

1
2
3
4
5
6 **Manuscript 4580, Revision 2**
7
8
9

10
11
12
13
14
15
16
17 **Visible to near-infrared optical properties of pure synthetic olivine**
18 **across the olivine solid solution**
19

20
21 Peter J. Isaacson^{1*}, Rachel L. Klima², Jessica M. Sunshine³, Leah C. Cheek⁴, Carle M. Pieters⁴,
22 Takahiro Hiroi⁴, M. Darby Dyar⁵, Melissa Lane⁶, and Janice Bishop⁷
23
24

25 ¹Hawaii Institute of Geophysics and Planetology, School of Ocean and Earth Sciences and
26 Technology, University of Hawai'i at Mānoa, Honolulu, HI 96822

27 ²Johns Hopkins University, Applied Physics Laboratory Laurel, MD 20742

28 ³Astronomy Department, University of Maryland, College Park, MD

29 ⁴Brown University, Department of Geological Sciences, Box 1846, Providence RI 02912

30 ⁵Mount Holyoke College

31 ⁶Planetary Science Institute

32 ⁷SETI Institute
33

34 *Correspondence author: isaacson@higp.hawaii.edu, (808) 956-6938
35

36 Revised for: American Mineralogist
37

38 Keywords: Crystal chemistry, Mineralogy, Spectroscopy, Spectroscopy/Infrared
39

40 Number of Figures: 7

41 Number of Tables: 8
42
43
44
45

46 **1 Abstract**

47 Olivine exhibits highly diagnostic absorption features across visible to near-infrared
48 (VNIR) wavelengths due to electronic transitions of Fe^{2+} in its crystal structure. The properties
49 of these absorptions vary with composition, enabling compositional analysis of olivine through
50 VNIR spectroscopy, both in the laboratory and through remote sensing. Previous analyses of
51 these trends have relied on natural olivine samples, which are influenced by the presence of
52 minor cations that can affect the diagnostic absorptions. We conduct a systematic analysis of a
53 suite of synthetic (pure Mg/Fe) olivine samples with VNIR (300-2600 nm) reflectance
54 spectroscopy and quantitative spectral deconvolutions. From the full suite of samples described
55 and characterized by Dyar et al. (2009), we identify a small suite of well-characterized and
56 chemically pure olivine samples that demonstrates consistent and reliable spectral reflectance
57 properties across visible to near-infrared wavelengths. This suite covers the stoichiometric
58 olivine solid solution from $x = \text{Mg}/(\text{Mg}+\text{Fe}) = 0$ to $x = 70$ (Fo_0 to Fo_{70}). Because of their tight
59 compositional control, these synthetic samples improve on previous analyses of natural samples.
60 The results of this study provide a new standard for spectral reflectance properties of olivine
61 across visible to near-infrared wavelengths for the compositions present in the suite. We present
62 updated data on the trends in olivine band position as a function of olivine composition, which
63 are the basis for remote compositional evaluation of olivine with visible to near-infrared
64 reflectance spectroscopy. For these reasons, these improved olivine band position trends are of
65 major importance to remote compositional analyses of terrestrial planets.

66 **2 Introduction**

67 Olivine is a major component of rocky solar system bodies. It plays a dominant role in
68 planetary mantles, and is common in mafic igneous rocks occurring on or near planetary

69 surfaces. Olivine composition is a valuable means for interpretation of the thermal and chemical
70 evolution of diverse planetary bodies. Because olivine is the first mineral to crystallize from
71 typical mafic melts, it is a common constituent of planetary surfaces and chondritic meteorites.
72 Olivine composition (major element chemistry, typically expressed as Mg#, or molar
73 $(\text{Mg}/(\text{Mg}+\text{Fe}))$) is an important petrologic indicator for properties such as the body or sampled
74 unit's degree of petrologic evolution and parent melt composition. Furthermore, olivine's VNIR
75 absorption features are sensitive to olivine composition; olivine composition can often be
76 evaluated through remote sensing using techniques that quantify these characteristic olivine
77 absorption features. Thus, determining olivine composition through remote sensing is an
78 important tool for evaluating planetary petrology.

79 Characterization of the olivine absorption properties using laboratory VNIR spectroscopy
80 for samples with known major element compositions provides the basis for the evaluation of
81 unknown olivines with VNIR spectroscopy. However, pure stoichiometric olivine
82 $[(\text{Mg},\text{Fe})_2\text{SiO}_4]$ is very rare in natural samples, and the substitution of minor cations like Mn^{2+}
83 (Burns, 1993; Cloutis, 1997), Ni^{2+} (Hu et al., 1990), and Cr^{2+} (Burns, 1975, 1993) can have
84 unexpected effects on olivine's VNIR reflectance properties. Minor cations can introduce
85 additional absorption features in olivine spectra. However, the most significant effects of these
86 minor cations (with respect to compositional evaluation of olivine through analysis of the $1\ \mu\text{m}$
87 Fe^{2+} feature) largely involve changes to the principal Fe^{2+} crystal field transition absorptions near
88 $1\ \mu\text{m}$ (including changes to band position, relative intensities, and shape). These changes result
89 from changes to the olivine structure resulting from the introduction of cations with different
90 sizes and site occupancy preferences from the stoichiometric cations (Mg^{2+} , Fe^{2+}). Previous
91 efforts to define olivine's absorption properties across VNIR wavelengths have relied on natural

92 samples covering a compositional range that is as broad as possible (Sunshine and Pieters, 1998).
93 These studies described general trends in olivine's spectral properties across VNIR wavelengths,
94 but the samples studied were not pure stoichiometric olivine of precise compositions because
95 they were natural samples, typically containing minor amounts of non-stoichiometric cations
96 (e.g., Ca^{2+} , Mn^{2+} , Cr^{2+} , Fe^{3+}). Furthermore, natural samples cannot sample the full range of
97 olivine compositions with equal levels of sample purity, because certain olivine compositions are
98 not common in natural samples. In particular, pure Fe^{2+} -rich olivine samples are relatively rare.
99 For all of these reasons, studies of systematic trends in olivine absorption properties are
100 challenging when relying on natural samples. While it is true that natural samples are unlikely to
101 be completely free from minor cations and thus not perfectly represented by these synthetic
102 samples, the definition of a reference standard is best based on the ideal case, as it represents an
103 unbiased framework in which "real" natural samples can be evaluated.

104 The present study conducts a systematic survey of olivine's optical properties across the
105 compositional range using synthetic samples, which do not suffer from the complexities and
106 uncertainties resulting from using natural samples. The samples analyzed here are part of a larger
107 suite of synthetic olivine samples described by Dyar et al. (2009) along the solid substitution
108 series between forsterite [Fo , $x = \text{Mg}/(\text{Mg}+\text{Fe}) = 1$] and fayalite [Fa , $x = \text{Mg}/(\text{Mg}+\text{Fe}) = 0$]. The
109 preparation of these samples is discussed by Dyar et al. (2009). Dyar et al. (2009) incorporated a
110 variety of analytical spectroscopic techniques to characterize the properties and systematics of
111 olivine with the high degree of fidelity afforded by the use of synthetic samples. Lane et al.
112 (2011) conducted a systematic survey of the group 1 subset of these samples (subset groups 1
113 and 2 are defined below) using a variety of spectroscopic techniques at mid-infrared
114 wavelengths. However, the full sample suite is not perfect, and minor levels of contamination

115 and the presence of minor amounts of Fe^{3+} (stoichiometric olivine would contain only Fe^{2+} and
116 Mg^{2+} cations) affect the VNIR spectral properties of some samples significantly. While many of
117 the samples are appropriate for other analysis techniques, those with significant optically active
118 contaminants in the VNIR are excluded because our goal is to produce the definitive
119 characterization of “ideal” olivine’s spectral characteristics. Thus the samples used in this study
120 were selected from the larger suite to describe most clearly the systematics of olivine absorptions
121 in the VNIR across the compositional range. We eliminated samples that exhibit obvious and
122 substantial variations from previously identified trends (based on natural samples) in olivine
123 absorption properties (e.g., Burns, 1970, 1974; King and Ridley, 1987; Sunshine and Pieters,
124 1998). The previous trends are assumed to be generally correct, so the goal of the present work is
125 to refine, rather than redefine, these trends. As such, these previously identified trends are a
126 reasonable standard to use in defining samples to exclude from the present analysis. The
127 remaining high-quality samples are used to refine the trends in olivine absorption properties
128 across VNIR wavelengths. The results should be considered the definitive standard for olivine
129 absorption properties in the VNIR.

130 **3 Background**

131 **3.1 Olivine structure and VNIR spectroscopy**

132 Olivine structure is discussed by Dyar et al. (2009), but is summarized briefly here.
133 Olivine is an orthosilicate with a nearly hexagonally close-packed structure. It has a general
134 formula of M1M2SiO_4 , where the cation sites (M1, M2) are six-coordinated (octahedral). The
135 dominant (stoichiometric) divalent cations (Fe^{2+} , Mg^{2+}) occupy the octahedral cation sites, and
136 the octahedra share edges, creating chains along the *c* axis. The M2 site is the larger and more
137 distorted of the sites, although size effects between the sites are relatively minor (e.g., Bish,

138 1981). Olivine does not exhibit strong ordering in the cation sites between Fe^{2+} and Mg^{2+} ,
139 although minor site ordering has been detected using optical absorption spectroscopy (e.g., Taran
140 and Matsyuk, 2013). While Fe^{2+} has a slight preference for the M1 site at relatively low
141 temperatures (below $\sim 600 \pm \sim 100^\circ\text{C}$, depending on the sample and analytical technique
142 employed), each cation will occupy both sites in pure stoichiometric olivine (Abdu et al., 2008).
143 The geometry of the two M06 octahedral sites is illustrated (for the forsterite endmember) in
144 Figure 1.

145 The asymmetry of the cation sites (i.e., the irregularity of the M06 octahedra, as
146 illustrated in Figure 1) leads to a splitting in the energy levels of the outer d orbitals in Fe^{2+} in the
147 cation sites. This splitting gives rise to the characteristic olivine absorption features near $1 \mu\text{m}$.
148 Example olivine spectra are illustrated in Figure 2. The position and shape of the absorption
149 features are controlled by the magnitude of the energy splitting between these orbitals. Cation-
150 ligand (i.e., $\text{Fe} - \text{O}$) bond length controls the splitting energy (also termed “crystal field
151 stabilization energy”, or Δ), and scales as $\sim r^{-5}$, where r is the bond length (Burns, 1993). As
152 incident photons with the proper energy (Δ) interact with the olivine structure, they are absorbed
153 and cause electrons to “jump” between the split energy levels. This photon absorption causes
154 absorption features in the wavelength-resolved optical properties of the mineral.

155 The characteristic olivine absorption near $1 \mu\text{m}$ is comprised of three components. Two
156 features (near $\sim 0.85 \mu\text{m}$ and $\sim 1.25 \mu\text{m}$) are caused by Fe^{2+} in the M1 site, and the central
157 absorption by Fe^{2+} in the M2 site (Burns, 1970, 1974). These features shift in relative strength,
158 width, and most substantially, position (energy), as a function of the olivine’s relative Fe^{2+} vs.
159 Mg^{2+} content (Mg# or Fo#), largely driven by the changing bond lengths due to the different
160 sizes of the Fe^{2+} and Mg^{2+} cations. The general trend is that with increasing Fe^{2+} content,

161 absorption features shift to longer wavelengths and become slightly broader, while the M1
162 absorptions become stronger relative to the central M2 absorption (Burns, 1970, 1974; Sunshine
163 and Pieters, 1998).

164 **3.2 The Modified Gaussian Model**

165 The reflectance properties of the olivine samples studied are quantified by analysis with
166 the Modified Gaussian Model (MGM). The MGM is a spectral deconvolution routine developed
167 by Sunshine et al. (1990) that has been widely validated as a means for accurately quantifying
168 crystal field absorption features caused by Fe²⁺ in mineral cation sites both in laboratory spectra
169 (e.g., Hiroi and Sasaki, 2001; Isaacson and Pieters, 2010; Sunshine and Pieters, 1993, 1998;
170 Noble et al., 2006; Klima et al., 2007; Mayne et al., 2010; Clénet et al., 2011; Isaacson et al.,
171 2011a) and in remotely-acquired spectra (e.g., Hiroi et al., 1995; Sunshine et al., 2007; Birlan et
172 al., 2011; Farrand et al., 2011; Isaacson et al., 2011b). We note that the scientific literature of
173 research with the MGM is extensive, and we have not attempted a thorough review of it here; the
174 reader is referred to references within those example works we have cited (which emphasize
175 analyses of olivine) for further background. A major strength of the MGM is its ability to
176 quantify absorption features linked to specific transitions in spectra exhibiting multiple
177 overlapping absorptions in a small wavelength region such as olivine and olivine-pyroxene
178 mixtures. The MGM approach is based on the fact that electronic transitions such as those in
179 olivine can accurately be described as “modified” Gaussian distributions in which the
180 fundamental variable is transformed from energy to bond length to better model the properties of
181 real absorption features (Sunshine et al., 1990). MGM deconvolutions are carried out in
182 wavelength and natural log reflectance, in which overlapping features add linearly and thus
183 linear inversions can be used in the deconvolution (Sunshine and Pieters, 1993; Sunshine et al.,

184 1990). The optimization method employed by the MGM is the approach of Tarantola and Valette
185 (1982). The MGM inversion technique is iterative in which the parameters controlling the
186 Gaussians and spectral continuum slope are varied until a pre-defined RMS residual error
187 threshold is met (Sunshine et al., 1990). Additionally, the Tarantola and Valette (1982) and
188 MGM approaches allow the introduction of *a priori* information as constraints on the resulting
189 parameters; for example, Gaussian centers can be prescribed to lie within a specified wavelength
190 range within a certain level of confidence. MGM fits are initiated by a startup file containing
191 initial values for all parameters to be considered in the optimization. By varying the starting
192 parameters, the initial conditions for the run can be changed. MGM results can be sensitive to
193 starting parameters (e.g., Clénet et al., 2011), so the choice of initial parameters is significant.
194 For the results presented here, standard initial conditions roughly consistent with olivine, based
195 on previous work by Sunshine and Pieters (1998), were used for all fits.

196 **4 Samples**

197 A full listing of the olivine samples synthesized for an integrated spectroscopic characterization
198 of olivine can be found in the overview publication of Dyar et al. (2009). Two different suites of
199 olivines were described in that publication. The first suite [described here as “Group 1” or
200 “SUNY olivines”, which is the source of most of the olivines evaluated here and of those studied
201 by Lane et al. (2011)] was synthesized by Donald Lindsley at SUNY Stony Brook. The second
202 suite (described here as “Group 2” or “Bristol olivines”) was synthesized by Richard Brooker at
203 Bristol University. Both cover the full range of stoichiometric olivine composition, although the
204 SUNY samples were found to contain fewer impurities due to the different synthesis process
205 employed (the SUNY samples were synthesized with multiple heating/drying steps, while the
206 Bristol samples experienced only a single cycle, leading to less complete reactions in some

207 cases). More complete descriptions of the sample suites, synthesis processes, and quality
208 assessments (including through Mössbauer spectroscopy and X-ray diffraction analyses) are
209 provided by Dyar et al. (2009), and the full listing of samples evaluated is provided in Table 1.
210 Subsequent tables identify the subset of samples used for further analyses.

211 **5 Methods**

212 **5.1 Spectral reflectance measurements**

213 Visible to near-infrared reflectance spectra of all of the synthetic olivine samples were
214 collected over a wavelength range of 300 nm – 2600 nm using the bidirectional reflectance
215 spectrometer in the Brown University RELAB. Reflectance measurements were collected as
216 “bidirectional reflectance”, using the standard RELAB viewing geometry of incidence (i) = 30°,
217 emergence (e) = 0° (Pieters, 1983; Pieters and Hiroi, 2004). The samples were in particulate
218 form for these measurements with a standard grain size of <45 μm for all measurements.

219 **5.2 Compositional Analyses**

220 All of the samples evaluated in the course of this study were synthesized under carefully
221 controlled conditions, and in general were found to be of high quality and purity based on x-ray
222 diffraction and Mössbauer spectroscopy analyses, although many of the Group 2 samples
223 exhibited higher levels of impurities (likely linked to incomplete synthesis reactions) and higher
224 degrees of compositional variability, manifested in both the abundance of Fe^{3+} and impurities
225 based primarily on Mössbauer spectroscopy analyses (Dyar et al., 2009). However, to ensure that
226 the samples evaluated in this study are consistent with their stated compositions, we performed
227 compositional “spot checks” with electron microprobe (EMP) analysis of grain mounts prepared
228 from selected samples using the Cameca SX-100 at Brown University. Compositional analyses
229 were performed on all six samples that passed the spectral quality assessment (section 5.4).

230 Many individual analyses were performed on each grain mount, and were averaged to determine
231 mean compositions for the samples. Only analyses that passed quality thresholds were included
232 in these averages. These thresholds were defined based on stoichiometry (moles of cations per
233 four moles of oxygen between 2.98 and 3.02) and oxide weight percent totals (between 98.0%
234 and 102%; the 102% upper cutoff is higher than might typically be used, but the olivine samples
235 tended to have systematically high totals for all analyses).

236 **5.3 MGM Modeling**

237 As discussed above, the absorption features in the acquired reflectance spectra were
238 quantified with the Modified Gaussian Model (MGM). Each spectrum was modeled with a series
239 of four standard “fits”, where a “fit” refers to a particular starting condition. Because MGM
240 model results can be sensitive to the initial conditions, as discussed above, we used the same
241 starting conditions for the absorptions in each fit, varying only the constraints on the degree to
242 which the parameters were allowed to vary in optimizing the result. Furthermore, because MGM
243 fits have some uncertainty in predicted band center, we included some fits in which the
244 absorption widths were highly constrained to values consistent with olivine, which generally
245 tended to improve the consistency of the band center results with previously-established trends.
246 Fits were run iteratively until the pre-defined RMS error threshold of 10^{-5} , consistent with the
247 value employed by Sunshine et al. (1990), was met (when the improvement in RMS residual
248 error between successive iterations was less than this threshold). The continuum used was a first-
249 order polynomial in wavelength (nm) and natural log reflectance units, with an initial offset of
250 0.8 and a slope of 0.0. The initial parameter values for the absorptions are provided in Table 2.
251 Bands 1-3 are short-wavelength bands required to accommodate short wavelength spectral
252 structure and enable a mathematically valid fit; the ultraviolet (UV) and visible wavelengths are

253 not analyzed thoroughly in this study, and results for these absorptions are not reported. While
254 Gaussians may not be the most accurate of the UV absorptions (which are largely caused by
255 various charge transfer processes), we employ this approach for consistency with previous
256 treatments of olivine spectra concentrating on the electronic transition absorptions that are the
257 focus of this work. Quantitative analysis of charge transfer absorptions in the UV and visible are
258 beyond the scope of this work. The initial values for the principal olivine absorptions (M1-1, M2,
259 M1-2) are intended to be consistent with an intermediate composition ($\sim\text{Fo}_{50}$). The parameter
260 constraints for each fit are listed in Table 3. Note that the other parameters not mentioned
261 explicitly in Table 3 (i.e., Gaussian centers and intensities) were allowed to vary independently
262 during fitting. The same initial offset and slope was used for each fit; the only parameters
263 changed were the constraints imposed on various parameters during fitting. The possibility of a
264 second absorption caused by Fe^{2+} in the M2 site has been raised previously (e.g., Runciman et
265 al., 1974; Hiroi and Takeda, 1992), and there is some support for this idea from analysis of Mg-
266 free orthorhombic olivine-type phosphates where Fe^{2+} is confined to the M2 site (Langer et al.,
267 2006). However, MGM analyses in this work and by previous authors (e.g., Sunshine and
268 Pieters, 1998) do not reveal evidence for this additional band.

269 **5.4 Quality assessment and filtering**

270 The “quality” of the resulting fits was determined by analyzing the consistency of the
271 MGM-derived parameters with previously-derived trends (Sunshine and Pieters, 1998). Because
272 the present study is intended to refine the spectral properties of olivine rather than re-define
273 them, we assume that the results of these previous studies represent valid guidelines for the
274 general reflectance properties of olivine spectra. In other words, we used these previous studies
275 to determine which spectra and fits from the present study were most consistent with olivine’s

276 known spectral properties. The criteria used most heavily in filtering the results were the band
277 positions relative to the known composition (Fo#) and the relative band strengths of the M1
278 absorptions.

279 We determined the deviation of the M1-1, M2, and M1-2 band positions from previously
280 established trends (Sunshine and Pieters, 1998). Fits with deviations near or above 10 nm for
281 M1-1, 7.5 nm for M2, and 12.5 nm for M1-2 were considered to be of lower quality. Fits with
282 extreme deviations (i.e., deviations far exceeding this “allowable” range) were rejected. While
283 these ranges may seem large, the previous results are only a guide, so we must consider results
284 that deviate from previous results in order not to bias the present study excessively. Previous
285 results indicate that while the strength of the M2 band relative to that of the M1-2 band varies as
286 a function of Fo#, the strength of the M1-1 band relative to the M1-2 band is approximately
287 constant with composition. This is fully consistent with theory, as Fe²⁺ in the same site causes
288 both of these (M1) absorptions. Thus, we rejected any fits in which the strength of the M1-1
289 band relative to that of the M1-2 band fell beyond a specified range of 0.5 to 0.65. Any fits
290 falling close to the extremes of this range (within ~0.02) were not excluded, but were considered
291 to be lower quality. Examples of excellent, good, mediocre, and poor/rejected results are given in
292 Table 4. The fits described in Table 4 are illustrated in Figure 3. The parameters causing the fit to
293 be downgraded are in boldface type. Quality assessment is according to shading: white /
294 excellent, light grey / good, dark grey / mediocre, and black / rejected. Generally, a deviation
295 indicating reduced quality according to the criteria given above for a single parameter would
296 indicate a “good” result (as opposed to “excellent”, which would exhibit no such deviations),
297 while more than one such deviation would indicate a “mediocre” result. In cases where multiple
298 fits produced results of comparable quality (e.g., two “excellent” fits for a single spectrum), we

399 chose the result derived with a flat continuum slope for consistency with previous analyses by
300 Sunshine and Pieters (1998). This situation was only observed for fits 2 and 4 (both with
301 constrained widths, with the only difference being the constraint on continuum slope), so we
302 never were forced to choose between results obtained with the same continuum slope constraint
303 (i.e., fit 1 vs. 2 or 3 vs. 4).

304 As evidenced by Figure 3, the differences between “good” and unacceptable spectra in
305 defining trends for olivine across VNIR wavelengths are difficult to determine from plots of the
306 spectra; all samples and spectra evaluated in this study are olivine with an extraordinarily high
307 purity level. The difference between spectra that pass our defined quality threshold and those
308 classified as unacceptable is likely a small amount of contamination or residual uncrystallized
309 reactants. Because the samples are all of very high purity, we must rely on such quality criteria
310 as those outlined above to identify those samples and spectra that best capture the properties of
311 pure olivine across VNIR wavelengths, and must also assume that those samples not passing the
312 quality test are slightly more affected by minor contamination.

313 **5.5 Trends in band position with composition**

314 The band positions derived from MGM analyses (section 5.3) were used along with the
315 compositions determined from EMP analysis (section 5.2) to establish trends in band position
316 with changing olivine composition (Mg#) for each of the three component absorptions. The
317 trends in band position with composition were assumed to vary linearly, as in previous
318 treatments (Burns, 1970; King and Ridley, 1987; Sunshine and Pieters, 1998; Burns et al., 1972).
319 Trend lines were produced for each individual absorption via least squares. In producing our
320 trend lines, we used the mean compositions for each sample determined from EMP analysis
321 rather than the nominal (stated) compositions, and did not use any weighting (e.g., by

322 measurement variance). However, we provide the band position (Table 5) and compositional
323 (Table 6) data that will enable the reader to produce trend lines using alternative approaches.

324 **6 Results**

325 The suite of spectra and fits that survived the quality-filtering step are provided in Table
326 5, along with the MGM results. Compositions for these samples are provided in Table 6; full
327 oxide abundances are provided in the supplemental material. These selected fits are shown
328 graphically in Figure 4. Shading in Table 5 is according to quality of the selected fit, and uses the
329 same scheme as Table 4 (excellent to poor from light to dark). Note that one case, due to
330 moderate deviations from previous trends either in band position or in relative M1 absorption
331 strengths, the best fit is only mediocre [C1DD96/Fo₂₀, due to deviations in M1-1 (~9 nm) and
332 M2 (~8 nm) band position]. As the difference between “excellent” and poorer fits is very
333 difficult to determine from plots such as those in Figure 4, the quantitative criteria given in Table
334 5 are critical for the quality assessment. Note that all results shown in Table 5 were produced
335 from either fit 2 or fit 4; fits 1 and 3 produced significantly poorer results in most parameters.
336 For example, C1DD98, which produced some of the “best” results based on consistency with
337 previously-established trends, exhibited deviations of ~15 nm in band position, up to 70 nm in
338 FWHM, as well as significant variations in relative M1-1 intensity in results from fits 1 and 3,
339 but produced excellent results in the constrained cases (fits 2 and 4). Results from other spectra
340 were often worse in the unconstrained cases, demonstrating the importance of the parameter
341 constraints in producing reliable results.

342 The suite of spectra and fit results given in Table 5 were used to prepare a series of
343 figures illustrating the reflectance properties of the “surviving” (highest-quality) synthetic olivine
344 compositional range (Fo₇₀ – Fo₀). The following series of plots is labeled by Fo#, with blue

345 colors (Fe-rich) trending to orange and red (Mg-rich). Note that the original suite includes a
346 wider compositional range, in particular more magnesian samples, but those samples did not
347 produce spectra or fits of sufficient quality for these quantitative VNIR analyses. Figure 2 shows
348 reflectance spectra of the synthetic olivine suite. Figure 5 shows the same spectra after
349 continuum removal (a) and after a normalization by band depth (b) that illustrates trends in
350 spectra property without the confusion of variable band strengths, following the approach of
351 Lucey (1998). The continuum slope removed was a straight line, with tangent points at 660 nm
352 and 1960 nm. Band minima and relative strength of the M1-2 (long-wavelength) absorption both
353 increase with increasing Fe content. The Fo₃₀ and Fo₄₀ spectra deviate slightly from the trend,
354 which is most apparent in Figure 5b, where the Fo₃₀ spectrum runs at shorter wavelength than the
355 Fo₄₀ spectrum in the M1-2 region. This deviation is discussed further below.

356 The MGM-derived band positions of the selected subset and fits are illustrated in Figure
357 6. Results from this study are in heavy lines and symbols, while results from Sunshine and
358 Pieters (1998) are in smaller filled symbols and dashed lines. The trend lines for the present
359 study were calculated with a linear fit to the results for each absorption (section 5.5). A “curve”
360 is observed in the M1-1 absorption at Fe-rich compositions, in which the band positions skew to
361 shorter wavelengths in a curved trend. This curve can be seen most clearly in the M1-1 band
362 centers for the Fo₁₀ and Fo₂₀ points, which fall at shorter wavelengths than the trend line in
363 Figure 6. A similar effect was observed previously (Sunshine and Pieters, 1998), also apparent in
364 the Sunshine and Pieters points in Figure 6. The slopes of the trend lines in this plot are
365 presented in Tables 7 and 8. Table 7 presents the slopes cast in the form of band positions as a
366 function of composition (Fo#), and Table 8 presents the same slopes cast in the form of Fo#

367 solved as a function of band position. The values in Table 8 are the bases of equations used to
368 predict olivine composition from MGM-derived band position (e.g., Isaacson et al., 2011b).

369 The MGM-derived widths and intensities of the selected subset are illustrated in Figure 7,
370 which plots these parameters against MGM-derived band strength. Absorption intensities (in
371 Figure 7b) are shown relative to the M1-2 absorption to eliminate variability resulting from
372 differences in absolute band intensity between the spectra, as overall band intensity is controlled
373 by a number of factors unrelated to the variation with composition that is of primary interest
374 here. As found in previous studies, the central M2 absorption weakens (relative to the M1-2
375 absorption) with increasing Fe content, whereas the intensity of the M1-1 absorption is roughly
376 constant, weakening slightly (relative to the M1-2 absorption) at Fe-rich compositions. While the
377 overall trend is of increasing M2 relative intensity with decreasing Fe content, the trend is rough,
378 as the Fo₁₀ and Fo₂₀ points are reversed, as are the Fo₃₀ and Fo₄₀ points. The absorption widths
379 are relatively constant for the M1-1 and M2 absorptions (they were highly constrained in these
380 fits), although the M1-2 absorption does get slightly wider at more Fe-rich compositions.

381 The outliers noted above in Figure 5b are notable outliers in Figure 7a as well. The Fo₄₀
382 and Fo₃₀ spectra are widely separated in M2 relative intensity, which likely contributes to their
383 apparent reversal in order in Figure 5b. This apparent reversal in Figure 5b could be caused by a
384 very strong M1-2 absorption relative to the M2 (i.e. a smaller value in Figure 7a) in the Fo₄₀
385 spectrum, which would cause the spectrum to appear at a longer wavelength in the 1250 nm M1-
386 2 region (scaled spectra in Figure 5b) as the scaling approach would tend to over-scale the M1-2
387 absorption slightly (the scaling factor is determined by the intensity at the band minimum, which
388 is driven primarily by the M2 absorption). Conversely, a weak M1-2 absorption relative to the
389 M2 absorption (as in the Fo₃₀ spectrum) would cause the spectrum to appear at shorter

390 wavelengths in the 1250-1300 nm region in Figure 5b). A similar effect to that observed in the
391 Fo_{30} and Fo_{40} spectra is observed in the Fo_{10} and Fo_{20} spectra; they appear reversed in the trends
392 with composition in Figure 7a, and also anomalous in the progression in scaled spectra in Figure
393 5b (although to a lesser degree than the Fo_{30} and Fo_{40} spectra; the Fo_{10} and Fo_{20} spectra merely
394 plot on top of each other, whereas the Fo_{30} and Fo_{40} spectra appear in the wrong order). Note that
395 this inconsistency between the Fo_{30} vs. Fo_{40} and Fo_{10} vs. Fo_{20} spectra does not mean that the
396 MGM-derived band positions fall in reverse order; as shown in Figure 6, the band centers for
397 these spectra are in the expected trend, with the Fo_{40} and Fo_{20} band centers at shorter
398 wavelengths (relative to the Fo_{30} and Fo_{10} points, respectively). The inconsistency is only
399 observed in apparent band position in the scaled spectra shown in Figure 5b.

400 **7 Discussion**

401 **7.1 Identification of best synthetic olivine spectra**

402 One of the major goals of this study is to identify, from the suite of olivines discussed by
403 Dyar et al. (2009), those spectra that best represent the spectral reflectance properties of olivine
404 across visible to near-infrared wavelengths for the largest compositional range possible. In
405 particular, this study seeks to improve coverage of Fe-rich olivine compositions, for which the
406 data used by Sunshine and Pieters (1998) were of lower purity. The quality of the spectra in the
407 visible to near-infrared is an important consideration, because the quality assessment of Dyar et
408 al. (2009) was based on Mössbauer analyses. While the quality assessments are generally
409 consistent between Mössbauer and VNIR spectroscopy (samples identified as being of poor
410 quality with Mössbauer analyses are unlikely to be good in the VNIR), the correlation is not
411 perfect, and samples that pass the Mössbauer quality test are not necessarily of high quality in
412 the VNIR.

413 As stated above, our basis for evaluating the “quality” of the VNIR olivine spectra was
414 previous work by Burns (1970; 1974), King and Ridley (1987), Sunshine and Pieters (1998), and
415 Burns et al. (1972). These previous studies represent suitable general guidelines; we assume that
416 these previous studies are not grossly incorrect. Thus, any spectra that were in serious
417 disagreement with results for olivine of similar composition from those previous studies were
418 considered to be of lower quality. The spectra that we consider to be of the best quality from the
419 full suite are listed in Table 5. Note that in order to cover as much of the compositional range as
420 possible, a few spectra of slightly lower quality (e.g., C1DD96/F_{O20}, based on inconsistency in
421 spectral characteristics with previous results) were included in this identified sub-suite. On the
422 flip side, several spectra that would pass the MGM-based quality assessment just based on their
423 spectral characteristics (C1DD92/F_{O55}, C1DD41/F_{O50}) were excluded because of their wide
424 variability in composition based on the EMP analyses.

425 *a Trends in band position*

426 The reflectance properties of our identified sub-suite are remarkably consistent with the
427 results of previous studies (Burns, 1970; Sunshine and Pieters, 1998). However, a few
428 differences are apparent. First, and perhaps most significantly, we observe a flatter slope for the
429 M1-1 absorption relative to the results of Sunshine and Pieters (1998) in the band position vs.
430 composition plot (Figure 6). These slopes can be found in numerical form in Tables 7 and 8.
431 While our M1-1 slope is flatter than that of Sunshine and Pieters (1998), our M1-1 and M1-2
432 slopes are more similar than those observed by Sunshine and Pieters (1998). The Sunshine and
433 Pieters (1998) study found a steeper M1-1 slope relative to the M1-2 slope. We observe a similar
434 pattern (steeper M1-1), but our values are more similar. In theory, because the same transition
435 causes both absorptions, the slopes for these lines should be similar (Burns, 1970; 1993; Burns et

436 al., 1972). Sunshine and Pieters (1998) used natural samples, which tend to exhibit greater
437 compositional variability than synthetic samples like the ones we analyzed, and this difference
438 could explain minor differences in the band position trends. However, the increased similarity
439 between our M1-1 and M1-2 slopes relative to the results of Sunshine and Pieters (1998) leads us
440 to trust the slope values from the present study.

441 *b Outliers in MGM-derived band properties*

442 We note a few outliers in Figures 4 and 6. More scatter is to be expected in the plots of
443 relative intensity and FWHM (Fig. 7) relative to the band position vs. composition plot (Fig. 6),
444 but insofar as the results can be interpreted directly, the Fo₄₀ and Fo₁₀ points in particular are
445 inconsistent with the overall trend of decreasing relative M2 strength with increasing Fe content
446 (Fo#), as noted previously in comparison to their close compositional neighbors (the Fo₃₀ and
447 Fo₂₀ points). Based on that simple trend, the Fo₄₀ spectrum appears to have an anomalously low
448 M2 relative intensity, and the Fo₁₀ spectrum to have an anomalously high M2 relative intensity.
449 The Fo₃₀ fit result (Figure 4) exhibits a slight mismatch in the ~1200 nm region, which may
450 contribute to the anomalous relative absorption intensities. The Fo₁₀ spectrum has the most
451 pronounced continuum slope as shown in Figure 4, and both the Fo₁₀ and Fo₂₀ fits (Figure 4)
452 required sloped continuum slopes, which may contribute to their inconsistency with the other
453 results. The C1DD96/Fo₂₀ fit result is more marginal, and was included in the interest of having
454 a more complete compositional series. Additionally, that sample has the highest standard
455 deviation in its composition of any sample included in our suite. While still quite low (1.72 in
456 Mg#, Table 6), this is another possible reason to treat this sample and its results with some
457 caution. Thus, care should be taken in using this spectrum, particularly if conclusions are heavily
458 influenced by these specific spectra or their MGM-derived properties. Overall, it is also

459 important to recall that despite the overall expected trend of relative intensity with composition
460 observed, substantial scatter is expected in the relative intensities based on previous results
461 (Sunshine and Pieters, 1998). This scatter, clearly observed in the results in Figure 7, combined
462 with the relatively small suite of input data points in the present study, caution against over-
463 interpretation of individual data points in the relative intensity results.

464 *c Features associated with contaminants*

465 We also note evidence for minor contamination in several of the olivine spectra. In
466 particular, several spectra exhibit weak features at longer wavelengths (e.g., C1DD39/Fo₇₀,
467 C1DD97/Fo₁₀, C1DD96/Fo₂₀). Several spectra also exhibit weak features attributable to H₂O/OH
468 contamination are also apparent across this wavelength range (>1500 nm). While these spectral
469 features are attributable to contamination rather than pure stoichiometric olivine (for example,
470 EMP analysis suggested that the Fo₇₀ sample in particular appeared to have modally minor
471 (several %) ~Mg₇₀ pyroxene), they do not affect the principal olivine absorptions substantially,
472 so they are not a major complication for the results derived from these spectra. This minor
473 pyroxene could, however, explain the very weak absorption in the C1DD₃₉/Fo₇₀ spectrum just
474 beyond 2000 nm. However, users of these data should not consider the weak absorption features
475 in the spectra beyond ~1500 nm as diagnostic or consistent with pure, ideal olivine.

476 **7.2 Choice of continuum slope (fitting approach)**

477 Our default choice of the best fit result was that of fit 4, which was a fit with constrained
478 band widths and an enforced flat continuum slope. This was done to enable the most consistency
479 with previous studies (Sunshine and Pieters, 1998). In practice, this meant choosing between fit 2
480 and fit 4, where the only difference was that fit 2 allowed the continuum slope to deviate from an
481 enforced “flatness” constraint, because, as mentioned above, valid results were obtained only

482 from the constrained fits (2 and 4). We only obtained two results in which both fits produced
483 equally good results: C1DD98/Fo₀ and C1DD42/Fo₄₀. In all other cases (for the results shown in
484 Figure 4 and presented in Table 4), the selected fit was clearly the best based on the band
485 positions and relative band strengths. While it would be desirable to have all of the MGM results
486 derive from the exact same starting condition, the goal of using only the best-quality fits is the
487 most important criterion in determining the best results for this study, and we do not believe that
488 the few instances in which a slightly different continuum slope was used impart a major error to
489 our results. For the spectra in which fit 2 was demonstrably superior, the reason was typically
490 that the long-wavelength portion of the spectrum exhibited spectral structure inconsistent with
491 pure olivine (i.e., contamination of some sort). This behavior can be observed in C1DD97/Fo₁₀ in
492 Figure 4, in which a positive slope can be observed in the spectrum beyond ~1800 nm. The small
493 sample size (6) of spectra surviving the quality assessment makes it challenging to determine if a
494 systematic effect results from the use of fit 2 vs. fit 4. The two spectra for which fit 2 produced
495 the best results are those for which the Fe-rich “curve” is observed (sections 5, 7.3). However, a
496 more pronounced “curve” for Fe-rich samples was observed by Sunshine and Pieters (1998), and
497 their results were produced using flat continuum slopes for all samples. Thus, we do not believe
498 that the use of fit 2 vs. fit 4 introduces a bias in the results, although the small sample size of our
499 suite of spectra surviving the quality assessment filtering makes this conclusion challenging to
500 test. A more complete suite of olivine compositions with spectra able to pass our strict quality
501 criteria would allow this conclusion to be tested, and would provide additional statistics for
502 refining the trends in olivine spectral properties with changing composition identified in this
503 study.

504 **7.3 Re-evaluation of Fe-rich samples**

505 After analysis with the standard series of fits described above, we observed a curve in the
506 M1-1 absorption on the band position vs. composition plot (Figure 6) for the Fe-rich
507 compositions. A similar, albeit more extreme, curve can be observed in the results of Sunshine
508 and Pieters (1998) shown in Figure 6. This trend is suspect, because no comparable curve on this
509 plot is observed for the M1-2 absorption, and transitions in the same cation site control both
510 absorptions. Sunshine and Pieters (1998) attributed this curve to noise in the spectra measured of
511 those compositions. We cannot envision a physical reason (with a basis in mineral physics) that
512 such a trend should be observed in the M1-1 absorption but not in the M1-2 absorption. In an
513 attempt to determine if this “curve” is due to a bias introduced by our use of initial conditions
514 more suited to intermediate compositions, we performed another series of fits using a new
515 starting point for the M1-1 absorption determined by extrapolating a linear trend from the band
516 positions of the more Mg-rich samples (Fo₃₀-Fo₇₀).

517 We found that in fits where we allowed the M1-1 position to deviate from its new
518 (extrapolated) initial value, the resulting positions were essentially identical to those obtained
519 from the standard series of fits (the results shown in Figures 3, 5, and 6, in which the curve in
520 M1-1 band positions is observed for the Fe-rich samples). We were not able to obtain
521 satisfactory fit results in which this “curve” in M1-1 band positions was not observed when we
522 allowed the band positions to deviate from these extrapolated initial positions. To obtain a result
523 in which the same “curve” was not observed, we had to constrain the position of the M1-1
524 absorption to remain at its new initial value. Relying on the constrained results would, in effect,
525 mean that the results for the Fe-rich samples would be determined by the results of the more Mg-
526 rich samples, and would carry an implicit *a priori* assumption that the Fe-rich samples should
527 have similar trends to the Mg-rich samples. While this is a reasonable assumption, some research

528 (using mid-infrared instead of VNIR data) suggests that Mg-rich olivines may behave slightly
529 differently from more intermediate and Fe-rich olivines (Hofmeister and Pitman, 2007; Pitman et
530 al., 2010). However, other mid-infrared studies of olivine [including work by Lane et al. (2011)
531 on some of the same samples studied in this work] did not find a need to define distinct trends
532 for Mg-rich samples (e.g., Koike et al., 2003; Lane et al., 2011). Thus, we do not present
533 alternate trend lines for this new series of fits, and our conclusion is that the best and most
534 testable results are those presented in Figures 5 and 6 and in the “Present Study” rows of Tables
535 5 and 6. Our tests did not suggest that this curve in the M1-1 band positions for Fe-rich samples
536 could be attributed to bias resulting from our choice of initial conditions. However, we note that
537 the deviation from the trend in the M1-1 band positions for the Fo₁₀ and Fo₂₀ samples is much
538 less extreme than that observed by Sunshine and Pieters (1998), and might not even be apparent
539 when not compared directly to the Sunshine and Pieters (1998) results. Furthermore, the data
540 point that deviates most strongly from the trend line (Fo₂₀) is one of the most suspect points in
541 the study, as discussed previously. We conclude that this slight “curve” in M1-1 band position of
542 the Fe-rich samples is likely an artifact of our fitting procedure and not attributable to a physical
543 property of Fe-rich olivine, effectively consistent with the conclusions of Sunshine and Pieters
544 (1998).

545 **7.4 Implications: Definitive standard for olivine optical properties in the VNIR**

546 Our analyses suggest that the results of this study can be considered the new standard in
547 the optical properties of olivine across VNIR wavelengths. Our results build on and refine
548 previous attempts to define these properties. In particular, we have refined the state-of-the-art in
549 understanding of the position and characteristics of the three diagnostic olivine absorption
550 features near 1 μm over the compositional range from Fo₀ to Fo₇₀. In the process of quantifying

551 these properties, we have identified a suite of synthetic olivine samples that can be considered
552 the best examples of the optical properties of pure stoichiometric olivine across these
553 wavelengths. These samples and their reflectance spectra represent the best examples of pure
554 olivine unaffected by the inclusion of minor cations, particularly across the critical 1 μm region.
555 While natural olivines are unlikely to be entirely pure and thus may not exhibit identical
556 properties to these samples, our results represent the behavior of the most “ideal” olivines, and
557 thus they are a standard to which natural samples can be compared that is not biased by minor
558 cations. However, the suite of olivine samples used to define these important trend lines is small
559 due to our strict quality criteria. A sample suite with a broader range of compositions
560 (particularly at intermediate and Mg-rich compositions) and spectra able to meet these quality
561 criteria would improve on the results of this study and should be a goal for future research.
562 Nevertheless, the synthesis of our results, captured in Tables 7 and 8, can be used to evaluate the
563 composition of olivine on planetary surfaces with VNIR reflectance spectroscopy.

564 **8 Acknowledgments**

565 We thank Donald Lindsley and Richard Brooker for synthesizing the samples used for this work.
566 Constructive reviews from Harold Clenet and Michail Taran greatly improved the quality of this
567 publication. This work was supported by NASA grants NNX07AR66G and NNA09DB34A. The
568 NASA RELAB is a multiuser facility supported through the NASA Lunar Science Institute
569 under NNA09DB34A. This is HIGP publication number 2015, SOEST publication number 8986,
570 and PSI contribution number 609.

571 **9 References**

572 Abdu, Y., Annersten, H., Ericsson, T., and Hawthorne, F. (2008) High-temperature cation
573 ordering in olivine: an *in situ* Mössbauer study of synthetic $(\text{Mg}_{0.55}\text{Fe}_{0.45})_2\text{SiO}_4$. Hyperfine
574 Interactions, 186(1), 99-103.

- 575 Birlan, M., Nedelcu, D.A., Descamps, P., Berthier, J., Marchis, F., Merouane, S., and Popescu,
576 M. (2011) Spectral properties of (854) Frostia, (1333) Cevenola and (3623) Chaplin.
577 Monthly Notices of the Royal Astronomical Society, 415(1), 587-595.
- 578 Bish, D.L. (1981) Cation ordering in synthetic and natural Ni-Mg olivine. American
579 Mineralogist, 66, 770-776.
- 580 Brown, G.E. (1970) The Crystal Chemistry of the Olivines, Ph.D., p. 121. Virginia Polytechnic
581 Institute and State University, Blacksburg, VA.
- 582 Burns, R.G. (1970) Crystal field spectra and evidence of cation ordering in olivine minerals.
583 American Mineralogist, 55, 1608-1632.
- 584 --- (1974) The polarized spectra of iron in silicates: Olivine. A discussion of neglected
585 contributions from Fe^{2+} ions in M(1) sites. American Mineralogist, 59, 625-629.
- 586 --- (1975) On the occurrence and stability of divalent chromium in olivines included in
587 diamonds. Contributions to Mineralogy and Petrology, 51(3), 213-221.
- 588 --- (1993) Mineralogical applications of crystal field theory. 551 p. Cambridge University Press,
589 New York.
- 590 Burns, R.G., Huggins, F.E., and Abu-Eid, R.M. (1972) Polarized absorption spectra of single
591 crystals of lunar pyroxenes and olivines. Moon, 4, 93-102.
- 592 Clénet, H., Pinet, P., Daydou, Y., Heuripeau, F., Rosemberg, C., Baratoux, D., and Chevrel, S.
593 (2011) A new systematic approach using the Modified Gaussian Model: Insight for the
594 characterization of chemical composition of olivines, pyroxenes and olivine-pyroxene
595 mixtures. Icarus, 213(1), 404-422.
- 596 Cloutis, E.A. (1997) Manganese-rich olivines: Identification from spectral reflectance properties.
597 Journal of Geophysical Research: Planets, 102(E11), 25575-25580.
- 598 Dyar, M.D., Sklute, E.C., Menzies, O.N., Bland, P.A., Lindsley, D., Glotch, T., Lane, M.D.,
599 Schaefer, M.W., Wopenka, B., Klima, R., Bishop, J.L., Hiroi, T., Pieters, C., and Sunshine,
600 J. (2009) Spectroscopic characteristics of synthetic olivine: An integrated multi-wavelength
601 and multi-technique approach. American Mineralogist, 94(7), 883-898.
- 602 Farrand, W.H., Lane, M.D., Edwards, B.R., and Aileen Yingst, R. (2011) Spectral evidence of
603 volcanic cryptodomes on the northern plains of Mars. Icarus, 211(1), 139-156.
- 604 Hiroi, T., and Sasaki, S. (2001) Importance of space weathering simulation products in
605 compositional modeling of asteroids: 349 Dembowska and 446 Aeternitas as examples.
606 Meteoritics and Planetary Science, 36, 1587-1596.
- 607 Hiroi, T., and Takeda, H. (1992) Crystal-field theory calculations for Fe^{2+} Ions in bronzite,
608 augite, and olivine. Physics and Chemistry of Minerals, 19(4), 229-235.
- 609 Hiroi, T., Binzel, R.P., Sunshine, J.M., Pieters, C.M., and Takeda, H. (1995) Grain sizes and
610 mineral compositions of surface regoliths of Vesta-like asteroids. Icarus, 115, 374-386.
- 611 Hofmeister, A., and Pitman, K. (2007) Evidence for kinks in structural and thermodynamic
612 properties across the forsterite-fayalite binary from thin-film IR absorption spectra. Physics
613 and Chemistry of Minerals, 34(5), 319-333.
- 614 Hu, X., Langer, K., and Bostroem, D. (1990) Polarized electronic absorption spectra and Ni-Mg
615 partitioning in olivines ($\text{Mg}_{1-x}\text{Ni}_x$)₂ (SiO_4). European Journal of Mineralogy, 2(1), 29-41.
- 616 Isaacson, P.J., and Pieters, C.M. (2010) Deconvolution of lunar olivine reflectance spectra:
617 Implications for remote compositional assessment. Icarus, 210(1), 8-13.
- 618 Isaacson, P.J., Basu Sarbadhikari, A., Pieters, C.M., Klima, R.L., Hiroi, T., Liu, Y., and Taylor,
619 L.A. (2011a) The Lunar Rock and Mineral Characterization Consortium: Deconstruction

- 620 and Integrated Mineralogical, Petrologic, and Spectroscopic Analyses of Mare Basalts.
621 Meteoritics and Planetary Science, 46, 228-251.
- 622 Isaacson, P.J., Pieters, C.M., Besse, S., Clark, R.N., Head, J.W., Klima, R.L., Mustard, J.F.,
623 Petro, N.E., Staid, M.I., Sunshine, J.M., Taylor, L.A., Thaisen, K.G., and Tompkins, S.
624 (2011b) Remote compositional analysis of lunar olivine-rich lithologies with Moon
625 Mineralogy Mapper (M³) spectra. Journal of Geophysical Research E: Planets,
626 116(E00G11).
- 627 King, T.V.V., and Ridley, W.I. (1987) Relation of the Spectroscopic Reflectance of Olivine to
628 Mineral Chemistry and Some Remote Sensing Implications. Journal of Geophysical
629 Research E: Planets, 92(B11), 11457-11469.
- 630 Klima, R.L., Pieters, C.M., and Dyar, M.D. (2007) Spectroscopy of synthetic Mg-Fe pyroxenes
631 I: Spin-allowed and spin-forbidden crystal field bands in the visible and near-infrared.
632 Meteoritics and Planetary Science, 42, 235-253.
- 633 Koike, C., Chihara, H., Tsuchiyama, A., Suto, H., Sogawa, H., and Okuda, H. (2003)
634 Compositional dependence of infrared absorption spectra of crystalline silicate. A&A,
635 399(3), 1101-1107.
- 636 Lane, M.D., Glotch, T.D., Dyar, M.D., Pieters, C.M., Klima, R., Hiroi, T., Bishop, J.L., and
637 Sunshine, J. (2011) Midinfrared spectroscopy of synthetic olivines: Thermal emission,
638 specular and diffuse reflectance, and attenuated total reflectance studies of forsterite to
639 fayalite. Journal of Geophysical Research E: Planets, 116(E8), E08010.
- 640 Langer, K., Taran, M.N., and Fransolet, A.-M. (2006) Electronic absorption spectra of phosphate
641 minerals with olivine-type structures: I. Members of the triphylite-lithiophilite series,
642 $M^{1[6]}Li^{M2[6]}(Fe_x^{2+}Mn_{1-x}^{2+})[PO_4]$. European Journal of Mineralogy, 18(3), 337-344.
- 643 Lucey, P.G. (1998) Model near-infrared optical constants of olivine and pyroxene as a function
644 of iron content. Journal of Geophysical Research E: Planets, 103(E1), 1703-1713.
- 645 Mayne, R.G., Sunshine, J.M., McSween, H.Y., McCoy, T.J., Corrigan, C.M., and Gale, A.
646 (2010) Petrologic insights from the spectra of the unbrecciated eucrites: Implications for
647 Vesta and basaltic asteroids. Meteoritics & Planetary Science.
- 648 Noble, S.K., Pieters, C.M., Hiroi, T., and Taylor, L.A. (2006) Using the modified Gaussian
649 model to extract quantitative data from lunar soils. Journal of Geophysical Research
650 (Planets), 111, 11009.
- 651 Pieters, C.M. (1983) Strength of Mineral Absorption Features in the Transmitted Component of
652 Near-Infrared Reflected Light: First Results From RELAB. Journal of Geophysical Research
653 E: Planets, 88(B11), 9534-9544.
- 654 Pieters, C.M., and Hiroi, T. (2004) RELAB (Reflectance Experiment Laboratory): A NASA
655 Multiuser Spectroscopy Facility. Lunar and Planetary Science Conference, 35, 1720.
- 656 Pitman, K.M., Dijkstra, C., Hofmeister, A.M., and Speck, A.K. (2010) Infrared laboratory
657 absorbance spectra of olivine: using classical dispersion analysis to extract peak parameters.
658 Monthly Notices of the Royal Astronomical Society, 406(1), 460-481.
- 659 Runciman, W.A., Sengupta, D., and Gourley, J.T. (1974) The polarized spectra of iron in
660 silicates. II. Olivine: A Reply. American Mineralogist, 59, 630-631.
- 661 Sunshine, J.M., and Pieters, C.M. (1993) Estimating modal abundances from the spectra of
662 natural and laboratory pyroxene mixtures using the modified Gaussian model. Journal of
663 Geophysical Research, 98, 9075-9087.
- 664 --- (1998) Determining the composition of olivine from reflectance spectroscopy. Journal of
665 Geophysical Research, 103, 13675-13688.

- 666 Sunshine, J.M., Pieters, C.M., and Pratt, S.F. (1990) Deconvolution of mineral absorption bands:
667 An improved approach. *Journal of Geophysical Research*, 95(B5), 6955-6966.
- 668 Sunshine, J.M., Bus, S.J., Corrigan, C.M., McCoy, T.J., and Burbine, T.H. (2007) Olivine-
669 dominated asteroids and meteorites: Distinguishing nebular and igneous histories.
670 *Meteoritics & Planetary Science*, 42(2), 155-170.
- 671 Taran, M., and Matsyuk, S. (2013) Fe²⁺, Mg-distribution among non-equivalent structural sites
672 M1 and M2 in natural olivines: an optical spectroscopy study. *Physics and Chemistry of*
673 *Minerals*, 40(4), 309-318.
- 674 Tarantola, A., and Valette, B. (1982) Generalized Nonlinear Inverse Problems Solved Using the
675 Least Squares Criterion (Paper 1R1855). *Reviews of Geophysics and Space Physics*, 20,
676 219-232.
- 677
678

679 **1 Figure Captions**

680

681 Figure 1: Illustration of geometry of the M106 and M206 octahedral sites in olivine. These
682 illustrations are viewed down the *a* axis. The specific geometries (i.e., sizes, bond lengths, and
683 deviation of the M06 symmetry from a regular octahedron) are shown for the forsterite [$x =$
684 $\text{Mg}/(\text{Mg}+\text{Fe}) = 1$] endmember. The illustration is based on data reported by Brown (1970).

685

686 Figure 2: Bidirectional reflectance spectra of synthetic olivine subset suite. The geometry for the
687 bidirectional reflectance measurements is discussed in the text. Legend is labeled by Fo# [$x =$
688 $\text{Mg}/(\text{Mg}+\text{Fe})$]. The principal olivine absorption near 1000 nm is composed of three overlapping
689 absorptions, caused by Fe^{2+} ions in the M1 (external absorptions) and M2 (central absorption)
690 sites. The component absorptions are labeled as referred to throughout the paper. The Fe-rich
691 samples (Fo₀ – Fo₃₀) are from Group 1, and the Mg-rich samples (Fo₄₀, Fo₇₀) are from Group 2.

692

693 Figure 3: Examples of MGM fits to synthetic olivine spectra. Fits correspond to the spectra
694 described in Table 4. Note the differing Y-axis scales between the plots, used to maximize the
695 contrast visible in each plot. The residual error in the fits (between model and spectrum) are
696 captured in the dotted RMS lines. Parts a-d are ordered by decreasing “quality” of the
697 spectrum/fit result, with (a) being excellent and (d) being poor. In order, the plots are C1DD98
698 (Fo₀) Fit 4, C1DD95 (Fo₃₀) Fit 2, C1DD96 (Fo₂₀) Fit 2, and C1DD115 (Fo₇₀) Fit 2. Note the
699 sloped continuum in b-d; these fits were produced with the Fit 2 setup, in which the slope is
700 allowed to deviate from the flat initial value. Part a was produced from Fit 4, which forces the
701 continuum slope to remain flat in optimizing the fit. The fit in part d is considered poor and
702 would be rejected from the final analysis for the parameters flagged in bold-face in Table 4 (the
703 M1-1 absorption is too strong, and the M1-2 band position deviates too much from previously-

704 established trends). The problems with the fit in part d are difficult to see on the plot, which
705 emphasizes the importance of quantitative quality threshold criteria.

706

707 Figure 4: MGM fit results for the identified sub-suite of spectra passing our quality criteria. The
708 results shown are presented numerically in Table 5. Note the differing Y-axis scales between the
709 plots, used to maximize the contrast visible in each plot. The residual error in the fits (between
710 model and spectrum) are captured in the dotted RMS lines. The fits are to the spectra illustrated
711 in Figures 2 and 5 and listed in Tables 5 and 6.

712

713 Figure 5: Reflectance spectra of the synthetic olivine subset suite after continuum removal (a) by
714 straight line tangent at 660 nm and 1960 nm and normalization to maximum band depth (b).
715 Legend is labeled by Fo#. The spectra are the same as those shown in Figures 2 and 4.

716

717 Figure 6: MGM-derived band position vs. composition for synthetic olivine subset, compared to
718 the results of Sunshine and Pieters (1998). Our results are in heavy open black symbols, whereas
719 the Sunshine and Pieters results are in small blue symbols. The trend lines calculated from the
720 results of this study are in heavy black lines, and the trend lines from Sunshine and Pieters
721 (1998) are in the dashed grey lines. Error bars are ± 1 standard deviation (σ) in measured olivine
722 composition (Table 6). The parameters for these trend lines are presented in Tables 7 and 8. Our
723 results suggest a slightly shallower slope in the M1-1 position with composition, and we observe
724 a curve in the M1-1 position for Fe-rich compositions, similar (but lesser in magnitude) to that
725 observed by Sunshine and Pieters (1998). Note that the Fo_{20} band positions (which deviated
726 slightly from the best-fit trend line) are determined from a more marginal fit (Table 5).

727

728 Figure 7: MGM-derived parameters for synthetic olivine subset, based on the fits shown in
729 Figure 4. (a) Relative absorption intensity (relative to the M1-2 absorption) vs. band position. (b)
730 Full width at half maximum intensity (FWHM) vs. band position. Points are labeled by Fo#/10,
731 using the same color scheme used for Figures 2 and 4. The overall trends are consistent with
732 those identified by previous studies, in which the M2 absorption weakens relative to the M1-2
733 with increasing olivine Fe content, and the M1-1 stay relatively constant (weakening slightly at
734 very Fe-rich compositions). The widths stay relatively constant, except for a slight increase in the
735 M1-2 feature's FWHM at very Fe-rich compositions.

736

737 Tables

738

739 Table 1: List of samples and compositions evaluated in this study

Sample	Nominal Composition (Fo#)	Suite (1/2)
C1DD37	90	2
C1DD38	80	2
C1DD39	70	2
C1DD40	60	2
C1DD41	50	2
C1DD42	40	2
C1DD43	30	2
C1DD44	20	2
C1DD45	10	2
C1DD46	90	2
C1DD85	100	1
C1DD86	89.5	1
C1DD87	80	1
C1DD88	75	1
C1DD89	70	1
C1DD90	65	1
C1DD91	60	1
C1DD92	55	1
C1DD93	50	1
C1DD94	40	1
C1DD95	30	1
C1DD96	20	1
C1DD97	10	1
C1DD98	0	1
C1DD115	70	1
C1DD116	89.5	1

740

741

742 Table 2: Initial values for individual bands used in MGM fits.

Absorption	Position (nm)	FWHM* (nm)	Intensity (ln reflectance)
1	300	375	-0.5
2	650	115	-0.01
3	750	115	-0.01
M1-1	879	228	-0.15
M2	1054	176	-0.13
M1-2	1254	424	-0.25

743

744

745

* Full width at half maximum strength.

746 Table 3: MGM model starting conditions

Fit Number	Offset Constraint	Slope Constraint	Width constraint
1	Free	Free	Free
2	Free	Free	Constrained*
3	Free	Fixed Flat	Free
4	Free	Fixed Flat	Constrained*

747 *Principal olivine bands (M1-1, M2, M1-2) allowed to vary slightly in width, but were highly
 748 constrained to the relations determined by previous analyses (Sunshine and Pieters, 1998).

749

750 Table 4: Example evaluations of MGM fit results illustrated in Figure 4

Sample/Fit	M1-1 Rel. Str.*	M2 Rel. Str.	M1-1 Deviation**	M2 Deviation	M1-2 Deviation
C1DD98 / 4	0.55	0.47	3.81	-4.93	-6.67
C1DD95 / 2	0.65	0.63	3.06	-0.42	9.40
C1DD96 / 2	0.54	0.48	-9.29	-7.97	-6.67
C1DD115 / 2	0.71	0.80	-7.35	-7.20	-25.36
Quality criterion	0.5-0.65		±10	±7.5	±12.5

751 Shading indicates the quality assessment: white / excellent, light grey / good, dark grey /
 752 mediocre, and black / rejected.

753 *Intensity of absorption (natural log reflectance) divided by intensity of M1-2 absorption

754 **Difference between projected band position at the composition of that sample from trends of
 755 Sunshine & Pieters (1998) and MGM-derived band position. This quantity is effectively the
 756 deviation of the result from the trend lines derived by Sunshine and Pieters (1998).

757

758

759 Table 5: Results of best MGM fits for spectra with MGM results passing the quality tests

Spectrum / Fit	Fo#	Position (nm)			FWHM (nm)			Intensity (ln reflectance)		
		M1-1	M2	M1-2	M1-1	M2	M1-2	M1-1	M2	M1-2
C1DD98 / 4	0	917.1	1073.7	1289.2	228.4	177.7	507.9	-0.49	-0.41	-0.89
C1DD97 / 2	10	899.6	1068.4	1287.2	228.5	178.0	493.8	-0.41	-0.39	-0.72
C1DD96 / 2	20	887.6	1061.6	1269.8	228.6	178.6	529.1	-0.48	-0.42	-0.88
C1DD95 / 4	30	889.2	1062.8	1268.1	228.3	177.5	478.1	-0.35	-0.33	-0.57
C1DD42 / 4	40	876.7	1056.0	1245.7	228.5	178.3	523.6	-0.53	-0.45	-0.97
C1DD39 / 4	70	854.8	1044.7	1232.3	228.2	177.7	493.1	-0.43	-0.42	-0.73

760

761 Table 6: Compositional data for samples passing the quality tests

Name	RELAB ID	Nominal Mg#	Measured Mg#	Mg# Standard Deviation	# Analyses (n) ²
C1DD98	DD-MDD-098	0	0.01	0.02	7
C1DD97	DD-MDD-097	10	10.47	0.19	8
C1DD96	DD-MDD-096	20	21.44	1.72	15
C1DD95	DD-MDD-095	30	30.40	0.29	13
C1DD42	DD-MDD-042	40	41.52	0.26	4
C1DD39	DD-MDD-039	70	65.42	1.42	14

762 ¹RELAB ID's are unique identifiers for these samples in the Brown University RELAB database.

763 ²Analyses reported only if they meet quality criteria (moles of cations per 4 moles oxygen
 764 between 2.98 and 3.02, oxide weight percent totals between 98.0% and 102.0%).

765

766 Table 7: Linear trend data for band position vs. Fo; Band position(Fo)

Sample Suite	Fo Range for Slope	Absorption	Slope	Offset
Sunshine & Pieters	01 - 90	M1-1	-0.64	900.9
		M2	-0.44	1076.4
		M1-2	-0.89	1290.5
Present study	0 - 70	M1-1	-0.89	912.4
		M2	-0.43	1073.2
		M1-2	-0.95	1292.1

767

768 Table 8: Linear trend data for band position vs. Fo; Fo(Band position)

Sample Suite	Fo Range for Slope	Absorption	Slope	Offset
Sunshine & Pieters	01 - 90	M1-1	-1.23	1139.9
		M2	-2.18	2347.8
		M1-2	-1.06	1376.2
Present study	0 - 70	M1-1	-1.09	993.8
		M2	-2.28	2442.9
		M1-2	-1.01	1305.6

769