1 **REVISION 1** 2 Chromo-alumino-povondraite, NaCr<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, a new 3 4 mineral species of the tourmaline supergroup 5 6 LEONID REZNITSKII<sup>1</sup>, CHRISTINE M. CLARK<sup>2</sup>, FRANK C. HAWTHORNE<sup>2,\*</sup>, JOEL D. GRICE<sup>3</sup>, 7 HENRIK SKOGBY<sup>4</sup>, ULF HÅLENIUS<sup>4</sup> AND FERDINANDO BOSI<sup>4,5,\*</sup> 8 9 10 11 <sup>1</sup>Institute of the Earth's Crust, Siberian Branch, Russian Academy of Science, Lermontova str., 128, 12 664033 Irkutsk, Russia 13 <sup>2</sup>Department of Geological Sciences, University of Manitoba, Winnipeg, Manitoba R3T 2N2, Canada 14 <sup>3</sup>Research Division, Canadian Museum of Nature, P.O. Box 3443A, Station "D", Ottawa, Ontario, 15 Canada, K1P 6P4 16 <sup>4</sup>Department of Geosciences, Swedish Museum of Natural History, Box 50007, SE-10405 Stockholm, 17 18 <sup>5</sup>Dipartimento di Scienze della Terra, Sapienza Università di Roma, Piazzale Aldo Moro, 5, I-00185 19 Rome, Italy 20 21 22 \* E-mail: frank hawthorne@umanitoba.ca; ferdinando.bosi@uniroma1.it 23 24 25 26 27 ABSTRACT 28 Chromo-alumino-povondraite, NaCr<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, is a new mineral 29 of the tourmaline supergroup. It is found in metaquartzites of the Pereval marble quarry 30 (Sludyanka, Lake Baikal, Russia) in association with dravite, oxy-chromium-dravite, oxy-31 dravite, quartz, calcite, chromphyllite, eskolaite, chromite, uvarovite, chromian phlogopite, and pyroxenes of the diopside-kosmochlor series, Cr-bearing tremolite, Cr-bearing titanite, Crbearing rutile, and pyrite.

Crystals are green and transparent with a vitreous luster, and exhibit a pale-green streak and conchoidal fracture. Chromo-alumino-povondraite has a Mohs hardness of approximately 7½, and a calculated density of 3.227 g/cm<sup>3</sup>. In plane-polarized light, chromo-aluminopovondraite is pleochroic (O = emerald green and E = pale yellowish green) and uniaxial negative:  $\omega = 1.745(5)$ ,  $\varepsilon = 1.685(5)$ . Chromo-alumino-povondraite is rhombohedral, space group R3m, with the unit-cell parameters a = 16.0277(2), c = 7.3085(1) Å, V = 1625.93(5) Å<sup>3</sup>, Z = 3. Crystal-chemical analysis resulted in the empirical structural formula:  ${}^{X}(Na_{0.87}Ca_{0.07}\square_{0.04}K_{0.02})_{\Sigma1.00}{}^{Y}(Cr^{3+}_{2.29}Mg_{0.71})_{\Sigma3.00}{}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{1.18}V^{3+}_{0.22}Fe^{3+}_{0.01})_{\Sigma6.00}$ 

 $^{T}[(Si_{5.96}Al_{0.04})O_{18}]^{B}(BO_{3})_{3}^{V}(OH)_{3}^{W}[O_{0.73}F_{0.25}(OH)_{0.02}]_{\Sigma1.00}$ 

The crystal structure of chromo-alumino-povondraite was refined to an R1 index of 1.68% using 1803 unique reflections collected with MoKα X-radiation. Ideally, chromoalumino-povondraite is related to oxy-dravite and oxy-chromium-dravite by the homovalent substitution  $Cr^{3+} \leftrightarrow Al^{3+}$ . Tourmaline with chemical compositions classified as chromoalumino-povondraite can be either Al-dominant or Cr-dominant as a result of the compositional boundaries along the solid solution between Al and Cr<sup>3+</sup> that are determined at  $^{Y+Z}(Cr_{1}, 5Al_{5}, 5),$ corresponding to  $Na^{Y}(Cr_{1.5}Al_{1.5})^{Z}(Al_4Mg_2)Si_6O_{18}(BO_3)_3(OH)_3O_5$  $^{Y+Z}(Cr_5Al_2)$ , corresponding to  $Na^Y(Cr_3)^Z(Cr_2Al_2Mg_2)Si_6O_{18}(BO_3)_3(OH)_3O$ .

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53 INTRODUCTION

The tourmaline supergroup minerals are widespread, occurring in sedimentary, igneous and metamorphic rocks (Dutrow and Henry 2011). They are important indicator minerals that can provide information on the compositional evolution of their host rocks, chiefly due to their ability to incorporate a large number of elements (e.g., Novák et al. 2004; Agrosì et al. 2006; Lussier et al. 2011a; Novák et al. 2011; van Hinsberg et al. 2011; Bačík et al. 2012). However, the chemical composition of tourmalines is also controlled by short-range and long-range constraints (e.g., Hawthorne 1996, 2002; Bosi and Lucchesi 2007; Bosi 2010, 2011; Henry and Dutrow 2011; Skogby et al. 2012; Bosi 2013). Tourmaline supergroup minerals are complex borosilicates and their crystal structure and crystal chemistry have been extensively studied (e.g., Foit 1989; Hawthorne and Henry 1999; Bosi and Lucchesi 2007; Lussier et al. 2008; Bosi 2008; Bosi et al. 2010; Lussier et al. 2011b; Filip et al. 2012). In accordance with

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Henry et al. (2011), the general formula of tourmaline may be written as: XY_3Z_6T_6O_{18}(BO_3)_3V_3W, where X (\equiv {}^{[9]}X) = Na^+, K^+, Ca^{2+}, \Box (=vacancy); Y (\equiv {}^{[6]}Y) = Al^{3+}, Fe^{3+}, Cr^{3+}, V^{3+}, Mg^{2+}, Fe^{2+}, Mn^{2+}, Li^+; Z (\equiv {}^{[6]}Z) = Al^{3+}, Fe^{3+}, Cr^{3+}, V^{3+}, Mg^{2+}, Fe^{2+}; T (\equiv {}^{[4]}T) = Si^{4+}, Al^{3+}, B^{3+}; B (\equiv {}^{[3]}B) = B^{3+}; W (\equiv {}^{[3]}O1) = OH^{1-}, F^{1-}, O^{2-}; V (\equiv {}^{[3]}O3) = OH^{1-}, O^{2-} and where, for example, T represents a group of cations (Si^{4+}, Al^{3+}, B^{3+}) accommodated at the [4]-coordinated T sites. The dominance of these ions at one or more sites of the structure gives rise to a range of distinct mineral species.

Recently, several new minerals of the tourmaline supergroup were approved by the
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Recently, several new minerals of the tourmaline supergroup were approved by the Commission on New Minerals, Nomenclature and Classification (CNMNC) of the International Mineralogical Association (IMA). Among these are a number of oxy-tourmalines related by complete solid solution in the Al<sup>3+</sup>-Cr<sup>3+</sup>-V<sup>3+</sup> subsystem: oxy-dravite, end-member formula NaAl<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2012-004a; Bosi and Skogby 2013), oxy-chromium-dravite, NaCr<sub>3</sub>(Cr<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2011-097; Bosi et al. 2012a); oxy-vanadium-dravite, NaV<sub>3</sub>(V<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 11-E; Bosi et al. 2013a); vanadio-oxy-dravite, NaV<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2012-074; Bosi et al. 2014a); vanadio-oxy-chromium-dravite, NaV<sub>3</sub>(Cr<sub>4</sub>Mg<sub>2</sub>)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O (IMA 2012-034; Bosi et al. 2014b).

Chromo-alumino-povondraite was originally approved as a new species by the IMA CNMMN in 2009 (no. 2009-088). Subsequent additional work connected with publication of the resultant paper showed that the chemical analysis associated with the original proposal was not reliable. Further work uncovered an appropriate sample with the requisite chemical composition for this species. Consultation with the Chair of CNMMN resulted in a new proposal which involved changing the holotype sample. The original approval of IMA 2009-088 was hence withdrawn on 30 January 2013 (Williams et al. 2013), and a new proposal was submitted and approved by the IMA-CNMNC (no. 2013-089). The holotype specimen (sample PR85v) is deposited in the collections of the Museum of Mineralogy, Earth Sciences Department, Sapienza University of Rome, Italy, catalogue number 33069/1. A formal description of the new species chromo-alumino-povondraite is presented here, including a full characterization of its physical, chemical and structural properties.

#### OCCURRENCE, APPEARANCE AND PHYSICAL AND OPTICAL PROPERTIES

The crystals of chromo-alumino-povondraite are green and occur in metaquartzites in the Pereval marble quarry, Sludyanka crystalline complex, Southern Baikal region, Russia (51°37'N 103°38'E). The Sludyanka complex underwent regional granulite facies metamorphism in the Early Ordovician (Kotov et al. 1997; Kovach et al. 2013). The Pereval quarry is the type locality (see Reznitsky et al. 2001 and Bosi et al. 2012a for a more detailed description) for natalyite, florensovite, kalininite, magnesiocoulsonite, oxy-vanadium-dravite, oxy-chromium-dravite, vanadio-oxy-chromium-dravite, vanadio-oxy-dravite, oxyvanite and cuprokalininite. Minerals associated with the holotype specimen are: dravite, oxy-chromium-dravite, oxy-dravite, quartz, calcite, chromphyllite, eskolaite, chromite, uvarovite, chromian phlogopite, and pyroxenes of the diopside-kosmochlor series, Cr-bearing tremolite, Cr-bearing titanite, Cr-bearing rutile, and pyrite, The host rocks (quartz-diopside) are Cr-V-bearing carbonate-siliceous sediments, metamorphosed to granulite facies and partly diaphtorized (Salnikova et al. 1998) to amphibolite facies (retrograde stage). Chromoalumino-povondraite probably formed during the prograde stage (granulite facies). The crystals are euhedral, reaching up to 0.3 mm in length, and may be chemically zoned (for details, see Figure 3 of Bosi et al. 2013a), but homogeneous crystals also occur.

The morphology of chromo-alumino-povondraite consists of elongated  $\{10\bar{1}0\}$  and  $\{11\bar{2}0\}$  prisms terminated by a prominent  $\{0001\}$  pedion and small, minor  $\{10\bar{1}1\}$  pyramidal faces. Crystals are green, with a pale-green streak, transparent and display vitreous luster. They are brittle and show conchoidal fracture. The Mohs hardness is approximately  $7\frac{1}{2}$  (Reznitsky et al. 2001). The calculated density is 3.227 g/cm<sup>3</sup>. In transmitted light, chromo-alumino-povondraite is pleochroic with O = emerald green and E = pale yellowish green. Chromo-alumino-povondraite is uniaxial negative with refractive indices, measured by the immersion method using white light from a tungsten source, of  $\omega = 1.745(5)$ ,  $\varepsilon = 1.685(5)$ . The mean index of refraction, density and chemical composition lead to superior compatibility indices  $(1 - K_p/K_c = 0.011)$  (Mandarino 1976, 1981)

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125 Methods

### **General comment**

The present crystal structure refinement and electron microprobe analyses were obtained from the same single crystal. However, complementary optical data, Fourier-

transform infrared and optical-absorption spectra, refractive indices and X-ray powder diffraction data were recorded on coexisting crystals. Small differences in composition are likely to occur between these crystals, as well as minor variations in chemistry within the crystal population.

# **Single-crystal structural refinement**

A representative fragment of the type specimen was selected for X-ray diffraction measurements on a Bruker KAPPA APEX-II single-crystal diffractometer, at Sapienza University of Rome (Earth Sciences Department), equipped with a CCD area detector (6.2 × 6.2 cm<sup>2</sup> active detection area,  $512 \times 512$  pixels) and a graphite crystal monochromator, using Mo $K\alpha$  radiation from a fine-focus sealed X-ray tube. The sample-to-detector distance was 4 cm. A total of ca. 3265 exposures (step = 0.2°, time/step = 20 s) covering a full reciprocal sphere with a redundancy of about 8 was used. Final unit-cell parameters were refined using the Bruker AXS SAINT program for reflections with I > 10  $\sigma(I)$  in the range  $5^{\circ} < 20 < 72^{\circ}$ . The intensity data were processed and corrected for Lorentz, polarization, and background effects with the APEX2 software program of Bruker AXS. The data were corrected for absorption using the multi-scan method (SADABS). The absorption correction led to a significant improvement in  $R_{\text{int}}$ . No violations of R3m symmetry were noted.

Structural refinement was done with the SHELXL-97 program (Sheldrick 2008). Starting coordinates were taken from Bosi et al. (2004). Variable parameters were: scale factor, extinction coefficient, atom coordinates, site-scattering values and atom-displacement factors. To obtain the best values of statistical indexes (*R*1, *wR*2), a fully ionized scattering curve for O was used, whereas neutral scattering curves were used for the other atoms. In detail, the occupancy of the *X* site was modeled by using the Na scattering factor, and both the *Y* and *Z* sites using Cr and Mg scattering factors. The *T* and *B* sites were modeled, respectively, with Si and B scattering factors and with a fixed occupancy of 1, because refinement with unconstrained occupancies showed no significant deviations from this value. Three full-matrix refinement cycles with isotropic-displacement parameters for all atoms were followed by anisotropic cycles until convergence was attained. No significant correlations over a value of 0.7 between the parameters were observed at the end of refinement. Table 1 lists crystal data, data-collection information and refinement details; Table 2 gives the fractional atom coordinates and site occupancies; Table 3 gives the displacement parameters; Table 4 gives selected bond distances.

X-ray powder diffraction

X-ray powder data were collected with a Bruker D8 Discover SuperSpeed micropowder diffractometer with multi-wire 2D detector using a modified Gandolfi attachment (Cu $K\alpha$  radiation,  $\lambda = 1.54178$  Å, 50 kV, 60 mA). Two 10 h frames were collected and merged; no internal standard was used. Data listed in Table 5 are indexed on a = 16.095(17), c = 7.341(9) Å.

# **Electron-Microprobe analysis**

Electron-microprobe analyses of the crystal used for X-ray diffraction refinement were obtained by wavelength-dispersive spectrometry (WDS mode) with a Cameca SX50 instrument at the "Istituto di Geologia Ambientale e Geoingegneria (Rome, Italy), CNR", operating at an accelerating potential of 15 kV and a sample current of 15 nA, 10 μm beam diameter. Minerals and synthetic compounds were used as standards: wollastonite (Si, Ca), magnetite (Fe), rutile (Ti), corundum (Al), vanadinite (V) fluorphlogopite (F), periclase (Mg), jadeite (Na), K-feldspar (K), sphalerite (Zn), metallic Cr, Mn and Cu. The overlap corrections and the PAP routine were applied (Pouchou and Pichoir 1991). The results (Table 6) represent mean values of 10 spot analyses. In accord with the very low concentration of Li in dravitic samples (e.g., Henry et al. 2011) and the crystallization environment of the studied sample (i.e., Cr-bearing metaquartzites), the Li<sub>2</sub>O content was assumed to be insignificant. Manganese, Zn, Cu and Ti were below the detection limits (0.03 wt%) in the studied sample.

# **Infrared spectroscopy**

A homogeneous chromo-alumino-povondraite crystal from the holotype material was measured by Fourier-transform infrared (FTIR) absorption spectroscopy in the wavenumber range 2000-5000 cm<sup>-1</sup> to characterize OH absorption bands, using a Bruker Equinox 55 spectrometer equipped with a tungsten-halogen source and a CaF<sub>2</sub> beam-splitter. Polarized spectra with a resolution of 4 cm<sup>-1</sup> were acquired parallel and perpendicular to the crystallographic **c**-axis direction using an IR microscope with a KRS-5 wire-grid polarizer and an MCT detector. The 26 μm thick doubly-polished crystal plate had been oriented parallel the **c**-axis by morphology and optical microscopy. As normally observed, it was not possible to

thin the sample sufficiently to avoid off-scale absorption intensity for the strongest band (Fig. 1).

**Optical-absorption spectroscopy** 

Polarized, room-temperature optical-absorption spectra (OAS) were recorded on the same 26  $\mu$ m thick crystal platelet used for the FTIR measurements. The spectra were measured in the range 300-1100 nm at a resolution of 1 nm using an AVASPEC-ULS2048X16 spectrometer attached via a 400  $\mu$ m UV optical fiber to a Zeiss Axiotron UV-microscope. A 75 W Xenon arc lamp served as illuminating source and Zeiss Ultrafluar 10x lenses served as objective and condenser. A UV-quality Glan-Thompson prism with a working range from 250 to 2700 nm (40000 to 3704 cm<sup>-1</sup>) was used as a polarizer. The circular aperture was 64  $\mu$ m in diameter. The wavelength scale of the spectrometer was calibrated against Ho<sub>2</sub>O<sub>3</sub> doped and Pr<sub>2</sub>O<sub>3</sub>/Nd<sub>2</sub>O<sub>3</sub> doped standards (Hellma glass filters 666F1 and 666F7) with an accuracy better than 15 cm<sup>-1</sup> in the wavelength range 300-1100 nm. Recorded spectra were fitted using the Jandel PeakFit 4.12 software assuming Gaussian peak shapes.

209 RESULTS AND DISCUSSION

# **Determination of atomic proportions**

In agreement with the structural-refinement results, the boron content was assumed to be stoichiometric in the chromo-alumino-povondraite sample ( $B^{3+} = 3.00$  apfu). Both the site-scattering results and the bond lengths of B and T are consistent with the B site fully occupied by boron and no amount of  $B^{3+}$  at the T site (e.g., Hawthorne 1996; Bosi and Lucchesi 2007). The small contents of Fe were calculated as Fe<sup>3+</sup>, on the basis of results of (Bosi et al. 2013b). The (OH) content can then be calculated by charge balance with the assumption T + Y + Z = 15.00. The atomic proportions were calculated on this assumption (Table 6). The excellent match between the number of electrons per formula unit (epfu) derived from chemical and structural analysis supports this procedure, respectively: 265.6 and 266.3 epfu.

### **Site populations**

The anion site populations in the studied sample follow the general preference suggested for tourmaline (e.g., Grice and Ercit 1993; Henry et al. 2011): the O3 site (V position in the general formula) is occupied by (OH), and the O1 site (W position in the general formula) is occupied by  $O^{2-}$ ,  $(OH)^{1-}$  and  $F^{1-}$ . The cation distribution at the T, Y and Z

226 sites was optimized by using a quadratic program to minimize the residuals between calculated 227 and observed data (based on the chemical and structural analysis). Site-scattering values, octahedral and tetrahedral mean bond-distances (i.e.,  $\langle Y-O \rangle$ ,  $\langle Z-O \rangle$  and  $\langle T-O \rangle$ ) were 228 calculated as the linear contribution of each cation multiplied by its ideal bond-distance (Table 229 230 7). More details about the ideal distances as well as the optimization procedure may be found 231 in Bosi et al. (2004) and Bosi and Lucchesi (2004; 2007). The robustness of this approach was 232 confirmed by another optimization procedure (Wright et al. 2000), which led to very similar 233 cation distributions (Table 7). This result represents another example of convergence of these two procedures to similar solutions for tourmaline (e.g., Bosi and Lucchesi 2007; Filip et al. 234 2012; Bosi et al. 2012a, 2013a; Bosi and Skogby 2013). 235

The final structural formula is as follows:

- $237 \qquad {}^{X}(Na_{0.87}Ca_{0.07}\square_{0.04}K_{0.02})_{\Sigma1.00} \\ {}^{Y}(Cr^{3+}_{2.29}Mg_{0.71})_{\Sigma3.00} \\ {}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{1.18}V^{3+}_{0.22}Fe^{3+}_{0.01})_{\Sigma6.00} \\ {}^{Y}(Cr^{3+}_{2.29}Mg_{0.71})_{\Sigma3.00} \\ {}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{1.18}V^{3+}_{0.22}Fe^{3+}_{0.01})_{\Sigma6.00} \\ {}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{0.22}Fe^{3+}_{0.01})_{\Sigma6.00} \\ {}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.01})_{\Sigma6.00} \\ {}^{Z}(Al_{3.04}Mg_{1.54}Cr^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}Fe^{3+}_{0.22}F$
- $238 \qquad {}^{T}[(Si_{5.96}Al_{0.04})O_{18}] \ {}^{B}(BO_{3})_{3} \ {}^{V}(OH)_{3} \ {}^{W}[O_{0.73}F_{0.25}(OH)_{0.02}]_{\Sigma1.00}.$
- 239 The bond-valence analysis is consistent with the optimized structural formulae. Bond-valence
- sums (BVS), using the formula and bond-valence parameters from Brown and Altermatt
- 241 (1985), are reported in Table 8. In this regard, note that the equation for the W(OH) calculation
- [reported by Bosi (2013),  $^{W}(OH) = 2 1.01 \text{ BVS}(O1) 0.21 \text{F}]$  yields an  $^{W}(OH)$  value very
- close to that of the empirical formula: 0.03 apfu and 0.02 apfu, respectively.

### Name and crystal chemistry

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246 The composition of the present sample PR85v is consistent with a tourmaline 247 belonging to the alkali group, oxy-subgroup 3 (Henry et al. 2011): Na-dominant at the X position of the general formula of tourmaline XY<sub>3</sub>Z<sub>6</sub>T<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>V<sub>3</sub>W and oxygen-dominant 248 at W with  $O^{2-} > (OH+F)$ . As  $Cr^{3+}$  is the dominant cation at Y and  $Al^{3+}$  is the dominant cation 249 at Z, its end-member composition may be represented as NaCr<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O. 250 251 Although tourmaline chemical compositions compatible with that of the studied sample were 252 already reported in the literature (samples TMt6b, TMt3c, TMpr79f and TM1p43e of Bosi et al. 2004), no samples have vet been formally defined as Cr<sup>3+</sup>- and Al<sup>3+</sup>-dominant at Y and Z 253 (respectively). Therefore, the present tourmaline can be classified as a new species. Its closest 254 255 end-member composition of a valid tourmaline species is that of oxy-dravite. The name chromo-alumino-povondraite is the same as the original name given in the 2009-088 proposal. 256

Although there exists a significant degree of  $Cr^{3+}$  and Mg disorder over the Y and Z sites, the structural formula of sample PR85v indicates a preference of  $Cr^{3+}$  for the Y site and

Mg for the Z site. Aluminum, on the other hand, is completely ordered at the Z site. The O1 site is dominated by  $O^{2-}$  with a significant concentration of F (0.25 apfu) and a minor concentration of (OH). The presence of only minor concentrations of (OH) at O1 (~0.02 apfu) are consistent with the observation of weak absorption bands at 3720 and 3760 cm<sup>-1</sup> in the infrared spectrum (Fig. 1), i.e. the area typically ascribed to the O1 site (see below).

# **Infrared spectroscopy**

Spectra recorded in polarized mode perpendicular and parallel to the crystallographic **c** axis show an intense broad band around 3530 cm<sup>-1</sup> and two weak bands at 3720 and 3760 cm<sup>-1</sup>, all strongly polarized in the **c** direction (Fig. 1). As it was not possible to thin the sample sufficiently to avoid off-scale absorption for the main band in the **c**-axis direction, any possible fine structure cannot be discerned. However, in line with previous studies (Bosi et al. 2012b; Bosi et al. 2013b), the band can be related to the local arrangement ( ${}^{Y}Cr^{3+} {}^{Z}R {}^{Z}R$ )-O3, i.e., to the occurrence of (OH) at the V position of the tourmaline general formula (O3 site in the structure). The two weak bands at ca. 3720 and 3760 cm<sup>-1</sup> are consistent with the minor concentrations of (OH) (ca. 0.02 apfu) assigned to the W position (O1 site in the structure), and may be related to the local arrangements  ${}^{Y}(Mg Mg Mg)$  and  ${}^{Y}(Mg Mg Rg)$  (cf. Gonzalez-Carreño et al. 1988; Bosi et al. 2012b; Bosi et al. 2013b).

#### **Optical-absorption spectroscopy**

The optical spectra of chromo-alumino-povondraite show two broad absorption bands at ca. 435 and 590 nm, superimposed on an intense UV absorption edge (Fig. 3). These absorption bands are ascribed to spin-allowed *d-d* transitions in octahedrally coordinated Cr<sup>3+</sup>, in accord with comparable bands at ca. 430 and 590 nm observed in spectra of Fe-bearing chromo-alumino-povondraite (Bosi et al. 2013b) as well as bands observed in spectra of other Cr-bearing tourmalines (Manning 1969; Taran et al. 1993; Ertl et al. 2008).

# RELATION TO THE OTHER SPECIES

Chromo-alumino-povondraite [NaCr<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O] is related to oxydravite [NaAl<sub>3</sub>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, in the ordered form] by the substitution  $3Cr^{3+} \rightarrow 3Al$ , and to oxy-chromium-dravite [NaCr<sub>3</sub>(Cr<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O] by the substitution  $4Al \rightarrow 4Cr^{3+}$ . The properties of these three tourmalines are compared in Table 9.

Figure 3 displays the ternary diagram of the  $V_{total}$ - $Cr_{total}$ - $Al_{total}$  system for oxy(O1)-tourmaline  $Na^Y(R^{3+})_3{}^Z(R^{3+}_4Mg_2)Si_6O_{18}(BO_3)_3(OH)_3O$  (where  $R^{3+}=V$ , Cr, Al) showing the position of the studied sample and that of ideal chromo-alumino-povondraite. This species occurs between the series oxy-chromium-dravite and oxy-dravite, its compositional limits were calculated by Bosi et al. (2013b) at  $^{Y+Z}(Cr_{1.5}Al_{5.5})$ , corresponding to  $Na^Y(Cr_{1.5}Al_{1.5})^Z(Al_4Mg_2)Si_6O_{18}(BO_3)_3(OH)_3O$  towards to oxy-dravite, and and  $^{Y+Z}(Cr_5Al_2)$ , corresponding to  $Na^Y(Cr_3)^Z(Cr_2Al_2Mg_2)Si_6O_{18}(BO_3)_3(OH)_3O$  toward to oxy-chromium-dravite. In a similar way, the compositional limits of vanadio-oxy-chromium-dravite and vanadio-oxy-dravite were inferred (Bosi et al. 2014a,b).

302 IMPLICATIONS

The discovery of the new mineral chromo-alumino-povondraite provides new information on the crystal chemistry of the tourmaline supergroup. The current chemical data supports complete solid-solution of V<sup>3+</sup>, Cr<sup>3+</sup> and Al in species of the tourmaline supergroup (Reznitsky et al. 2001; Bosi et al. 2004, 2013a,b, Bosi 2014a,b) showing that tourmaline with chemical compositions classified as chromo-alumino-povondraite can be either Al-dominant or Cr-dominant as a result of the compositional boundaries along the solid solution between Al and Cr.

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481	$Na^{Y}(R^{3+})_{3}^{Z}(R^{3+}_{4}Mg_{2})Si_{6}O_{18}(BO_{3})_{3}(OH)_{3}O$ . The full black circle is the studied
482	sample.
483	

484

Table 1. Single-crystal X-ray diffraction data details for chromo-alumino-povondraite

Sample	PR85v
Crystal sizes (mm)	$0.10\times0.10\times0.20$
a (Å)	16.0277(2)
c (Å)	7.3085(1)
$V(\mathring{A}^3)$	1625.93(5)
Range for data collection, 2θ (°)	5-72
Reciprocal space range hkl	–26 ≤ <i>h</i> ≤ 15
	$-26 \le k \le 26$
	<b>–</b> 11 ≤ <i>I</i> ≤ 10
Total number of frames	3265
Set of measured reflections	7819
Unique reflections, $R_{int}$ (%)	1803, 2.02
Redundancy	8
Absorption correction method	SADABS
Refinement method	Full-matrix last-squares on F <sup>2</sup>
Structural refinement program	SHELXL-97
Extinction coefficient	0.00057(8)
Flack parameter	0.05(2)
wR2 (%)	3.43
R1 (%) all data	1.68
$R1 \ (\%) \text{ for } I > 2\sigma(I)$	1.58
GooF	1.055
Largest diff. peak and hole (±e <sup>-</sup> /Å <sup>3</sup> )	-0.39 and 0.34

*Notes*:  $R_{\text{int}}$  = merging residual value; R1 = discrepancy index, calculated from F-data; wR2 = weighted discrepancy index, calculated from F-data; GooF = goodness of fit; Diff. Peaks = maximum and minimum residual electron density. Radiation, Mo $K\alpha$  = 0.71073 Å. Data collection temperature = 293 K. Space group R3m; Z = 3.

TABLE 2. Fractional atom coordinates and site occupancy for chromo-alumino-povondraite

Site	X	У	z	Site occupancy
Χ	0	0	0.22801(18)	Na <sub>1.044(6)</sub>
Υ	0.12380(2)	0.061898(10)	0.63723(5)	Cr <sub>0.767(4)</sub> Mg <sub>0.233(4)</sub>
Z	0.297903(19)	0.261913(19)	0.60938(6)	Cr <sub>0.280(2)</sub> Mg <sub>0.720(2)</sub>
В	0.10993(6)	0.21986(12)	0.4549(2)	B <sub>1.00</sub>
Τ	0.190426(19)	0.18870(2)	0	Si <sub>1.00</sub>
O1 (≡W)	0	0	0.7661(3)	O <sub>1.00</sub>
O2	0.06037(4)	0.12074(8)	0.48900(15)	O <sub>1.00</sub>
O3 (≡V)	0.25735(9)	0.12868(4)	0.51027(17)	O <sub>1.00</sub>
04	0.09256(4)	0.18511(9)	0.07144(16)	O <sub>1.00</sub>
O5	0.18285(9)	0.09142(4)	0.09074(15)	O <sub>1.00</sub>
O6	0.19226(6)	0.18294(6)	0.78002(12)	O <sub>1.00</sub>
O7	0.28334(5)	0.28296(5)	0.07444(12)	O <sub>1.00</sub>
O8	0.20763(6)	0.26830(6)	0.43859(13)	O <sub>1.00</sub>
H3	0.2575(16)	0.1287(8)	0.395(3)	H <sub>1.00</sub>

TABLE 3. Displacement parameters (Å<sup>2</sup>) for chromo-alumino-povondraite

Site	U <sup>11</sup>	U <sup>22</sup>	U 33	U <sup>23</sup>	U <sup>13</sup>	U 12	$U_{\rm eq}/U_{\rm iso}^*$
X	0.0212(5)	0.0212(5)	0.0186(7)	0	0	0.0106(2)	0.0203(4)
Y	0.00518(13)	0.00495(10)	0.00650(12)	-0.00037(4)	-0.00074(8)	0.00259(6)	0.00552(7)
Z	0.00424(12)	0.00464(12)	0.00442(11)	0.00026(8)	0.00002(8)	0.00207(10)	0.00450(7)
В	0.0065(5)	0.0073(6)	0.0083(6)	0.0005(5)	0.0003(2)	0.0036(3)	0.0072(3)
T	0.00567(13)	0.00524(12)	0.00609(13)	-0.00050(9)	-0.00042(10)	0.00266(10)	0.00570(7)
O1	0.0071(5)	0.0071(5)	0.0073(8)	0	0	0.0035(2)	0.0072(3)
O2	0.0066(3)	0.0048(4)	0.0083(5)	0.0006(3)	0.00029(17)	0.0024(2)	0.0068(2)
О3	0.0138(5)	0.0128(4)	0.0054(4)	0.00036(19)	0.0007(4)	0.0069(3)	0.0105(2)
04	0.0085(3)	0.0174(6)	0.0087(5)	-0.0004(4)	-0.0002(2)	0.0087(3)	0.0106(2)
O5	0.0161(6)	0.0079(3)	0.0087(5)	0.0005(2)	0.0011(4)	0.0081(3)	0.0100(2)
O6	0.0104(3)	0.0080(3)	0.0055(3)	-0.0009(2)	-0.0004(3)	0.0052(3)	0.00773(14)
07	0.0084(3)	0.0072(3)	0.0092(3)	-0.0005(3)	-0.0024(3)	0.0015(3)	0.00938(15)
08	0.0055(3)	0.0094(4)	0.0173(4)	0.0031(3)	0.0010(3)	0.0033(3)	0.01093(16)
H3	. ,		. ,	. ,	` ,	. ,	0.016*

*Notes:* Equivalent ( $U_{eq}$ ) and isotropic ( $U_{iso}$ ) displacement parameters; H-atom was constrained to have a  $U_{iso}$  1.5 times the  $U_{eq}$  value of the O3 oxygen.

TABLE 4. Selected bond distances (Å) for chromo-alumino-povondraite

TABLE 4. OCICCI	TABLE 4. Ociocica bona distances (7) for enforme diamino povendrala					
<i>B</i> -O2	1.3614(12)	Y-O1	1.9594(9)			
<i>B</i> -O8 <sup>A</sup> (× 2)	1.398(2)	Y-O2 <sup>B</sup> (× 2)	2.0138(7)			
< <i>B</i> -O>	1.374	Y-O3	2.0731(12)			
		Y-O6 <sup>C</sup> (× 2)	1.9818(8)			
<i>T</i> -O4	1.6267(5)	< Y-O>	2.004			
<i>T</i> -O5	1.6420(5)					
<i>T</i> -O7	1.5957(7)	<i>Z</i> -O3	2.0296(6)			
<i>T</i> *-O6	1.6115(9)	<i>Z</i> -O6	1.9701(8)			
< <i>T</i> -O>	1.619	<i>Z</i> -O8 <sup>E</sup>	1.9380(8)			
		<i>Z</i> -O7 <sup>E</sup>	1.9839(8)			
X-O2 <sup>B,F</sup> (× 3)	2.5390(14)	<i>Z</i> -O7 <sup>D</sup>	1.9504(8)			
X-O4 <sup>B,F</sup> (× 3)	2.8126(13)	<i>Z</i> -08	1.9519(9)			
$X$ -O5 <sup>B,F</sup> ( $\times$ 3)	2.8127(13)	< <i>Z</i> -O>	1.971			
<x-o></x-o>	2.694	О3-Н	0.84(2)			

*Notes*: Standard uncertainty in parentheses. Superscript letters: A = (y - x, y, z); B = (y - x, -x, z); C = (x, x - y, z); D = (y - x + 1/3, -x + 2/3, z + 2/3); E = (-y + 2/3, x - y + 1/3, z + 1/3); F = (-y, x - y, z). Transformations relate coordinates to those of Table 2.

<sup>\*</sup> Positioned in adjacent unit cell.

TABLE 5. Powder X	TABLE 5. Powder X-ray data for chromo-alumino-povondraite			
I <sub>(meas.)</sub> %	d <sub>(meas.)</sub> Å	d <sub>(calc.)</sub> Å	hkl	
47	6.496	6.496	101	
24	5.058	5.055	0 2 1	
18	4.643	4.646	300	
42	4.279	4.280	2 1 1	
55	4.019	4.024	220	
44	3.548	3.550	0 1 2	
6	3.417	3.421	1 3 1	
51	3.010	3.012	122	
6	2.934	2.932	3 2 1	
3	2.659	2.662	3 1 2	
100	2.601	2.606	0 5 1	
4	2.477	2.479	2 4 1	
13	2.415	2.411	232	
12	2.370	2.370	5 1 1	
5	2.327	2.323	600	
11	2.225	2.220	502	
12	2.161	2.165	0 3 3	
46	2.006	2.068	1 5 2	
7	2.039	2.042	161	
2	2.015	2.012	4 4 0	
18	1.943	1.944	3 4 2	
6	1.913	1.907	1 4 3	
4	1.872	1.869	6 2 1	
14	1.685	1.685	063	
12	1.658	1.659	271	
19	1.608	1.610	5 5 0	
2	1.563	1.562	4 6 1	
8	1.533	1.534	7 3 1	

Note: Estimated errors in  $d_{\text{(meas.)}}$ -spacing range from 0.01 Å for large d-values to 0.001 Å for small d-values

**TABLE 6.** Chemical composition of chromo-alumino-povondraite

	Weight %		Apfu
SiO <sub>2</sub>	34.06(08)	Si	5.96(6)
B <sub>2</sub> O <sub>3</sub> *	9.93	В	3.00
$Al_2O_3$	14.94(10)	Al	3.08(3)
$Cr_2O_3$	25.09(21)	Cr <sup>3+</sup>	3.47(4)
$V_2O_3$	1.56(37)	V <sup>3+</sup>	0.22(5)
Fe <sub>2</sub> O <sub>3</sub> <sup>†</sup>	0.10(2)	Fe <sup>3+</sup>	0.014(3)
MgO	8.65(14)	Mg	2.26(4)
CaO	0.37(9)	Ca	0.07(2)
Na <sub>2</sub> O	2.57(14)	Na	0.87(5)
K <sub>2</sub> O	0.08(4)	K	0.017(8)
F	0.45(7)	F	0.25(4)
H <sub>2</sub> O*	2.59	ОН	3.02
-O=F	-0.19		
Total	100.20		

<sup>\*</sup> Calculated by stoichiometry.

*Notes:* Errors for oxides are standard deviations (in brackets) of 10 spot analyses. Standard errors for the atomic proportions (in brackets) were calculated by error-propagation theory; apfu = atoms per formula unit.

<sup>†</sup> Calculated as Fe<sup>3+</sup>.

TABLE 7. Cation site populations (apfu), mean atomic numbers and mean bond lengths (Å) for chromo-alumino-povondraite

oito	site population	mean ator	nic number	mean bond length	
site	site population	refined	calculated	refined	calculated*
X	0.87 Na + 0.07 Ca + 0.04 □ + 0.02 K	11.49(7)	11.30		
Υ	2.29 Cr <sup>3+</sup> + 0.71 Mg	21.2(1)	21.2	2.004(1)	2.003
	$(2.06 \text{ Cr}^{3+} + 0.66 \text{ Mg} + 0.23 \text{ V}^{3+} + 0.03 \text{ Fe}^{3+} + 0.03 \text{ Al})^{\dagger}$				
Z	$3.04 \text{ AI} + 1.54 \text{ Mg} + 1.18 \text{ Cr}^{3+} + 0.22 \text{ V}^{3+} + 0.01 \text{ Fe}^{3+}$	15.4(1)	15.3	1.971(1)	1.964
	$(2.99 \text{ Al} + 1.59 \text{ Mg} + 1.42 \text{ Cr}^{3+})^{\dagger}$				
Τ	5.96 Si + 0.04 Al	14 <sup>‡</sup>	14.00	1.619(1)	1.620
В	3 B	5 <sup>‡</sup>	5		

*Notes*: apfu = atoms per formula unit;

<sup>\*</sup> Calculated using the ionic radii of Bosi and Lucchesi (2007).

† Site populations optimized by the procedure of Wright et al. (2000).

‡ Fixed in the final stages of refinement

**TABLE 8.** Bond valence calculations (valence units) for chromo-alumino-povondraite.

<u> </u>						
Site	X	Υ	Z	T	В	BVS
O1(W)		0.50 <sup>×3</sup> →				1.50
O2	0.14 <sup>×3</sup> ↓	$0.45^{\times 2} \downarrow \rightarrow$			0.93	1.97
O3(V)		0.38	$0.39^{\times 2} \rightarrow$			1.16
O4	$0.07^{\times3} \downarrow$			$0.99^{\!\times\!2}\!\!\rightarrow$		2.05
O5	0.09 <sup>×3</sup> ↓			$0.95^{\!\times\!2}\!\!\rightarrow$		1.99
O6		$0.49^{\times 2} \downarrow$	0.46	1.03		1.98
07			0.480.44	1.08		2.00
O8			0.480.50		1.02 <sup>×2</sup> ↓	2.00
BVS	0.89	2.75	2.74	4.06	2.98	
MFV	1.03	2.76	2.74	3.99	3.00	

*Note*: BVS = bond valence sum; MFV = mean formal valence from site populations

 TABLE 9. Selected properties of oxy-dravite, chromo-alumino-povondraite and oxy-chromium-dravite

	Oxy-dravite	Chromo-alumino-povondraite	Oxy-chromium-dravite
a (Å)	15.9273(2)	16.0277(2)	16.0539(7)
С	7.2001(1)	7.3085(1)	7.3247(5)
$V(Å^3)$	1581.81(4)	1625.93(5)	1634.86(15)
Space group	R3m	R3m	R3m
Optic sign	Uniaxial (-)	Uniaxial (-)	Uniaxial (-)
ω	1.650(5)	1.745(5)	1.765(5)
3	1.620(5)	1.685(5)	1.715(5)
Color	Dark red	Green	Emerald Green
Pleochroism	O = orange	O = emerald green	O = dark green
	E = pink	E = pale yellowish green	E = yellow green
Ref.*	(1)	(2)	(3)

<sup>\*(1)</sup> Bosi and Skogby (2013); (2) this study; (3) Bosi et al. (2012a).

Figure 1

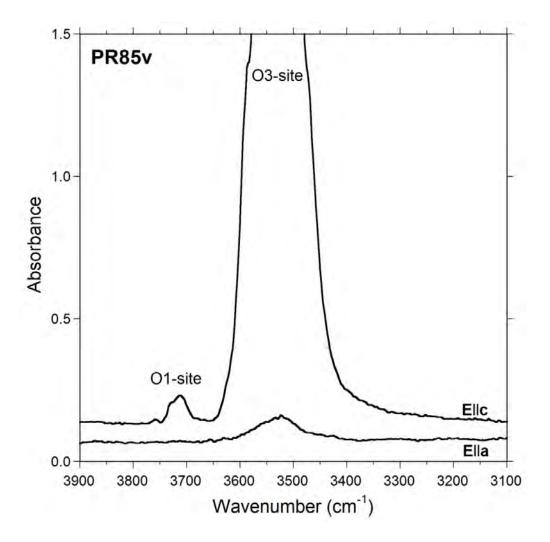


Figure 2

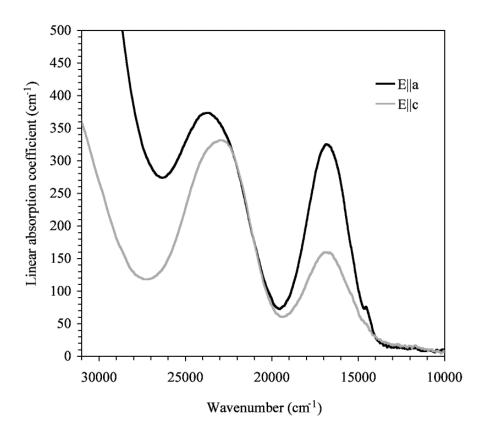


Figure 3

