

24 contribution of the less intense band at 3581 cm^{-1} is likely related to the partial protonation of a silicon vacancy (protonation of the O3-type oxygen) associated with the migration of the silicon cation to the Si2 site. This assignment is consistent with several experimental constraints: wavenumber and pleochroism of infrared OH-stretching bands, pressure-dependence of the bands wavenumber, evidence from X-ray diffraction of magnesium vacancies in M3 site, and increase of the *b*/*a* axial ratio with water content. The integrated absorption coefficients of the corresponding OH-defects are also calculated and thus complement the set of data obtained previously for forsterite and ringwoodite. Absorption coefficients of wadsleyite computed at 0 and 15 GPa indicate that for a precise quantification of the hydrogen content in *in situ* experiments, one must consider higher absorption coefficients than those determined at 0 GPa after quench. It is 35 also shown that a single theoretical relation can account for the three Mg_2SiO_4 36 polymorphs at 0 GPa: $K_{int} = 278.7 \pm 18.1$ (3810 ± 465 - x), where K_{int} is the integrated molar absorption coefficient of the OH stretching modes and x is the average 38 wavenumber in cm⁻¹. Absorption coefficients are significantly lower than the general calibrations, the use of which would lead to an underestimation of the water concentrations.

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- **Keywords:** wadsleyite, OH-defect, IR spectroscopy, *ab initio* calculations, DFT
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1. INTRODUCTION

 The interest for OH-defects in nominally anhydrous minerals (NAMs) is essentially two-fold. First, given the abundance of the NAMs in the mantle, these OH- defects may contribute significantly to the water budget of the Earth. Second, the presence of OH-defects within the structure of NAMs modifies the bulk properties of minerals. These hydroxyl groups may affect the global mantle dynamics because hydrogen incorporation is likely to change melting temperatures, atomic diffusivities, and rheological properties**.** In addition, it is essential to know the effect of OH-defects on electrical conductivity and elasticity of minerals for interpreting accurately the geophysical properties that are used to probe the Earth interior (Keppler and Smyth, 2006).

 Wadsleyite corresponds to a high-pressure polymorph of forsterite. It is considered to be the dominant mineral in the upper part of the transition zone between 410 and 520 km depth, before its transformation into ringwoodite at higher pressure. The 59 structure of hydrogen-free wadsleyite $(\beta \text{-Mg}_2 \text{SiO}_4)$ can be refined in the orthorhombic *Imma* space group, with the following cell parameters: $a = 5.698 \text{ Å}, b = 11.438 \text{ Å}, \text{and } c$ $61 = 8.257$ Å (Horiuchi and Sawamoto, 1981). Silicon atoms in tetrahedral sites (Wyckoff 62 position $8h$) form $Si₂O₇$ groups (sorosilicate). There are three distinct octahedral sites occupied by magnesium atoms, M1 (*4a*), M2 (*4e*), and M3 (*8g*), with twice as many M3 sites as M1 or M2 sites in each unit cell. M1 and M2 sites share edges and form a single octahedral chain, which runs parallel to the *b*-axis. M3 octahedra form a double chain running along the *a*-axis. Finally, there are four distinct oxygen sites in the structure.

67 Among them, O2 is the bridging oxygen of the $Si₂O₇$ group, and O1 is a non-silicate oxygen atom (not bonded to silicon atoms). This structural feature makes the O1 atoms strongly underbonded with a valence of 1.82 valence units (v.u.) using the method of Donnay and Allmann (1970), and for this reason O1 atoms are expected to be very favourable sites for protonation (Smyth, 1987). The ability of wadsleyite to incorporate significant amount of hydrogen has been confirmed experimentally (e.g. McMillan et al. 1991, Young et al. 1993, Inoue et al. 1995, Kohlstedt et al. 1996, Kohn et al. 2002, Demouchy et al. 2005, Ye et al. 2010). A review of the water content in wadsleyite quantified by secondary ion mass spectrometry can be found in Demouchy et al. (2005), and highlights the variability of water contents obtained experimentally. Demouchy et al. (2005) have shown that the water solubility in Fe-free wadsleyite decreases with

78 increasing temperature, from ~2.2 wt % H_2O at 900 °C down to ~0.9 wt % H_2O at 1400 °C

(15 GPa). On the other hand, no significant effect of pressure was observed.

 The knowledge of the location of hydrogen within the crystal structure is a prerequisite to understand the mechanisms controlling the changes of mineral bulk properties. Because of the low hydrogen concentrations, it is not possible to determine experimentally the protons location using for instance diffraction techniques. In this field, modeling techniques proved to be very useful. Several theoretical studies have investigated the mechanisms of hydrogen incorporation via the formation of OH-defects within the wadsleyite structure (Winkler et al. 1995, Wright and Catlow 1996, Haiber et al. 1997, Ross et al. 2003, Parker et al. 2004, Walker et al. 2006, Tsuchiya and Tsuchiya 2009, Verma and Karki 2009). Based on total energy calculations and on vibrational frequencies, the proposed models imply the protonation of magnesium and silicon

 vacancies. In the present work, we propose to test different models of OH-defects with the calculation of the relative intensities and the polarization of the infrared (IR) absorption bands of OH-stretching modes at 0 and 15 GPa. This approach will enable a direct comparison with experiment, and so the interpretation of the major IR absorption bands of wadsleyite, which can be considered as the most discriminating signatures of OH-defects in this mineral.

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2. THEORETICAL METHODS

99 Calculations are done on a $2 \times 1 \times 1$ supercell of wadsleyite (i.e. 112 atoms) containing up to four hydrogen atoms. This corresponds to a maximum hydrogen 101 concentration of 1.65 wt % H₂O. With \sim 10 Å between the neutral OH-defect and its periodic images in the three directions, the OH-defect can be considered as sufficiently isolated to account for experimental samples displaying lower hydrogen concentrations. Calculations are performed within the density functional theory framework, using the generalized gradient approximation (GGA) to the exchange-correlation functional with the PBE parametrization (Perdew et al. 1996) and periodic boundary conditions. The ionic cores are described by the ultrasoft pseudopotentials Si.pbe-rrkjus.UPF, O.pbe- rrkjus.UPF, and H.pbe-rrkjus.UPF from http://www.quantum-espresso.org and the magnesium norm-conserving pseudopotential previously used to model forsterite, ringwoodite and enstatite (Blanchard et al. 2009, Balan et al. 2011, 2012). The electronic wave-functions and the charge density are expanded in plane-waves with 60 and 600 Ry cutoffs, respectively. For the electronic integration, the Brillouin zone is sampled using a

 $2 \times 2 \times 2$ k-point grid according to the Monkhorst-Pack scheme (Monkhorst and Pack 114 1976). It is checked that increasing the wave-functions cutoff to 100 Ry and using a 4×4 \times 4 k-point grid does not change the total energy by more than 1.1 meV/atom and does not modify significantly the residual forces on atoms. Atomic relaxations are performed with the PWscf code (Giannozzi et al., 2009; http://www.quantum-espresso.org) until the 118 residual forces on atoms are less than 10^{-4} Ry/a.u.. The optimized cell parameters are 119 only slightly smaller than the experimental ones, i.e. $a = 5.691 \text{ Å}$, $b = 11.428 \text{ Å}$, and $c =$ 120 8.240 Å (this work) versus $a = 5.698$ Å, $b = 11.438$ Å, and $c = 8.257$ Å (Horiuchi and Sawamoto, 1981). Defective structures are fully relaxed at constant pressure (at 0 and 15 GPa), by optimizing the atomic positions as well as the cell parameters. This approach allows assessing the evolution of the cell parameters associated with the protonation mechanisms.

 The dynamical matrix and dielectric quantities, such as the Born effective charges 126 and the electronic dielectric tensor, are calculated at the Brillouin zone center $(\Gamma \text{ point})$, within the harmonic approximation using the linear response theory (Baroni et al. 2001) as implemented in the PHonon code (Giannozzi et al., 2009; http://www.quantum- espresso.org). The high-frequency OH-stretching modes can be accurately calculated by only considering the displacement of the oxygen and hydrogen atoms involved in OH groups (Balan et al. 2005). We note that the anharmonicity of OH modes is not explicitly considered here. However, a reasonable agreement with experimental frequencies (i.e., 133 within \sim 1 rel.%) can be expected due to the cancelation of the error related to the neglect of anharmonicity with that related to the use of the GGA approximation (Balan et al. 2007; Finocchi et al. 2008). The complex low-frequency dielectric permittivity tensor is

3. EXPERIMENTAL CONSTRAINTS ON OH-DEFECTS IN WADSLEYITE

 Diffraction techniques and vibrational spectroscopies are not able to determine unambiguously the atomic configuration of OH-defects, but they provide crucial constraints on the nature of OH-defects. For wadsleyite, the proposed atomic models will have to take into account or reproduce the four following experimental constraints:

 (i) IR spectroscopy is the most widely used technique to investigate the concentration and the nature of OH-defects in minerals. Several IR spectra, most of the time unpolarized, of protonated wadsleyite are reported in the literature (e.g. Young et al. 1993, Kohlstedt et al. 1996, Bolfan-Casanova et al. 2000, Kohn et al. 2002, Jacobsen et al. 2005, Deon et al. 2010). We will refer here to the polarized IR spectra measured by Jacobsen et al. (2005) on wadsleyite samples containing between ~100 and ~10000 wt ppm H2O (Fig. 1). In the absorption range of OH-stretching modes, we can distinguish two main groups of bands; the more intense group being the one at lower wavenumber. Each group is composed by at least two absorption bands. The first group displays two bands at and 3360 cm⁻¹ (here referred to as bands A and B, respectively) that both show a minimal pleochroism even if the maximum intensities occur for a polarization along the *c*-axis. The second

159 group has a band at 3581 cm⁻¹ (referred to as band C) that is polarized in the *bc* plane and 160 another band at 3614 cm⁻¹ of smaller intensity (referred to as band D), polarized in the *ac* 161 plane. It is important to note that at very low concentration (\sim 150 wt ppm H₂O) the OH 162 spectrum is relatively simple. Only one absorption band is present in each region, i.e. the band D and the band A, which is in fact a doublet with a contribution at 3317 cm⁻¹ 163 164 polarized along the *b*-axis and a contribution at 3324 cm⁻¹ polarized in the *ac* plane. In 165 addition to these main bands, we can note a broad absorption band centered around 3000 166 cm⁻¹ with a maximum intensity for the polarization along the *a*-axis, as well as small 167 features between the two main groups of bands and shoulders to these main bands. Some 168 studies have also reported extra bands at high wavenumbers, \sim 3660 cm⁻¹, but their 169 intensities are relatively weak (Bolfan-Casanova et al. 2000, Kohn et al. 2002, Deon et al. 170 2010). The observation of IR spectra thus suggests the presence of several types of OH-171 defects with a more diverse population when hydrogen content increases.

172 (ii) Deon et al. (2010) measured the IR spectrum of a synthetic wadsleyite containing 173 \sim 8000 wt ppm H₂O, from ambient conditions up to 15.4 GPa. Essentially, the low-174 wavenumber group of bands shifts strongly toward lower wavenumbers $(\sim -8 \text{ cm}^{-1}/\text{GPa})$, 175 whereas the position of the high-wavenumber group of bands remains about constant $(\sim$ 176 +0.3 cm⁻¹/GPa). In their Raman spectroscopic study, Kleppe et al. (2001) found very 177 similar pressure dependences for a sample containing 1.65 wt $\%$ H₂O.

 (iii) Structure refinements from single-crystal X-ray diffraction measurements show that most of the cation vacancies are located on the M3 sites (Kudoh and Inoue 1999, Deon et al. 2010, Ye et al. 2010), and that this amount of M3 vacancies increases with increasing water content (Jacobsen et al. 2005). This is consistent with theoretical results, which

 We focus here on both M2 and M3 vacancies. In total, about twenty starting defect configurations are investigated. In agreement with previous theoretical data (e.g. 203 Walker et al. 2006), the most stable defect, i.e. V_{M3} O1_O1, corresponds to a M3 vacancy with protonation of O1-type oxygens, which are the non-silicate oxygens already

205 strongly underbonded in the non-defective structure. The two OH bonds point toward the 206 apical O4 atoms of this vacancy and have similar bond lengths (Fig. 2, Table 1). This 207 leads to a coupling of OH motions with an in-phase stretching mode at 3367 cm^{-1} that is 208 polarized along the *b*-axis and an out-of-phase stretching mode at 3374 cm^{-1} , mainly 209 polarized in the *ac* plane. Overall this doublet displays its strongest contribution along the 210 *c*-axis.

 The second most stable configuration for a protonated magnesium vacancy still 212 involves a M3 site and corresponds to V_{M3} O1 O3. It differs from the previous OH- defect by one of the two protons that is now located between the two O3 oxygens and is 0.37 eV less stable. This energy difference implies that, even at high temperatures relevant to the mantle transition zone, the probability of occurrence of this configuration is small (less than 10 %). Bond valence analysis using the method of Donnay and Allmann (1970) evidences that the creation of a M3 vacancy leaves O1 and O3 oxygens with residual valence of 1.46 and 1.67 v.u., respectively. The loosely bound proton (O3- 219 H = 1.022 Å) gives rise to an absorption band at 2837 cm⁻¹ polarized along the *a*-axis. In this OH-defect, the small relaxation of the O1-H…O4 hydrogen bond and the decoupling 221 of the H motions shifts the corresponding band to 3441 cm^{-1} .

222 The first protonated M2 vacancy $(V_{M2}$ O1 O4) is 0.47 eV less stable than 223 V_{M3} O1 O1 (Table 1). When a M2 vacancy is created, the first proton necessarily goes 224 to the O1 oxygen that is significantly underbonded compared to others. This proton 225 points toward the centre of the vacancy and undergoes repulsion from the second proton. 226 Therefore the corresponding band at 3431 cm^{-1} displays a polarization in the *ac* plane (Fig. 3). The second proton is close to the O4-O4 octahedral edge, and with a long O4-H 228 bond (1.017 Å) leads to a band at 2909 cm⁻¹ polarized along the *b*-axis.

 Finally, the OH-defect corresponding to the building block of the most common structural model for hydrous wadsleyite (Smyth 1994) is found to be less favorable 231 energetically (i.e. V_{M2} O1 O1 is 0.69 eV less stable than V_{M3} O1 O1). The two O1-H groups are aligned along the *c*-axis with no obvious hydrogen bond (Fig. 3). One points 233 toward the centre of the M2 vacancy (3454 cm^{-1}) while the other points toward the centre 234 of a normally vacant octahedral site (3515 cm^{-1}) . This OH-defect corresponds to the one considered for wadsleyite in the theoretical studies of Winkler et al. (1995) and Balan et al. (2008).

4.2. Hydrogarnet-type OH-defects

 In wadsleyite, all silicon sites are equivalent. In the case of a silicon vacancy, a bond valence analysis (Donnay and Allmann 1970) indicates the following residual valences 0.96, 1.00 and 1.22 v.u. for the O3, the two O4 and the O2 oxygens, respectively. Seven starting configurations are investigated. All optimized configurations are similar in energy, within 0.1 eV. The two most favorable are shown in Fig. 4 and Table 2.

245 V_{Si}A OH-defect can be described as a cycle $O4...O3...O4$ where each oxygen atom is donor and acceptor of proton. The fourth proton bound to the O2 oxygen points toward a O4 oxygen. The corresponding spectrum displays four bands. The band at the 248 lowest wavenumber (3409 cm^{-1}) and with contributions in the three directions is related to the longest group, O2-H. The O4-H…O4 bond is parallel to the *a*-axis and corresponds

250 to the band at 3550 cm^{-1} . The two other groups, O4-H and O3-H, lead to a doublet at 251 higher wavenumber, 3673 and 3708 cm^{-1} , that are mainly polarized along the *a* and *c*-252 axis, respectively.

253 In the second hydrogarnet-type defect, $V_{Si}B$, the cycle $O4...O3...O4$ is also 254 present, leading to the same polarization of the three highest absorption bands. Because of slightly shorter OH bonds, these three bands are observed between 3701 and 3798 cm-255 256 ¹. The main difference with respect to the $V_{Si}A$ defect is the O2-H group, which switches 257 from the O4 oxygen to the O2 oxygen. This latter OH group corresponds to the band at 258 3375 cm⁻¹ polarized along the *a* axis. It must also be noticed that because of the similar 259 stability of these two OH-defects, $V_{Si}B$ configuration is not retained at high-pressure and 260 turns back to a $V_{Si}A$ configuration.

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262 **4.3. Silicon vacancies compensated by a magnesium cation and two protons**

 This type of OH-defects involving more complex charge-compensating mechanisms has already been invoked in other nominally anhydrous minerals (e.g. Blanchard et al. 2009, Balan et al. 2012, Panero 2010). Here the magnesium cation is placed either in the vacant tetrahedral site or in a normally vacant octahedral sites adjacent to the silicon vacancy. In other words, magnesium is in interstitial site. Eight starting configurations are investigated. The most stable configuration found corresponds to the situation where the extra magnesium is located outside the silicon vacancy (V_{Si} M_i) but the relative stabilization of this defect is small. All optimized OH-defects 271 are within 0.28 eV. In this V_{Si} M_i defect as well as in the defect with magnesium in 272 silicon vacancy (M_{Si}) , the same oxygens, O3 and O4, are protonated (Fig. 5 and Table 3).

 The orientation of the OH groups is however very different leading to contrasted IR 274 spectra. In V_{Si} M_i, the O4-H group points toward the O3 oxygen and the O3-H group points toward the other O4 oxygen of the vacancy. This leads to the bands at 3468 and 276 3541 cm⁻¹, respectively, with components in the three directions. A small coupling of hydrogen motions explains why the high-wavenumber band is more intense than the other, in contrast with the general relation between OH absorption coefficients and wavenumber (e.g. Libowitzky and Rossman 1997). This general relation relies on the fact that in the case of OH-stretching modes, the IR absorptivity is directly related to the magnitude of the hydrogen Born effective charge along the OH bond (Balan et al. 2008). 282 In the M_{Si} defect, the vacancy being occupied by the magnesium cation, the O3-H group 283 points toward the O3 oxygen of another $Si₂O₇$ group giving rise to the band at 2995 cm⁻¹, mainly polarized along the *a*-axis. The O4-H group is mainly oriented along the *b*-axis 285 with a non-linear hydrogen bond (O4-H...O4 angle \sim 135°) that leads to a IR band at 3301 cm⁻¹.

4.4. OH-defects associated with the migration of a silicon cation

 In line with the interpretation of Kudoh and Inoue (1999) about the partial occupancy by silicon of a normally vacant tetrahedral site (Si2), we consider a last type of OH-defects formed by a silicon cation displaced from its site to the Si2 site adjacent to either one or two magnesium vacancies in M3 site. Two or four protons are needed to compensate the charge and must be bound to the oxygen atoms of the silicon and/or magnesium vacancies. In contrast, the model of Smyth et al. (1997) assuming that the Si2 silicon cation comes from a M3 site, does not necessitate protons to compensate the charge, which would be balanced instead by an extra magnesium vacancy or a magnesium cation in silicon site. In our case, bond valence analysis evidences that the most underbonded oxygen atoms are the O3 of silicon vacancy followed equally by the O2 of the silicon vacancy and the O1 of the magnesium vacancies. Calculations confirm 300 this trend. The defect with only one magnesium vacancy (V_{Si} Si2 V_{M3}) is formed by a O3-H group pointing toward the centre of the silicon vacancy and a O1-H group pointing toward the apical O4 oxygen of the magnesium vacancy (Fig. 6 and Table 4). The O3-H 303 group with no obvious hydrogen bond leads to an absorption band at 3711 cm^{-1} that is 304 polarized in the *bc* plane, and the O1-H group is related to the band at 3365 cm⁻¹. In the case with two magnesium vacancies, another O1-H group is present in the second magnesium vacancy. Their hydrogen motions are coupled in in-phase and out-of-phase 307 stretching vibrations, giving rise to the overlapped bands at 3372 and 3377 cm^{-1} . The fourth hydrogen protonates the O2 oxygen of the silicon vacancy and points toward a O4 309 oxygen, leading to the band at 3245 cm^{-1} . The presence of this fourth hydrogen slightly 310 disturbs the O3-H group and shifts the corresponding band to 3796 cm^{-1} while the polarization is almost unchanged.

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313 **5. DISCUSSION**

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315 **5.1. Interpretation of the major group of OH-stretching bands**

316 As described in introduction, the major group of bands is formed by a doublet 317 with bands A and B that have an independent behavior as a function of the water content 318 (Fig. 1). At low water content (\sim 150 wt ppm H₂O), only the band A is visible and is itself 319 a thin doublet with a contribution at 3317 cm^{-1} polarized along the *b*-axis and a 320 contribution at 3326 cm⁻¹ polarized in the *ac* plane. In addition the band B appears with a 321 minimal pleochroism (slightly more intense along the *c*-axis) in samples containing more 322 than \sim 3000 wt ppm H₂O (Jacobsen et al. 2005).

323 The OH-defect V_{M3} O1 O1 is in good agreement with the position and pleochroism of the thin doublet A (Fig. 2), which is therefore assigned to OH groups vibrating along the O1…O4 unshared edges of the M3 octahedron. The calculated 326 pressure derivative of the OH-stretching frequencies (\sim -10 cm⁻¹/GPa, Table 1) is also in excellent agreement with the experimental measurements made on a sample containing 328 8000 wt ppm H₂O (\sim -9 cm⁻¹/GPa, Deon et al. 2010). Calculations suggest that the same O1-H…O4 group is also responsible for the band B, when it belongs to OH-defects like V_{Si} Si2 V_{M3} (Fig. 6).

 This assignment of the major group of bands is in general agreement with previous studies (Kohn et al. 2002, Jacobsen et al. 2005, Walker et al. 2006, Tsuchiya and Tsuchiya 2009, Deon et al. 2010). It corresponds to vacancies in M3 sites compensated by protons located along the O1…O4 edges. Here, we propose a modeling of the OH-stretching IR spectra, which allows us to specify the incorporation mechanisms. It is clearly shown that the evolution of the spectrum as a function of the water content necessarily involves at least two types of well-defined OH-defects.

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339 **5.2. Interpretation of the minor group of OH-stretching bands**

340 The situation is more complex for the minor group of bands that is made of the 341 band C polarized in the *bc* plane and another band of smaller intensity, band D, polarized

 in the *ac* plane (Fig. 1). Here again these two bands are independent and at low water content, only the band D is present.

 Experimentally, Jacobsen et al. (2005) assigned this minor doublet to protonation of other edges of a M3 vacancy (i.e. O1…O3 and O1…O1). This assignment that would fit the pleochroic behavior of these bands, has been ruled out by Deon et al. (2010) because it would not reproduce their high wavenumbers and their pressure dependence. Here, calculations show that it is impossible to stabilize OH groups along these octahedral edges. Deon et al. (2010) suggested instead the involvement of a non-linear O1-H…O2 hydrogen bond. The only possibility we found to deviate the O1-H group 351 from the c-axis in a fully protonated M2 vacancy is the V_{M2} O1 O4 OH-defect (Fig. 3). 352 The polarization is in good agreement with the observed band D, but the V_{M2} O1 O4 353 defect would also give rise to a band at \sim 3000 cm⁻¹, polarized along the *b*-axis. This latter band is not observed experimentally since only a contribution along the *a*-axis is present 355 at ~3000 cm⁻¹ (Fig. 1).

 From their theoretical investigation, Walker et al. (2006) would rather assign the observed minor doublet to a partial protonation of silicon vacancies. In the present study, we obtained several configurations for the hydrogarnet-type defects with similar energies, leading to variable IR signatures. One could find bands in the right wavenumber range and a similar pleochroism as the minor doublet (Fig. 4). But hydrogarnet-type defects generate too many bands and more importantly could not explain the independent behavior of the two bands as a function of the water content. That is why partial protonation has been evoked, like for instance when the charge of the silicon vacancy is compensated by a magnesium cation and two protons (Fig. 5). But in this case, the

 models do not display OH-stretching bands with the right polarizations. A partial 366 protonation of the silicon vacancy is also present in the V_{Si} Si2 V_{M3} defect already mentioned for explaining the band B of the major doublet. With such a defect-structure, the O3-H group would lead to a band polarized in the *bc* plane like the observed band C 369 (Fig. 6). With \sim -4 cm⁻¹/GPa for the O3-H group, this OH-defect shows the weakest pressure dependence after the low-wavenuber band of the M_{Si} OH-defect. This is also in agreement with its experimental counterpart that is almost pressure insensitive.

 The present results thus suggest that the main contribution of the minor doublet (band C) can be interpreted as the partial protonation of a silicon vacancy coupled with the migration of a silicon cation to the normally vacant tetrahedral site (Si2). The less intense contribution (band D) may be related to non-linear O1-H…O2 hydrogen bonds in M2 vacancies, but none of the investigated OH-defects fully satisfies the experimental constraints.

5.3. Absorption coefficient of OH-defects in wadsleyite

 The quantification of the water content by IR spectroscopy relies on the knowledge of molar absorption coefficients of OH-stretching modes involved. General calibrations have been derived showing a negative relation between the absorption coefficient and the mean wavenumber of the corresponding vibrational mode (Paterson 1982, Libowitzky and Rossman 1997). However for nominally anhydrous minerals, mineral-specific calibrations are needed to obtain accurate concentrations (e.g. Balan et al. 2008, Koch-Müller and Rhede 2010). To summarize, the OH-defects that account for the main absorption bands of IR spectra of quenched wadsleyite samples are those 388 involving M3 vacancies, i.e. the fully protonated M3 vacancies, V_{M3} O1 O1, and OH-389 defects associated with the migration of a silicon cation, V_{Si} Si2 V_{M3} and V_{Si} Si2 $2V_{M3}$. Only the theoretical integrated molar absorption coefficients of these likely OH-defects are reported in Figure 7a. We can distinguish two sets of data, one for 0 GPa and the other for 15 GPa. As already pointed out in Blanchard et al. (2009) for ringwoodite, the evolution of the wavenumber and the absorption coefficient with pressure depends on the local environment of the OH group but globally absorption coefficients increase with pressure while the average wavenumbers decrease. The distinction of two data sets means that for a precise quantification of the hydrogen content in *in situ* experiments, one must consider a higher absorption coefficient than the one determined after quench at 0 GPa (Deon et al. 2010).

 Absorption coefficients of OH-defects in forsterite and ringwoodite have already been computed using the same theoretical method (Blanchard et al. 2009, Balan et al. 401 2011). We can then compile the theoretical data for the three Mg_2SiO_4 polymorphs, considering only the most-likely OH-defects at 0 GPa (Fig. 7b). In agreement with experimental studies (Thomas et al. 2009, Koch-Müller et al. 2010, Withers et al. 2012), we find that all data are below the general theoretical trend derived by Balan et al. (2008) using mainly hydrous minerals. Considering the small number of data points for each mineral and considering the relatively small data scattering, it is possible to fit the whole 407 data of the three polymorphs by the single following relation: $K_{int} = 278.7 \pm 18.1$ (3810 \pm 465 - x). This trend suggests that an absorption coefficient obtained from a general 409 calibration would underestimate the water content of the Mg_2SiO_4 minerals. This underestimation depends on the position of the main absorption bands. It is almost

411 negligible for wavenumbers around 3600 cm⁻¹, but becomes important for smaller 412 wavenumbers. For instance for the major band of wadsleyite, observed at \sim 3350 cm⁻¹, the 413 absorption coefficient obtained from the Mg₂SiO₄ calibration is ~23 % lower than the one 414 obtained from the general calibration. This is consistent with the experimental 415 determination by Deon et al. (2010) of an integrated molar absorption coefficient of 73 416 000 cm⁻² per mol H₂0/L that is well below the general calibration curves. For instance, 417 the calibration of Libowitzky and Rossman (1997) gives an absorption coefficient of 91 418 512 cm⁻² per mol H₂0/L (~25 % higher) for the same wavenumber. To compare the three 419 Mg2SiO⁴ polymorphs, one must keep in mind the positions of the main absorption bands 420 in each mineral. For the range of the main absorption bands in wadsleyite $(\sim 3350 \text{ cm}^{-1})$ 421 and ringwoodite $({\sim}3150 \text{ cm}^{-1})$, the calculated absorption coefficient for ringwoodite is 44 422 % larger than the one for wadsleyite, which is in good agreement with the results of 423 Koch-Müller et al. (2010) who obtained an absorption coefficient for the Mg end-424 member of ringwoodite 43 % larger than the wadsleyite value of Deon et al. (2010). 425 Forsterite displays a large variety of OH-stretching IR spectra with bands ranging from 426 3200 to 3650 cm⁻¹. For IR spectra dominated by high-frequency bands, Thomas et al. 427 (2009) and Withers et al. (2012) found absorption coefficients between 35 000 and 50 428 000 cm⁻² per mol H₂0/L that is significantly lower than the wadsleyite coefficient of Deon 429 et al. (2010), i.e. 73 000 cm⁻² per mol H₂0/L. All these trends are supported by our 430 theoretical results (Fig. 8). It is also interesting to note that the relation between the OH 431 bond length and the stretching wavenumber determined for forsterite, ringwoodite and 432 enstatite (Balan et al. 2012) also apply to wadsleyite.

 In conclusion, a specific calibration for the Mg2SiO4 polymorphs must be used to quantify accurately the water content of these minerals from IR spectroscopy. Results also suggest that the calibration determined at 0 GPa from quenched samples does not apply with the same accuracy to *in situ* IR measurements.

5.4. Concluding remarks about the bands assignment

 Jacobsen et al. (2005) pointed out that the *b*/*a* axial ratio of wadsleyite increases with increasing water content. Figure 8 shows this empirical relation as well as the calculated OH-defects. All OH-defects involving M3 vacancies and proposed here to interpret the IR spectra, follow a similar trend as the one seen experimentally (i.e. the 443 fully protonated M3 vacancies, V_{M3} O1 O1, and OH-defects associated with the 444 migration of a silicon cation, V_{Si} Si2 V_{M3} and V_{Si} Si2 $2V_{M3}$). Note that the V_{M3} O1 O3 defect, which falls in the right range, was ruled out because of its relative energy with 446 respect to the V_{M3} O1 O1 defect (0.37 eV less stable). On the other hand, the protonated M2 vacancies lead to a smaller *b*/*a* ratio than that observed for anhydrous wadsleyite and the creation of the hydrogarnet-type defects does not change much the axial ratio.

 Based on results inferred from X-ray diffraction studies and from the confrontation of experimental and theoretical IR spectra, we propose the following mechanisms of hydrogen incorporation in wadsleyite. At low water content, protons are 452 associated with M3 vacancies, thus forming preferentially V_{M3} O1 O1 OH-defects, which are the energetically most favorable OH-defects. This yields the thin doublet A (Fig. 1a). In addition, the small band D indicates that a small amount of non-linear O1- H...O2 hydrogen bonds is formed in M2 vacancies. As experimentally seen, the number 456 of M3 vacancies increases with the water content. More V_{M3} O1 O1 OH-defects are formed. The high concentration of M3 vacancies facilitates the migration of silicon cation to Si2 sites that is accompanied by a redistribution of protons. This leads to the formation 459 of OH-defects like V_{Si} Si2_ V_{M3} and V_{Si} Si2_2 V_{M3} that are responsible of the absorption bands B and C (Fig. 1b). Other OH-defects may also be involved in a smaller extent as reflected by the additional weak bands observed in wadsleyite spectra.

 If the migration of silicon cation to Si2 sites occurs during the decompression, as suggested by Smyth et al. (1997) and Kudoh and Inoue (1999), it means that the IR spectra collected at 0 GPa and even the high-pressure IR spectra collected on quenched samples do not reflect the population of OH-defects present in mantle conditions. In this case, calculations predict that the wadsleyite IR spectra would be dominated by the main 467 group of absorption bands reported around 3350 cm^{-1} at 0 GPa (these bands would be shifted to lower wavenumbers under pressure), and associated with fully protonated M3 vacancies. IR spectra collected *in situ* and during the sample decompression would allow tackling this question.

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 Figure 7. a. Theoretical integrated molar absorption coefficients of the stretching modes of OH-defects in wadsleyite as a function of their average wavenumber. Circles and squares represent the likely OH-defects in wadsleyite at 0 and 15 GPa, respectively. The general theoretical trend of Balan et al. (2008) is also reported for comparison. b. Same graph for the likely OH-defects in forsterite, triangles (Balan et al. 2011), in wadsleyite at 0 GPa, circles (this study) and in ringwoodite at 0 GPa, squares (Blanchard et al. 2009). 624 A single linear fit for the three Mg_2SiO_4 polymorphs leads to the following equation: K_{int} 625 = 278.7 \pm 18.1 (3810 \pm 465 - x).

 Figure 8. Variation of the (*b*/*a*) axial ratio of wadsleyite (at 0 GPa) with water content. Solid square is for anhydrous wadsleyite. The bold line corresponds to the empirical 629 relation derived by Jacobsen et al. (2005), $(b/a) = 2.008(1) + 1.25(3) \times 10^{-6} \cdot C_{H_2O}$.

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631 **TABLES**

632

633 **Table 1**. Properties of protonated Mg vacancies

634 Dimensions of the $2\times1\times1$ supercell are converted in unit-cell parameters.

 635 $\frac{b}{b}$ No obvious single hydrogen bond.

636 **Table 2**. Properties of hydrogarnet-type protonated Si vacancies

649 **Table 3**. Properties of Si vacancies compensated by a magnesium cation and two protons

- 662 **Table 4**. Properties of OH-defects formed by a silicon cation moved from its site to a
- 663 normally vacant tetrahedral site (Si2) adjacent to either one or two magnesium vacancies
- 664 in M3 site

665 a Dimensions of the $2\times1\times1$ supercell are converted in unit-cell parameters.

666 ^b The relative energy is not reported since these OH-defects do not have the same 667 stoichiometry.

- 668 ^c No obvious single hydrogen bond.
- 669 d Modes involving a coupled motion of OH groups.
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673 Fig. 1

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696 Fig. 8