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2 **Revision 1**

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4 **Pressure-induced phase transitions in coesite**

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18 **ABSTRACT**

19 High-pressure behavior of coesite was studied on single crystals using diamond-anvil cells with
20 neon as the pressure transmitting medium by means of in situ Raman spectroscopy up to

21 pressures of ~51 GPa. The experimental observations were complemented with theoretical
22 computations of the Raman spectra under similar pressure conditions. We find that coesite
23 undergoes two phase transitions and does not become amorphous at least up to ~51 GPa. The first
24 phase transition (coesite I to coesite II) is reversible and occurs around 23 GPa. The second
25 transition (coesite II to coesite III) at about 35 GPa is also reversible but involves a large
26 hysteresis. Samples recovered from the highest pressure achieved, ~51 GPa, show Raman spectra
27 of the initial coesite. The ab initio calculations gave insight into the initiation mechanism of the
28 first phase transition, implying, from the analysis of unstable phonon modes, that it is probably a
29 displacive phase transition due to shearing of the four-membered rings of SiO₄ tetrahedra upon
30 compression. The transition to the lowest-symmetry phase, coesite III, is possibly a first order
31 phase transition which leads to a very distinct structure. None of the metastable high-pressure
32 phases of coesite has been previously studied and it was widely accepted that coesite undergoes
33 pressure-induced amorphization at significantly lower pressures (30 GPa). The study of the high-
34 pressure behavior of coesite is important to better constrain the metastable phase diagram of
35 silica. Further crystallographic investigations are necessary for characterizing the structures of
36 these metastable coesite forms. Crystalline or amorphous metastable phases derived from coesite
37 under high-pressure conditions are of particular interest because they can be used as potential
38 tracers of peak transient pressures (stress) reached in processes such as impacts or faulting.

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40 **Keywords:** silica, coesite, high-pressure polymorph, metastable, diamond-anvil cell, high-
41 pressure Raman spectra

42

INTRODUCTION

43 Silica, SiO₂, exhibits a very rich polymorphism with more than 30 stable or metastable phases,
44 most of which occur at ambient to moderate pressures (< 9 GPa). Some of these phases, such as
45 quartz, tridymite and cristobalite consist of frameworks of SiO₄ tetrahedra and are abundant in
46 nature. Phases of silica stable at higher pressures and temperatures are mostly composed of
47 octahedrally coordinated silicon atoms. Coesite, thermodynamically stable above ca. 2.5 GPa and
48 at temperatures in excess of 500 °C, is the densest known polymorph with silicon atom
49 tetrahedrally coordinated to oxygen. This polymorph was first observed experimentally by Coes
50 in 1953 and later discovered in nature by Chao and coworkers, in impact breccia from Meteor
51 Crater (Chao et al. 1960). Nowadays coesite is widely accepted as a high-pressure indicator in
52 rocks related to meteorite impact sites, where it forms due to the high-pressure and high-
53 temperature regime of the impact and to the very rapid quenching conditions which prevent a
54 reversion from coesite to quartz (e.g., Xiao et al. 2011). Shock-induced formation of coesite has
55 been also observed in meteorites, where it can be found along with other shock minerals (e.g.,
56 Ohtani et al. 2011). Terrestrial occurrence of coesite was reported from deeply (>100 km)
57 subducted crustal rocks (ultra-high pressure metamorphic rocks) or mantle derived rocks
58 (kimberlites), which have been exhumed under conditions that prevent retrograde transition to
59 quartz (e.g., Smyth and Hatton 1977; Chopin 1984; Mosenfelder and Bohlen 1997; Parkinson
60 2000; Korsakov et al. 2007; Ruiz-Cruz et al. 2012). In these rocks coesite commonly occurs as an
61 inclusion within clinopyroxene, olivine, garnet or diamond. Coesite-in-diamond assemblage was
62 recently introduced as an important high-pressure barometer (Sobolev et al. 2000).

63 Because of the relatively strong Si-O bonding in silica, there are high kinetic barriers associated
64 with the transitions to stable high-pressure phases consisting of SiO₆ octahedra (Haines et al.
65 2001). This results in complicated metastable phenomena at high pressures and ambient
66 temperature such as the persistence of low-pressure phases far out of their stability fields and

67 transitions to metastable crystalline and amorphous phases. Slow kinetics, enhanced
68 metastability, formation of poorly crystallized or structurally disordered materials usually give
69 rise to weak X-ray diffraction patterns which are difficult to interpret, resulting in contradictory
70 interpretations. Another major complexity in studying the behavior of silica also arises from the
71 fact that the transition mechanisms, the pathways through intermediate metastable phases and the
72 onset of pressure-induced amorphization strongly depend on the starting polymorph, as well as
73 on the hydrostaticity and stress anisotropy present in the sample during experiments. For
74 example, experimental studies on polycrystalline α -quartz in neon pressure medium (Kingma et
75 al. 1993) evidenced pressure-induced phase transitions above 21 GPa to a crystalline metastable
76 phase of reduced symmetry, followed by pressure-induced amorphization above 30 GPa. In
77 helium pressure medium, which provides higher hydrostaticity, polycrystalline quartz is found to
78 remain crystalline up to 45 GPa and to transform to a monoclinic phase at such pressure (Haines
79 et al. 2001). A transformation path through several metastable phases has been reported in
80 cristobalite, as well (e.g., Yagi and Yamakata 2000; Prokopenko et al. 2001; Dera et al. 2011).
81 Numerical calculations also stressed the fact that stable polymorphs of SiO_2 may undergo a
82 number of phase transitions to metastable phases with increasing pressure before reaching
83 amorphization (Teter and Hemley 1998; Wentzcovitch et al. 1998).

84 In an experimental study on the high-pressure behavior of coesite, Hemley (1987) observed
85 changes in the Raman spectra of coesite collected at ambient temperature and at the pressures of
86 22-25 GPa in argon pressure medium, and reported that coesite becomes amorphous above 30
87 GPa. The authors attributed such changes to a high-pressure phase transformation occurring in
88 coesite; however, no further investigation has been reported to better describe such behavior.

89 In this study we aim at giving insight into the high-pressure behavior of coesite. We measure in
90 situ Raman spectra at pressures up to 51 GPa and complement them with theoretical

91 computations of Raman spectra under similar pressure conditions. We find two phase transitions
92 clearly distinguishable by Raman spectroscopy and explain the mechanism of the first transition
93 based on the theoretical calculations.

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EXPERIMENTAL METHODS

96 **Synthesis**

97 The starting material for the coesite synthesis was SiO₂ glass powder with very low trace
98 elements content, as analyzed at the BGI using LA-ICP-MS: Al 20 ppm, Ge 1.3 ppm, Na 1.0
99 ppm, Li 0.8 ppm, and B, Ti, Fe, Ga, Rb and Sn below the detection limits. Coesite single crystals
100 were synthesized by mixing the starting powder with ~5 wt% distilled water inside a platinum
101 capsule, which was then welded shut. The capsule was first placed into pyrophyllite sleeves and
102 then in a 0.5" talc-pyrex piston-cylinder assembly containing internal, tapered graphite resistance
103 furnaces (Bromiley and Keppler 2004). The mixture was pressurized to 3.5 GPa and slowly
104 heated up to 1250 °C, kept at this temperature for ~15 hours, then cooled down to 1100 °C in 5
105 hours, and finally quenched. Slow cooling procedure and water-saturated conditions resulted in
106 growth of relatively large (above 100 μm in linear dimensions) crystals. No Raman peaks were
107 observed in the spectra of synthesized coesite in the O-H vibration region (2800-3400 cm⁻¹). This
108 is in agreement with the study on pressure dependence of hydrogen solubility in coesite
109 (Mosenfelder 2000; Koch-Müller et al. 2001), according to which no water should be present in
110 coesite synthesized at 3.5 GPa and 1250 °C.

111

112 **High-pressure Raman spectroscopy**

113 Piston-cylinder type diamond anvil cells (DAC) developed at BGI (Kantor et al. 2012) made of
114 high-temperature resistant alloy were used for the high-pressure experiments. Experiments in
115 quasi-hydrostatic conditions were carried out using 250 μm culet size diamonds, a cylindrical
116 pressure chamber of 30-40 μm height and ~ 125 μm diameter drilled in a pre-indented rhenium
117 gasket. We used neon as pressure transmitting medium, loaded using the BGI gas loading system
118 (Kurnosov et al. 2008). Selected crystals were in the form of plates of typical size 15 x 30 x 50
119 μm^3 and they were loaded together with a ~ 5 μm in diameter ruby sphere for pressure
120 determination. The pressure was determined using the ruby luminescence line (Mao et al. 1986).

121 Raman spectroscopy measurements were performed using a LabRam system (Horiba Scientific
122 Inc.) with a He-Ne-laser (excitation wavelength 632.8 nm, output power 0.15 W, spectral
123 resolution 2 cm^{-1}) as well as Dilor XY Raman spectrometer with Ar⁺ ion laser (514.5 nm,
124 Coherent Innova 300, spectral resolution 1 cm^{-1}). The output power was varied according to the
125 sample signal between 0.3 and 1.0 W when using the Dilor instrument. The spectrometers were
126 calibrated using either the silicon peak at 520 cm^{-1} or referencing the value of the absolute
127 wavelength of the ruby R₁ fluorescence line. Raman spectra were collected between 200 and
128 1200 cm^{-1} . This region is optimal because the LabRam spectrometer used in our experiments is
129 equipped with a notch filter which prevents collection of the frequencies lower than 150 cm^{-1} ,
130 whereas above 1200 cm^{-1} low intensity peaks of coesite are suppressed by strong diamond peak
131 at centered near 1300 cm^{-1} . We performed five different DAC experiments at room temperature
132 by pressurizing the coesite crystals in steps of several GPa up to a maximum of ~ 51 GPa. In two
133 of these experiments, Raman spectra were also collected during decompression. Peak positions in
134 the Raman spectra were determined by fitting to Lorentzian peak shape using Igor Pro v. 6.22
135 software.

136 **External electrical heating in DAC**

137 Heating experiments up to ~400 °C were performed using platinum-wire resistive whole-cell
138 heater. The heater has an external diameter of 50 mm, it is made of fired pyrophyllite and 0.5 mm
139 in diameter platinum wire is folded over the inner surface of the heater. For the purpose of
140 protecting the Raman spectrometer from heating over several hours, we used a thermal insulator
141 that is made of a square-shaped double-layered stainless steel box. Stainless steel of “type
142 1.4571” is oxidation-resistant at elevated temperatures and shows no oxidation effect to at least
143 600 °C (Fig. 1). The DAC and the whole-cell heater were fixed on a stainless steel holder, which
144 was then placed inside the steel box with openings at the bottom and the top, to allow optical
145 access. Temperature was measured using a Pt Pt/Rh S-type thermocouple placed in the vicinity of
146 the gasket.

147 In the first experiment, the coesite sample was pressurized at room temperature up to 29 GPa and
148 then slowly heated up to 400 °C. During heating up to 200 °C the pressure dropped down to 27
149 GPa and was readjusted before further heating. In a second experiment, the coesite sample was
150 pressurized from room pressure up to 34 GPa and then slowly heated up to 200 °C. At this
151 temperature, however, the Raman spectrum was very poor and therefore the experiment was
152 interrupted. Pressure was monitored from a position of the ruby fluorescence peaks corrected for
153 temperature after Rekhi et al. (1999).

154

155 **Transmission Electron Microscopy**

156 Conventional TEM characterizations in bright-field and dark-field imaging and selected electron
157 diffraction techniques were carried out with a Philips CM20FEG equipped with an energy
158 dispersive X-ray (EDX) analyzer, operated at 200 kV. Analyses were performed on several chips

159 of a crystal which was subjected to a maximum pressure of 35 GPa. The recovered material was
160 first characterized by Raman spectroscopy and then further crushed between two tungsten carbide
161 plates. The ethanol dispersion was placed on a lacey carbon film on Cu-grid (300 mesh). Selected
162 area electron diffraction (SAED) patterns from different chips were obtained to distinguish
163 between crystalline and amorphous domains or particles. The chemistry of the grains was also
164 cross-checked using EDX spectroscopy.

165 **X-ray diffraction**

166 We selected coesite single crystals at BGI using a rotating anode high-brilliance Rigaku
167 diffractometer with Mo K α radiation, equipped with Osmic focusing X-ray optics and Bruker
168 Apex CCD detector.

169 The crystal recovered after a repeated experiment at 34 GPa and \sim 300 °C was analyzed by means
170 of single-crystal X-ray diffraction at the beamline 13IDD of the GSECARS at the Advanced
171 Photon Source, Argonne National Laboratory, USA. An incident monochromatic beam with
172 wavelength of 0.3344 Å was focused to a spot of 4x8 μm^2 . Diffraction images were collected
173 using a MAR165 CCD detector with sample-to-detector distance of approximately 250 mm. The
174 sample was also analyzed by X-ray diffraction mapping along a 5 x 5 points, covering an area of
175 12 x 12 μm^2 (step size 3 μm), with exposure time of 10 s at each point. Diffraction images were
176 integrated using the Fit2d software (Hammersley 1996).

177 **Ab initio calculations**

178 We analyze the pressure dependence of the phonon spectra also from the first-principles
179 calculations. We determine the ground-state properties using standard density-functional theory
180 (Kohn and Sham 1965; Payne et al. 1992; Martin 2003) in the ABINIT implementation, based on

181 planewaves and pseudopotentials (Gonze et al. 2002, 2009). Starting from the crystal structure
182 refined by Angel et al. (2003) we determine the theoretical structure of coesite up to 40 GPa in 10
183 GPa steps. Then we compute the energy derivatives to build the dynamical matrices and the
184 Raman tensors in the framework of the density-functional perturbation theory (Baroni et al. 2001;
185 Gonze et al. 2005; Veithen et al. 2005). We employ a 4x4x4 grid of special k points (Monkhorst
186 and Pack) to sample the electron density in the reciprocal space and a kinetic energy cutoff of 38
187 Hartrees (1 Hartree = 27.2116 eV). With these parameters the precision of the calculation is
188 typical on the order of 0.001 Hartree in energy and better than 1 GPa in pressure. We store all the
189 Raman spectra computed under pressure on the WURM website (<http://www.wurm.info>). More
190 details of the Raman calculations can be found in the original WURM paper (Caracas and
191 Bobocioiu, 2011).

192 RESULTS

194 **High-pressure, room temperature Raman spectroscopy**

195 Selected Raman spectra collected during compression are shown in Figure 2. At ambient pressure
196 coesite shows prominent peaks at 202, 261, 322, 352, 371, 425, 469 and 519 cm^{-1} , with the last
197 peak being the most intense. The intensities and positions of these peaks are in good agreement
198 with previous studies (Sharma et al. 1981; Boyer et al. 1985; Hemley 1987). The peaks that are
199 expected to be centered around 665 cm^{-1} and 785 cm^{-1} at ambient pressure have only been
200 observed at elevated pressures and not in all experiments. A peak located at 815 cm^{-1} was
201 observed starting from ambient conditions, but was not detected in all experiments. Another
202 peak, not previously reported in high-pressure Raman spectroscopy studies of coesite, was
203 observed in two experiments. It appeared at and above 4.5 GPa, and, according to its pressure

204 shift, it can be expected to occur around 860 cm^{-1} at ambient conditions. Raman bands at 1036
205 and 1164 cm^{-1} were observed at ambient conditions outside the DACs, but not under compression
206 due to their very low intensities. All the vibration bands are significantly weaker than the main
207 519 cm^{-1} band, they all show continuous positive pressure shift during compression and can be
208 followed up to $\sim 28\text{ GPa}$ (Fig. 2a). The bands located at 519, 665, 785 and 815 cm^{-1} have the
209 largest pressure shifts, in good agreement with the previous study by Hemley (1987).

210

211 The most intense 519 cm^{-1} band was described as $\nu_s(\text{Si-O-Si})$ symmetric stretching mode with A_g
212 symmetry (Sharma et al. 1985, and references therein). It shows continuous positive pressure
213 increase up to 23 GPa . Above 23 GPa a new strong vibrational band appears at a frequency of
214 about 20 cm^{-1} lower than the main Raman peak, which at this pressure has a frequency of 593 cm^{-1} ,
215 as was also observed by Hemley (1987). Several other new modes become visible at lower
216 frequencies, confirming the results reported by Hemley (1987) and their interpretation of a
217 possible phase transition to a phase with reduced symmetry or enlarged unit cell. This post-
218 coesite phase, hereafter referred as coesite II, can be followed upon compression up to $\sim 35\text{ GPa}$
219 (Fig. 2b). The A_g mode persists up to $\sim 35\text{ GPa}$, although its intensity decreases significantly.
220 Above $\sim 35\text{ GPa}$, the two high-frequency peaks disappear; only one new peak centered around
221 $\sim 600\text{ cm}^{-1}$ can be observed (Fig. 2b). This sudden change in the Raman spectrum indicates a
222 possible discontinuous transition to a new phase, hereinafter referred as to coesite III. On further
223 compression up to $\sim 51\text{ GPa}$, the peak centered around $\sim 600\text{ cm}^{-1}$ decreases in intensity and
224 exhibits a very small pressure increase (Fig. 2b). Above $\sim 35\text{ GPa}$, in addition to the peak centered
225 around $\sim 600\text{ cm}^{-1}$, new bands appear as two triplets of equally spaced peaks between $\sim 250\text{-}300$
226 cm^{-1} and between $\sim 620\text{-}720\text{ cm}^{-1}$. Appearance of a large number of new vibrational modes may

227 indicate that coesite III has either an enlarged unit cell or lower symmetry with respect to coesite
228 I and II structures. At ~51 GPa no vibrational modes can be observed. The pressure dependence
229 of the main 519 cm^{-1} Raman mode (dv/dP , where v is the ambient-pressure phonon frequency
230 and P is the pressure) was derived from a series of experiments and it is, on average, 3.0 ± 0.4
231 $\text{cm}^{-1}/\text{GPa}$ (Fig. 3). Based on the zero-pressure bulk modulus $K_{T0} = 100.8 \pm 0.5 \text{ GPa}$ reported by
232 Angel et al. (2001) and using the formulation of the mode-Grüneisen parameter as $\gamma_v =$
233 $(K_{T0}/v)(dv/dP)$, we obtain a value of 0.58 ± 0.08 for the A_g mode, only slightly higher than the
234 values reported in Hemley (1987).

235 At the highest pressure reached in these experiments, about 51 GPa, the Raman spectrum of
236 coesite almost vanishes. Under decompression the vibrational modes of coesite III regain
237 intensity, with the most prominent peaks visible at ~250, ~540 and ~580 cm^{-1} . Their frequencies
238 decrease continuously down to ~24 GPa without any abrupt change, suggesting that the structure
239 of coesite III is preserved down to 24 GPa and that the transition to the coesite II phase is not
240 reversible or has a very large hysteresis. A strong single band centered around 600 cm^{-1} appears
241 at ~24 GPa and its intensity increases with further decompression, until it dominates the Raman
242 spectra below ~20 GPa, suggesting that the structure of coesite has been retrieved. Indeed, the
243 Raman spectrum of the recovered crystal at room pressure is identical to that of the starting
244 material (Fig. 4). The fact that the pressure-quenched material is crystalline coesite strongly
245 suggests that the high-pressure phase coesite III was also crystalline at 51 GPa even if the Raman
246 vibrations could not be observed.

247 *Raman active modes from ab initio calculations*

248 We compute the phonons in the Brillouin zone-center, with both frequencies and atomic
249 displacement patterns, for coesite I up to 40 GPa. We equally determine the Raman tensors and

250 compute the Raman intensities for the Raman-active modes. All the spectra are reported on the
251 WURM website (<http://www.wurm.info>), but for sake of clarity only the spectrum at ambient
252 pressure is represented here, in Figure 5a. Figure 5b shows pressure dependence of selected most
253 dominant modes. The Raman spectra are dominated by the A_g mode situated at $\sim 510\text{ cm}^{-1}$ at
254 ambient pressure. A series of low intensity peaks are observed at frequencies higher than 600 cm^{-1} ,
255 1 , in excellent agreement with the experimental data.

256 The strongest A_g mode, located at $\sim 510\text{ cm}^{-1}$ at ambient pressure (at 541 cm^{-1} at 10 GPa) is a
257 breathing mode of the four-membered rings comprised of SiO_4 tetrahedra (Fig. 6a) and it arises
258 from bending of four different Si-O-Si bonds that vibrate in phase, namely Si-O2-Si, Si-O3-Si,
259 Si-O4-Si and Si-O5-Si (nomenclature after Angel et al. 2003, see Figures 2 and 4 in the cited
260 reference). Except for the Si-O4-Si bond angle, which changes by only a few degrees throughout
261 the pressure range examined by Angel et al. (2003), other three bond angles show strong pressure
262 dependence. In contrast to these four angles, Si-O1-Si is rigid due to symmetry constrains and
263 remains 180° throughout the investigated range. Atomic movements of this breathing A_g mode
264 are illustrated in a representative four-membered ring in Figure 6a seen during contraction. The
265 largest displacement is seen in O3 and O4 atoms, which are moving symmetrically inward, and in
266 O5 atoms, that are moving parallel to each other along a-axes, but in the opposite direction.

267 A B_g mode that appears at 77 cm^{-1} at ambient conditions (also 77 cm^{-1} at 10 GPa and 56 cm^{-1} at
268 20GPa) is dominant in the lower frequency region and shows softening with increasing pressure
269 (Fig. 5b), similar to previous experimental studies (Hemley 1987). This mode corresponds to
270 shearing of the tetrahedra rings (Fig. 6b and c) and may have a significant influence on the
271 structural distortion at elevated pressures. Shearing of the four-membered rings is illustrated by
272 the two adjacent four-membered rings shown in Figure 6b. Within the same ring O3 atoms are

273 displaced in the same direction, but this displacement direction is opposite between the two
274 adjacent rings. O2 atoms vibrate opposite to the O3 displacement direction. Within the ring, one
275 of the O4 atoms moves downwards and the other upwards from the plane of the ring. All O5
276 atoms of the structure move in the same direction, in the plane of the SiO₄ rings. Similar intensity
277 variation is observed in an A_u mode that appears at 115 cm⁻¹ at ambient pressure (Fig. 5b). Above
278 20 GPa this infrared active mode becomes even more unstable than the Raman active 77 cm⁻¹
279 mode; this enhances the displacement of the O1 atom from the rigid 180° Si-O1-Si bond (Fig.
280 6d). Then this displacement could initiate the structural changes that we observe above 24 GPa.
281 However, an IR study is required in order to confirm this prediction. For a better visualization of
282 the atomic motions, please visit <http://www.wurm.info>.

283 Other two modes in the 400-500 cm⁻¹ region (Fig. 5a) are correlated with the motion of the SiO₄
284 tetrahedra rings, although they have low intensity. One is a complex B_g mode at 430 cm⁻¹,
285 corresponding to both breathing and shearing of the SiO₄ tetrahedra rings, and the other one is an
286 A_g mode at 455 cm⁻¹, which arises from stretching of the rings. The 430 cm⁻¹ mode has not been
287 observed experimentally, but the 455 cm⁻¹ mode, albeit very weak, was present in some of our
288 spectra.

289 **High-temperature, high-pressure Raman spectra**

290 The Raman spectra collected at 27-29 GPa up to 400 °C and at 34 GPa up to 200 °C are shown in
291 Figure 7a and b, respectively. The first heating experiment was carried out in the pressure range
292 where coesite II phase (Figs. 2a and 2b) was observed, and, as expected, the main feature in the
293 Raman spectra is the peak doublet located ~600 cm⁻¹ (Fig. 7a). The intensity of the vibrational
294 modes decreases with increasing temperature, and at 27 GPa, as expected, they shift to lower
295 frequency. However, at 29 GPa the higher frequency peak shifted 4 cm⁻¹ towards higher

296 frequency between room temperature and 425 °C, whereas the lower frequency band softened of
297 $\sim 1 \text{ cm}^{-1}$, increasing in this way the total spread of the doublet. At this temperature an intermediate
298 peak appears in the spectra but due to the very poor quality of the spectra it is hard to tell if it is a
299 signal from the sample or a spike. At higher temperatures the intensity of the doublet decreased
300 significantly and the doublet peak was not observable.

301 At 34 GPa and 100 °C, the peak doublet of coesite II is still present and decreases both in
302 intensity and in frequency with increasing temperature. No signal was observed at 200 °C (Fig.
303 7b). Raman spectra were collected during decompression at elevated temperature, but no Raman
304 spectra of coesite were observed. At room pressure and after cooling down to room temperature,
305 a broad peak close in position to the main vibration mode of coesite was observed (Fig. 7b). The
306 recovered crystal was also analyzed by single-crystal X-ray diffraction (see further).

307 **TEM observations**

308 One of the recovered samples was investigated by transmission electron microscopy. This sample
309 was selected for the TEM analysis because the central, thickest part of the crystal was bridged
310 between diamond anvils at $\sim 35 \text{ GPa}$ during compression, whereas the thinner edges of the crystal
311 remained largely undisturbed by the diamond culets. Raman spectra collected on the material
312 recovered from the thin edge of the sample show a weak peak around 425 and 514 cm^{-1} , and the
313 TEM analysis of the same material revealed that very thin grain domains are indexed with coesite
314 in the SAED pattern (Fig. 8). On the other hand, the material recovered from the bridged part of
315 the crystal does not show any Raman active modes and appears entirely amorphous at the TEM.
316 Our observations agree qualitatively with the previous studies on high-pressure silica behavior: in
317 experiments with stiff pressure media where amorphization is initiated at lower pressures with
318 respect to experiments in quasi-hydrostatic conditions (e.g., Hemley 1987; Haines et al. 2001).

319

320 **X-ray diffraction**

321 X-ray mapping along 25 spots of the sample quenched from ~300 °C and 34 GPa indicates that
322 the material is well-crystallized. All observed diffraction peaks belong to coesite (Fig. 9). Phase
323 identification was carried out using the Match! Software v 1.11 (with Crystallography Open
324 Database), considering only the region of d-spacing between 4.5 and 1.0 Å (2theta 5 – 19°). The
325 suggested possible phases by the program are coesite, diamond and quartz, however all of the
326 peaks assigned to quartz have d-spacing which can also be assigned to coesite, and no peak
327 belonging solely to quartz has been observed (Fig. 9).

328

DISCUSSION

329 Coesite appears to undergo at least two distinct transitions to metastable phases upon
330 compression, as shown in DAC experiments with quasi-hydrostatic medium (Ne) up to ~51 GPa.
331 A similar high-pressure behavior, i.e., complex path of transformations to metastable phases
332 before amorphization, have been reported for quartz (Kingma et al. 1993; Haines et al. 2001) and
333 it is in agreement with predictions from numerical models, which suggest that a number of
334 metastable phases can appear on compression of SiO₂ materials before amorphization
335 (Wentzcovitch et al. 1998; Dubrovinsky et al. 2004). Most certainly these phases appear as a
336 consequence of the lack of energy at room temperature to overcome the considerable potential
337 barrier needed to increase the silicon coordination by oxygen from 4 to 6.

338 The transition from coesite I to coesite II at around 23 GPa was initially reported by Hemley
339 (1987). The study focused on the phase transition, showing its relation to phonon softening, but
340 did not go further into clarifying the eventual splitting of the most intense Raman doublet above

15

341 ~23 GPa. The author also reports one single broad peak above 30 GPa centered around 600 cm^{-1} ,
342 interpreted as a feature similar to that reported for silica glass, and therefore concludes that above
343 such pressure coesite becomes amorphous. Hemley (1987) also reported that the splitting at 23
344 GPa was obscured in all experiments except in those with argon, which was providing the best
345 quasi-hydrostatic conditions in their experiments at higher pressures. Neon is shown to be a softer
346 pressure-transmitting medium and to give quasi-hydrostatic conditions at higher pressures than
347 argon (Klotz et al. 2009). In our experiments, therefore, we are able not only to clearly follow the
348 evolution of the splitting of the main Raman mode above 23 GPa, but also to observe another
349 phase transition at ~35 GPa. Non-hydrostatic compression, as we observed on the example of
350 partially bridged coesite crystal as well as reported in the experiments with Ar or stiffer pressure
351 media (Hemley 1987), leads to amorphization of coesite at lower pressures at room temperature.

352 Our experimental and theoretical results are in very good agreement, with a discrepancy in
353 frequencies of about 10 cm^{-1} between the calculated and the measured spectra. The main Raman
354 mode is at 519 cm^{-1} in the experimental spectra and at 510 cm^{-1} in the calculated; another intense
355 A_g mode which is observed at 426 cm^{-1} is calculated at 415 cm^{-1} . Moreover, the calculated
356 frequency shifts with pressure are very similar to those observed experimentally, at least for the
357 strongest Raman modes.

358 At 8 GPa, a small shoulder at 228 cm^{-1} (Fig. 2a) can be observed overlying the 220 cm^{-1} peak
359 (202 cm^{-1} at ambient pressure). From a previous study (Hemley 1987) it appears that the mode at
360 176 cm^{-1} intersects the 202 cm^{-1} mode at ~8 GPa, and therefore it could be the cause of the
361 shoulder in the spectra we recorded at this pressure. However, it should be noted that the results
362 of our ab initio calculations are in contrast with such an interpretation since these two calculated

363 modes do not cross at any pressure (calculated up to 40 GPa), although their separation decreases
364 with increasing pressure.

365 In light of our experiments and ab initio calculation we can give an insight on the mechanism of
366 transformation from the coesite I to coesite II phase in terms of a complex phonon softening. The
367 crystal structure of coesite II has a lower symmetry than coesite I, as it appears from the
368 increasing number of the vibrational modes observed in this study. The calculations show that
369 three phonon modes, one A_g , one A_u and one B_u soften under compression (Fig. 4b), the former
370 two becoming imaginary around 27 GPa. They all correspond to some form of shear of the SiO_4
371 tetrahedra rings, but the A_u and B_u modes are only infrared active. Freezing of any or all of these
372 modes could enhance a displacive transition and may be associated with a breaking of the
373 symmetry of the four-membered rings. This can also give rise to two slightly different breathing
374 modes for two differently sheared rings, a consequence of which is the splitting of the main A_g
375 mode (at 510 cm^{-1} at room pressure) associated with the breathing of the four-membered rings.

376 An alternative explanation is that the second Raman peak at 560 cm^{-1} in the coesite II structure
377 arises from the coupling between the soft B_g mode, situated at 77 cm^{-1} at ambient conditions and
378 imaginary at the transition, with either the B_g mode at 430 cm^{-1} (at ambient conditions, 450 cm^{-1}
379 at 20 GPa) or the A_g mode at 455 cm^{-1} (at ambient conditions, 468 cm^{-1} at 20 GPa). In particular
380 the coupling and further energy transfer between the two B_g modes is very likely due to their
381 identical symmetry. This transfer can also enhance the Raman intensity of the higher-energy
382 mode. Coupling of the unstable B_g and the stable A_g modes is possible because of similarities
383 between their eigendisplacements. The coupling could also enhance the polarizability of the
384 higher-energy mode and hence its Raman signature.

385 The transition from coesite I to coesite II is likely to be second order in character because of the
386 phonon softening, continuous behavior of the main Raman peak, and because of the absence of
387 hysteresis. Decrease in intensity and softening of the main peak doublet during heating, support
388 the observation that coesite II is a metastable phase, which exists only at room temperatures far
389 outside the pressure stability field of coesite. The heating which is insufficient to transform
390 coesite to the thermodynamically stable phase – stishovite – enhances the thermal motion of the
391 metastable structure and causes instability but, most probably, does not lead to amorphization.
392 Hence, the pressure- and temperature-quenched sample retransforms to coesite. The anomalous
393 spreading of the doublet and shifting to the higher frequency at 29 GPa and 400 °C requires
394 further understanding and structural analysis of coesite II.

395 The transition of coesite II to coesite III is possibly first order in character, as it appears from the
396 abrupt change in Raman spectra above ~35 GPa (Fig. 2b). As expected in a first order transition,
397 the frequency of the main mode, situated at ~600 cm⁻¹, does not vary significantly with pressure
398 (Fig. 4) and a hysteresis is probably hindering reversible transformation from coesite III to
399 coesite II.

400 The mechanisms of this transition cannot be revealed from ab initio calculation, since the
401 structure of the two polymorphs is still unknown. If we assume that the major peak observed in
402 the Raman spectra is still correlated to the O-Si-O motions as the A_g mode, the major decrease in
403 frequency may suggest a softer Si-O bond and, hence an increase in coordination number of the
404 Si. Coesite III would therefore have at least some of Si in octahedral coordination. However, to
405 confirm such hypothesis, a structural study is required.

406

IMPLICATIONS

407 The study of the high-pressure behavior of coesite reveals two high-pressure polymorphs that
408 remain crystalline far above the previously reported amorphization pressure. This significant
409 discrepancy between previous and our current results implies that the metastable phase diagram
410 of silica phases requires a substantial revision. Crystalline or amorphous metastable phases
411 derived from coesite under high-pressure conditions can be of interest in geological application as
412 potential tracers of peak transient pressure conditions reached in processes such as impacts or
413 faulting.

414

415

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427 **REFERENCES CITED**

- 428 Angel, R.J., Mosenfelder, J.L., and Shaw, C.S.J. (2001) Anomalous compression and equation of
429 state of coesite. *Physics of the Earth and Planetary Interiors*, 124, 71–79.
- 430 Angel, R.J., Shaw, C.S.J., and Gibbs, G.V. (2003) Compression mechanisms of coesite. *Physics*
431 *and Chemistry of Minerals*, 30, 167–176.
- 432 Araki, T. and Zoltai, T. (1969) Refinement of a coesite structure. *Zeitschrift für Kristallographie*,
433 129, 381-387.
- 434 Baroni, S., de Gironcoli, S., Dal Corso, A., and Giannozzi, P. (2001) Phonons and related crystal
435 properties from density-functional perturbation theory. *Reviews in Modern Physics*, 73,
436 515–562.

- 437 Bromiley, G.D. and Keppler, H. (2004) An experimental investigation of hydroxyl solubility in
438 jadeite and Na-rich clinopyroxenes. *Contributions to Mineralogy and Petrology*, 147, 189–
439 200.
- 440 Boyer, H., Smith, D.C., Chopin, C., and Lasnier, B. (1985) Raman microprobe (RMP)
441 determinations of natural and synthetic coesite. *Physics and Chemistry of Minerals*, 12,
442 45–48.
- 443 Chao, E.C.T., Shoemaker, E.M., and Madsen, B.M. (1960) First natural occurrence of coesite.
444 *Science*, 132, 220-222.
- 445 Chopin, C. (1984) Coesite and pure pyrope in high-grade blueschists of the Western Alps: a first
446 record and some consequences. *Contributions to Mineralogy and Petrology*, 86, 107–118.
- 447 Coes, L. (1953) A New Dense Crystalline Silica. *Science*, 118, 131–132.
- 448 Dean, D.W., Wentzcovitch, R.M., Keskar, N., Chelikowsky, J.R. and Binggeli, N. (2000)
449 Pressure-induced amorphization in crystalline silica: Soft phonon modes and shear
450 instabilities in coesite. *Physical Review B*, 61, 3303–3309.
- 451 Dubrovinskaia, N. and Dubrovinsky, L. (2001) High-pressure silica polymorphs as hardest
452 known oxides. *Materials Chemistry and Physics*, 68, 77–79.
- 453 Dubrovinsky, L., Dubrovinskaia, N., Prakapenka, V., Seifert, F., Langenhorst, F., Dmitriev, V.,
454 Weber, H.-P., and Le Bihan, T. (2004) A class of new high-pressure silica polymorphs.
455 *Physics of the Earth and Planetary Interiors*, 143–144, 231–240.

- 456 Glinnemann, J., King, H.E., Schulz, H., Hahn, T., La Placa, S.J., and Dacol, F. (1992) Crystal
457 structures of the low-temperature quartz-type phases of SiO₂ and GeO₂ at elevated
458 pressure. *Zeitschrift für Kristallographie*, 198, 177-212
- 459 Gonze, X., Beuken, J.-M., Caracas, R., Detraux, F., Fuchs, M., Rignanese, G.-M., Sindic, L.,
460 Verstraete, M., Zerah, G., Jollet, F., Torrent, M., Roy, A., Mikami, M., Ghosez, Ph., Raty,
461 J.-Y., and Allan, D.C. (2002) First-principle computation of material properties the
462 ABINIT software project. *Computational Materials Science*, 25, 478–492.
463 [<http://www.abinit.org>]
- 464 Gonze, X., Rignanese, G.-M., and Caracas, R. (2005) First-principles studies of the lattice
465 dynamics of crystals, and related properties. *Zeitschrift für Kristallographie*, 220, 458–
466 472.
- 467 Gonze, X., Amadon, B., Anglade, P.-M., Beuken, J.-M., Bottin, F., Boulanger, P., Bruneval, F.,
468 Caliste, D., Caracas, R., M. Côté, Deutsch, T., Genovesi, L., Ghosez, Ph., Giantomassi,
469 M., Goedecker, S., Hamann, D.R., Hermet, P., Jollet, F., Jomard, G., Leroux, S., Mancini,
470 M., Mazevet, S., Oliveira, M.J.T., Onida, G., Pouillon, Y., Rangel, T., Rignanese, G.-M.,
471 Sangalli, D., Shaltaf, R., Torrent, M., Verstraete, M.J., Zerah G., and Zwanziger, J.W.
472 (2009) ABINIT: First-principles approach to material and nanosystem properties.
473 *Computer Physics Communications*, 180, 2582-2615.
- 474 Haines, J., Léger, J.M., Gorelli, F., and Hanfland, M. (2001) Crystalline Post-Quartz Phase in
475 Silica at High Pressure. *Physical Review Letters*, 87, 155503.
- 476 Hammersley (1996), Publication No. ESRF98HAO1T, ESRF.

- 477 Hazen, R.M., Finger, L.W., Hemley, R.J., and Mao, H.K. (1989) High-pressure crystal chemistry
478 and amorphization of alpha-quartz. *Solid State Communications*, 72, 507-511.
- 479 Hemley, R.J. (1987) Pressure dependence of Raman spectra of SiO₂ polymorphs: α -quartz,
480 coesite, and stishovite. In M.H. Manghnani and Y. Syono, Eds., *High-pressure Research*
481 *in Mineral Physics*, 347-359.
- 482 Hemley, R.J., Jephcoat, A.P., Mao, H.K., Ming, L.C., and Manghnani, M.H. (1988) Pressure-
483 induced amorphization of crystalline silica. *Nature*, 334, 52–54.
- 484 Kantor, I., Prakapenka, V., Kantor, A., Dera, P., Kurnosov, A., Sinogeikin, S., Dubrovinskaia, N.,
485 and Dubrovinsky, L. (2012) BX90: A new diamond anvil cell design for X-ray diffraction
486 and optical measurements. *Review of Scientific Instruments*, 83, 125102.
- 487 Kingma, K.J., Hemley, R.J., Mao, H., and Veblen, D.R. (1993) New high-pressure
488 transformation in α -quartz. *Physical Review Letters*, 70, 3927–3930.
- 489 Klotz, S., Chervin, J.-C., Munsch, P. and Le Marchand, G. (2009) Hydrostatic limits of 11
490 pressure transmitting media. *Journal of Physics D-Applied Physics*, 42, 075413-075419.
- 491 Koch-Muller, M., Fei, Y., Hauri, E., and Liu, Z. (2001) Location and quantitative analysis of OH
492 in coesite. *Physics and Chemistry of Minerals*, 28, 693–705.
- 493 Kohn, W. and Sham, L.J. (1965) Self-consistent equations including exchange and correlation
494 effects. *Physical Reviews*, 140, A1133–A1138.

- 495 Kurnosov, A., Kantor, I., Boffa-Ballaran, T., Lindhardt, S., Dubrovinsky, L., Kuznetsov, A., and
496 Zehnder, B.H. (2008) A novel gas-loading system for mechanically closing of various
497 types of diamond anvil cells. *Review of Scientific Instruments*, 79, 045110.
- 498 Korsakov, A.V., Hutsebaut, D., Theunissen, K., Vandenabeele, P., and Stepanov, A.S. (2007)
499 Raman mapping of coesite inclusions in garnet from the Kokchetav Massif (Northern
500 Kazakhstan). *Spectrochimica Acta Part A: Molecular and Biomolecular Spectroscopy*, 68,
501 1046–1052.
- 502 Martin, R.M. (2003) *Electronic structure. Basic theory and practical methods*. Cambridge
503 University Press. 596 pp.
- 504 Mao, H.K., Xu, J., and Bell, P.M. (1986) Calibration of the ruby pressure gauge to 800 kbar
505 under quasi-hydrostatic conditions. *Journal of Geophysical Research*, 91, 4673-4676.
- 506 Monkhorst, H.J. and Pack, J.D. (1976) Special points for Brillouin-zone integrations. *Physical*
507 *Reviews B*, 13, 5188-5192.
- 508 Mosenfelder, J. and Bohlen, S. (1997) Kinetics of the coesite to quartz transformation. *Earth and*
509 *Planetary Science Letters*, 153, 133–147.
- 510 Ohtani, E., Ozawa, S., Miyahara, M., Ito, Y., Mikouchi, T., Kimura, M., Arai, T., Sato, K., and
511 Hiraga, K. (2011) Coesite and stishovite in a shocked lunar meteorite, Asuka-881757, and
512 impact events in lunar surface. *Proceedings of the National Academy of Sciences U.S.A.*,
513 108, 463–466.
- 514 Oxford Diffraction, CrysAlis CCD, Oxford Diffraction Ltd, Abingdon, England (2006)

- 515 Parkinson, C.D. (2000) Coesite inclusions and prograde compositional zonation of garnet in
516 whiteschist of the HP-UHPM Kokchetav massif, Kazakhstan: a record of progressive
517 UHP metamorphism. *Lithos*, 52, 215–233.
- 518 Payne, M.C., M.P. Teter, D.C. Allan, T.A. Arias, and J.D. Joannopoulos (1992) Iterative
519 minimization techniques for ab initio total-energy calculations: molecular dynamics and
520 conjugate gradients. *Reviews of Modern Physics*, 64, 1045-1097.
- 521 Prokopenko, V.B., Dubrovinsky, L.S., Dmitriev, V., and Weber, H.P. (2001) In situ
522 characterization of phase transitions in cristobalite under high-pressure by Raman
523 spectroscopy and X-ray diffraction. *Journal of Alloys and Compounds*, 327, 87–95.
- 524 Rekhi, S., Dubrovinsky L. S., and Saxena S. (1999) Temperature-induced ruby fluorescence
525 shifts up to a pressure of 15 GPa in an externally heated diamond anvil cell. *High
526 Temperatures - High Pressures*, 31, 299–305.
- 527 Ruiz-Cruz, M.D. and Sanz de Galdeano, C. (2012) Diamond and coesite in ultrahigh-pressure-
528 ultrahigh-temperature granulites from Ceuta, Northern Rif, northwest Africa.
529 *Mineralogical Magazine*, 76, 683–705.
- 530 Smyth, J., Hatton, C., (1977). Coesite-Sanidine Grosopydite from Roberts-Victor Kimberlite.
531 *Earth and Planetary Science Letters*, 34, 284–290.
- 532 Sobolev, N.V., Fursenko, B.A., Goryainov, S.V., Shu, J.F., Hemley, R.J., Mao, H.K., and Boyd,
533 F.R. (2000) Fossilized high pressure from the Earth's deep interior: The coesite-in-
534 diamond barometer. *Proceedings of the National Academy of Sciences. U. S. A.*, 97,
535 11875–11879.

- 536 Teter, D.M., Hemley, R.J. (1998). High Pressure Polymorphism in Silica. Physical Reviews
537 Letters, 80, 2145-2148.
- 538 Veithen, M., Gonze, X., and Ghosez, Ph. (2005) Non-linear optical susceptibilities, Raman
539 efficiencies and electrooptic tensors from first-principles density functional perturbation
540 theory. Physical Reviews B, 71, 125107.
- 541 Wentzcovitch, R.M., Da Silva, C., Chelikowsky, J.R., and Binggeli, N. (1998) A New Phase and
542 Pressure Induced Amorphization in Silica. Physical Review Letters, 80, 2149–2152.
- 543 Wyckoff, R.W.G. (1963) Crystal Structures 1, 2nd edition. Interscience Publishers, New York,
544 New York, 7-83
- 545 Xiao, W., Chen, M., and Xie, X. (2011). Shock-Produced Coesite in the Xiuyan Crater, China.
546 Meteoritics & Planetary Science, 46, A256–A256.
- 547 Yagi, T. and Yamakata, M. (2000). Effect of hydrostaticity on the phase transformations of
548 cristobalite. In H. Aoki et al. Eds., Physics Meets Mineralogy, 242-255.
- 549

550 FIGURE CAPTIONS

551 **FIGURE 1.** External heating assembly, consisting of a pyrophyllite whole cell heater surrounding
552 DAC placed on a stainless steel holder and double-layered stainless-steel insulation box. The coin
553 is used as a scale marker.

554 **FIGURE 2.** (a) Raman spectra collected up to 30 GPa during compression. The star located above
555 the doublet peak at 24.2 GPa indicates appearance of coesite II phase. (b) Raman spectra
556 collected between 30 and 51 GPa during compression. The star above the spectrum at 38.3 GPa
557 indicates disappearance of the main doublet peak and the transition to coesite III.

558 **FIGURE 3.** Pressure dependence of the main Raman active mode obtained from several different
559 experiments at ambient temperature. Phase transitions at ~23 and ~35 GPa are indicated by solid
560 lines, and the thermodynamic stability field of coesite is indicated by dotted lines.

561 **FIGURE 4.** Raman spectra collected during decompression from ~51 GPa to ambient pressure.
562 Note that the power of the incident laser beam was decreased below 20 GPa from 1.0 to 0.3 W.

563 **FIGURE 5.** (a) Theoretical Raman spectrum at ambient conditions. The spectrum is dominated by
564 the strong A_g mode at 510 cm^{-1} . Several peaks are visible at low frequencies and the most intense
565 is the B_g mode at 77 cm^{-1} , while all peaks above 600 cm^{-1} are very weak. A few peaks are
566 symmetry-labeled. The inset shows the peaks occurring in the $400\text{-}500\text{ cm}^{-1}$ range, with intensity
567 magnified, with the B_g mode at 430 and the A_g mode at 455 cm^{-1} . (b) Pressure variation of
568 selected modes in coesite. Note that the lowest frequency modes become unstable with increasing
569 pressure. The A modes bear symbols; the g modes are represented with solid line; the dashed line
570 represent modes that are only infrared active. Three modes become unstable between 20 and 30

571 GPa, including a B_g Raman active mode at 77cm⁻¹ and an A_u infrared active at 111 cm⁻¹ (ambient
572 conditions). These two modes are displayed in Figure 6c and d at elevated pressure.

573 **FIGURE 6.** Atomic displacement pattern of some of the most representative Raman modes of
574 coesite at 10 GPa. They dominate the Raman spectra and their behavior is likely responsible for
575 the structural changes exhibited under pressures. White corresponds to silicon atom, grey is
576 oxygen; atomic nomenclature is after Angel et al. (2003). Ellipses are surrounding representative
577 SiO₄ rings and grey arrows placed on oxygen atoms are indicating displacement directions. (a)
578 The strongest A_g mode at 541 cm⁻¹ (located at 510 cm⁻¹ at ambient pressure) is a breathing mode
579 of the four-membered tetrahedral rings. (b) The B_g mode at 77 cm⁻¹ (also 77 cm⁻¹ at 10 GPa)
580 dominates the spectra at low frequencies and it corresponds to shearing of the four-membered
581 silica rings. The sketch in the upper left corner on the figure is for simplifying the structure of one
582 ring. (c) The same B_g mode at 20 GPa (56 cm⁻¹) in different crystallographic orientation. (d)
583 Infrared active A_u mode at 20 GPa (78 cm⁻¹) showing significant displacement of the O1 atom
584 from the rigid 180 °Si-O1-Si bond.

585 **FIGURE 7.** (a) Raman spectra collected at 27-29 GPa during heating up to 400 °C. At both
586 pressures one can notice a decrease in intensity and an increase in the splitting of the doublet. (b)
587 Raman spectra collected at 34 GPa during heating up to 200 °C and after quenching to ambient
588 conditions. Only a decrease in intensity can be observed at this pressure, without an increase in
589 the splitting of the doublet.

590 **FIGURE 8.** TEM analysis of one of the grains recovered from the thin edge of the crystal bridged
591 at 35 GPa, showing diffraction profiles that can be attributed to coesite. The material from which
592 this grain was selected shows a weak Raman peak around 520 cm⁻¹.

593 **FIGURE 9.** Representative X-ray diffraction data showing that the crystalline phase quenched
594 from ~34 GPa and ~300 °C is coesite. Reference phases are taken from Crystallography Open
595 Database: coesite - Araki and Zoltai (1969), synthetic sample at ambient conditions; quartz -
596 Hazen et al. (1989), synthetic sample at pressure of 1bar; and diamond – Wyckoff (1963).

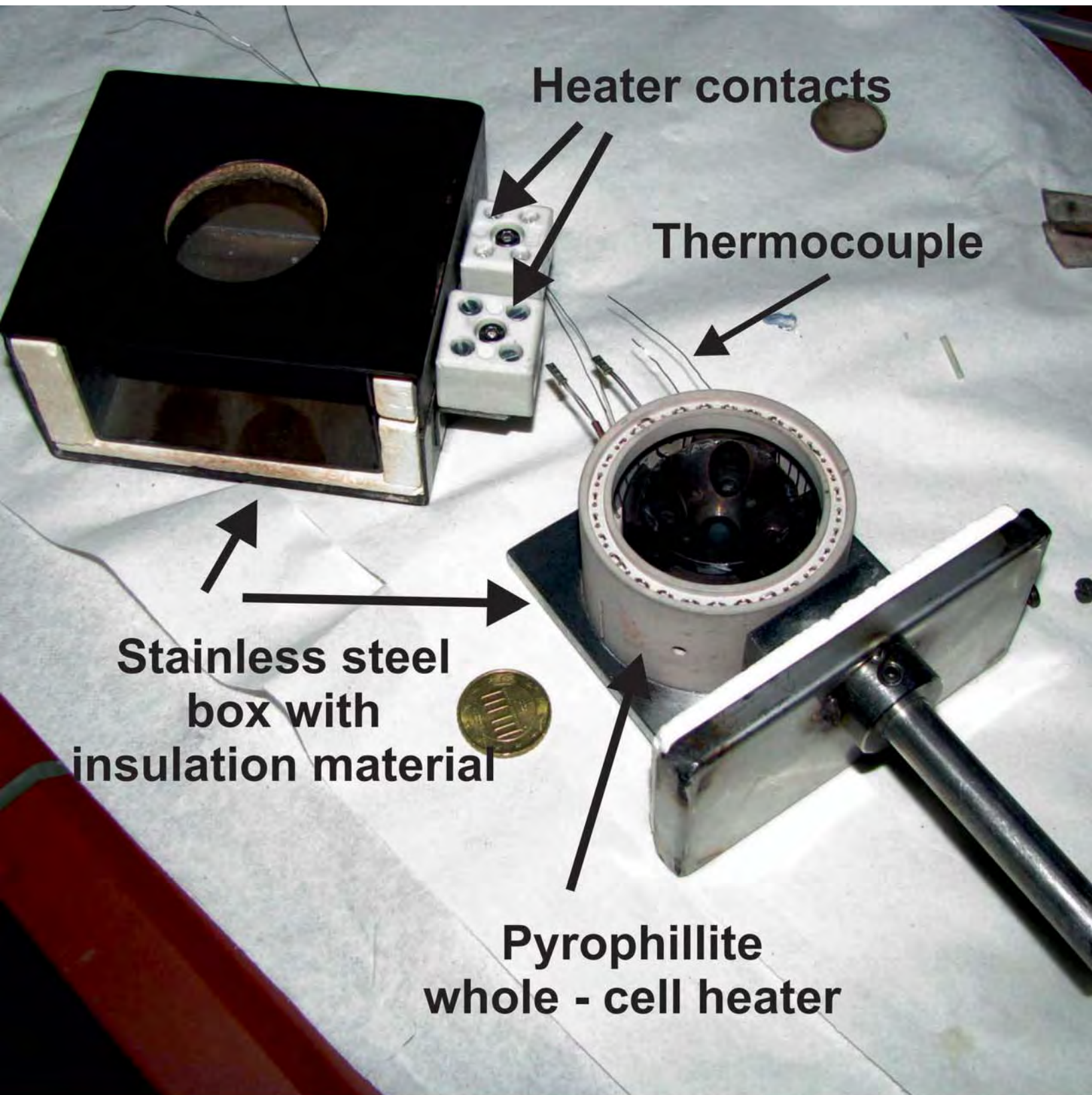


Figure 1.

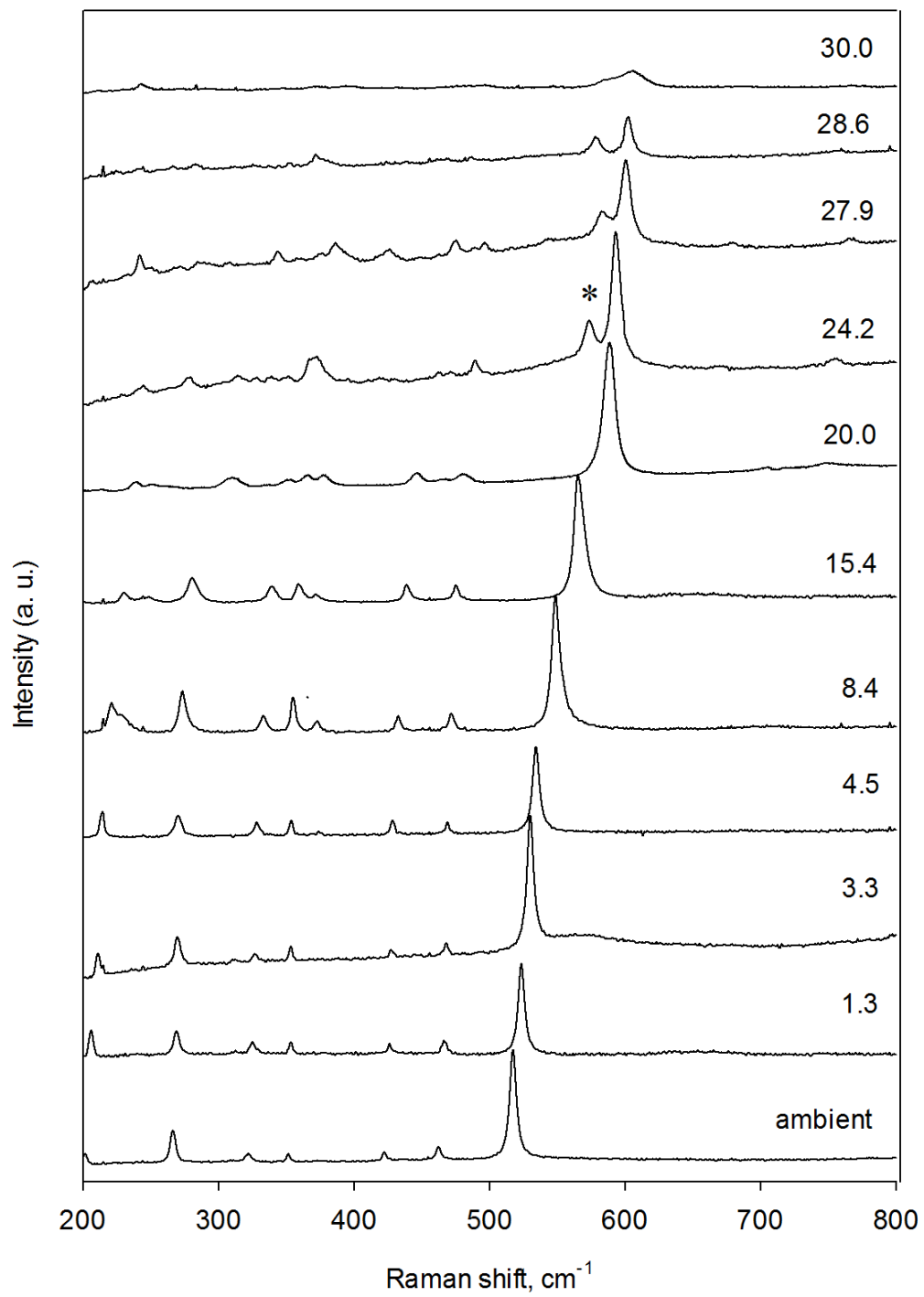


Figure 2a

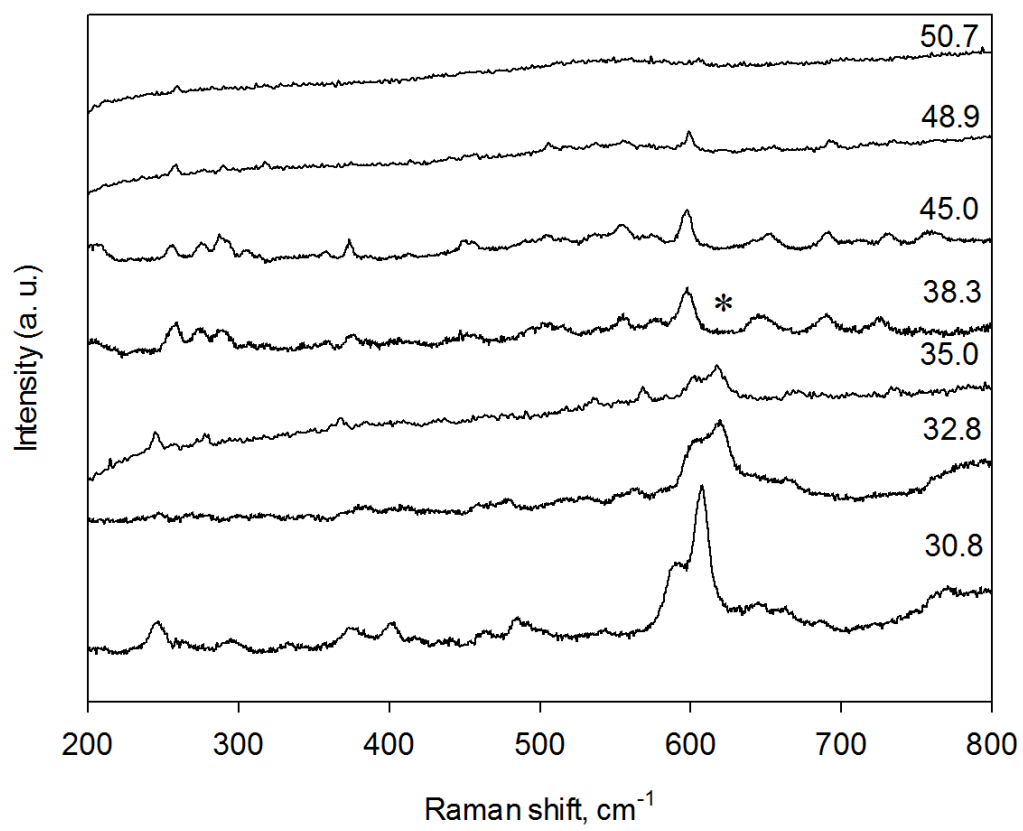


Figure 2b.

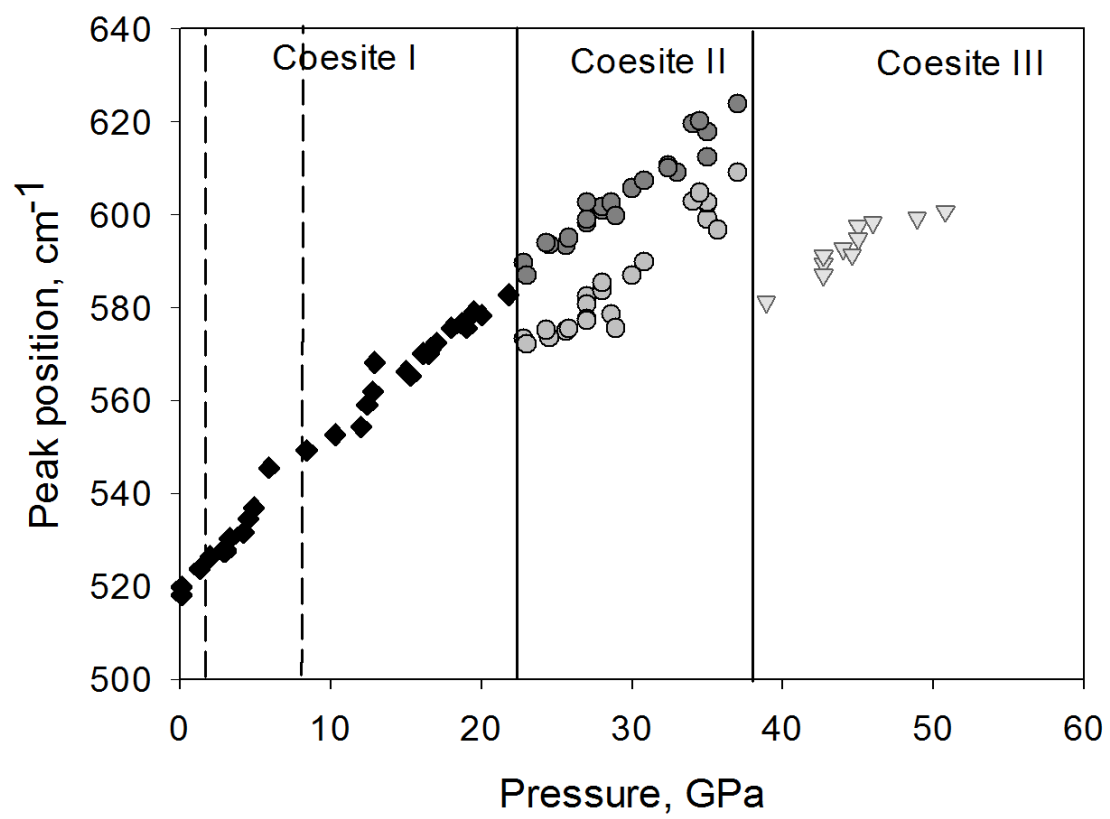


Figure 3.

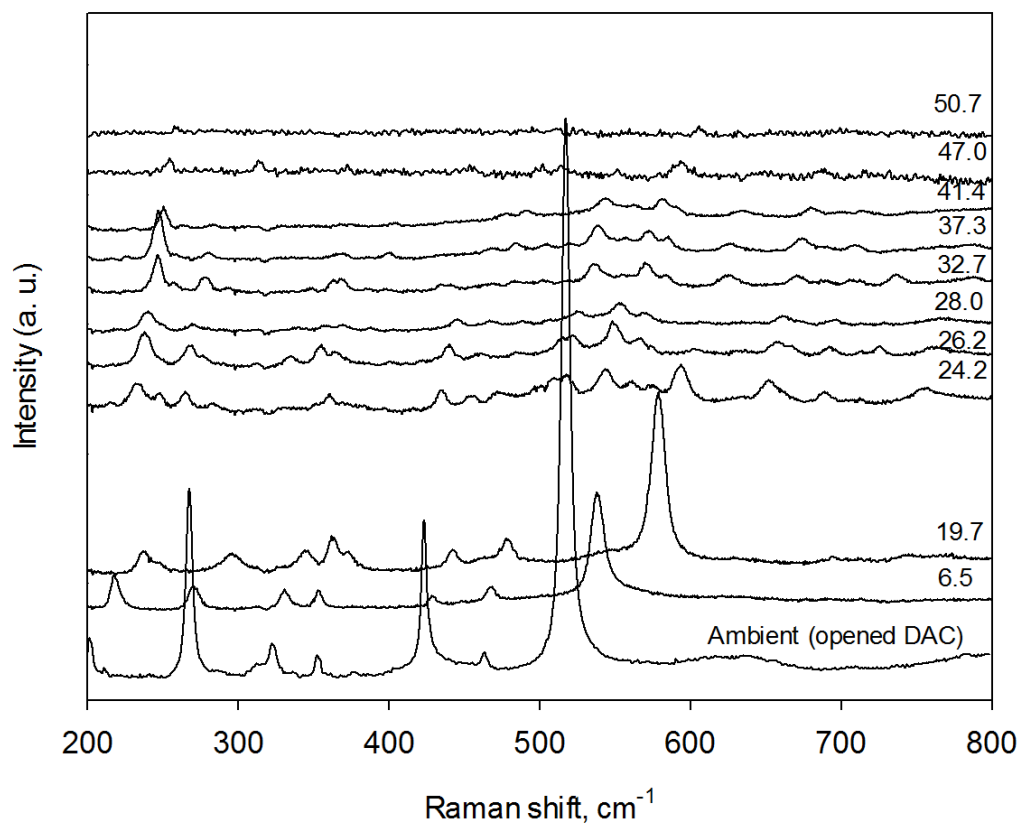


Figure 4.

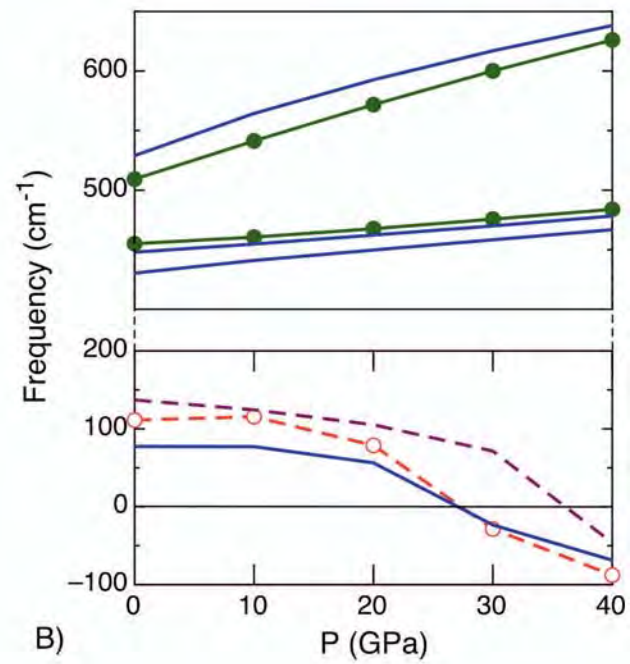
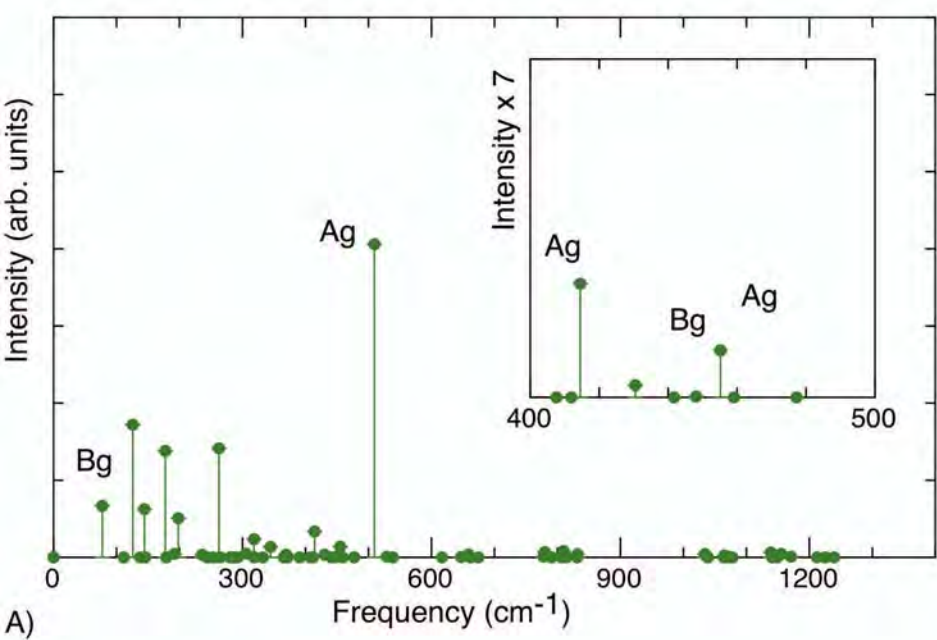
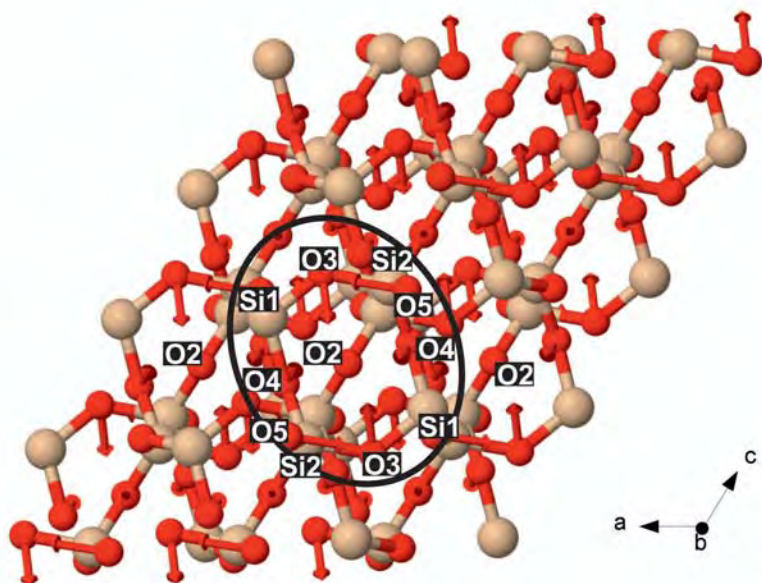
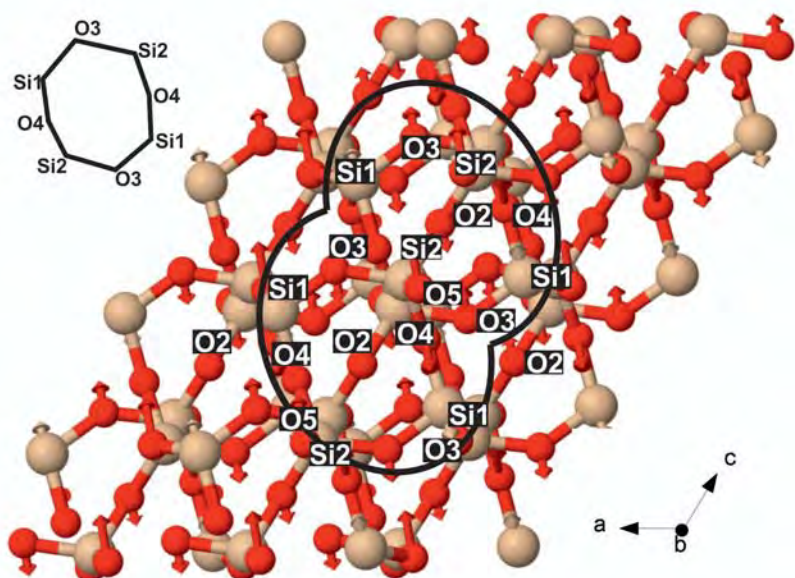


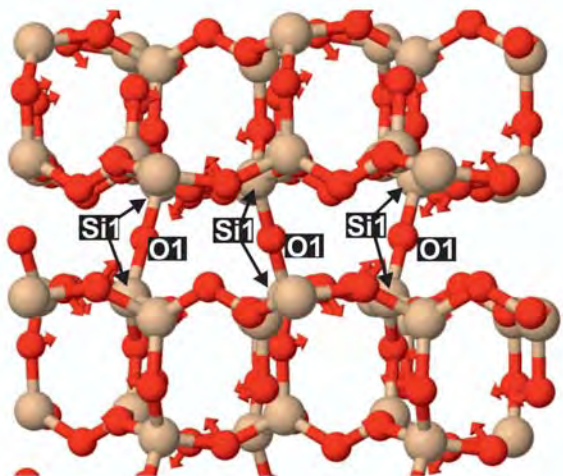
Figure 5.



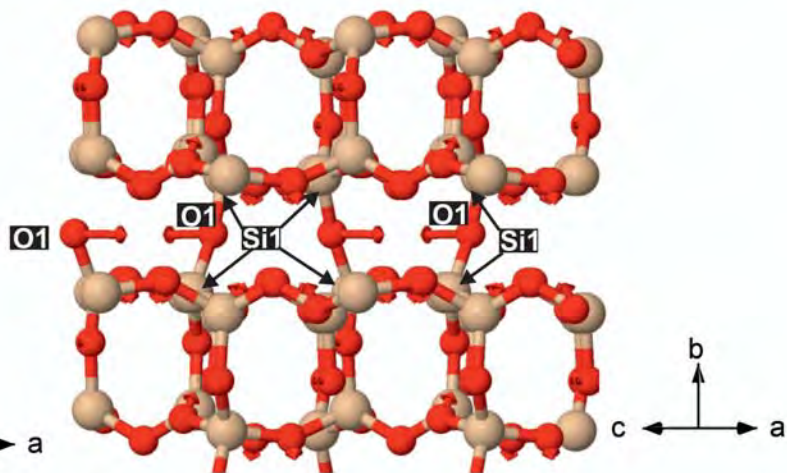
A)



B)



C)



D)

Figure 6.

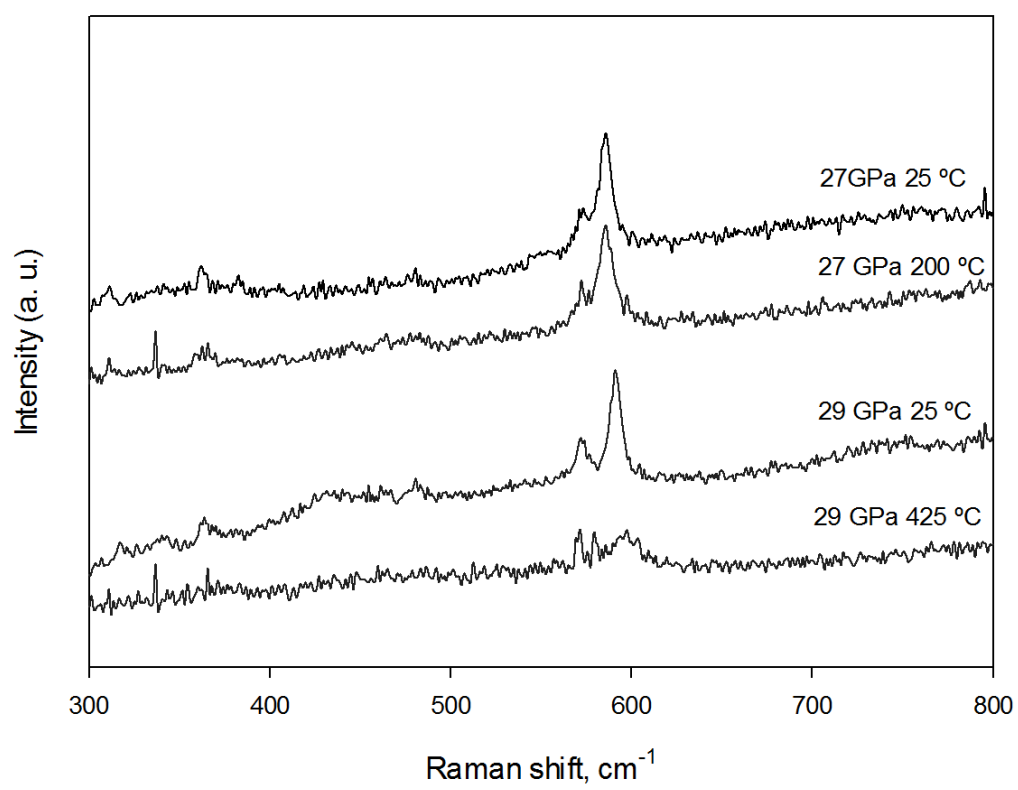


Figure 7a.

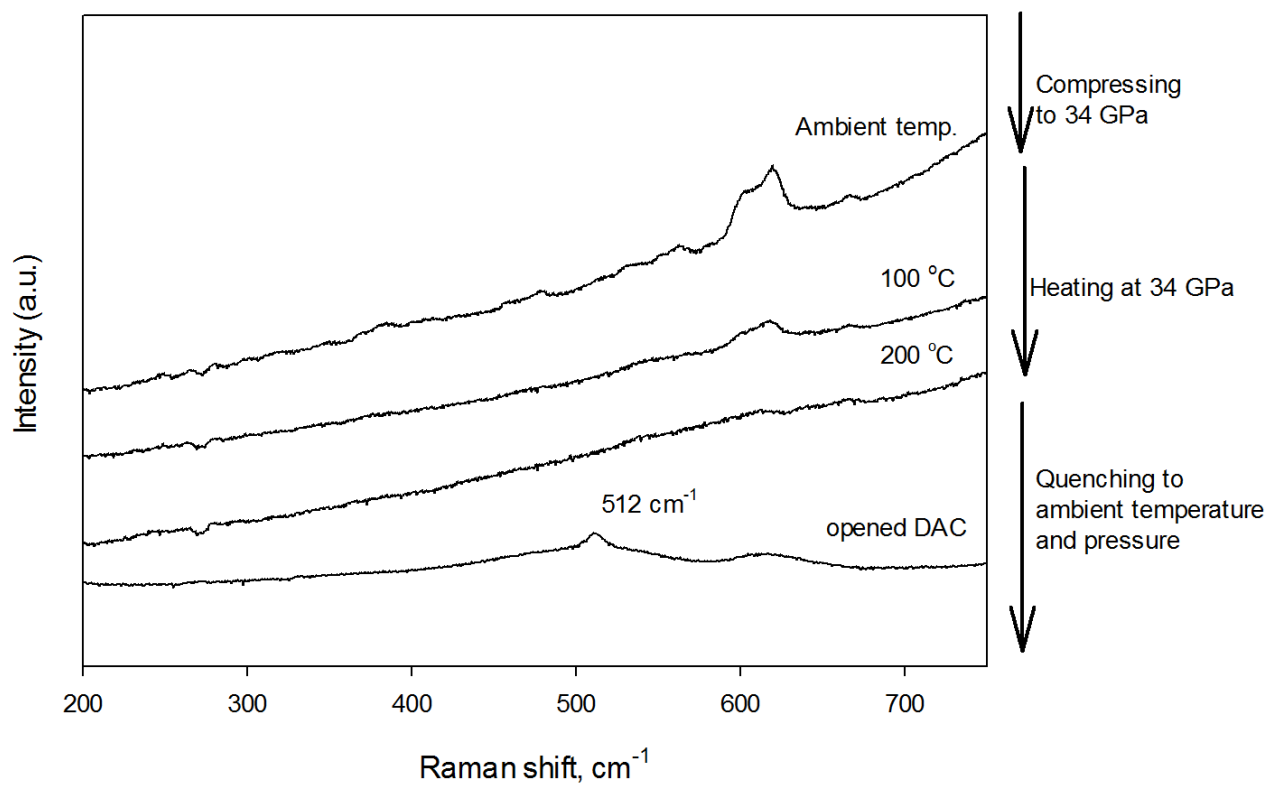


Figure 7b.

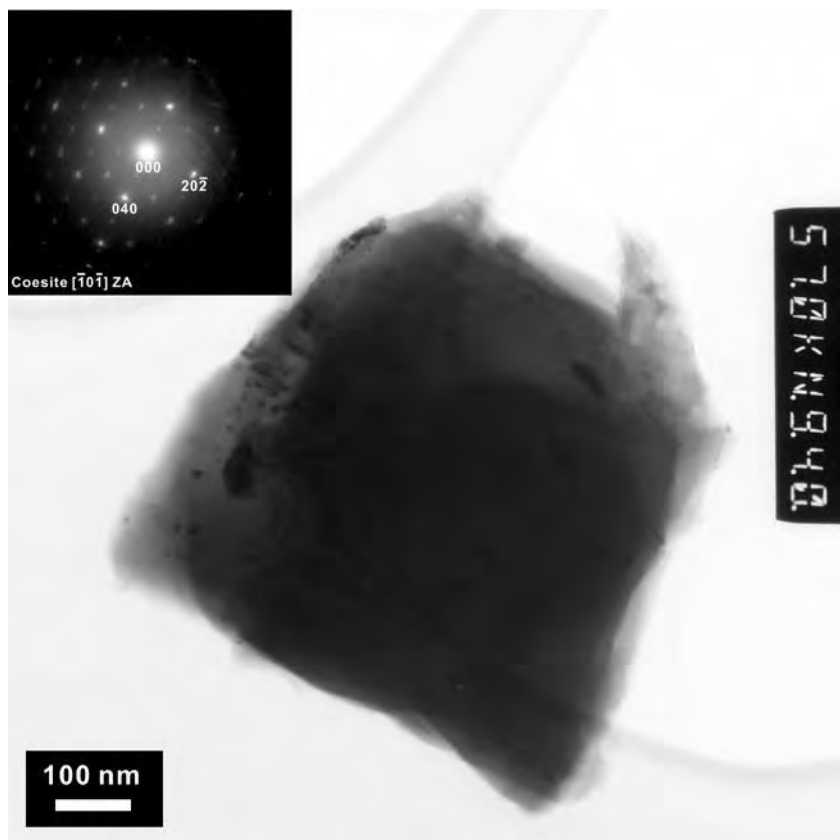


Figure 8.

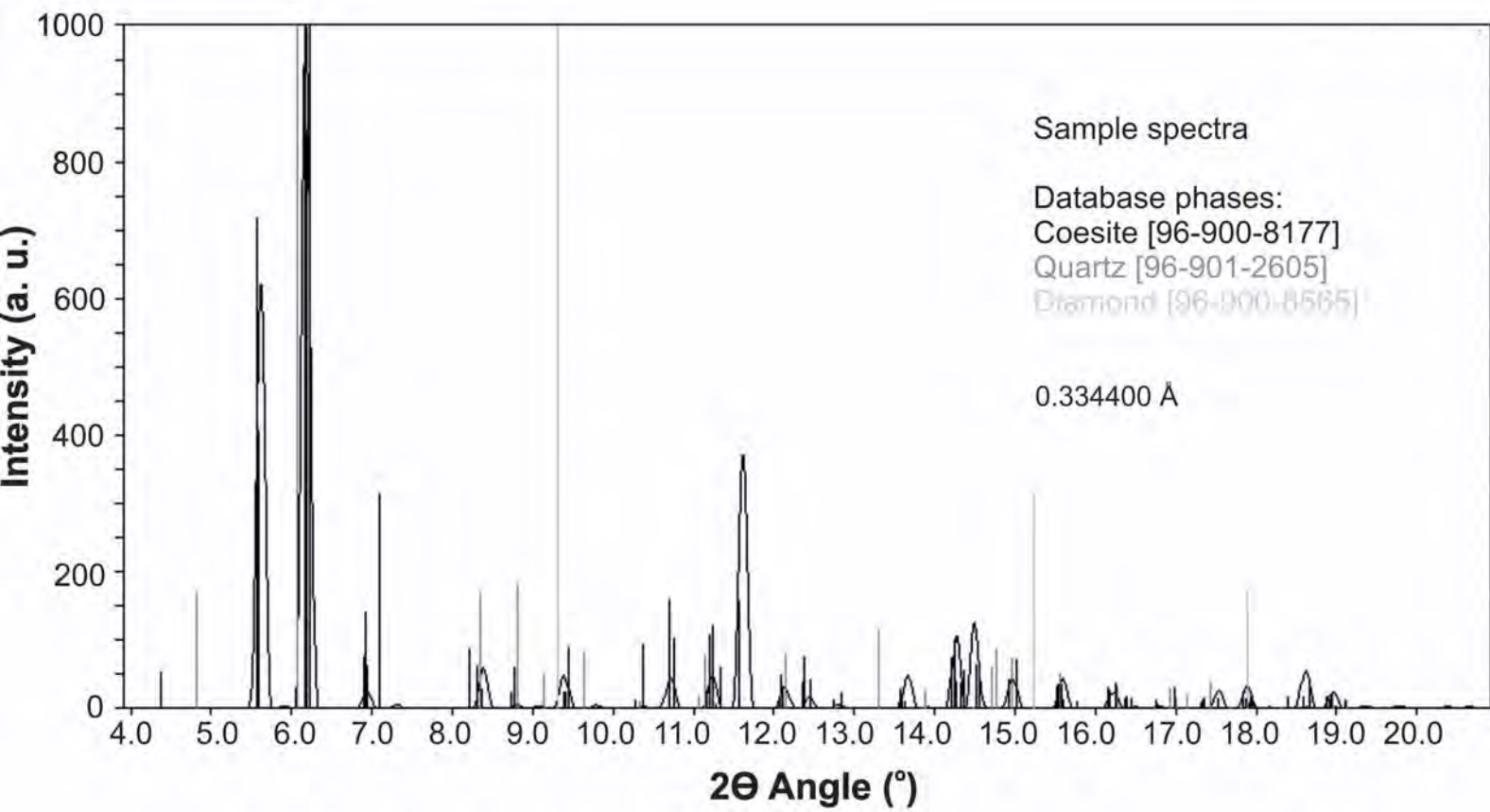


Figure 9.