6

7 ¹Center for Quantum Science and Technology under Extreme Conditions, Osaka
8 University, Toyonaka, Osaka 560-8531, Japan

9 ²National Metrology Institute of Japan, National Institute of Advanced Industrial
10 Science and Technology, Tsukuba, Ibaraki 305-8563, Japan.

11 ³Earth-Life Science Institute, Tokyo Institute of Technology, Meguro, Tokyo 152-8551,
12 Japan.

13

14 *E-mail: ohta@hpr.osaka-u.ac.jp

15

16

17

ABSTRACT

18 Thermal diffusivities of MgSiO₃ perovskite (MgPv) and MgSiO₃ perovskite
19 containing 2 wt% Al₂O₃ (Al-MgPv) were measured at ambient conditions using the
20 micro-spot heating Ångström method. The obtained values of thermal diffusivities of
21 MgPv and Al-MgPv are 2.6 ± 0.1 and 2.4 ± 0.1 mm²/s, respectively. Present result for
22 MgPv is much higher than previously reported value of 1.7 mm²/s. Substitution of
23 aluminum into MgPv has little effect on its thermal diffusivity at ambient conditions,
24 and such impurity effect would remain insignificant at high pressures and high
25 temperatures corresponding to the Earth's lower mantle.

26 **Keywords:** thermal diffusivity, thermal conductivity, MgSiO₃ perovskite (MgPv),
27 Al-bearing MgSiO₃ perovskite (Al-MgPv)

28

29

INTRODUCTION

30 To understand the thermal structure and the thermal evolution of the Earth, it is
31 indispensable to know the transport properties of the material that constitutes the
32 interior of the Earth. Since MgSiO₃ perovskite (MgPv) with some amount of iron and
33 aluminum is accepted to be a primary mineral in the Earth's lower mantle (e.g., Hirose

34 2002), thermophysical properties of the (Al,Fe)-bearing MgPv are of great importance
35 to comprehend the heat transfer system in the deep mantle. Osako and Ito (1991)
36 reported lattice thermal diffusivity (D) of MgPv to be $1.7 \text{ mm}^2/\text{s}$ at ambient conditions,
37 from which lattice thermal conductivity ($\kappa = D\rho C_p$, where ρ is density and C_p is specific
38 heat at constant pressure) was calculated to be 5.1 W/m/K . However, as Hofmeister and
39 Branlund (2007) claimed, the conventional Ångström method involving multiple
40 physical contacts, which was employed in Osako and Ito (1991) often underestimates
41 thermal diffusivity due to the contribution of the contact resistance between the sample
42 and heater or thermocouple. Thus the value of thermal diffusivity of MgPv at ambient
43 conditions needs revisit by using a contact free method.

44 Recent technical progress both in the experiment and the theoretical calculation
45 enables us to reveal high pressure and high temperature behavior of lattice thermal
46 diffusivity (and conductivity) of lower mantle minerals, MgSiO_3 perovskite and MgO
47 periclase (de Koker 2010; Stackhouse et al. 2010; Tang and Dong 2010; Manthilake et
48 al. 2011; Haigis et al. 2012; Ohta et al. 2012; Dekura et al. 2013). However, there is
49 only one report regarding the effect of chemical impurity on the lattice conductivity of
50 the lower mantle minerals (Manthilake et al. 2011). They reported dissolution of 3
51 mol% FeSiO_3 or 2 mol% Al_2O_3 into MgPv induces $\sim 70\%$ decrease of the thermal
52 diffusivity at 26 GPa and 300 K, which seems to be the significant impurity effect on
53 the diffusivity relative to other mantle minerals. Measurements of iron-bearing olivine
54 and orthopyroxene have yielded only 8% reduction in conductivity in the presence of 10
55 mol% iron (Horai 1971). Enrichment in aluminum into magnesium silicate is expected
56 to have a smaller effect on the diffusivity because of the similarity in atomic mass of
57 aluminum to magnesium and silicon relative to iron. Here we report the thermal
58 diffusivities of perovskites on MgSiO_3 end member composition and on its Al-bearing
59 solid solution at ambient conditions determined by the micro-spot heating Ångström
60 method that is a contact free technique. We found that thermal diffusivity of MgPv is
61 $2.6 \pm 0.1 \text{ mm}^2/\text{s}$, 50% higher than the value reported by Osako and Ito (1991), and there
62 is no measurable difference between the diffusivities for MgPv and Al-MgPv with 2
63 wt% Al_2O_3 at ambient conditions.

64

65 **EXPERIMENTAL METHODS**

66 A polycrystalline sample of MgPv has been synthesized from orthoenstatite (En) at

67 25 GPa and 2073 K for 1 hour in a multi-anvil apparatus. Starting material of En + 2
68 wt% Al₂O₃ in a rhenium capsule was also converted to Al-MgPv sample at 23 GPa and
69 2273 K for 1 hour. The perovskite structure was confirmed by a combination of
70 micro-focused (50-100 μm) X-ray diffraction measurements and Raman spectroscopy.

71 The micro-spot heating Ångström method was employed to measure the thermal
72 diffusivities of those perovskites (Fig. 1). This is a contact-free technique, and thus we
73 can eliminate the contribution of contact resistance. A pump laser beam (808 nm
74 wavelength, 3.7-7.0 mW laser power, modulated frequency f of 4 kHz, and 15 μm in
75 diameter) periodically heats a 100 nm-thick Mo film deposited on the polished samples.
76 During steady state heating, the heat inside the sample hemispherically propagates from
77 the heated spot and oscillated with the frequency f . The reflectivity of the Mo film
78 changes with temperature (i.e., thermorefectance effect) (Weaver et al. 1975; Wang et
79 al. 2010). The thermorefectance effect of the Mo film induced by the temperature
80 oscillation was detected using a continuous probe laser (782 nm wavelength, 1.5 mW
81 laser power, and 5 μm in diameter). The temperature phase (θ) is negatively
82 proportional to the distance from the heated spot (r), and the thermal diffusivity (D) of
83 the sample can be determined by the following equation,

$$84 \quad \theta = -\sqrt{\frac{\pi f}{D}} r + A \quad (1)$$

85 where A is a constant. The probe laser scanned the sample surface by a 1-μm step to
86 obtain a θ map around the heated spot with an area of 80 μm².

87 Artificially synthesized materials often contain micro size cracks. Such
88 discontinuities inside the material potentially lead underestimation of the
89 thermophysical properties of a bulk sample. In order to inspect our samples for the
90 cracks, we obtained a widely scanned temperature phase (θ) maps for samples using a
91 lower modulation frequency of 172 Hz. For MgPv, some discontinuities of heat
92 conduction were observed which were highlighted by white broken line (Fig. 2a). The
93 highlighted areas were almost overlapped with the cracks in the sample surface (Fig.
94 2b). This result indicates that cracks inside the sample inhibit heat conduction. We
95 avoided such cracks in samples for accurate thermal diffusivity measurements, using the
96 widely scanned map as a reference.

97 It is well known that MgPv undergoes crystal to amorphous transition when heated
98 above 400 K at atmospheric pressure (Durben and Wolf 1992). This is also true for

99 Al-MgPv (Liu et al. 1995). Since the temperature increase at the heated spot is
100 estimated to be about 50 K, and the temperature drops rapidly as r increases,
101 crystallinity of both MgPv and Al-MgPv should be kept during experiments. After the
102 diffusivity measurements, we also confirmed the crystallinity of the samples by means
103 of the Raman spectroscopy.

104

105

RESULTS AND DISCUSSION

106 We measured thermal diffusivity of the MgPv sample at four areas (Fig. 3a). The
107 pump beam of 4 kHz modulations heated the center of each measurement area. The
108 spherical distribution of the θ indicates that the measured area has no cracks and
109 homogeneous thermal diffusivity (Fig. 3b). The θ was plotted against the distance from
110 the heated spot (r) in all directions (Fig. 3c). The amplitude of temperature oscillation
111 decays with increasing the r , resulting the scatter of the θ in the r larger than 20 μm . We
112 fitted the Equation 1 to the data in the r range between 10 and 20 μm to calculate
113 thermal diffusivity (yellow line in Fig. 3c). The obtained values of the diffusivity of
114 MgPv in each measurement area are summarized in Table 1. Averaged value of thermal
115 diffusivity of MgPv is $2.6 \pm 0.1 \text{ mm}^2/\text{s}$, from which the thermal conductivity is
116 evaluated to be 8.1 W/m/K combining the reported values of heat capacity and density
117 of MgPv (Akaogi and Ito 1993; Fiquet et al. 2000).

118 The obtained thermal diffusivity of MgPv is 50% higher than that reported by
119 Osako and Ito (1991). As shown in Figure 2, cracks in the sample lower its thermal
120 diffusivity. Measured sample in Osako and Ito (1991) is a bulk sample with a cylindrical
121 shape, which could have many micro cracks, leading to substantial phonon scattering at
122 the cracks and a reduction in thermal diffusivity. In addition, Hofmeister and Branlund
123 (2007) claimed that the conventional Ångström method for minerals often
124 underestimates by about 20% due to interface resistance between the sample and the
125 heater or thermocouple. On the other hand, our method is a contact free technique and
126 enables us to select the measurement area with no cracks. Thus we can avoid
127 underestimation of thermal diffusivity. We have measured thermal diffusivity of MgPv
128 in a pressure range between 11 and 144 GPa and 300 K, and have revealed pressure
129 dependence of the diffusivity of MgPv ($\partial(\ln D_{\text{MgPv}})/\partial P$) of 1.2%/GPa by fitting the
130 obtained data (Ohta et al. 2012). Combining the present result for MgPv at 1 bar,
131 $\partial(\ln D_{\text{MgPv}})/\partial P$ is updated to be 1.1%/GPa, which is almost same to that we have

132 determined before (Ohta et al. 2012).

133 We also conducted similar thermal diffusivity measurements on the Al-MgPv
134 sample (Fig. 4a). As well as the measurements for MgPv, we selected measurement
135 areas with no cracks and homogeneous diffusivity (Fig. 4b), and determined thermal
136 diffusivities of selected areas with fitting obtained data via Equation 1 (Fig. 4c and
137 Table 1). The averaged thermal diffusivity of Al-MgPv is $2.4 \pm 0.1 \text{ mm}^2/\text{s}$, which is a
138 comparable value to that of MgPv within the experimental uncertainty. The present
139 results indicate that substitution of aluminum into MgPv has little effect to the thermal
140 diffusivity, contradictory to the results of Manthilake et al. (2011). Thermal diffusivity
141 of MgPv at 26 GPa determined by Manthilake et al. (2011) is quite higher than other
142 high-pressure data for MgPv (Ohta et al. 2012; Dekura et al. 2013). This possible
143 overestimation of the diffusivity on MgPv could induce an apparently large effect of Al
144 substitution on the diffusivity of MgPv.

145 The effect of scattering by solute aluminum on the lattice thermal conductivity of
146 MgPv can be estimated as follows (Klemens 1960; Pature and Klemens 1997):

$$147 \quad \kappa = \kappa_i \left(\frac{\omega_0}{\omega_M} \right) \arctan \left(\frac{\omega_M}{\omega_0} \right) \quad (2)$$

148 with

$$149 \quad \left(\frac{\omega_0}{\omega_M} \right)^2 = \frac{\chi T}{C(1-C)} \quad (3)$$

150 where ω_M is the phonon frequency corresponding to the maximum of the acoustic
151 branch of the phonon spectrum, ω_0 is that phonon frequency where the intrinsic mean
152 free path is equal to that due to solute atoms, χ is a constant, and C is the concentration
153 of the solute atoms. κ_i is the solid-solution thermal conductivity without the solute-atom
154 phonon scattering and is given by

$$155 \quad \kappa_i = C\kappa_A + (1-C)\kappa_B \quad (4)$$

156 where κ_A and κ_B are the thermal conductivities of solids with end member compositions
157 of $C = 0$ and $C = 1$, respectively, at a given temperature. At high temperature conditions
158 corresponding to the lower mantle, as (ω_M/ω_0) is very small, $\arctan(\omega_M/\omega_0)$ will be
159 comparable to (ω_M/ω_0) . Then, the scattering effect will be negligible at high temperature.
160 In addition, according to Hofmeister (1999), a pressure derivative of thermal
161 conductivity of Al_2O_3 corundum is similar to that of MgPv, and thus the difference in

162 the conductivity between MgPv and corundum will not change appreciably even at high
163 pressures. Hence the lower mantle impurity effect of solute aluminum in silicate
164 perovskite is predicted to be weak. The effect of iron on the conductivity of MgPv also
165 needs to be investigated. Direct measurements of thermal diffusivity of (Mg,Fe)O
166 ferropericlase, (Al,Fe)-bearing silicate perovskite, and post-perovskite under actual
167 Earth's lower mantle condition will provide tighter constraints on thermal transport
168 properties of Earth's mantle.

169

170

ACKNOWLEDGEMENTS

171 Comments from two anonymous reviewers were helpful to improve the manuscript.
172 K.O. is supported by the Japan Society for Promotion of Science.

173

174

REFERENCE CITED

- 175 Akaogi, M., and Ito, E. (1993) Heat capacity of MgSiO₃ perovskite. *Geophys. Res. Lett.*,
176 20, 105-108.
- 177 de Koker, N. (2010) Thermal conductivity of MgO periclase at high pressure:
178 Implications for the D'' region. *Earth Planet. Sci. Lett.*, 292, 392-398.
- 179 Dekura, H., Tsuchiya, T., and Tsuchiya, J. (2013) *Ab initio* lattice thermal conductivity
180 of MgSiO₃ perovskite as found in Earth's lower mantle. *Phys. Rev. Lett.*, 110,
181 025904.
- 182 Durben, D.J., and Wolf, G.H. (1992) High-temperature behavior of metastable MgSiO₃
183 perovskite: a Raman spectroscopic study. *Am. Miner.*, 77, 890-893.
- 184 Fiquet, G., Dewaele, A., Andrault, D., Kunz, M., and Bihan, T. (2000) Thermoelastic
185 properties and crystal structure of MgSiO₃ perovskite at lower mantle pressure and
186 temperature conditions. *Geophys. Res. Lett.*, 27, 21-24.
- 187 Haigis, V., Salanne, M., and Jahn, S. (2012) Thermal conductivity of MgO, MgSiO₃
188 perovskite and post-perovskite in the Earth's deep mantle. *Earth Planet. Sci. Lett.*,
189 355-356, 102-108.
- 190 Hirose K. (2002) Phase transitions in pyrolitic mantle around 670-km depth:
191 Implications for upwelling of plumes from the lower mantle. *J. Geophys. Res.*, 107,
192 2078, 10.1029/2001JB000597.
- 193 Hofmeister, A.M. (1999) Mantle values of thermal conductivity and the geotherm from
194 phonon lifetimes. *Science*, 283, 1699-1706.

- 195 Hofmeister, A.M., and Branlund, J.M. (2007) Properties of rocks and minerals- Thermal
196 conductivity of the Earth. *Treat. Geophys.*, 1, 543-577.
- 197 Horai, K. (1971) Thermal conductivity of rock-forming minerals. *J. Geophys. Res.*, 76,
198 1278-1308.
- 199 Klemens, P. G. (1960) Thermal resistance due to point defects at high temperatures.
200 *Phys. Rev.*, 119, 507-509.
- 201 Liu, L., Mernagh, T.P., and Irifune, T. (1995) Raman spectra of MgSiO_3 -10%
202 Al_2O_3 -perovskite at various pressures and temperatures. *Phys. Chem. Miner.*, 22,
203 511-516.
- 204 Manthilake, G.M., de Koker, N., Frost, D.J., and McCammon, C.A. (2011) Lattice
205 thermal conductivity of lower mantle minerals and heat flux from Earth's core. *Proc.*
206 *Natl. Acad. Sci. U.S.A.*, 108, 17901-17904.
- 207 Ohta, K., Yagi, T., Taketoshi, N., Hirose, K., Komabayashi, T., Baba, T., Ohishi, Y., and
208 Hernlund, J. (2012) Lattice thermal conductivity of MgSiO_3 perovskite and
209 post-perovskite at the core-mantle boundary. *Earth Planet. Sci. Lett.*, 349-350,
210 109-115.
- 211 Osako, M., and Ito, E. (1991) Thermal diffusivity of MgSiO_3 perovskite. *Geophys. Res.*
212 *Lett.*, 18, 239-242.
- 213 Padture, N., and Klemens, P. (1997) Low thermal conductivity in garnets. *J. Am. Ceram.*
214 *Soc.*, 80, 1018-1020.
- 215 Stackhouse, S., Stixrude, L., and Karki, B. (2010) Thermal Conductivity of Periclase
216 (MgO) from First Principles. *Phys. Rev. Lett.*, 104, 208501.
- 217 Tang, X., and Dong, J. (2010) Lattice thermal conductivity of MgO at conditions of
218 Earth's interior. *Proc. Nat. Acad. Sci. USA*, 107, 4539-4543.
- 219 Wang, Y., Park, J.-Y., Koh, Y.K., and Cahill, D.G. (2010) Thermoreflectance of metal
220 transducers for time-domain thermoreflectance. *J. Appl. Phys.*, 108, 043507.
- 221 Weaver, J.H., Olson, C.G., Lynch, D.W., Piancentini, M. (1975) Thermoreflectance of
222 Mo from 0.5 to 35 eV. *Solid State Comm.*, 16, 163-166.

223

224 **FIGURE CAPTIONS**

225 **Figure 1.** Schematic of thermal diffusivity measurement by the micro-spot heating
226 Ångström method. Thermal diffusivity of the sample can be calculated by using
227 Equation 1.

228

229 **Figure 2.** Photographs of the MgPv sample before Mo coating (a) with and (b) without
230 a scanned temperature phase (θ) map. Areas with low θ almost overlap cracks in the
231 sample indicated by white lines.

232

233 **Figure 3.** (a) Photograph of the MgPv sample before Mo coating. Four squares indicate
234 measurement areas. (b) Distribution of the θ for the area 1 (yellow square in Fig. 2a). (c)
235 The θ as a function of the distance from the heated spot (r) for the area 1. Thermal
236 diffusivity of MgPv was calculated in the range of r between 10 and 20 μm (yellow
237 line) by using Equation 1.

238

239 **Figure 4.** (a) Photograph of the Al-MgPv sample before Mo coating (Upper room of
240 rhenium capsule). Yellow and blue squares are measurement areas. (b) Distribution of
241 the θ for the area 1 (Yellow square in Fig. 4a). (c) The θ as a function of the r for the
242 area 1. Yellow line shows a fitting range of r to calculate the diffusivity of Al-MgPv by
243 using Equation 1.

TABLE 1. Thermal diffusivities of MgPv and Al-MgPv at each measurement point.

Sample	Measurement point	θ/r (rad/m)	Thermal diffusivity, D (mm^2/s)
MgPv	1	71600	2.45
	2	71300	2.47
	3	67900	2.73
	4	68200	2.70
Al-MgPv	1	71300	2.47
	2	74000	2.30

Modulation frequency (f) of the pump laser is 4 kHz.

The values of θ/r were derived from the data of $r = 10 \sim 20 \mu\text{m}$.

Thermal diffusivity (D) was calculated using Equation 1.

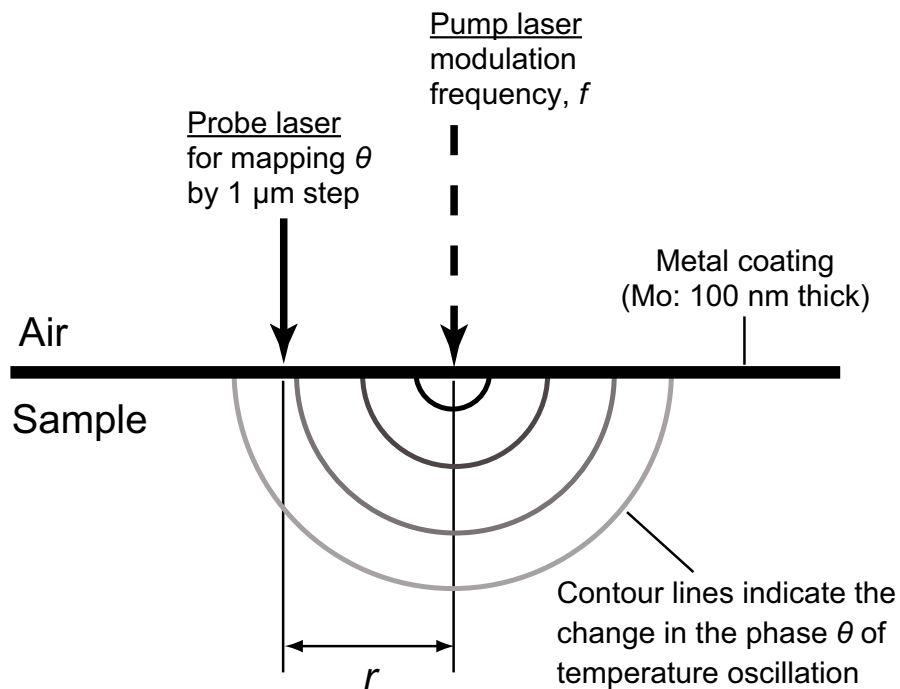
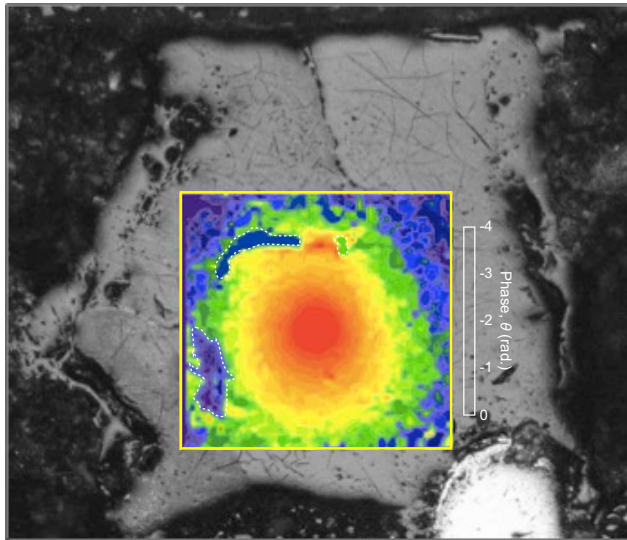


FIGURE 1

(a)



(b)

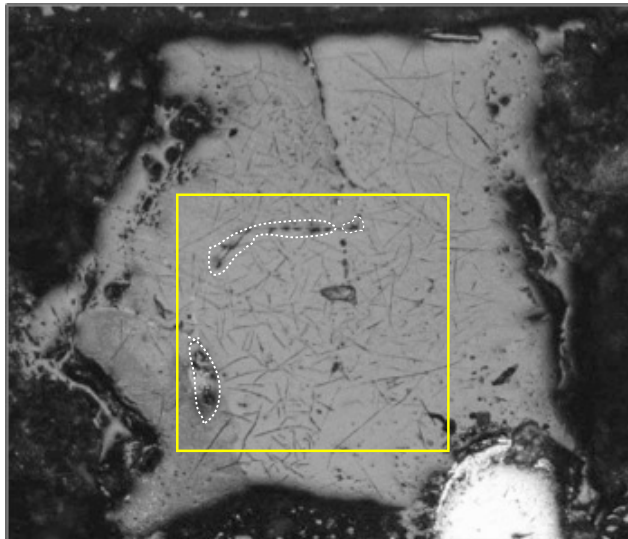
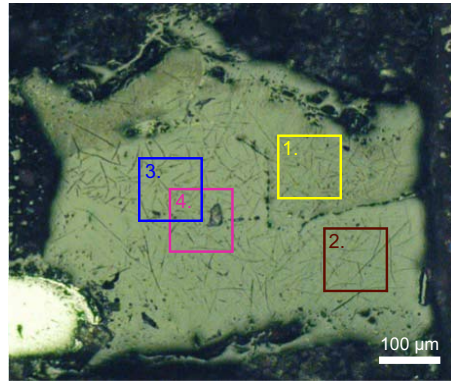
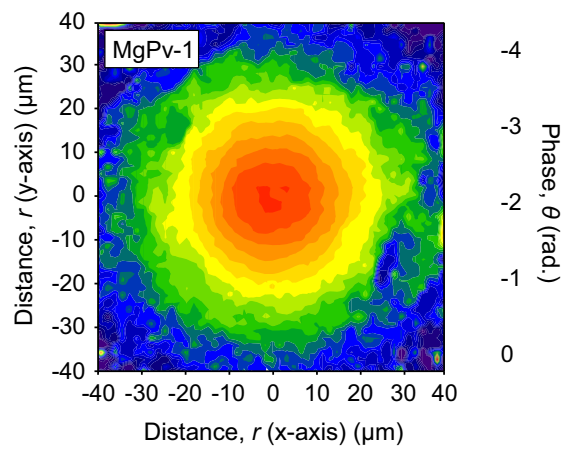


FIGURE 2

(a)



(b)



(c)

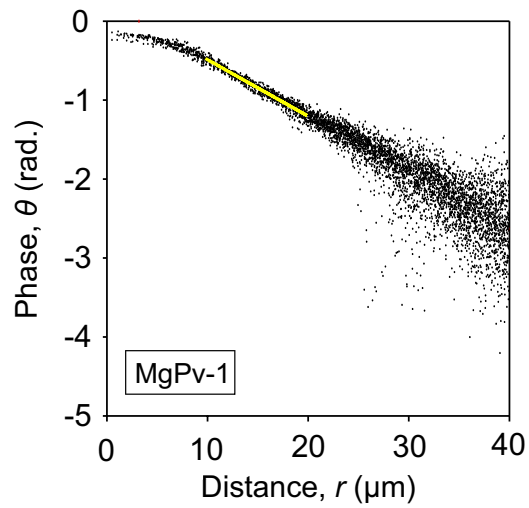
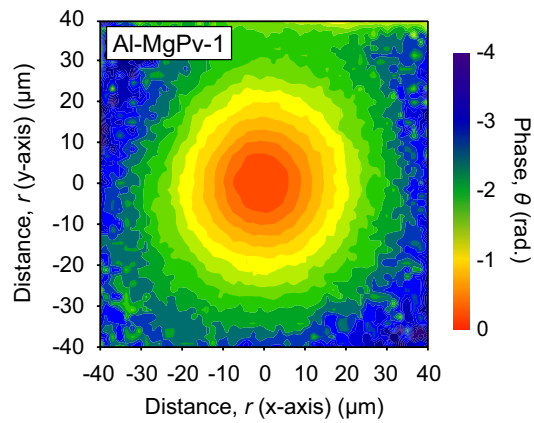


FIGURE 3

(a)



(b)



(c)

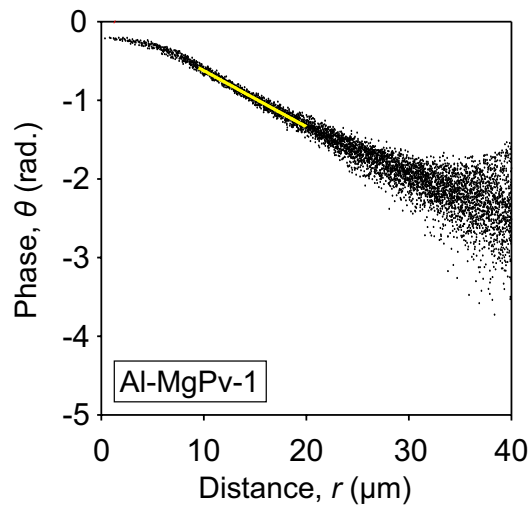


FIGURE 4