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1	Revision 1
2	High-pressure phase transitions in $FeCr_2O_4$ and structure analysis of new post-spinel $FeCr_2O_4$ and
3	Fe ₂ Cr ₂ O ₅ phases with meteoritical and petrological implications
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12	
13	Abstract
14	We determined phase relations in FeCr ₂ O ₄ at 12-28 GPa and 800-1600°C using a multianvil apparatus. At
15	12-16 GPa, FeCr ₂ O ₄ spinel (chromite) first dissociates into two phases: a new Fe ₂ Cr ₂ O ₅ phase + Cr ₂ O ₃ with the
16	corundum structure. At 17-18 GPa, the two phases combine into CaFe ₂ O ₄ -type and CaTi ₂ O ₄ -type FeCr ₂ O ₄
17	below and above 1300°C, respectively. Structure refinements using synchrotron X-ray powder diffraction data
18	confirmed the CaTi ₂ O ₄ -structured FeCr ₂ O ₄ (<i>Cmcm</i>), and indicated that the Fe ₂ Cr ₂ O ₅ phase is isostructural to a
19	modified ludwigite-type Mg ₂ Al ₂ O ₅ (Pbam). In situ high-pressure high-temperature X-ray diffraction
20	experiments showed that CaFe ₂ O ₄ -type FeCr ₂ O ₄ is unquenchable and is converted into another FeCr ₂ O ₄ phase

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21	on decompression. Structural analysis based on synchrotron X-ray powder diffraction data with transmission
22	electron microscopic observation clarified that the recovered FeCr ₂ O ₄ phase has a new structure related to
23	CaFe ₂ O ₄ -type. The high-pressure phase relations in FeCr ₂ O ₄ reveal that natural FeCr ₂ O ₄ -rich phases of
24	CaFe ₂ O ₄ - and CaTi ₂ O ₄ -type structures found in the shocked Suizhou meteorite were formed above about 18
25	GPa at temperature below and above 1300 °C, respectively. The phase relations also suggest that the natural
26	chromitites in the Luobusa ophiolite previously interpreted as formed in the deep-mantle were formed at
27	pressure below 12-16 GPa.
28	Keywords: Post-spinel, Rietveld analysis, crystal structure, high pressure, phase transition, shocked meteorite,
29	FeCr ₂ O ₄ , ophiolite

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phases in the subducted crust into the deep mantle. CaFe₂O₄-, CaTi₂O₄- and CaMn₂O₄-structures have been 35regarded as the $A^{2+}B^{3+}O_4$ post-spinel structures (Ringwood, 1975). We hereafter abbreviate CaFe₂O₄, CaTi₂O₄ 36 and CaMn₂O₄ as CF, CT and CM respectively. The $A^{2+}B^{3+}_{2}O_4$ phases with the CF- and CT-structures (space 37groups *Pnma* and *Cmcm*, respectively) consist of double chains of edge-shared $B^{3+}O_6$ octahedra running 38 parallel to one of the orthorhombic cell axes, and A^{2+} ions occupy tunnel spaces surrounded by corner-sharing 3940of four double chains (Decker et al., 1957; Rogge et al., 1998; Yamanaka et al., 2008). The CM-structure (*Pbcm*) is interpreted to be formed by Jahn-Teller distortion of $B^{3+}O_6$ octahedra from the CT-structure (Geisber 41et al., 2001; Yamanaka et al., 2008). FeCr₂O₄ chromite with the spinel structure forms spinel solid solutions 42with other end-members, MgCr₂O₄, MgAl₂O₄, FeAl₂O₄ etc., in various crustal and upper-mantle rocks. 4344Chromite is also a major chromium ore mineral. It was recently found that by compression at room temperature pure FeCr₂O₄ spinel undergoes a cubic-to-tetragonal transition due to Jahn-Teller effect in tetrahedrally 45coordinated Fe²⁺ (Kyono et al., 2012). However, compared with other $A^{2+}B^{3+}_{2}O_{4}$ spinels, particularly MgAl₂O₄, 46

little is known about high-pressure transitions in FeCr₂O₄. 47

FeCr₂O₄-rich spinel and its high-pressure polymorphs were found in a shock-metamorphosed L6-chondrite 48named Suizhou which contained shock-induced melt veins and high-pressure minerals including ringwoodite 49and majorite garnet (Chen et al., 2003a, b). Chen et al. (2003a, b) synthesized CF- and CT-structured chromite 50phases with the same composition as in the L6-chondrite in laser-heated diamond anvil cell experiments at 12.5 51

52	GPa and above 20 GPa, respectively. They interpreted that the FeCr ₂ O ₄ -rich high-pressure polymorphs in the
53	meteorite were the CF- and CT-phases, on the basis of powder X-ray diffraction patterns. However, detailed
54	structural analysis of the CF- and CT-phases and determination of their stability relations with chromite spinel
55	have not yet been made. Chromitites composed mostly of FeCr ₂ O ₄ -rich spinel which occured in the Luobusa
56	ophiolite in Tibet included high-pressure minerals such as diamond and coesite, which suggest possible
57	deep-mantle origin of the chromitites (Yang et al., 2007; Arai, 2010; Yamamoto et al., 2009).
58	At 1600 °C, MgAl ₂ O ₄ spinel transforms to a mixture of MgO + Al ₂ O ₃ at 16 GPa and to a CF-type
59	phase at 27 GPa (Akaogi et al., 1999). The CF-type MgAl ₂ O ₄ transforms to a CT-type phase at about 40 GPa
60	(Funamori et al., 1998). Above about 2000°C, however, we recently found that MgAl ₂ O ₄ spinel first dissociates
61	into an $Mg_2Al_2O_5$ phase + Al_2O_3 at 20-22 GPa, and changes to CF phase at 26-28 GPa via an unknown
62	MgAl ₂ O ₄ phase (Kojitani et al. 2010). Our structure analysis indicated that the Mg ₂ Al ₂ O ₅ phase has a modified
63	ludwigite structure (Pbam) consisting of edge- and corner-shared (Mg,Al)O ₆ octahedra running parallel to
64	c-axis, and tunnel spaces surrounded by the octahedra are occupied by Mg^{2+} (Enomoto et al., 2009). We
65	abbreviate the modified ludwigite structure as mLd. The mLd structure of Mg ₂ Al ₂ O ₅ is similar to that of
66	recently found high-pressure Fe ₄ O ₅ phase (Lavina et al., 2011), but the manner of octahedral connection is
67	different between the two structures. To our best knowledge, the Mg ₂ Al ₂ O ₅ phase has been the only compound
68	having the mLd structure.
69	In this study, we have investigated in detail the post-spinel transitions of FeCr ₂ O ₄ chromite at high

into a mixture of Fe₂Cr₂O₅ and Cr₂O₃. The two phases combine into a FeCr₂O₄ phase with the CF structure at

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pressure and high temperature. We have found that with increasing pressure FeCr₂O₄ chromite first dissociates

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80	Experimental methods
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78	chromitites in the Luobusa ophiolite in the deep mantle.
77	discuss on P,T conditions of the shock event which the Suizhou meteorite suffered and on the origin of the
76	FeCr ₂ O ₄ phase have been analyzed and refined, together with that of CT-type FeCr ₂ O ₄ . Based on the data, we
75	phases have been precisely determined. The structures of the Fe ₂ Cr ₂ O ₅ phase and the CF-related structured
74	CF-related structure on release of pressure. The high-pressure phase relations in FeCr ₂ O ₄ involving these new
73	in situ X-ray diffraction experiments that the CF-type FeCr ₂ O ₄ is changed into a FeCr ₂ O ₄ phase with a new
72	relatively low temperature and into the CT-type FeCr ₂ O ₄ at relatively high temperature. We have also found by

81 1. Starting materials

FeCr₂O₄ spinel was prepared from a mixture of reagent-grade Fe₂O₃ and Cr₂O₃ with 1:2 molar ratio. 82 The mixture was heated at 1200°C for 24 h in a controlled oxygen fugacity using a mixture of H₂, CO₂ and Ar 8384with volume ratios of 1:1:2. After heating, the product was confirmed to be single-phase of FeCr₂O₄ with spinel 85 structure by a powder X-ray diffractometer ($CrK\alpha$) and a scanning electron microscope with an energy dispersive X-ray spectrometer (SEM-EDS). The FeCr₂O₄ spinel powder mixed with 5 wt% Fe metal powder 86 was used as the starting material for phase equilibrium experiments. Fe_xO wüstite was made by heating a 87 88 reagent-grade Fe_2O_3 at the same conditions as $FeCr_2O_4$ spinel. The composition of Fe_xO was estimated as x =0.933 from the relationship between the lattice parameter and x by McCammon (1993). The $Fe_{0.933}O$ and Cr_2O_3 89 powders were mixed with atomic ratio Fe : Cr = 1:1 with additional 5 wt% Fe metal powder, and the mixture 90 91was used as the starting material to synthesize $Fe_2Cr_2O_5$ phase. The addition of 5 wt% Fe metal in both the

92 starting materials was to keep iron in the phases ferrous state.

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94 2. High-pressure high-temperature experiments by quench method

95	The high-pressure high-temperature quench experiments were carried out at 11.7-28 GPa and
96	800-1600°C with a Kawai-type 6-8 multianvil high-pressure apparatus at Gakushuin University. Tungsten
97	carbide anvils with truncated edge lengths (TEL) of 2.5 mm were used with a pressure medium of Cr ₂ O ₃ -doped
98	MgO octahedron with 7.0 mm edge length. A cylindrical rhenium furnace was placed in the center of the
99	magnesia octahedron. For thermal insulation, a LaCrO3 sleeve was put outside of the rhenium heater and
00	LaCrO ₃ end-plugs were placed at both ends of the heater. Temperature was measured at the central part of
01	outer-surface of the furnace with a Pt/Pt-13%Rh thermocouple of 0.1 mm in diameter. No correction was made
02	on pressure effect of the thermocouple emf. Phase relations of FeCr ₂ O ₄ were examined with the multi-sample
03	cell technique (Fei and Bertka, 1997) described by Ishii et al. (2011, 2012). The starting material of FeCr ₂ O ₄
04	spinel + 5 wt% Fe and a pressure marker (Mg ₂ SiO ₄ , MgSiO ₃ or MgO+Al ₂ O ₃) were packed in two holes of 0.2
05	mm diameter in a Re capsule which was 1.0 mm in diameter and 0.7 mm in thickness. Two Re discs of 1.0 mm
06	in diameter and 0.1 mm thickness were placed on both sides of the Re capsule. A boron nitride sleeve and discs
07	were inserted between the Re capsule and the heater for electrical insulation.
08	The EeCr ₂ Ω_{i} high-pressure phases with the CE-related and CT structures used for Rietveld refinements

08	The FeCr ₂ O ₄ high-pressure phases with the CF-related and CT structures used for Rietveld refinements
09	were synthesized from FeCr ₂ O ₄ spinel as the starting material at 25 GPa, 1000°C with a Fe capsule and at 25
10	GPa, 1400°C with an Au-Pd capsule, respectively. A BN sleeve was placed between the Fe or Au-Pd capsule
11	and the tubular rhenium heater for electrical insulation. The CF-related FeCr ₂ O ₄ phase for TEM observation

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and the Fe₂Cr₂O₅ high-pressure phase for structure refinement were synthesized from sample materials packed
directly in the rhenium heater and held at 25 GPa, 1000°C and 16 GPa, 1300°C, respectively.
Pressure calibration was made in a similar manner to those of Ishii et al. (2011, 2012). The pressure
was calibrated at room temperature using transition points of ZnS (15.5 GPa), GaAs (18.3 GPa) and GaP (23)

- 16 GPa) (Dunn and Bundy, 1978, Ito, 2007). Temperature effect on pressure was corrected at 1600°C by Mg₂SiO₄
- 17 forsterite-wadsleyite transition (15.1 GPa) (Morishima et al., 1994), Mg₂SiO₄ wadsleyite-ringwoodite transition
- 18 (21.3 GPa) (Suzuki et al., 2000), MgSiO₃ akimotoite-perovskite transition (22.3 GPa) (Fei et al., 2004), and
- 19 transition of Al₂O₃ corundum + MgO periclase to MgAl₂O₄ calcium ferrite (24.9 GPa) (Irifune et al., 2002),
- using the pressure markers (MgSiO₃, Mg₂SiO₄ and MgO + Al_2O_3) packed in one of the holes of the Re capsule.
- The pressure calibration curve at 1200 °C was very close to that at 1600°C. The relative uncertainty of pressure
- of the quench experiments in this study was estimated to be about ± 0.2 GPa (Ishii et al., 2011, 2012).

In each high-pressure run, pressure was increased to a targeted pressure at an almost constant rate during about 2-4 h, and then temperature was increased to 800-1600°C at a rate of about 100°C/min. The sample assembly was kept for 1-3 h at the pressure-temperature conditions, then quenched, and decompressed for 5-12 h, and finally recovered to ambient conditions. The recovered samples in the capsules were mounted on slide-glass plates with epoxy resin, and were polished into flat for phase identification and compositional analysis.

The phases in the recovered samples were identified using a microfocus X-ray diffractometer (Rigaku RINT 2500V, MDG) with an X-ray beam collimated to 50 μ m. The powder X-ray diffractometer (Rigaku RINT 2500V) was used with the step-scan mode (step size of 0.02°) in the 2 θ range of 10-140° to determine

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32	lattice parameters of the samples. Both the X-ray diffraction measurements were conducted using CrKa
33	radiated from a rotating anode at 45 kV and 250 mA. The scanning electron microscope (SEM, JEOL
34	JMS-6360) with the energy dispersive X-ray spectrometer (EDS, Oxford INCA energy 300) was used for
35	composition analysis and also for phase identification in the recovered samples. The SEM was operated with
36	acceleration voltage of 15 kV and probe current of 0.43 nA. Natural fayalite and synthetic Cr ₂ O ₃ eskolaite were
37	used as standard materials for Fe and Cr, respectively.
38	
39	3. In-situ X-ray observations at high pressures and high temperatures
40	High-pressure high-temperature in-situ X-ray observations were conducted using a high-pressure
41	apparatus SPEED-1500 with a Kawai-type multianvil system with intense white X-ray of synchrotron radiation
42	at the beam line BL04B1 at SPring-8. An X-ray beam collimated to $50 \times 150 \ \mu\text{m}$ was emitted to the sample
43	chamber in the pressure medium via horizontal and vertical slits through the anvil gap of the first and
44	second-stage anvils. Energy dispersive X-ray diffraction was measured in the range between 3 and 150 keV
45	with a pure Ge solid-state detector connected to a multichannel analyzer with 4096 channels, and radiographic

46 images of the sample chamber were recorded using a CCD camera. The detector was calibrated using known

47 X-ray emission lines of a set of metals (Au, Cu, Mo, Ag, Ta, Pt and Pb). The diffraction angle (2θ) was fixed to

48 about ~6° using diffraction peaks of MgO at ambient conditions. We used sintered MgO as the pressure maker,

and pressure was estimated on the basis of equation of state of MgO by Speziale et al. (2001).

50 The mixture of $FeCr_2O_4$ spinel + 5 wt% Fe was used as the starting material for the *in-situ* 51 high-pressure high-temperature experiments. Tungsten carbide anvils with 4 mm TEL were used in

52	combination with a Cr ₂ O ₃ -doped MgO octahedron of 10.0 mm edge length as the pressure medium. A
53	cylindrical TiB_2 -BN was put in the central part of the magnesia octahedron as a heating element. A LaCrO ₃
54	sleeve and end-plugs were placed between the heater and the pressure medium. Mo electrodes were inserted in
55	the center of both the LaCrO ₃ end-plugs. Temperature was measured in the central part of the furnace with a
56	W-3%Re/W-25%Re thermocouple of 0.05 mm in diameter. The pressure effect of the thermocouple emf was
57	ignored. The MgO pressure maker was put next to the thermocouple junction on both sides, and the samples
58	were put between the pressure marker and MgO end-plugs. Mo foils were put at both ends of the sample. The
59	graphite and MgO rods were placed in the pressure medium and the LaCrO ₃ sleeve as a window for the X-ray
60	beam.
61	The press load was raised to the desired values up to 1000 ton, and then temperature was gradually
62	increased to 1000°C in about 30 min. The samples were kept at the targeted temperature for 1 h. The diffraction
63	patterns of the sample and the pressure marker were measured for 120 and 300 sec, respectively. Using the
64	diffraction patterns at high pressures and room temperature, lattice parameters of CF-type FeCr ₂ O ₄ were refined
65	by Le Bail method (Le Bail, 1988) using the GSAS-EXPGUI software package (Larson and Von Dreele, 2000;
66	Toby, 2001; Sanehira et al., 2008). Reduced Chi-square (χ^2) and R_P values (Larson and Von Dreele, 2000)
67	estimated by Le Bail fitting were used to evaluate goodness of fit after the refinement. χ^2 values were converged
68	less than \sim 3, suggesting that the fits were of high quality.
69	
70	4. Transmission electron microscopic observations and Rietveld structure analyses

71 Transmission electron microscopic (TEM) observation of CF-related FeCr₂O₄ phase was performed to

72 constrain its space group and to determine approximate cell parameters. For the TEM observation, a small chip of $\sim 0.1 \,\mu m$ thickness was prepared from the central part of the sintered sample synthesized at 25 GPa and 731000°C. The thin film was observed by the TEM (JEOL, JEM-2010) using an accelerating voltage 200 kV at 74Geodynamics Research Center, Ehime University. To examine possible existence of centrosymmetry of the 7576FeCr₂O₄ phase, the optical second harmonic generation (SHG) response was measured with a Continuum Minilite YAG:Nd laser (λ =1064 nm) (Rabin and Tang, 1975, Inaguma et al., 2012). 7778For structural determination and refinement, angle-dispersive synchrotron powder X-ray diffraction measurements of CF-related and CT-type $FeCr_2O_4$ phases and the mLd-type $Fe_2Cr_2O_5$ phase were conducted 7980 using an imaging plate as the detector and a Debye-Scherrer camera in a 2θ angle of 0-75° with an angle resolution of 0.01° in the beam line BL02B2 at SPring-8. The wavelength of the X-ray was determined using 81fluorite-type CeO₂ as 0.41927 Å for CT-type FeCr₂O₄ and Fe₂Cr₂O₅ and as 0.49973 Å for CF-related FeCr₂O₄. 82Each polycrystalline sample was finely ground in an agate mortar, and was packed in a Lindemann glass 83 84capillary. The X-ray diffraction patterns were measured at ambient conditions on rotating the samples along the capillary axis. 85

Rietveld analysis was performed using the RIETAN-FP/VENUS package (Izumi and Momma, 2007). The initial structure models of CT-type FeCr₂O₄ and mLd-type Fe₂Cr₂O₅ were CaTi₂O₄ (Bertaut and Blum, 1956) and Mg₂Al₂O₅ (Enomoto et al., 2009), respectively. The starting structure model of CF-related FeCr₂O₄ was constructed by the powder charge-flipping method using SUPERFLIP and EDMA softwares (Baerlocher et al., 2007, Palatinus et al., 2007, 2012). Observed integrated intensities of synchrotron XRD pattern (λ = 0.49973 Å), |F_{obs}|², were extracted by the Le Bail analysis method (Le Bail, 1988). The initial values of lattice

parameters in the Rietveld refinement of each sample were determined from the powder X-ray diffraction patterns (CrK α) with DICVOL06 software (Loüer and Boultif, 2007). Corundum-type Cr₂O₃ was included as the second phase in the Rietveld analysis of all the samples, and Re metal as the third phase in the analysis of mLd-type Fe₂Cr₂O₅.

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Results and discussion

- 98 1. High-pressure phase transitions in FeCr₂O₄ by quench and *in situ* X-ray diffraction experiments.
- Table 1 summarizes the results of the quench experiments in $FeCr_2O_4$ that indicate phases in the 99 00 recovered run products identified by microfocus and powder X-ray diffractometers and SEM-EDS. At 800-1600 °C and about 12-16 GPa, FeCr₂O₄ spinel first dissociates into a mixture of two phases. Powder X-ray 01diffraction patterns of the dissociation products consisted of corundum-type Cr₂O₃ and another phase whose 02diffraction pattern was very similar to that of mLd-type Mg₂Al₂O₅ (Enomoto et al., 2009). The SEM-EDS 0304analysis indicated that the former phase had Cr₂O₃ composition and the latter Fe₂Cr₂O₅ with analyzed Fe : Cr ratio of 1.98(1) : 2.01(2). As described below, the structural refinement confirmed that the Fe₂Cr₂O₅ phase has 05the mLd structure. At temperature above 1300-1400 °C at 16-18 GPa, the mixture of mLd-type Fe₂Cr₂O₅ and 06 corundum-type Cr₂O₃ combine into a single phase whose diffraction pattern was very similar to that of CaTi₂O₄. $\mathbf{07}$ $\mathbf{08}$ The Rietveld refinement and SEM-EDS analysis confirmed that the phase has the CT structure with $FeCr_2O_4$ composition (Fe : Cr = 0.99(1) : 2.01(1)), as shown below. At 16-18 GPa below 1300-1400 °C, on the other 0910 hand, the mixture of $Fe_2Cr_2O_5$ and Cr_2O_3 changes into another phase with the composition of $Fe_2Cr_2O_4$ (Fe : Cr = 0.99(1) : 2.00(1)). Figure 1 shows that, though the diffraction pattern of this phase resembles that of the 11

12	CF-type structure, relative intensities of some diffraction lines are very different from those of the simulated
13	pattern for CF-type FeCr ₂ O ₄ . The structural analysis described below clarified that this recovered FeCr ₂ O ₄
14	phase has a different from the CF-type structure but related to it. Therefore, we call it a "modified
15	CaFe ₂ O ₄ -type" (mCF) FeCr ₂ O ₄ phase Although detailed results of structure analyses of CT-type FeCr ₂ O ₄ ,
16	mCF-type FeCr ₂ O ₄ and mLd-type Fe ₂ Cr ₂ O ₅ are described in the following section, we use here lattice
17	parameters determined by the structure analyses and densities of the phases listed in Table 2. Figure 2 illustrates
18	the phase relations in FeCr ₂ O ₄ up to 28 GPa and 1600°C. The densities of FeCr ₂ O ₄ spinel (Lenaz et al., 2004),
19	Fe ₂ Cr ₂ O ₅ mLd + Cr ₂ O ₃ corundum (Belokoneva and Shcherbakova, 2003), FeCr ₂ O ₄ mCF, and FeCr ₂ O ₄ CT are
20	5.059(1), 5.48(4), 5.213(1), and 5.604(1) g/cm ³ , respectively, at ambient conditions. The lower density of
21	mCF-type $FeCr_2O_4$ compared to that of mLd-type $Fe_2Cr_2O_5 + Cr_2O_3$ corundum is inconsistent with the fact that
22	mCF was synthesized at higher pressure than mLd + Cr_2O_3 . This strongly suggests that mCF is not the stable
23	phase at the P,T conditions of our experiments but is a retrograde transformation product. Therefore, we
24	conducted in situ X-ray diffraction experiments in the P,T field from which mCF was recovered.
25	Figure 3 shows <i>in-situ</i> X-ray diffraction patterns of FeCr ₂ O ₄ at high pressures and high temperatures.
26	By pressurization to 26.2 GPa at room temperature, diffraction pattern of $FeCr_2O_4$ cubic spinel + α -Fe (I in Fig.
27	3a) changed into a broad diffraction pattern consisting of tetragonal spinel, cubic spinel and ϵ -Fe (II in Fig. 3a).
28	The cubic-tetragonal transition of FeCr ₂ O ₄ spinel at room temperature is consistent with the results by Kyono et
29	al. (2012). By heating to 1000 °C at 26.2 GPa, the diffraction pattern of FeCr ₂ O ₄ transformed into a different
30	one (III) with a small pressure decrease to 22.7 GPa. After quenching, the diffraction patterns collected in
31	decompression at 20.1 GPa (IV), at 16.5 GPa (V) and down to at least 9.3 GPa are consistent with (III). It

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am.2014.4736 4/2 32should be noted that peaks of ε -Fe changed into those of γ -FeH_x + FeO by heating to 1000 °C, probably due to reaction of Fe with H₂O adsorbed in the sample, as reported by Ohtani et al. (2005). To examine these patterns 33in more detail, we simulated diffraction pattern at 16.5 GPa and 27°C for CF-type FeCr₂O₄ and compared it 34with the observed one at the P, T conditions in Figure 3b. Except for the peaks of γ -FeH_x and FeO, the observed 3536 diffraction pattern is quite consistent with the simulated pattern of CF-type $FeCr_2O_4$. After decompression to ambient conditions (VI in Fig. 3a), however, the diffraction pattern was changed to that of mCF-type. These in 37situ X-ray diffraction profiles indicate that the stable phase is not mCF but the CF phase at the relevant P, T 38conditions and that CF changes to mCF by decompression. 39We determined the cell volumes of CF-type FeCr₂O₄, 248.44(3) and 245.16(2) Å³, at 16.5 and 20.1 40GPa, respectively, at 25 °C by in situ X-ray diffraction measurements, and those per formula unit are shown in 41Figure 3c. By fitting the third-order Birth-Murnaghan equation of state (BM) to the cell volume data, assuming 42K' = 4, bulk modulus (K_0) and zero-pressure volume (V_0) were determined to be 199 GPa and 266.95 Å³, 43respectively, in which the latter gives molar volume (V_m) of 40.19 cm³/mol. The obtained K_0 and V_m are 44consistent with an empirical relationship, K_0V_m = constant, for CF-type oxides, when we compare our K_0 and 45 V_m of CF-type FeCr₂O₄ with those of other CF-type oxides, MgAl₂O₄ (213 GPa and 36.22 cm³/mol) by Irifune 46 et al. (2002) and NaAlSiO₄ (220 GPa and 36.58 cm³/mol) by Dubrovinsky et al. (2002). The V_0 of CF phase is 47smaller by about 7 % than that of mCF, smaller by ~5 % than zero-pressure volume of $Fe_2Cr_2O_5$ mLd + Cr_2O_3 48corundum, and very close to that of CT-type FeCr₂O₄. As shown in Fig. 2, CF-type FeCr₂O₄ transforms to CT at 49

50 27.5 GPa at 1300 °C. The slightly smaller V_0 of CF than CT is probably derived from relatively large

uncertainties in the K_0 and V_0 of CF determined using only the two volume data. As a future study, it is

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52	necessary to measure more precise compression data of CF. Based on all the results on mCF and CF, we
53	conclude that CF phase is stable above about 18-19 GPa below about 1300 °C, and the CF phase is converted to
54	mCF phase on release of pressure.
55	In Fig. 2, phase boundaries between $FeCr_2O_4$ spinel and mLd-type $Fe_2Cr_2O_5 + Cr_2O_3$ and between
56	mLd-type $Fe_2Cr_2O_5 + Cr_2O_3$ and CT-type $FeCr_2O_4$ have small negative slopes. While the boundary between
57	mLd-type $Fe_2Cr_2O_5 + Cr_2O_3$ and CF-type $FeCr_2O_4$ has a small positive slope, that between CF-type and
58	CT-type has a strongly negative slope probably due to the small volume difference. The high-pressure phase
59	transition behaviors of FeCr ₂ O ₄ shown here have some similarities to those in MgAl ₂ O ₄ by Enomoto et al.
60	(2009) and Kojitani et al. (2010).
61	Chen et al. (2003b) reported that CF-type and CT-type phases were recovered after compression of natural
62	FeCr ₂ O ₄ -rich spinels to maximum pressures of 12.5 GPa and 20 GPa, respectively, at about 2000 °C. These
63	results exhibit some inconsistency with our study with respect to the P,T conditions of CF-type FeCr ₂ O ₄ . The
64	differences in synthesis pressure and temperature for CF phase may be due to uncounted thermal pressure
65	produced during laser-heating (in fact they measured pressure only before laser-heating). Minor components
66	other than FeCr ₂ O ₄ may also affect stability fields of the CF-type and CT-type phases.
67	
68	2. Crystal structures of modified CaFe ₂ O ₄ (mCF)-type FeCr ₂ O ₄ , CaTi ₂ O ₄ (CT)-type FeCr ₂ O ₄ , and modified
69	ludwigite(mLd)-type Fe ₂ Cr ₂ O ₅ .
70	Electron diffraction patterns of mCF by TEM are shown in Figure 4. The results of Rietveld analysis using

synchrotron X-ray diffraction pattern of mCF, CT and mLd are shown in Figures 5-8 with the data in Tables 3

72	and 4. The crystal structures illustrated in Figures 5-8 were drawn by VESTA (Momma and Izumi, 2008). The
73	structural parameters and reliability factors (R_{wp} , R_B and R_F) and goodness-of-fit indicator (S) of mCF, CT and
74	mLd are shown in Table 3. We confirmed that R_{wp} , R_B , R_F and S for the mCF, CT and mLd and the impurity
75	phase converged to sufficiently small values. The interatomic distances, bond angles, effective coordination
76	numbers (n_c) (Nespolo et al., 2001) and the bond valence sum (BVS) (Brown and Altermatt, 1985) values of
77	mCF, CT and mLd are shown in Table 4. In the following sections, we describe detailed results on structure
78	analyses of the three phases.

79

80 a) mCF-type $FeCr_2O_4$

As shown in Figures 1 and 3, the powder X-ray diffraction pattern of FeCr₂O₄ mCF is similar to CF. But 81the intensity of that is different from CF. To determine the crystal structure of FeCr₂O₄ mCF, we first examined 82the mCF sample by TEM. The synthesized sample, run no. 40, was used for TEM observation. Figure 4 8384illustrates electron diffraction patterns ((a) hk0 plane, (b) h0l plane and (c) 0kl plane) taken from different grains of the sample. In Figure 4a, diffraction spots of hk0 with h = 2n + 1 had weaker intensities than those 85with h = 2n, and the spots disappeared by declining the hk0 plane from the Bragg condition. Therefore, the 86 spots of hk0 with h = 2n + 1 appeared by double diffractions from hk1 plane, showing the extinction rule of h =87 2n. The spots of 0k0 with k = 2n + 1 were also double diffractions in hk0 plane, because they vanished by 88 declining the hk0 plane. In a similar manner, Figure 4c shows the systematic absence of k + l = 2n + 1 in the 89spots 0kl, respectively. No extinction rule was observed in Figure 4b except for h00 with h = 2n + 1 and 00l with 90 l = 2n + 1, which appeared by double diffractions in h0l plane. The reflection conditions derived from TEM 91

92	observation are consistent with those of X-ray diffraction. From these results of TEM observations, we
93	conclude that mCF-type $FeCr_2O_4$ has an orthorhombic unit cell with space group <i>Pnma</i> or <i>Pn</i> 2 ₁ <i>a</i> . Next, we
94	performed the SHG measurement, but no SHG signal was observed for the run no. 40 sample. This indicated
95	that the structure of mCF has centrosymmetry and thus its space group is limited to <i>Pnma</i> .
96	The lattice parameters of mCF were obtained as the initial values for Rietveld refinement, using 33
97	reflections of the powder X-ray diffraction pattern (CrKa) with DICVOL06. Figure 5a shows synchrotron
98	powder X-ray diffraction patterns of mCF. $ F_{obs} ^2$ values of 210 reflections in the region d > 0.95 Å were
99	extracted by the Le Bail method. We used the powder charge flipping method to obtain the initial structure of
00	mCF. The cation positions (Fe and Cr) determined using EDMA software were close to those in the CF-type
01	structure, while the oxygen positions could not be determined because oxygen has relatively low electron
02	number. Therefore, we performed Rietveld refinement from the initial structure model in which oxygens were
03	at the same positions as those of CF-type CaCr ₂ O ₄ (Arévalo-López et al. 2010) and Fe and Cr at the positions
04	determined by EDMA.
05	The results of Rietveld analysis of mCF in Figures 5a and 6 indicate mCF has edge-shared CrO ₆
06	octahedral chains running along b-axis, and the two octahedral chains form double chains by edge-sharing.
07	Furthermore, by sharing corners of these four double chains, tunnel structure is formed. Although the structure

08 of mCF is very similar to CF-type structure, positions of Fe are moved by about half of b-edge length along the

09 b-axis of CF, while Cr and O positions in mCF are almost the same as in the CF structure. Thus, the

- 10 coordination polyhedron of Fe in mCF is FeO₅ (FeO₃ plane triangle + two oxygens), while that in CF is FeO₈
- 11 bicapped trigonal prism (FeO₆ prism + two oxygens). A similar asymmetric coordination environment of Fe to

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that of mCF is found in the structure of Fe₃(PO₄)₂ ($P2_1/c$) (Kosiner and Rea, 1974). The reason why Fe²⁺ ions prefer such the special coordination in the mCF structure may be the relatively small ionic radius of ^{VIII}Fe²⁺ (0.91 Å), compared with those of A^{2+} in other CF-structured phases such as CaCr₂O₄ and CdCr₂O₄ (^{VIII}Ca: 1.12 Å, ^{VIII}Cd: 1.10 Å) in the tunnels formed by CrO₆ double-octahedral chains (Shannon, 1976; Hill et al., 1956; Arévalo-López et al. 2010).

- 17
- 18 (b) CT-type $FeCr_2O_4$

Figures 5b and 7 show, together with Tables 3 and 4, the results of Rietveld refinement using 1920synchrotron X-ray diffraction pattern of CT-type FeCr₂O₄. Although the double chains of mCF are linked to have a glide plane corresponding to the plane and direction of corner-sharing between the octahedra, those of 2122CT are linked to have a mirror plane in the direction of those, indicating that both the structures have different types of tunnels. The coordination number of Fe in the tunnel of CT is 6 + 2 (bicapped trigonal prism). We 2324compared CT-type FeCr₂O₄ with the structure of CaTi₂O₄ (Bertaut and Blum, 1956). The Cr-O distances in the CrO₆ octahedron are 1.96-2.03 Å, with the average distance, 1.997 Å, which is close to 2.015 Å by summation 25of effective ionic radii of Cr^{3+} (0.615 Å for six-fold coordination) and O^{2-} (1.40 Å) (Shannon, 1976). The bond 26angles of O1-Cr1-O3 and O2-Cr1-O3 are 171.3° and 178.7°, respectively, and those of O1-Ti1-O3 and 27O2-Ti1-O3 of CaTi₂O₄ are 171.2° and 166.2°, respectively. Therefore, bond angles (124.1° of Cr1-O2-Cr1 and 28141.6° of Ti1-O2-Ti1) of two corner-shared double chains are largely different. This may depend on the sizes of 29ion radii of A-site cations (Fe²⁺ and Ca²⁺). The effective coordination number (n_c) (Nespolo et al., 2001) of Fe 30 31was calculated to be 4.87, and that of Ca in CaTi₂O₄ 6.86. The Fe-O distances in the FeO₈ bicapped trigonal This is a preprint, the final version is subject to change, of the American Mineralogist (MSA)

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32	prism are 2.01-2.31 Å (FeO ₆ prism) and 2.66 Å (two longest Fe-O bonds), but the Ca-O distances in the CaO ₈
33	bicapped trigonal prism of CaTi ₂ O ₄ are 2.32-2.46 Å (CaO ₆ prism) and 2.74 Å (two longest Ca-O bonds). Two
34	longest Fe-O bonds are too long, compared with summation of effective ion radii Fe ²⁺ (0.92 Å for eight-fold
35	coordination) and O ²⁻ (1.40 Å) (Shannon, 1976). This indicates the coordination number of Fe in tunnel spaces
36	is 6 rather than 8.
37	
38	(c) mLd-type $Fe_2Cr_2O_5$.
39	Figures 5c and 8 show the results of Rietveld analysis using synchrotron X-ray diffraction pattern of
40	mLd-type Fe ₂ Cr ₂ O ₅ , together with Tables 3 and 4. The isotropic atomic displacement parameters of oxygens of
41	mLd-type Fe ₂ Cr ₂ O ₅ were fixed to the same value as done in CT-type FeCr ₂ O ₄ . The site occupancies of Fe and
42	Cr in mLd-type Fe ₂ Cr ₂ O ₅ were estimated from bond distances between cations and oxygens. The crystal
43	structure of mLd-type $Fe_2Cr_2O_5$ in Figure 8 has five non-equivalent cation sites (M1 – M5), four of which are
44	(Fe, Cr)O ₆ octahedra with randomly distributed Fe and Cr, and the other a FeO ₆ prism. These edge-shared
45	octahedra construct a zigzag framework, and the tunnels formed in this framework accommodate Fe. The Fe-O
46	distances of FeO ₆ prism (M5 site) are 2.18-2.22 Å, and the average distance, 2.200 Å, is close to 2.18 Å by
47	summation of effective ionic radii of Fe^{2+} (0.78Å for six-fold coordination) and O^{2-} (1.40 Å) (Shannon, 1976).
48	The Fe ₂ Cr ₂ O ₅ phase is the second compound with the mLd structure reported so far after the Mg ₂ Al ₂ O ₅ phase.

49

3. Implications to FeCr₂O₄-rich high-pressure phases in the shocked Suizhou meteorite and to ultra-high 50pressure chromitites in the Luobusa ophiolite complex. 51

52	The natural observation of CF-type and CT-type FeCr ₂ O ₄ -rich phases near the shock melt veins of the
53	Suizhou meteorite was first reported by Chen et al. (2003b). The CT-type FeCr ₂ O ₄ -rich phase was named xieite
54	(Chen et al., 2008). Our study indicates that at pressure above about 18 GPa CT-type FeCr ₂ O ₄ and CF-type
55	FeCr ₂ O ₄ are stable, but that CF-type FeCr ₂ O ₄ is not quenched but changed into mCF-type which has a very
56	similar powder X-ray diffraction pattern to CF-type. Because of the very similar X-ray diffraction pattern of
57	mCF, it would be possible that Chen et al's CF-phase was not CF but mCF. Alternatively, minor components
58	such as Mg^{2+} , Al^{3+} etc. in the natural FeCr ₂ O ₄ -rich phase might stabilize the CF-type structure. Therefore, it is
59	desirable to directly analyze the structure of the natural CF-type FeCr ₂ O ₄ -rich phase in the meteorite. Figure 2
60	strongly suggests that the natural CF-type and CT-type FeCr ₂ O ₄ rich-phases in the meteorite were formed above
61	about 18 GPa at temperatures lower and higher than 1300°C, respectively. It is interesting that the CT-type
62	FeCr ₂ O ₄ -rich phase was found within and in contact with the shock melt veins in the meteorite, and the CF-type
63	FeCr ₂ O ₄ -rich phase was found between the CT-type FeCr ₂ O ₄ -rich phase contacting with the shock melt veins
64	and FeCr ₂ O ₄ -rich chromite (Chen et al., 2003b, Chen et al., 2008). These observations are consistent with our
65	experimental results in Figure 2 in which the CT-type $FeCr_2O_4$ is stable at higher temperature than the CF-type.
66	The decomposition phases, $Fe_2Cr_2O_5$ mLd + Cr_2O_3 corundum, synthesized at 13-18 GPa in this study were not
67	observed between FeCr ₂ O ₄ -rich chromite and those of the CF-type and CT-type phases in the shocked meteorite.
68	The high-pressure polymorphs including CT-type FeCr ₂ O ₄ rich-phase (xieite), CF-type FeCr ₂ O ₄ rich-phase,
69	majorite garnet, Fe-bearing ringwoodite and NaAlSi ₃ O ₈ -rich lingunite were found in the meteorite (Xie et al.,
70	2011a, b). Considering the stability fields of these high-pressure phases (Kubo and Akaogi, 2000, Ito and
71	Takahashi, 1989, Yagi et al., 1994, Liu, 2006, Akaogi et al., 2010), it is inferred that these phases were

produced at pressure up to about 23 GPa and temperature up to liquidus temperature in various portions in the meteorite during the shock event. Based on Figure 2, we suggest that the CF-type and CT-type $FeCr_2O_4$ rich-phases were formed in the portions of temperature below and above 1300 °C, respectively, at pressure around 18-23 GPa.

Finally, we discuss the ultra-high pressure chromitites in the Luobusa ophiolite. High-pressure 7677minerals including diamond and coesite were found in the chromitites in the Luobusa ophiolite, Tibet (Yang et al., 2007). Yang et al. (2007) suggested that coesite was formed by pseudomorphic replacement of stishovite 78which was derived from the deep upper-mantle (>9 GPa). Furthermore, exsolution lamellaes of coesite and 7980 clinopyroxene were found in chromite in the podiform chromitites of the Luobusa ophiolite (Yamamoto et al., 2009). Yamamoto et al. (2009) suggested that these exsolution lamellaes in chromite were formed by the 81process of the back-transformation from FeCr₂O₄-rich CF phase which could include other components such as 82CaO and SiO₂ during mantle upwelling, and that ultrahigh-pressure chromitites in the Luobusa ophiolite were 83 84originated from deep mantle at pressure above 12.5 GPa. We show that, in high-pressure phase relations of FeCr₂O₄ in Figure 2, the stability field of Fe₂Cr₂O₅ mLd + Cr₂O₃ corundum exists at about 12-18 GPa between 85those of FeCr₂O₄ spinel and CT-type or CF-type phase. However, the decomposed phases (Fe₂Cr₂O₅ + Cr₂O₃) 86 have not been reported in the natural chromitites in the Luobusa ophiolite. Based on our results, we suggest that 87 88 the chromitites have not undergone the mantle condition deeper than 12-16 GPa, because the decomposed phases would have been likely preserved in the chromitites if the chromitites had really been subducted to the 8990 deeper conditions.

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99	
00	Appendix
01	The synchrotron X-ray diffraction patterns of the whole 2θ range for modified CaFe ₂ O ₄ -type FeCr ₂ O ₄ ,
02	$CaTi_2O_4$ -type $FeCr_2O_4$ and modified ludwigite-type $Fe_2Cr_2O_5$ are deposited as Figure 1(a), (b) and (c),
03	respectively, at the MSA website.
04	
05	References cited
06	Akaogi, M., Hamada, Y., Suzuki, T., Kobayashi, M., Okada, M. (1999) High-pressure transitions in the system
07	MgAl ₂ O ₄ -CaAl ₂ O ₄ : a new hexagonal aluminous phase with implication for the lower mantle. Physics of the
08	Earth and Planetary Interiors, 115, 67–77.
09	Akaogi, M., Haraguchi, M., Nakanishi, K., Ajiro, H., Kojitani, H. (2010) High-pressure phase relations in the
10	system CaAl ₄ Si ₂ O ₁₁ -NaAl ₃ Si ₃ O ₁₁ with implication for Na-rich CAS phase in shocked Martian meteorites,
11	Earth and Planetary Science Letters, 289, 503-508.

4/2

- 12 Arai, S. (2010) Possible recycled origin for ultrahigh-pressure chromitites in ophiolite. Journal of Mineralogical
- 13 and Petrological Sciences, 105, 280-285, doi: 10.2465/jmps.100622a.
- 14 Arévalo-López, Á.M., Dos santos-García, A.J., Castillo-Martínez, E., Durán, A., Alario-Franco, M.Á. (2010)
- 15 Spinel to CaFe₂O₄ Transformation: Mechanism and Properties of β -CdCr₂O₄. Inorganic Chemistry, 49,
- 16 2827-2833, doi:10.1021/ic902228h.
- 17 Baerlocher, Ch., McCusker, L.B., Palatinus, L. (2007) Charge flipping combined with histogram matching to
- 18 solve complex crystal structures from powder diffraction data, Zeitschrift für Kristallographie, 222(2), 47-53.
- 19 Belokoneva, E.L., Shcherbakova, Y.K. (2003) Electron density in synthetic escolaite Cr₂O₃ with a corundum
- structure and its relation to antiferromagnetic properties, Russian journal of inorganic chemistry, 48,
- 21 **861-869**.
- 22 Bertaut, E.F. and Blum, P. (1956) Détermination de la Structure Ti₂CaO₄ par la Méthode Self-Consistence
- d'Approche Directe. Acta Crystallographyca, 9, 121-125.
- Brown, I.D., Altermatt, D. (1985) Bond-valence parameters obtained from a systematic analysis of the
- 25 Inorganic Crystal Structure Database. Acta Crystallographica, B41, 244–247.
- 26 Chen, M., Shu, J., Xie, X., Mao, H.K. (2003a) Natural CaTi₂O₄-structure FeCr₂O₄ polymorph in the Suizhou
- meteorite and its significance in mantle mineralogy. Geochimica et Cosmochimica Acta, 67, 3937-3942, doi:
- 28 10.1016/S0016-7037(03)00175-3.
- 29 Chen, M., Shu, J., Mao, H.K., Xie, X., Hemley, R.J. (2003b) Natural occurrence and synthesis of two new
- 30 postspinel polymorphs of chromite. PNAS, 100, 14651-14654, doi/10.1073/pnas.2136599100.
- 31 Chen, M., Shu, J., Mao, H.K. (2008) Xieite, a new mineral of high-pressure FeCr₂O₄ polymorph. Chinese

32 Science Bulletin, 53, 3341-3345, doi: 10.1007/s11434-008-0407-1.

- 33 Decker, B.F., Kasper, J.S. (1957) The structure of calcium ferrite. Acta Crystallographyca, 10, 332-337.
- 34 Dubrovinsky, L.S., Dubrovinskaia, N.A., Prokopenko, V.B., Le Bihan, T. (2002) Equation of state and crystal
- 35 structure of NaAlSiO₄ with calcium-ferrite type structure in the conditions of the lower mantle. High
- 36 Pressure Research, 22, 495-499.
- 37 Dunn, K.J., Bundy, F.P., (1978) Materials and techniques for pressure calibration by resistance-jump transitions
- up to 500 kilobars. Review of Scientific Instruments, 49, 365-370.
- 39 Enomoto, A., Kojitani, H., Akaogi, M., Yusa, H. (2009) High-pressure transitons in MgAl₂O₄ and a new
- 40 high-pressure phase of Mg₂Al₂O₅. Journal of Solid State Chemistry, 182, 389-395,
 41 doi:10.1016/j.jssc.2008.11.015.
- 42 Fei, Y. and Bertka, C.M. (1999) Phase transitions in the Earth's mantle and mantle mineralogy. In: Fei, Y.,
- 43 Bertka, C.M., Mysen, B.O. (Eds.), Mantle Petrology: Field Observations and High Pressure Experimentation:
- 44 The Geochemistry Society, pp. 189–207.
- 45 Fei, Y., Orman, J. Van., Li, J., Western, W. Van., Sanloup, C., Minarik, W., Hirose, K., Komabayashi, T. (2004)
- 46 Experimentally determined post-spinel transformation boundary Mg₂SiO₄ using MgO as an internal pressure
- 47 standard and its geophysical implications. Journal of Geophysical Research, 109, doi:
- 48 10.1029/2003JB002562.
- 49 Funamori, N., Jeanloz, R., Nguyen, J.H. Karner, A., Caldwell, W.A. (1998) High-pressure transforms in
- 50 MgAl₂O₄. Journal of Geophysical Research, 103, 20813-20818, doi: 101029/98JB01575.
- 51 Geisber, H.G., Pennington W.T., Kolis, J.W. (2001) Redetermination of CaMn₂O₄. Acta Crystallographyca, C57,

52 **329-330**.

- 53 Hill P.M., Peiser, H.S., Rait, J.R. (1956) The crystal structure of calcium ferrite and β calcium chromite. Acta
- 54 Crystallographyca, 9, 981-986, doi:10.101107/S0365110X56002862.
- 55 Inaguma, Y., Sakurai, D., Aimi, A., Yoshida, M., Katsumata, Y., Mori, D., Yeon, J., Halasyamani, P.S. (2012)
- 56 Dielectric properties of a polar ZnSnO₃ with LiNbO₃-type structure. Journal of Solid State Chemistry, 195,
- 57 115-119.
- 58 Irifune, T., Naka, H., Sanehira, T., Inoue T., Funakoshi, K., (2002) In situ X-ray observations of phase
- transitions in MgAl₂O₄ spinel to 40 GPa using multianvil apparatus with sintered diamond anvils. Physics
- and Chemistry of Minerals, 29, 645-654.
- 61 Ishii, T., Kojitani, H., Akaogi, M., (2011) Post-spinel transitions in pyrolite and Mg₂SiO₄ and
- 62 akimotoite-perovskite transition in MgSiO₃: Precise comparison by high-pressure high-temperature
- experiments with multi-sample cell technique. Earth and Planetary Science Letters, 309, 185-197, doi:
- 64 10.1016/j.epsl.2011.06.023.
- 65 Ishii, T., Kojitani, H., Akaogi, M., (2012) High-pressure phase transitions and subduction behavior of
- 66 continental crust a pressure-temperature conditions up to the upper part of the lower mantle. Earth and
- 67 Planetary Science Letters, 357-358, 31-41, doi: 10.1016/j.epsl.2012.09.019.
- Ito, E., 2007. Theory and Practice Multianvil cells and high-pressure experimental methods, in: G. Schubert,
- B. Romanowicz, A. Dziewonski (Eds.), Mineral Physics, 2, pp. 197-230. Treatise on Geophysics, Elsevier,
 Amsterdam.
- 71 Ito, E. and Takahashi, E. (1989) Post-spinel transformation in the system Mg₂SiO₄-Fe₂SiO₄ and some

- geophysical implications. Journal of Geophysical Research, 94, 10637-10646.
- Izumi, F., and Momma, K. (2007) Three-dimensional visualization in powder diffraction. Solid State
 Phenomena, 130, 15-20.
- 75 Kojitani, H., Enomoto, A., Tsukamoto, S., Akaogi, M., Miura, H., Yusa, H. (2010) High-pressure
- high-temperature phase relations in MgAl₂O₄. Journal of physics: Conference Series, 215, 012098,
- 77 doi:10.1088/1742-6596/215/1/012098.
- Kosiner, K. and Rea, J.R. (1974) Crystal structure of ferrous phosphate, Fe₃(PO₄)₂, Inorganic Chemistry, 13,
- 79 2876-2880, doi: 10.1021/ic50142a021.
- 80 Kubo, A. and Akaogi, M. (2000) Post-garnet transitions in the system Mg₄Si₄O₁₂-Mg₃Al₂Si₃O₁₂ up to 28 GPa:
- 81 phase relations of garnet, ilmenite and perovskite. Earth and Planetary Interiors, 121, 85-102.
- 82 Kyono, A., Gramsch S.A., Yamanaka, T., Ikuta, D., Ahart, M., Mysen, B.O., Mao, H.K. Hemley, R.J. (2012)
- 83 The influence of the Jahn-Teller effect at Fe2+ on the structure of chromite at high pressure. Physics and
- 84 Chemistry of Minerals, 39, 131-141, doi:10.1007/s00269-011-0468-6.
- 85 Larson, A.C. and Von Dreele, R.B. (2000) General structure analysis system (GSAS). Los Alamos National
- Laboratory Report, LAUR, 86–748.
- 87 Lavina, B., Dera, P., Kim, E., Meng, Y., Downs, R.T., Weck, P.F., Sutton, S.R., Zhao, Y. (2011) Discovery of the
- recoverable high-pressure iron oxide Fe₄O₅. PNAS, 108, 17281-17275, doi/10.1073/pnas.1107573108.
- 89 Le Bail, A., Duroy, H., Fourquet, J. L., (1988) Ab-initio structure determination of LiSbWO₆ by X-ray powder
- 90 diffraction, Materials Research Bulletin, 23, 447.
- 91 Lenaz, D., Skogby, H., Princivalle, F., Halenius, U. (2004) Structural changes and valence states in the MgCr₂

92 O₄-FeCr₂O₄ solid solution series. Physics and Chemistry of Minerals, 31, 633-642, doi:

- 93 10.1007/s00269-004-0420-0.
- 94 Liu, X. (2006) Phase relations in the system KAlSi₃O₈-NaAlSi₃O₈ at high pressure high temperature
- 95 conditions and their implications for the petrogenesis of lingunite. Earth and Planetary Science Letters, 246,

96 317-325.

- 97 Louër, D. and Boultif, A. (2007) Powder pattern indexing and the dichotomy algorithm. Zeitschrift für
- 98 Kristallographie Supplements, 26, 191-196, doi: 10.1524/zksu.2007.2007.suppl_26.191.
- 99 McCammon, C. (1993) Composition limits of Fe_xO and the earth's lower mantle. Science, 261, 923-925.
- 00 Momma, K. and Izumi, F. (2008) VESTA: a three-dimensional visualization system for electronic and structural
- analysis. Journal of Applied Crystallography, 41, 653-658.
- 02 Morishima, H., Kato, T., Suto, M., Ohtani, E., Urakawa, U., Shimomura, O., Kikegawa, T. (1994) The phase
- boundary between alpha-Mg₂SiO₄ and beta-Mg₂SiO₄ determined by in-situ x-Ray observation. Science, 265,

04 1202-1203.

- 05 Nespolo, M., Ferraris, G., Hoppe, R. (2001) Charge distribution analysis of ceramic materials. Journal of
- 06 Ceramic Processing Research, 2, 38-44.
- 07 Ohtani, E., Hirao, N., Kondo, T., Ito, M., Kikegawa, T. (2005) Iron-water reaction at high-pressure and
- temperature, and hydrogen transport into the core. Physics and Chemistry of Minerals, 32, 77-82, doi:
 10.1007/s00269-004-0443-6.
- 10 Ono, S., Kikegawa, T., Ohishi, Y. (2006) The stability and compressibility of MgAl₂O₄ high-pressure
- polymorphs. Physics and Chemistry of Minerals, 33, 200-206, doi: 10.1007/s00269-006-0068-z.

- 12 Palatinus, L. and Chapuis, G. (2007) Superflip a computer program for the solution of crystal structures by
- 13 charge flipping in arbitrary dimensions. Journal of Applied Crystallography, 40, 786-790.
- 14 Palatinus, L., Prathapa, S. J., van Smaalen, S. (2012): EDMA: a computer program for topological analysis of
- discrete electron densities. Journal of Applied Crystallography, 45, 575-580.
- 16 Rabin, H. and Tang, C.L. (1975) Physics; Nonlinear Optical Susceptibilities; Nonlinear Optical Process;
- 17 Parametric; Raman; Spectroscopy. Quantum Electronics, 1.
- 18 Ringwood, A.E. (1975) Composition and Petrology of the Earth's Mantle. McGraw-Hill, New York, NY,
- 19 618pp.
- 20 Rogge, M.P., Caldwell, J.H., Ingram, D.R., Green, C.E., Geselbracht, M.J., and Siegrist, T. (1998) A new
- synthetic route to peseudo-brookite-type CaTi₂O₄. Journal of Solid State Chemistry, 141, 338–342.
- 22 Sanehira, T., Irifune, T., Shinmei, T., Ohfuji, H., Brunet, F., Funakoshi, K.-I. (2008) Density profiles of pyrolite
- and MORB compositions across the 660-km seismic discontinuity. High Pressure Research, 28, 335-349,
- doi:10.1080/08957950802251357.
- 25 Shannon, R.D. (1976) Revised effective ionic radii and systematic studies of interatomic distances in halides
- and chalcogenides. Acta Crystallographica, A32, 751-767.
- 27 Speziale, S., Zha, C.S., Duffy, T.S., Hemley, R.J., Mao, H.K. (2001) Quasi-hydrostatic compression of
- magnesium oxide to 52 GPa: implications for the pressure-volume-temperature equation of state. Journal of
- 29 Geophysical Research, 106, 515-528.
- 30 Suzuki, A., Ohtani, E., Morishima, H., Kubo, T., Kanbe, Y., Kondo, T. (2000). In situ determination of the phase
- boundary between wadsleyite and ringwoodite in Mg₂SiO₄, Geophysical Research Letters, 27, 803-806.

32 Toby, B.H. (2001) EXPGUI, a graphical user interface for GSAS. Journal of Applied Crystallography, 34,

- 33 **210–213**.
- 34 Xie, X., Chen, M., Wang, C.Y. (2011a) Occurrence and mineral chemistry of chromite and xieite in the Suizhou
- ³⁵ L6 chondrite. Science China Earth Science, 54, 998-1010, doi: 10.1007/s11430-011-4199-9.
- 36 Xie, X., Sun, Z., Chen, M. (2011b) The distinct morphological and petrological features of shock melt veins in
- 37 the Suizhou L6 chondrite. Meteoritics and Planetary Science, 46, 459-469, doi:
 38 10.1111/j.1945-5100.2100.01168.x.
- 39 Yagi, A., Suzuki, T., Akaogi, M. (1994) High pressure transitions in the system KAlSi₃O₈-NaAlSi₃O₈. Physics
- 40 and Chemistry of Minerals, 21, 12-17.
- 41 Yamamoto, S., Kojima, T., Hirose, K., Maruyama, S. (2009) Coesite and clinopyroxene exsolution lamella in
- 42 chromites: *In-situ* ultrahigh-pressure evidence from podiform chromitites in the Luobusa ophiolite, southern
- 43 Tibet. Lithos, 109, 314-322.
- 44 Yamanaka, T., Uchida, A., Nakamoto, Y. (2009) Structural transition of post-spinel phases CaMn₂O₄, CaFe₂O₄
- and CaTi₂O₄ under high-pressures up to 80 GPa. American Mineralogist, 93, 1874-1881.
- 46 Yang, J.S., Dobrzhinetskaya, L., Bai, W.J., Fang, Q.S., Robinson, P.T., Zhang, J., Green, H.W. (2007)
- 47 Diamond– and coesite-bearing chromitites from the Luobusa ophiolite, Tibet. Geology, 35, 875-878, doi:
- 48 10.1130/G23766A
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50**Figure captions**

51

52	Figure 1. Synchrotron X-ray diffraction pattern of FeCr ₂ O ₄ phase synthesized at 25 GP and 1000°C. The
53	pattern was taken at ambeint conditions. Miller indices are based on the unit cell of modified CaFe2O4-type
54	FeCr ₂ O ₄ in Table 2. Es represents diffraction peaks of Cr ₂ O ₃ escolaite (upper figure). The blue pattern
55	represents simulated pattern of CaFe ₂ O ₄ -type FeCr ₂ O ₄ (lower figure). Note that intensities of 202, 104, 211,401
56	and 113 peaks in the observed XRD pattern are different from the coresponding peaks in the simulated pattern.
57	
58	Figure 2. Phase relations of FeCr ₂ O ₄ at high pressure and high temperature. Solid circle, Sp; half closed circle,
59	Sp + mLd + Es; solid square, mLd + Es; half closed square, mLd + Es + CT; solid triangle, CF; solid inverse
60	triangle, CT; half closed inverse triangle, CF + CT. Solid lines represent phase boundaries. Sp, spinel-type
61	FeCr ₂ O ₄ ; mLd, modified ludwigite-type Fe ₂ Cr ₂ O ₅ ; Es, Cr ₂ O ₃ eskolaite; CF, CaFe ₂ O ₄ -type FeCr ₂ O ₄ ; CT,
62	CaTi ₂ O ₄ -type FeCr ₂ O ₄ .
63	
64	Figure 3. (a) <i>In-situ</i> X-ray diffraction profiles of FeCr ₂ O ₄ as a function of pressure and temperature. Small red
65	circles indicate diffraction peaks whose intensities are different from those of CaFe ₂ O ₄ -type FeCr ₂ O ₄ . c-Sp:
66	cubic spinel, t-Sp: tetragonal spinel. (b) <i>In-situ</i> X-ray diffraction pattern of CaFe ₂ O ₄ -type FeCr ₂ O ₄ with small

- amounts of γ -FeH_x and FeO (upper figure). Simulated pattern of CaFe₂O₄-type FeCr₂O₄ (lower figure). Both 67
- patterns are at 16.5 GPa and 27 °C. (c) Compression curve of CaFe₂O₄-type FeCr₂O₄ up to 25 GPa at 27 °C. 68
- 69 The solid line represents the Birch-Murnaghan fitting curve.

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Figure 4. Electron diffraction patterns of (a) <i>hk</i> 0 plane, (b) <i>h</i> 0 <i>l</i> plane, and (c) 0 <i>kl</i> plane of the recovered sample
(Run no. 40). The electron diffraction patterns were taken from different grains of the same run product.
Figure 5. Synchrotron X-ray diffraction patterns of (a) modified CaFe ₂ O ₄ -type FeCr ₂ O ₄ , (b) CaTi ₂ O ₄ -type
$FeCr_2O_4$ and (c) modified ludwigite-type $Fe_2Cr_2O_5$ at ambient conditions. The diffraction patterns in (a), (b)
and (c) show those in 2 θ ranges below 35, 45 and 40°, respectively. The diffraction patterns in the whole 2 θ
range up to 75° (deposited in Appendix) were used for Rietveld refinements. Data points and solid lines show
the observed and the calculated profiles, respectively, and the difference between them is shown at the bottom.
Bragg peak positions are marked by small ticks. These lines are for main phases, Cr ₂ O ₃ eskolaite and Re from
top to bottom. The refined crystal structure is shown with each profile.
Figure 6. (a) and (b) Crystal structure of modified CaFe ₂ O ₄ -type FeCr ₂ O ₄ . (c) Coordination environments of Fe
and Cr. (d) Fe-O distances in the coordination environment of Fe.
Figure 7. (a) and (b) Crystal structure of CaTi ₂ O ₄ -type FeCr ₂ O ₄ . (c) Coordination environments of Fe and Cr.
(d) Fe-O distances in the coordination environment of Fe.
Figure 8. (a) and (b) Crystal structure of modified ludwigite-type Fe ₂ Cr ₂ O ₅ . (c) Coordination environments of
]

90 Table 1

91 Results of high-pressure high-temperature quench experiments

92	Run no.	Pressure	Temperature	Duration	[†] Phases
93		(GPa)	(°C)	(min)	
94	36	15.9	800	180	Sp
95	35	17.6	800	180	mCF
96	34	19.8	800	180	mCF
97	32	25.0	800	180	mCF
98	7	14.7	1000	120	Sp
99	12	15.9	1000	120	mLd+Es
00	14	16.5	1000	120	mLd+Es
01	22	17.6	1000	120	mCF
02	25	19.8	1000	120	mCF
03	40	25.0	1000	120	mCF
04	37	28.0	1000	120	mCF
05	39	15.9	1100	120	mLd+Es
06	38	28.0	1100	120	mCF
07	19	13.5	1200	120	Sp
08	18	14.7	1200	120	mLd+Es
09	10	17.6	1200	120	mLd+Es

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10	24	18.6	1200	120	mCF			
11	17	22.8	1200	120	mCF			
12	23	25.0	1200	120	mCF			
13	26	28.0	1200	120	mCF			
14	27	22.8	1300	60	mCF			
15	30	27.1	1300	60	mCF			
16	31	28.0	1300	60	mCF+CT			
17	13	13.0	1400	60	Sp			
18	5	13.9	1400	60	Sp+mLd+Es			
19	15	14.7	1400	60	mLd+Es			
20	16	17.6	1400	60	mLd+Es			
21	20	18.6	1400	60	mLd+Es+CT			
22	28	19.8	1400	60	CT+mLd(tr)+Es(tr)			
23	33	28.0	1400	60	СТ			
24	1	19.8	1500	60	СТ			
25	11	11.7	1600	60	Sp			
26	8	13.0	1600	60	mLd+Es			
27	4	14.7	1600	60	mLd+Es			
28	9	15.9	1600	60	mLd+Es			
29	6	16.6	1600	60	СТ			

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0	2	23.0	1600	60	СТ	
1	[†] Phases	in the recovered s	amples. A small a	mount of metallic	iron in the run products derived	d from the starting
2	material,	, the mixture of Fe	eCr_2O_4 spinel + 5 v	vt% Fe, is not list	ed.	
3	Abbrevia	ations: Sp, spinel	-type FeCr ₂ O ₄ ; mL	d, modified ludw	vigite-type Fe ₂ Cr ₂ O ₅ ; Es, Cr ₂ O ₃	corundum; mCF,
4	modified	l CaFe ₂ O ₄ -type Fe	eCr ₂ O ₄ ; CT, CaTi ₂ O	O ₄ -type FeCr ₂ O ₄ ;	tr, trace.	
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50 Table 2

51 Lattice parameters of modified CaFe₂O₄(mCF)-type FeCr₂O₄, CaTi₂O₄(CT)-type FeCr₂O₄ and modified

52 ludwigite(mLd)-type Fe₂Cr₂O₅.

53	Phase	mCF-type FeCr ₂ O ₄	CT-type FeCr ₂ O ₄	mLd-type Fe ₂ Cr ₂ O ₅
54	Space group	<i>Pnma</i> (no. 62)	<i>Cmcm</i> (no. 63)	<i>Pbam</i> (no. 55)
55	<i>a</i> (Å)	9.0633(3)	2.8845(1)	9.6642(2)
56	<i>b</i> (Å)	2.9579(1)	9.5207(2)	12.5000(3)
57	<i>c</i> (Å)	10.6391(3)	9.7532(2)	2.9023(1)
58	$V(\text{\AA}^3)$	285.22(1)	267.84(1)	350.59(1)
59	Ζ	4	4	4
60	V_m (cm ³ /mol)	42.941(2)	40.323(2)	52.781(2)
61	$D (g/cm^3)$	5.213(1)	5.553(1)	5.604(1)

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Table 3

71 Structure parameters of modified CaFe₂O₄(mCF)-type FeCr₂O₄, CaTi₂O₄(CT)-type FeCr₂O₄ and modified

72 ludwigite(mLd)-type Fe₂Cr₂O₅.

73	Atom	Wyckoff site	<i>g</i> (Fe)	<i>g</i> (Cr)	x	у	Ζ	B_{iso} (Å ²)
74	mCF-type F	eCr ₂ O ₄						
75	Fe	4c	1.0	0.0	0.2560(2)	0.25	0.3300(2)	1.72(4)
76	Cr1	4c	0.0	1.0	0.0511(2)	0.25	0.6164(2)	0.50(3)
77	Cr2	4c	0.0	1.0	0.0904(2)	0.25	0.0974(2)	0.42(3)
78	01	4c	-	-	0.3083(6)	0.25	0.1526(6)	0.84(6)
79	O2	4c	-	-	0.3895(6)	0.25	0.4759(6)	0.84(6)
80	O3	4c	-	-	0.4481(6)	0.25	0.7160(5)	0.84(6)
81	04	4c	-	-	0.0803(6)	0.25	0.4323(6)	0.84(6)
82								
83	CT-type Fe	Cr_2O_4						
84	Fe	4c	1.0	0.0	0	0.1128(2)	0.25	1.76(4)
85	Cr	8f	0.0	1.0	0	0.3686(2)	0.0714(1)	0.25(2)
86	01	4b	-	-	0	0	0	0.48(10)
87	O2	4c	-	-	0	0.4656(6)	0.25	0.65(11)
88	03	8f	-	-	0	0.2640(4)	0.6109(3)	0.60(9)

4/2 90 mLd-type Fe₂Cr₂O₅ 91M1 2a 0.3 0.7 0 0 0 0.51(4)0 0.5 M2 2d 0.8 0.2 0.5 920.43(4)0.35 93 M3 4g 0.65 0.0221(2)0.2824(1)0 0.72(3)0.1 0.9 0.3827(1)94 M4 4h 0.2730(1) 0.5 0.13(3) M5 1.0 0.0 0.2432(2)0.1306(2) 0 954g 1.64(4)96 01 4h 0.1423(6) 0.0237(5)0.5 0.68(4)O2 0.4115(5) 0.3510(5) 0 0.68(4)974g 98 O3 4h 0.4103(5) 0.1424(5)0.5 0.68(4)_ O4 0.4296(5) 0 0.68(4)99 4g 0.1373(6) 05 4h 0.1598(6) 0.5 00 0.2488(4) 0.68(4)_ _

01 The reliability factors and goodness-of-fit indicator for the mCF.

 $R_{wp} = 5.910\%, S = 0.9867$

- 03 mCF-type FeCr₂O₄: $R_{\rm B} = 2.873\%$, $R_{\rm F} = 1.512\%$
- 04 Corundum-type Cr_2O_3 : $R_B = 4.245\%$, $R_F = 2.447\%$
- 05 The reliability factors and goodness-of-fit indicator for the CT.
- 06 $R_{wp} = 6.380\%, S = 1.0663$
- 07 CT-type FeCr₂O₄: $R_{\rm B} = 1.969\%$, $R_{\rm F} = 2.036\%$
- 08 Corundum-type Cr_2O_3 : $R_B = 3.006\%$, $R_F = 2.773\%$
- 09 The reliability factors and goodness-of-fit indicator for the mLd.

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10 $R_{wp} = 3.905\%, S = 0.6443$

- 11 mLd-type Fe₂Cr₂O₅: $R_{\rm B} = 1.247\%$, $R_{\rm F} = 0.935\%$
- 12 Corundum-type Cr_2O_3 : $R_B = 2.420\%$, $R_F = 2.040\%$
- 13 Re: $R_{\rm B} = 0.826\%$, $R_{\rm F} = 0.532\%$

$$14 \qquad R_{wp} = \left\{ \frac{\sum_{l} w_{l} [y_{l} - f_{l}(x)]^{2}}{\sum_{l} w_{l} y_{l}^{2}} \right\}^{1/2}, \ R_{B} = \frac{\sum_{K} |l_{0}(h_{K}) - l(h_{K})|}{\sum_{K} l_{0}(h_{K})}, \ R_{F} = \frac{\sum_{K} ||F_{0}(h_{K})| - |F(h_{K})||}{\sum_{K} |F_{0}(h_{K})|}, \ S = \left\{ \frac{\sum_{l} w_{l} [y_{l} - f_{l}(x)]^{2}}{N - F} \right\}^{1/2}$$

15 where y_i , w_i and $f_i(x)$ are the observed intensity at step *i*, the statistical weight and calculated intensity, $I_0(\mathbf{h}_K)$,

16 $I(\mathbf{h}_K)$, $F_0(\mathbf{h}_K)$ and $F(\mathbf{h}_K)$ are the observed and calculated intensity and structure factor for reflection K, N and P

- are number of all data points and refined parameter, respectively.
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30 Table 4

31 Interatomic distances and angles in the structures of modified CaFe₂O₄(mCF)-type FeCr₂O₄, CaTi₂O₄(CT)-type

32 FeCr₂O₄ and modified ludwigite(mLd)-type Fe₂Cr₂O₅.

33	mCF-type FeCr ₂ O ₄		CT-type FeCr ₂ O ₄		mLd-type Fe ₂ Cr ₂ O ₅		
34	Bond length (Å)		Bond length (Å)		Bond length (Å)		
35	Fe-O1	1.945(6)	Fe–O1 × 2	2.6644(7)	$M1-O1 \times 4$	2.021(4)	
36	Fe-O2	1.968(6)	Fe–O2 × 2	2.011(5)	$M1-O2 \times 2$	2.049(6)	
37	$Fe-O3^i \times 2$	2.661(5)	Fe–O3 × 4	2.301(3)	Average	2.030	
38	Fe-O4	1.928(5)	Average	2.319	n _c	5.99	
39	Average	2.233	n _c	4.87	BVS	2.64	
40	n _c	3.01	BVS	1.97	M2-O3 × 2	1.980(6)	
41	BVS	1.85	$Cr1-O1^i \times 2$	2.0323(9)	$M2-O4 \times 4$	2.155(4)	
42	$Cr1-O1^i \times 2$	1.991(4)	Cr1–O2	1.972(3)	Average	2.097	
43	Cr1–O3 ⁱⁱ	2.013(6)	$Cr1-O3 \times 2$	1.955(3)	n _c	5.55	
44	$Cr1-O4^{iii} \times 2$	1.968(4)	Cr1–O3	2.038(3)	BVS	2.30	
45	Cr1–O4	1.977(6)	Average	1.997	M3-O2	1.980(6)	
46	Average	1.985	<i>n</i> _c	5.92	M3-O3 × 2	2.039(4)	
47	n _c	5.99	BVS	2.88	M3-O4	2.151(6)	
48	BVS	2.97	Bond angles (°)		M3-O5 × 2	2.013(4)	
49	Cr2-01	2.060(6)	Cr1 ^{iv} -O1-Cr1 ^{iv}	90.42(5)	Average	2.039	

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50	Cr2–O2 ⁱⁱ	1.981(5)	Cr1 ^v -O1 ^{vi} -Cr ^v	89.58(5)	<i>n</i> _c	5.85
51	$Cr2-O2^i \times 2$	1.973(4)	Cr1 ^{vii} –O3 ^{vi} –Cr1 ^{vii}	95.1(2)	BVS	2.61
52	$Cr2-O3^{i} \times 2$	1.975(4)	Cr1 ^{vii} –O3i ^{vii} –Cr1 ^{viii}	98.2(1)	M4-01	1.943(6)
53	Average	1.990	Cr1 ^{vii} -O2 ^{vi} -Cr1 ^{viii}	124.1(3)	M4-O2 × 2	2.014(4)
54	n _c	5.95	$O3^{v}$ - $Cr1^{vi}$ - $O1^{vii}$	171.3(1)	M4-O4 × 2	2.042(4)
55	BVS	2.94	O2 ^{vi} -Cr1 ^{vi} -O3 ^{viii}	178.7(2)	M4-O5	2.000(6)
56	Bond angles (°)				Average	2.009
57	Cr1 ⁱ -O1-Cr1 ⁱ	96.0(3)			n _c	5.93
58	Cr1 ⁱⁱⁱ –O4–Cr1 ⁱⁱⁱ	97.4(3)			BVS	2.80
59	Cr1 ⁱⁱⁱ –O4–Cr1	100.3(2)			M5-O1 × 2	2.201(5)
60	Cr2 ⁱ -O2-Cr2 ⁱ	97.1(2)			M5-O3 × 2	2.176(4)
61	Cr2 ⁱ -O2-Cr2 ⁱⁱ	100.0(2)			M5-O5 × 2	2.223(4)
62	Cr2 ⁱ -O3-Cr2 ⁱ	97.0(2)			Average	2.200
63	Cr1 ⁱ -O1-Cr2	124.0 (2)			n _c	5.98
64	Cr2 ⁱ -O3-Cr1 ⁱⁱ	130.4(1)			BVS	1.71
65	O4-Cr1-O3 ⁱⁱ	160.1(3)			Bond angles (°)	
66	O4 ⁱⁱⁱ –Cr1–O1 ⁱ	175.5(3)			01-M1-01 ⁱⁱⁱ	89.1(3)
67	O2 ⁱⁱ -Cr2-O1	95.2(2)			O1-M1-O2 ^{ix}	98.8(2)
68	O2 ⁱⁱ -Cr2-O3 ⁱ	173.4(3)			O4-M2-O4 ⁱⁱⁱ	173.4(3)
69					O3 ^{ix} -M2-O4 ⁱⁱⁱ	84.2(2)

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70	O5 ^{vi} –M3–O3 ^{ix}	163.7(3)	
71	O2 ^{ix} -M3-O4	178.3(3)	
72	O2-M4-O4 ^{vi}	174.8(3)	
73	O1 ^x -M4-O5	171.0(3)	
74	M3 ^{ix} -O2 ^{vi} -M1 ^x	122.2(3)	
75	M1-O1-M1 ^{vi}	90.9(3)	
76	M4 ^{vi} -O2-M1	95.9(3)	
77	M4 ^x -O1-M1 ^{vi}	98.2(3)	
78	M2-O4-M2 ^{vi}	84.9(2)	
79	M3-O4-M2 ^{vi}	92.0(2)	
80	M2-O3-M3 ^{ix}	100.4(2)	
81	M4-O4-M2	91.98(4)	
82	M3 ^{ix} -O3-M3 ^{ix}	90.7(3)	
83	M3-O5-M3 ^{iv}	92.1(3)	
84	M4 ^{vi} –O4–M3	95.2(2)	
85	M4-O5-M3	100.7(2)	
86	M3 ^{ix} -O2-M4	121.5(2)	
87	M4-O2-M4 ^{vi}	92.8(3)	
88	M4-O4-M4 ^{vi}	90.3(3)	

89 Symmetry codes: (i) 1/2-x, -y, 1/2+z. (ii) 1/2+x, y, 1/2-z. (iii) -x, -y, -z. (iv) 1/2-x, 1/2-y, -z. (v) 1/2-x,

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90 1/2-y, 1/2+z. (vi) x, y, z. (vii) 1/2+x, 1/2+y, z. (viii) x, y, 1/2-z. (ix) 1/2+x, 1/2-y, z. (x) 1/2-x, 1/2+y, -z.

- 91 n_c : effective coordination number
- 92 BVS: bond valence sum value

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Figure 1.







Figure 3.



Figure 4.



Figure 5.



Figure 6.



Figure 7.



Figure 8.