

1  
2  
3  
4  
5  
6  
7  
8  
9  
10  
11  
12  
13  
14  
15  
16  
17  
18  
19  
20  
21  
22  
23  
24  
25  
26

## Structure of Prismatic Halloysite

TOSHIHIRO KOGURE,<sup>1\*</sup> KIYOFUMI MORI,<sup>2</sup> VICTOR A DRITS,<sup>3</sup> AND  
YOSHIZO TAKAI<sup>2</sup>

<sup>1</sup>Department of Earth and Planetary Science, Graduate School of Science, The University of  
Tokyo, 7-3-1 Hongo, Bunkyo-ku, Tokyo, 113-0033, Japan

<sup>2</sup>Department of Material and Life Science, Graduate School of Engineering, Osaka University,  
2-1 Yamada-Oka, Suita, Osaka, 565-0871, Japan

<sup>3</sup>Geological Institute of the Russian Academy of Sciences, Pyzhevsky per 7, Moscow, Russia

\*E-mail: [kogure@eps.s.u-tokyo.ac.jp](mailto:kogure@eps.s.u-tokyo.ac.jp)

---

**Running title proposed:** Structure of prismatic halloysite

**Corresponding author:**

Toshihiro Kogure

Department of Earth and Planetary Science, Graduate School of Science,

The University of Tokyo

7-3-1 Hongo, Bunkyo-ku, Tokyo 113-0033, Japan

E-mail: [kogure@eps.s.u-tokyo.ac.jp](mailto:kogure@eps.s.u-tokyo.ac.jp)

**Revision 1**  
December 2012

27

## ABSTRACT

28 Halloysite from Olkhon Island, Lake Baikal, Russia has been investigated using X-ray  
29 diffraction (XRD), scanning electron microscopy (SEM), selected-area electron diffraction  
30 (SAED) and, in particular, high-resolution transmission electron microscopy (HRTEM), to  
31 reveal its atomic structure and formation process. XRD analysis indicated a basal spacing of ca.  
32 7.2 Å and two characteristic peaks with  $d = 4.28$  and  $4.03$  Å on the tail of the 02, 11 band. The  
33 halloysite grain cross-sections displayed various prismatic forms, ranging from a rectangle to a  
34 regular eighteen-sectored polygon in SEM and TEM. The SAED pattern from a sector of the  
35 polygonal prisms with the incident beam parallel to the prism axes showed a regular one-layer  
36 oblique reciprocal lattice, similar to that of kaolinite along  $Y_1$ -directions. HRTEM imaging  
37 performed with the new computer-assisted minimal-dose system and an incident beam  
38 perpendicular to the prism axis showed stacking of most dioctahedral 1:1 layers with their  
39 pseudo-mirror plane perpendicular to the prism axis and an almost random, or rather,  
40 alternating lateral stagger to the right or left from the preceding layer, which corresponds to  
41 interlayer displacements of  $\tau_+$  and  $\tau_-$  in Zvyagin symbols, or layer displacements of  $t_1$  and  $t_2$   
42 used to describe the stacking in kaolinite. This stacking feature explains the SAED pattern from  
43 the side of the prismatic grains and the two characteristic peaks on the tail of the 02, 11 band in  
44 the XRD pattern. Based on these results, it is proposed that tubular halloysite initially forms as  
45 a hydrated one with the pseudo-mirror plane of the kaolinite layers perpendicular to the tube  
46 axis, then dehydrates with, possibly, partially hydrogen-bonded interlayers, and finally  
47 transforms to a prismatic one consisting of sectored flat layers with the complete  
48 hydrogen-bonded interlayers. During this transformation, stacking with comparable ratio and  
49 frequent alternation of  $\tau_+$  and  $\tau_-$  is formed, to minimize morphological change of the tubes.  
50 **Keywords:** halloysite, stacking sequence, prismatic form, hydrogen bonding, HRETM, SAED

51

52

## INTRODUCTION

53

54

55

56

57

58

59

60

Halloysite,  $\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4 \cdot n\text{H}_2\text{O}$ , is one of the most common and ubiquitous clay minerals on the terrestrial surface. However, it is also true that the real structure of halloysite has yet to be fully understood despite the amount of research devoted to this mineral. Actually, no three-dimensional periodicity has been reported for halloysite because of very limited information from conventional diffractometry, which leads to the possibility that halloysite contains intense structural disorder. Moreover, from the standpoint of electron microscopy, halloysite is too beam-sensitive to apply recent high-resolution structure imaging in order to analyze its atomic structure.

61

62

63

64

65

66

67

68

69

70

71

72

73

74

75

Halloysite is one of the kaolin group minerals consisting of dioctahedral 1:1 layers which are termed “kaolinite layers” in this study, and hydrated/dehydrated interlayers. In cases involving halloysite with a tubular morphology, the kaolinite layers are rolled to compensate for the misfit between the tetrahedral and octahedral sheets (Bates et al., 1950). Bailey (1990) suggested ditrigonal rotation in the tetrahedral sheet, which reduces the misfit in the other kaolin group minerals may not be realized in halloysite due to water molecules and exchangeable cations possibly existing at the interlayer region. On the contrary, Singh (1996) proposed that the rolling mechanism of the kaolinite layers is preferable to ditrigonal rotation, if Si-Si repulsion in the tetrahedral sheet is considered. With respect to the stacking structure of halloysite, much research has suggested a two-layer periodicity for the dehydrated form, halloysite (7 Å), based on selected-area electron diffraction (SAED) taken with the incident beam perpendicular to the tube-axis (Honjo and Mihara, 1954; Honjo et al., 1954; Chukhrov and Zvyagin, 1966; Kohyama et al., 1978; Singh and Gilkes, 1992). Among them, probably only Honjo et al. (1954) and Chukhrov and Zvyagin (1966) considered the stacking structure to form the two-layer periodicity. Based on the intensity distribution in the SAED pattern, Honjo et al. (1954)

76 proposed a stacking sequence expressed as  $\sigma_1\tau_+\sigma_1\tau_0\sigma_1$  or  $\sigma_5\tau_-\sigma_5\tau_0\sigma_5$  with a triclinic symmetry,  
77 using the Zvyagin symbols (Chukhrov and Zvyagin, 1966). However, Chukhrov and Zvyagin  
78 (1966) suggested that the sequence expressed as  $\sigma_3\tau_+\sigma_3\tau_-\sigma_3$  with monoclinic *Cc* symmetry  
79 could also reproduce the intensity distribution in the SAED pattern and insisted that this  
80 sequence was more realistic than that proposed by Honjo et al. (1954). In this notation,  $\sigma_3$   
81 corresponds to the intralayer displacement of  $\mathbf{a}_o/3$  in the kaolinite layer having the orthogonal  
82 unit cell with  $\mathbf{a}_o$  and  $\mathbf{b}_o$  cell edges, and  $\tau_+$  and  $\tau_-$  correspond to the interlayer displacements of  
83  $+\mathbf{b}_o/3$  and  $-\mathbf{b}_o/3$ , respectively. This sequence is also described as the regular alternation of the  
84 layer displacements (Guggenheim et al., 2009) with  $\mathbf{a}_o/3 + \mathbf{b}_o/3$  ( $\sigma_3\tau_+$ ) and  $\mathbf{a}_o/3 - \mathbf{b}_o/3$  ( $\sigma_3\tau_-$ ).  
85 Bookin et al. (1989) showed that two-dimensional layer periodicity of defect-free *1Tc*  
86 kaolinite can be described equally well by the orthogonal cell  $\{a_o, b_o, \gamma_o\}$  ( $\gamma_o = 90^\circ$ ) as well as by  
87 the enantiomorphic oblique cells  $\{a_1, b_1, \gamma_1\}$  and  $\{a_2, b_2, \gamma_2\}$  which are related to each other by a  
88 mirror plane passing through the octahedral vacant site and the center of the ditrigonal ring of  
89 the tetrahedral sheet in kaolinite layer, and containing the  $\mathbf{a}_o$  axis. In particular, the  $\mathbf{a}_1$  and  $\mathbf{a}_2$   
90 axes of the enantiomorphic oblique cells are rotated clockwise and counterclockwise by  $\sim 120^\circ$   
91 with respect to the  $\mathbf{a}_o$  axis. The numerical values for the oblique cells coincide with those  
92 determined for kaolinite sample from Keokuk which is almost defect-free in the stacking  
93 sequence (Sutch and Young, 1983; Bish and von Dreele, 1989), indicating that the obliquity ( $\gamma_1$ ,  
94  $\gamma_2 \neq 90^\circ$ ) of the refined cell of kaolinite is just the choice of the unit cell and interlayer stacking  
95 sequence does not disturb the layer symmetry. On the other hand, the layer displacements  $\mathbf{t}_1$  and  
96  $\mathbf{t}_2$  which are corresponding to the individual enantiomorphs and related to each other by the  
97 mirror plane mentioned above form a defect-free *1Tc* kaolinite structure (Bookin et al., 1989).  
98 In terms of the conventional cell  $(a_1, b_1, \gamma_1)$  chosen for description of a defect-free kaolinite  
99 structure, the components of the  $\mathbf{t}_1$  and  $\mathbf{t}_2$  are  $(-0.369\mathbf{a}_1, -0.024\mathbf{b}_1)$  and  $(-0.352\mathbf{a}_1, 0.304\mathbf{b}_1)$ ,  
100 respectively. Bookin et al. (1989) also proposed that the stacking sequence in halloysite is

101 expressed as the regular interstratification of  $t_1$  and  $t_2$ . The components of the  $t_1$  and  $t_2$  expressed  
102 with the orthogonal cell are  $0.282a_o + 0.327b_o$  and  $0.282a_o - 0.327b_o$ , respectively. These  
103 expressions indicate the similarity and difference between the model by Chukhrov and Zvyagin  
104 (1966) and Bookin et al. (1989). The first authors used the layer structure with the idealized  
105 intralayer and interlayer displacements whereas the other assumed that the interlayer  
106 displacements in halloysite are identical to those in a periodic  $1Tc$  kaolinite despite the quite  
107 different mutual arrangement of the adjacent layers in these structures.

108 Recently, Kogure et al. (2011) succeeded in recording clear two-dimensional high-resolution  
109 transmission electron microscopy (HRTEM) images from tubular halloysite, using a new TEM  
110 with a computer-assisted minimal-dose system (Hayashida et al. 2007). These images showed  
111 no symptom of a two-layer periodicity which was suggested in the previous works. They did,  
112 however, show a more likely one-layer periodicity with high density of stacking disorder. The  
113 discrepancy between two-layer periodicity and one-layer periodicity with a high density of  
114 stacking disorder should be looked into during future halloysite research.

115 One of the characteristic features of halloysite is its various morphologies, typical of which  
116 are tubes and spheres (see the summarized table in Joussein et al. 2005). Furthermore, Bailey  
117 (1990) described the “prismatic” form as a category of halloysite morphologies. Chukhrov and  
118 Zvyagin (1966) and Dixon and McKee (1974) showed TEM replica images in which the  
119 elongate halloysite particles had regular and flat faces parallel to the elongation of the particles.  
120 It is noteworthy that the stacking sequence reported by Chukhrov and Zvyagin (1966) as  
121 mentioned above was derived from a specimen with prismatic morphology. Beside the TEM  
122 replica images, cross sectional TEM (Bates and Comer, 1958; Singh and Gilkes, 1992,  
123 Churchman et al., 1995) and scanning electron microscopy (SEM) (Diamond and Bloor, 1970)  
124 indicated prismatic shapes for halloysite. Such prismatic morphology was proposed to be

125 formed during dehydration (Kirkman, 1981) or transformation from kaolinite to halloysite  
126 (Robertson and Eggleton, 1991).

127 Following our previous work in investigating the atomic structure of tubular halloysite using  
128 advanced electron microscopy (Kogure et al., 2011), we have investigated a different halloysite  
129 specimen which apparently shows “higher crystallinity” in X-ray diffraction (XRD). SEM  
130 observation indicated that the morphology of most grains is prismatic, similar to those reported  
131 by Chukhrov and Zvyagin (1966) and Dixon and McKee (1974). HRTEM images were  
132 successfully obtained and showed a different stacking structure from that in tubular halloysite  
133 reported in Kogure et al. (2011). Finally, experimental XRD and SAED patterns were  
134 successfully reproduced in simulation using the recorded stacking structure in these HRTEM  
135 images. These results provide new insights into the real structure of halloysite and its formation  
136 process.

## 137 **SAMPLE AND METHODS**

138 The halloysite specimen investigated was from Olkhon Island, Lake Baikal, Russia, and  
139 formed by weathering of oligoclase (Chekin et al., 1972). It was reported that the halloysite  
140 crystallites initially grew epitaxially on the (001) surface of oligoclase with a platy and  
141 dendritic shape elongated along the [100], [110], [010] and [310] directions. When the  
142 crystallites became rather thick, they peeled away from the oligoclase surface and grew as a  
143 tubular crystal along the *b*-axis with a spiral-cylindrical structure (Chekin et al., 1972, 1976).

144 The powder XRD pattern was measured at an ambient condition using a Rigaku  
145 RINT-Ultima<sup>+</sup> diffractometer with monochromatized CuK $\alpha$  radiation, a 0.3 mm receiving slit,  
146 and 0.5° divergence and anti-scatter slits. A continuous scan rate of 0.1° (2 $\theta$ ) min<sup>-1</sup> and a  
147 sampling interval of 0.02° were adopted. High-resolution SEM was conducted using a Hitachi  
148 S-4500 SEM with a cold type field-emission gun operated at 2 kV. The specimen was coated  
149 with Pt-Pd of 5 nm in thickness using a sputter coater to provide electron conductivity.

150 TEM specimens for cross-sectional views of the prismatic halloysite were prepared using  
151 ion-milling. The halloysite grains were mixed with epoxy resin. After hardening, the mixture  
152 was cut into slices and mechanically thinned to less than 100  $\mu\text{m}$ , then made  
153 electron-transparent using an ion-milling instrument (Gatan Dual Ion Mill Model 600) with  
154 argon beams of  $\sim 4$  KeV and an incident angle of  $15^\circ$ . Images and electron diffraction patterns  
155 with the prism-axis perpendicular to the beam were acquired using halloysite particles, and  
156 dispersed on holey carbon microgrids through aqueous suspension. TEM images with low  
157 magnifications and SAED patterns were acquired using JEOL JEM-2010UHR with a LaB<sub>6</sub>  
158 filament operated at 200 kV. The images were recorded with films or a Gatan MSC 794  
159 bottom-mounted CCD camera, while SAED patterns were recorded with a Gatan ES-500W  
160 side-mounted CCD camera on the phosphor screen.

161 HRTEM images for analyzing the atomic structure in halloysite were acquired using a Hitachi  
162 HF-1000BC, with a cold field emission gun operated at 100 kV and a newly developed minimal  
163 dose system (MDS) for reducing the electron dose. The MDS consists of a CCD camera, an  
164 electric-field-type electron beam blanker that is installed beneath the condenser lens, and a  
165 computer for controlling the entire TEM system, including a CCD camera and a high speed  
166 blanker. The details for practical operation were described in our previous paper (Kogure et al.,  
167 2011). Noise contrast in the HRTEM images was removed using a Wiener-filter (Marks 1996;  
168 Kilaas 1998) developed by K. Ishizuka (HREM Research Inc.) and implemented with a Gatan  
169 DigitalMicrograph, version 3.1 0.0 (Kogure et al., 2008).

170 Finally, simulations of powder XRD and SAED patterns were performed using DIFFaX  
171 (Treacy et al., 1991), which calculates diffraction patterns from layered materials with stacking  
172 disorder.

## 173 RESULTS AND DISCUSSION

### 174 Powder XRD

175 The XRD pattern (Fig. 1a) of the specimen indicates a basal spacing of ca. 7.2 Å,  
176 suggesting that the halloysite is almost dehydrated, or halloysite-(7 Å). Apart from the basal  
177 reflections, several peaks are distinctly identified in the patterns, deviating from the XRD  
178 patterns of halloysite in other research (e.g. Joussein et al., 2005; Kogure et al., 2011). These  
179 peaks are compared with those calculated from the crystallographic parameters for kaolinite  
180 (Neder et al., 1999) in Figure 1. In the calculated patterns, the pattern for reflections with  $k = 3n$   
181 is drawn with a solid line, while that for all reflections is drawn with a broken line. First, the  
182 experimental pattern profile higher than  $2\theta$  of  $35^\circ$  corresponds, approximately, to the calculated  
183 pattern for kaolinite. In particular, calculated reflections with  $k = 3n$  also almost make an  
184 appearance in the experimental pattern. The reflections  $h \neq 0$  and  $k = 3n$  are called family  
185 reflections (Đurovič and Weiss, 1986) and their intensity distributions are distinct for the four  
186 polytypic groups (or structure types) of 1:1 phyllosilicates (Bailey, 1969; Zvyagin, 1967). The  
187 rough correspondence of the family reflections between the experimental and calculated  
188 patterns indicates that the stacking sequence in the present halloysite sample belongs to group  
189 A with a layer displacement of approximately  $\pm a_1/3$  and/or mutual rotations between adjacent  
190 layers with  $0^\circ$  or  $\pm 120^\circ$ . On the other hand, the experimental pattern shows two characteristic  
191 peaks on the tail of the 02, 11 band, as indicated with arrows one and two. The  $d$ -values are 4.28  
192 and 4.03 Å, which are in complete disagreement with the peaks for kaolinite. These  $d$ -values  
193 can be indexed as 021 and  $\bar{1}12$  with a two-layer unit cell (Chukhrov and Zvyagin, 1966; Bailey,  
194 1990). However, these reflections cannot be ascribed to dickite, because those for dickite have  
195 very weak intensities when compared to other reflections such as 022 or 111. On the contrary,  
196 these reflections have the strongest intensities in the 02 $l$ , 11 $l$ , and  $\bar{1}1l$  reciprocal lattice rows  
197 for the stacking model proposed by Chukhrov and Zvyagin (1966). Another intense reflection,  
198 112 with  $d \sim 3.49$  Å, as expected from their model, is overlapped with 004 peaks in the powder



199 XRD pattern. Hence, the experimental XRD pattern in Figure 1 seems consistent with the  
200 Chukhrov and Zvyagin model.

201

## 202 **Morphology analyses by SEM and TEM**

203         Several secondary electron images taken by the FE-SEM are shown in Figure 2. Most  
204 halloysite grains are elongated and appear roughly cylindrical, however their cross-section is  
205 often not circular but polygonal, as indicated by the white arrows. These morphologies are very  
206 close to those reported by Chukhrov and Zvyagin (1966) and Dixon and McKee (1974),  
207 although their images are TEM replicas. In Figure 2a, the arrowed grain has a rectangular  
208 prismatic form. Such a prismatic morphology, with a small number of faces and acute angles  
209 between the adjacent faces, is commonly observed. Cross sectional TEM images of the  
210 elongated grains thinned by ion-milling are presented in Figure 3. Some grains show smoothly  
211 curved layers and appear cylindrical rather than prismatic (Fig. 3a). On the contrary, Figure 3b  
212 shows the cross section of a prismatic halloysite with five major sectors. At the boundaries of  
213 the sectors with an acute corner angle, a section of the layers are connected with a small  
214 curvature radius. It is not certain whether the void-like regions between the connected layers at  
215 the boundaries are the original structure or formed by radiation damage before the recording  
216 was completed. In Figure 3c, the cross-section is almost circular, however, a close examination  
217 of the 7 Å lattice fringes reveals that the halloysite grain also consists of many sectors. This is  
218 clearly confirmed by the SAED pattern from the grain (Fig. 3d). Eighteen intensity maximums  
219 can be counted, with an equal interval on the rings corresponding to 7.2 and 3.6 Å. Hence, the  
220 grain consists of eighteen sectors with a center angle of 20°. This structure is very similar to that  
221 of a polygonal serpentine as reported in other research, in which fifteen and thirty sectors are  
222 common (e.g., Cressey and Zussman, 1976; Yada and Wei, 1987). If we return to the SEM  
223 images in Figure 2 and examine the contrast on the surface of relatively large halloysite grains

224 with cylindrical forms, it can be seen that the surface consists of facets elongated parallel to the  
225 cylinder axis, as indicated with the thin black arrows in Figure 2c.

226

### 227 **Stacking analyses using SAED and HRTEM**

228 Figure 4 shows the cross-section of a prismatic halloysite consisting of three sectors. One half  
229 of the prism disappeared during sample preparation by ion-milling. Moreover, the areas around  
230 the center and outside of the prism were amorphized by beam radiation. The corresponding  
231 SAED pattern from the entire prism is comprised of three superimposed reciprocal lattice nets,  
232 each of which corresponds to the individual sector. The three reciprocal lattice nets are  
233 identical; one-layer periodicity, the oblique angle, and no streak, indicating the typical pattern  
234 for 1:1 phyllosilicates of group A (Bailey, 1988), which corresponds to the XRD analysis result  
235 in Figure 1. On the contrary, two-dimensional HRTEM images recorded using the TEM with a  
236 minimal dose system are shown in Figure 5. As shown in our previous paper (Kogure et al.,  
237 2011), the contrast for each kaolinite layer consists of two-tiered dark dots separated by  $b/2$  (~  
238 4.5 Å) along the layer. Each dark dot can correspond to a pair of SiO<sub>4</sub> tetrahedra (or an SiO<sub>4</sub>  
239 chain running parallel to the incident beam) in the tetrahedral sheet and a pair of AlO<sub>2</sub>(OH)<sub>4</sub>  
240 octahedra in the dioctahedral sheet (see Fig. 1 in Kogure and Inoue, 2005). In Figure 5, the  
241 contrast of each layer shows vertically aligned dots with no lateral shift within the layer forming  
242 a bar-like shape (indicated with a square for each layer), except for three layers in Figure 5b,  
243 which are indicated with triangles. Such contrast should appear in the HRTEM images of  
244 kaolinite viewed along [110] or  $[\bar{1}\bar{1}0]$  (see Fig. 3 in Kogure and Inoue, 2005). More universally,  
245 these directions are expressed as those parallel to the pseudo-mirror plane described by Bookin  
246 et al. (1989), passing through both the octahedral vacant site and center of the tetrahedral  
247 six-membered ring in the kaolinite layer. Moreover, such contrast for each kaolinite layer is  
248 laterally shifted to the right or left against that for the lower and upper layers in the amount of

249  $b/6$  (only two exceptions are indicated with the white arrows in Figure 5c, where no lateral shift  
250 is observed). As described in Kogure and Inoue (2005), these right and left shifts correspond to  
251 the projections of the layer displacements for  $t_1$  and  $t_2$  (In actuality, they are related to each  
252 other through the pseudo-mirror plane), or directly correspond to the Zvyagin's interlayer shifts  
253 of  $\tau_+$  and  $\tau_-$ . Hence, the two-layer stacking model proposed by Chukhrov and Zvyagin (1966)  
254 should appear as the alternating shift of the contrast for the kaolinite layer in Figure 5. However,  
255 such alternation is only observed locally (indicated with square brackets in the figure) with six  
256 or seven layers, and not in the entire stacking.

257 As mentioned above, two-layer periodicity was proposed for the stacking structure of  
258 halloysite, based on electron diffraction from tubular grains in previous researches pertaining to  
259 halloysite (Honjo and Mihara, 1954; Honjo et al., 1954; Chukhrov and Zvyagin, 1966,  
260 Kohyama et al., 1978; Singh and Gilkes, 1992). Previous research found the intensity  
261 maximum at a position parallel to the tube axis on the  $02l$  row, which can be indexed 020. On  
262 the  $02l$  row, the next intensity peak to 020 has a distance that nearly corresponds to two-layer  
263 periodicity, which is perhaps the reason why the two-layer structure was proposed. However,  
264 the  $02l$  diffraction rows in their SAED patterns are considerably streaked, except for that  
265 presented by Chukhrov and Zvyagin (1966). It is possible that these patterns were obtained with  
266 a selected-area aperture including the whole diameter of the tubular form. In this case, the  
267 intensity maximum will be formed at the 020 position by the streaked  $02l$  diffraction row from  
268 the top and/or bottom parts of the tubular wall normal to the electron beam. Figures 6b and c  
269 show examples of the SAED patterns from the present specimen, obtained by including the  
270 whole diameter (Fig. 6b) and only the side region of the prism (Fig. 6c), using the selected-area  
271 apertures drawn in Figure 6a. The intensity maximum at the 020 position as indicated with the  
272 white arrows in Figure 6b does not appear in Figure 6c. Most halloysite prisms which are thick  
273 enough to set the aperture as in Figure 6a showed similar results; they did not indicate two-layer

274 periodicity. However, occasionally the intensity maximum appeared at the 020 position in the  
275 SAED pattern from the side region of a prism, implying that the possibility of a two-layer  
276 structure could not be fully excluded.

277

### 278 **Simulation of XRD and SAED pattern**

279 Another issue to be considered is the origin of the two characteristic peaks at the 02, 11 band  
280 with  $d = 4.28$  and  $4.03 \text{ \AA}$  in the XRD pattern (Fig. 1), which were indexed as 021 and  $\bar{1}12$  with  
281 a two-layer structure. In Figure 5, the HRTEM images indicate almost the same intralayer  
282 displacement ( $\sigma_3$  of the Zvyagin symbol) and heavily disordered interlayer displacement with  
283  $\tau_+$  and  $\tau_-$  (or layer displacement with  $t_1$  and  $t_2$ ). Powder XRD and SAED patterns from this  
284 structure model have been simulated using DIFFaX (Treacy et al., 1991). Such a simulation for  
285 XRD was already performed by Bookin et al. (1989), but their result (Fig. 6 in their paper) is  
286 confusing owing to an editorial mistake. The same result presented in Drits and Tchoubar  
287 (1990) is preferable as a reference.

288 The calculation procedure is similar to our previous research (Kogure et al., 2006a; Kogure et  
289 al., 2010). The lattice constants and atomic coordinates for the kaolinite layer were derived  
290 from the results by Neder et al. (1999) and coordinates for  $t_1$  and  $t_2$  are those described in  
291 Bookin et al. (1989) as described above. To make the calculation more simple, we assume equal  
292 proportions for  $t_1$  and  $t_2$  ( $W_1 = W_2 = 0.5$ ). Accordingly, all of the condition probabilities ( $P_{11}$ ,  
293  $P_{12}$ ,  $P_{21}$ ,  $P_{22}$ ) are determined if one of them, e.g.,  $P_{11}$ , is specified (Bookin et al., 1989), where  
294  $P_{11}$  refers to the probability that  $t_1$  comes next to  $t_1$  in the stacking. Figure 7 shows the calculated  
295 XRD patterns with various  $P_{11}$  values, and the experimental one for comparison. We adopted a  
296 pseudo-Voigt function with the same ratio of Gaussian and Lorentz shapes and a half-width of  
297  $0.1^\circ$  for the peak profile. In Figure 7,  $P_{11} = 0$  corresponds to the stacking model by Chukhrov  
298 and Zvyagin (1966) and Bookin et al. (1989) with two-layer periodicity, and  $P_{11} = 1$  to the

299 segregated two enantiomeric kaolinite crystals with one-layer periodicity. Based on the figure,  
300 it is understood that the two characteristic peaks at the 02, 11 band in the experimental pattern  
301 are reproduced in the wide range of  $P_{11}$ , from  $P_{11} = 0$  to even the most disordered state of  $P_{11} =$   
302 0.5. The peaks corresponding to those for kaolinite appear from  $P_{11} = 0.7$ . Hence, the two peaks  
303 in the experimental pattern do not necessarily indicate ordered two-layer periodicity, but only a  
304 tendency towards it.

305 The calculated intensity distributions along the reciprocal row next to 00 $l$  in the SAED  
306 patterns in Figure 6c were also calculated using DIFFaX and shown in Figure 8. As the disorder  
307 increases, the peaks are broadened and those with  $l = 2n$  disappear, losing any evidence of  
308 two-layer periodicity. The twin maximums at  $l = \pm 1$  correspond to those in the SAED pattern in  
309 Figure 6c and they also correspond to the peak with  $d = 4.28 \text{ \AA}$  in the XRD pattern. The  
310 intensity maximums corresponding to the peak with  $d = 4.03 \text{ \AA}$  in the XRD pattern are on the  
311 other reciprocal rows. As  $P_{11}$  exceeds 0.5, these peaks move and split and finally ( $P_{11} = 1.0$ ) the  
312 pattern becomes that of the 11 $l$  diffraction row from the two enantiomeric kaolinite crystals.

313

#### 314 **Formation and ripening process of halloysite**

315 When summarizing the result of the analysis described above, the crystal structure of  
316 prismatic halloysite is expressed as an almost-ordered layer orientation (intralayer  
317 displacement) and heavily-disordered layer displacement with  $t_1$  and  $t_2$ . This expression is  
318 almost the same as the conclusions for disordered kaolinite by Plançon et al. (1989) derived  
319 using XRD and Kogure and Inoue (2005) using HRTEM. However, to the best of our  
320 knowledge, similar peaks on the 02, 11 band observed in the XRD pattern of the present  
321 specimen were not reported for disordered kaolinite. As the XRD simulation indicated, the  
322 origin of these peaks is of comparable proportion to  $t_1$  and  $t_2$ , and their frequent alternation ( $P_{11}$   
323  $\leq 0.5$ ). Hence, this is characteristic of the structure of prismatic halloysite. Although there are

324 not a significant number of successful HRTEM images, it is probable that the pseudo-mirror  
325 plane in the kaolinite layer being perpendicular to tube or prismatic axis, or the kaolinite layer  
326 rolling around the axis normal to the pseudo-mirror plane could represent other structural  
327 characteristics of halloysite. The origin of these features will be considered in the following  
328 paragraph (Fig. 9).

329 It seems natural and undisputed that initially halloysite grows as a hydrated form, halloysite  
330 ( $10\text{\AA}$ ), without hydrogen, bonds between the basal oxygen and hydroxyl across the interlayer. It  
331 is probable that the roll direction of the layer at this initial stage is the same as that found in the  
332 present specimen, about the normal of the pseudo-mirror plane in the kaolinite layer (Fig. 9a).  
333 Although we have no evidence, it is suspected that the dioctahedral 1:1 layer can bend more  
334 easily along this direction to relax the lateral misfit between the tetrahedral and octahedral  
335 sheets. Next, the water molecules are removed from the interlayer to form halloysite ( $7\text{\AA}$ ) (Fig.  
336 9b). At this stage, because the layers are not flat but cylindrical, and cannot realize the specific  
337 positional relation between adjacent layers all over the circumference, hydrogen bonding across  
338 the interlayer must be limited in a circuit. Although the number of successful HRTEM images  
339 was limited, Kogure et al. (2011) reported the stacking structure of cylindrical halloysite with a  
340 pseudo-mirror plane normal to the tube axis and no lateral stagger between adjacent layers, as  
341 illustrated in Figure 9b. In this configuration,  $t_1$  or  $t_2$  layer displacement, which forms stable  
342 hydrogen bonding in kaolinite, does not exist. During the probable long-term ripening of  
343 halloysite cylindrical forms are converted to prismatic ones with sector domains in which the  
344 kaolinite layers are flat. During this deformation, each layer moves relative to the adjacent layer  
345 by  $\pm b/3$  (or  $\pm b/6$ ) along the tube-axis (the black arrows in Figure 9b), to form an interlayer  
346 configuration with hydrogen bonding such as that in kaolinite (Fig. 9c). At this time, in order to  
347 minimize deformation of the external form of halloysite, the proportion of the displacements of  
348  $+ b/3$  ( $= \tau_+$ ) and  $- b/3$  ( $= \tau_-$ ) should be comparable and they should make frequent exchanges

349 along the stacking as can be seen via the black arrows in Figure 9b, which results in a  $P_{11}$  value  
350 less than 0.5.

## 351 **CONCLUDING REMARKS**

352 If our model for the stacking structure and its formation pathway is accurate, the structural  
353 difference between kaolinite and halloysite is related to their formation processes. For kaolinite,  
354 the crystals grow *via* layer by layer or spiral growth mechanisms (e.g., Kogure et al., 2010).  
355 During these growth processes, a new kaolinite layer and hydrogen bonding at the interlayer is  
356 formed simultaneously. Our previous HRTEM research (Kogure and Inoue, 2005; Kogure et al.,  
357 2010) reported that stacking structure regarded as “growth faults” (Bookin et al., 1989) or  
358 “enantiometric twins” (Kogure et al., 2010) which are formed with a condition probability  $P_{11}$   
359 more than 0.5 are common in kaolinite. This implies that, in kaolinite, stacking with the same  
360 layer displacement to the adjacent layer is thermodynamically more favorable than stacking  
361 with the different layer displacement. In prismatic halloysite, by contrast, our model suggests  
362 that the hydrogen bonds in the interlayer are generated after the formation of concentric or  
363 spiraled kaolinite layers. In this case, it is probable that the kinetic factor to minimize the  
364 deformation of the external shape to form hydrogen bonding promotes the tendency to form a  
365 regular interstratification of the different layer displacements.

## 366 **ACKNOWLEDGMENTS**

367 The authors are grateful to E. Fujii for the XRD measurements and preparation of the  
368 specimens for SEM and TEM. The authors also thank to M. Mellini and G.J. Churchman for  
369 their reviews of the manuscript and valuable suggestions. This work was supported by  
370 Grants-in-Aid for Scientific Research 24340133 and 21107005 from the Japan Society for the  
371 Promotion of Science (JSPS) and Ministry of Education, Culture, Sports, Science and  
372 Technology (MEXT), respectively. V.A. Drits acknowledges the Russian Foundation for Basic  
373 Research, Grant 2012-05-00381.

374

## REFERENCES CITED

- 375 Bailey, S.W. (1969) Polytypism of trioctahedral 1:1 layer silicates. *Clays and Clay Minerals*,  
376 **17**, 355-371.
- 377 Bailey, S.W. (1988) Polytypism of 1:1 layer silicates. In S.W. Bailey, Ed., *Hydrous*  
378 *Phyllosilicates (Exclusive of Micas)*, p. 9-27. Reviews in Mineralogy, Mineralogical  
379 Society of America, Washington, D.C.
- 380 Bailey, S.W. (1990) Halloysite – A critical assessment. In: Farmer, V.C and Tardy, Y. (Eds),  
381 *Proceedings of the 9<sup>th</sup> International Clay Conference, Strasbourg, France. Sci. Geol. Mem.*  
382 **86**: 89-98.
- 383 Bates, T.F. and Comer, J.J. (1958) Further observations on the morphology of chrysotile and  
384 halloysite. *Clays and Clay Minerals*, **6**, 237-248.
- 385 Bates, T.F., Hildebrand, F.A., and Swineford, A. (1950) Morphology and structure of endellite  
386 and halloysite. *American Mineralogist*, **35**, 463-484.
- 387 Bish, D.L. and von Dreele, R.B. (1989) Rietveld refinement of non-hydrogen atomic positions  
388 in kaolinite. *Clays and Clay Minerals*, **37**, 289-296.
- 389 Bookin, A.S., Drits, V.A., Plançon, A., and Tchoubar, C. (1989) Stacking faults in kaolin-group  
390 minerals in the light of real structural features. *Clays and Clay Minerals*, **37**, 297-307.
- 391 Chekin S.S., Samptoin N.D., and Finko, V.I. (1972) Formation of halloysite by weathering of  
392 oligoclase. *Izvestiya Akademii Nauk SSSR, Seriya geologicheskaya*, N11, 88-114 (in  
393 Russian).
- 394 Chekin, S.S., Samotoin, N.D., and Finko, V.I. (1976) Spiral-cylindrical shape and growth of  
395 halloysite crystallites. *Izvestiya Akademii Nauk SSSR, Seriya geologicheskaya*, N 6,  
396 111-124 (in Russian).



- 397 Chukhrov, F.V. and Zvyagin, B.B. (1966) Halloysite, a crystallochemically and  
398 mineralogically distinct species. Proceedings of the International Clay Conference, 1966 (L.  
399 Heller and A. Weiss, editors). Israel Program for Scientific Translation, Jerusalem, Israel.
- 400 Churchman, G.J., Davy, T.J., Aylmore, L.A.G, Gilkes, R.J. and Self, P.G. (1995)  
401 Characteristics of fine pores in some halloysites. Clay Minerals, 30, 89-98.
- 402 Cressey, B.A. and Zussman, J. (1976) Electron microscopic studies of serpentinites. Canadian  
403 Mineralogist, 14, 307-213.
- 404 Diamond, S. and Bloor, J.W. (1970) Globular cluster microstructure of endellite (hydrated  
405 halloysite) from Bedford, Indiana. Clays and Clay Minerals, 18, 309-312.
- 406 Dixon, J.B. and McKee, T.R. (1974) Internal and external morphology of tubular and  
407 spheroidal halloysite particles. Clays and Clay Minerals, 22, 127-137.
- 408 Drits, V.A. and Tchoubar, C. (1990) X Ray Diffraction by Disordered Lamellar Structures:  
409 Theory and Applications to Microdivided Silicates and Carbons. Berlin; Tokyo:  
410 Springer-Verlag, XVII+648 pp.
- 411 Āuroviĉ, S. and Weiss, Z. (1986) OD structures and polytypes. Bulletin de Minéralogie, 109,  
412 15-29.
- 413 Guggenheim, S., Adams, J.M., Bergaya, F., Brigatti, M.F., Drits, V.A., Fromoso, M.L.L.,  
414 Galan, E., Kogure, T., Stanjek, H., and Stucki, J.W. (2009) Nomenclature for stacking in  
415 phyllosilicates: Report of the Association Internationale pour l'Etude des Argiles (AIPEA)  
416 nomenclature committee for 2008, Clay Minerals, 44, 157-159.
- 417 Hayashida, M., Nomaguchi, T., Kimura, Y., and Takai, Y. (2007) Development of  
418 computer-assisted minimal dose system with beam blanker for TEM. Micron, 38, 505-512.
- 419 Honjo, G. and Mihara, K. (1954) A study of clay minerals by electron-diffraction diagrams due  
420 to individual crystallites. Acta Crystallographica, 7, 511-513.

- 421 Honjo, G., Kitamura, N., and Mihara, K. (1954) A study of clay minerals by means of  
422 single-crystal diagrams – the structure of tubular kaolin. *Clay Minerals Bulletin*, 2, 133-141.
- 423 Joussein, E., Petit, S., Churchman, J., Theng, B., Righi, D., and Delvaux, B. (2005) Halloysite  
424 clay minerals – a review. *Clay Minerals*, 40, 383-426.
- 425 Kilaas, R. (1998) Optimal and near-optimal filters in high-resolution electron microscopy.  
426 *Journal of Microscopy*, 190, 45-51.
- 427 Kirkman, J.H. (1981) Morphology and structure of halloysite in New Zealand tephras. *Clays  
428 and Clay Minerals*, 29, 1-9.
- 429 Kogure, T. and Inoue, A. (2005a) Determination of defect structures in kaolin minerals by  
430 high-resolution transmission electron microscopy (HRTEM). *American Mineralogist*, 90,  
431 85-89.
- 432 Kogure, T., Jige, M., Kameda, J., Yamagishi, A., Miyawaki, R., and Kitagawa, R. (2006a)  
433 Stacking structures in pyrophyllite revealed by high-resolution transmission electron  
434 microscopy (HRTEM). *American Mineralogist*, 91, 1293-1299.
- 435 Kogure, T., Kameda, J., Matsui T., and Miyawaki, R. (2006b) Stacking structure in disordered  
436 talc: interpretation of its X-ray diffraction pattern by using pattern simulation and  
437 high-resolution transmission electron microscopy. *American Mineralogist*, 91, 1363-1370.
- 438 Kogure, T., Eilers, P.H.C., and Ishizuka, K. (2008) Application of optimum HRTEM noise  
439 filters in mineralogy and related sciences. *Microscopy and Analysis*, 22, S11-S14.
- 440 Kogure, T., Johnston, C.T., Kogel, J.E. and Bish, D. (2010) Stacking disorder in a sedimentary  
441 kaolinite. *Clays and Clay Minerals*, 58, 63-72.
- 442 Kogure, T., Mori, K., Kimura, Y. and Takai, Y. (2011) Unraveling the stacking structure in  
443 tubular halloysite using a new TEM with computer-assisted minimal-dose system, *American  
444 Mineralogist*, 96, 1776-1780
- 445 Kohyama, N., Fukushima, K. and Fukami, A. (1978) Observation of hydrated form of tubular

- 446 halloysite by an electron microscope equipped with an environmental cell. *Clays and Clay*  
447 *Minerals*, 26, 25-40.
- 448 Marks, L.D. (1996) Wiener-filter enhancement of noisy HREM images. *Ultramicroscopy*, 62,  
449 43-52.
- 450 Neder, R.B., Burghammer, M., Grasl, T., Schulz, H., Bram, A., and Fiedler, S. (1999)  
451 Refinement of the kaolinite structure from single-crystal synchrotron data. *Clays and Clay*  
452 *Minerals*, 47, 487-494.
- 453 Plançon, A., Giese, R.F., Snyder, R., Drits, V.A., and Bookin, A.S. (1989) Stacking faults in the  
454 kaolin-group minerals - defect structures of kaolinite. *Clays and Clay Minerals*, 37, 203-210.
- 455 Robertson, I.D.M. and Eggleton, R.A. (1991) Weathering of granitic muscovite to kaolinite and  
456 halloysite and of plagioclase-derived kaolinite and halloysite. *Clays and Clay Minerals*, 39,  
457 113-126.
- 458 Singh, B. (1996) Why does halloysite roll? A new model. *Clays and Clay Minerals*, 44,  
459 191-196.
- 460 Singh, B. and Gilkes, R.J. (1992) An electron optical investigation of the alteration of kaolinite  
461 to halloysite. *Clays and Clay Minerals*, 40, 212-219.
- 462 Switch, O.R. and Young, R.A. (1983) Atomic position in highly ordered kaolinite. *Clays and*  
463 *Clay Minerals*, 31, 357-366.
- 464 Treacy, M.M.J., Newsam, J.M. and Deem, M.W. (1991) A general recursion method for  
465 calculating diffracted intensities from crystals containing planar faults. *Proceedings of the*  
466 *Royal Society of London A*, 433, 499-520.
- 467 Yada, K. and Wei, L. (1987) Polygonal microstructures of Povlen chrysotile observed by  
468 high-resolution electron microscopy. Abstracts Sixth Meeting of the European Clay Groups  
469 (Seville), 596-597.

- 470 Zvyagin, B.B. (1967) Electron diffraction analysis of clay mineral structures. New York:  
471 Plenum Press, XVI+364 pp.  
472  
473  
474

475 **Figure captions.**

476

477 **FIGURE 1.** XRD pattern of halloysite from Olkhon Island, Lake Baikal, Russia, taken at  
478 ambient conditions. The calculated patterns for kaolinite including all reflections and those  
479 with  $k = 3n$  are also plotted for comparison. Although they are broadened, most halloysite peaks  
480 are also identified in the calculated kaolinite pattern with  $k = 3n$ . However, those indicated with  
481 the arrows 1 and 2 on the 02, 11 band do not correspond to kaolinite.

482

483 **FIGURE 2.** FE-SEM images of the halloysite, showing polygonal cross-sections as indicated  
484 with the white arrows. The thin black arrows in (c) indicate that the surface of the cylindrical  
485 halloysite actually consists of elongated faces parallel to the cylindrical axis (see also Fig. 3c).

486

487 **FIGURE 3.** (a-c) Cross-sectional TEM images of several prismatic halloysites. (d) SAED  
488 pattern from the cross-section in (c). The two arrows indicate the ring corresponding to 7.2 and  
489 3.6 Å. Eighteen intensity maximums with an almost equal interval are observed in the rings,  
490 indicating that the prism is actually a regular octadecagon with eighteen sectors.

491

492 **FIGURE 4.** (left) Cross-section of a prismatic halloysite with three major sectors. One-half of  
493 the prism in the top-left disappeared during the ion-milling sample preparation. In addition, the  
494 crystals around the center and outer surface amorphized before TEM recording. (Right) The  
495 SAED pattern from the prism is comprised of a superposition of three reciprocal lattices. Three  
496 sectors in the image correspond to the numbered reciprocal lattices.

497

498

499 **FIGURE 5.** (a) Low-magnified image of a halloysite grain, showing the area of the HRTEM  
500 image in (b). Filtered HRTEM image in the rectangle in (a). The squares and triangles at the  
501 individual kaolinite layers indicate the types of contrast relating to the layer orientation defined  
502 in Kogure and Inoue (2005). The square bracket indicates a position where the stacking is  
503 regarded as the two-layer structure proposed by Chukhrov and Zvyagin (1966). (c) Filtered  
504 HRTEM image of the side of another prismatic grain, showing a similar contrast. The  
505 interlayers indicated with white arrows show no lateral shift between the contrasts at the two  
506 layers across the interlayer.

507

508 **FIGURE 6.** (a) Bright-field TEM image of a thick halloysite prism, showing the position and  
509 size of selected-area aperture used to acquire the SAED patterns in (b) and (c). (b) SAED  
510 pattern from the area selected by the larger aperture with “b” in (a). (c) SAED pattern from the  
511 side of the prism, using the smaller aperture with “c” in (a). Note that the intensity maximums  
512 indicated with the arrows in (b), which were previously considered the 020 peak, have almost  
513 disappeared in (c). The white bracket in (c) indicates the diffraction row and its range for which  
514 the intensity distribution is simulated in Figure 8.

515

516 **FIGURE 7.** Simulated XRD patterns for halloysite as a function of  $P_{11}$ , the condition  
517 probability to continue the layer displacement of  $t_1$  over  $t_1$ . The experimental pattern is also  
518 shown at the top.

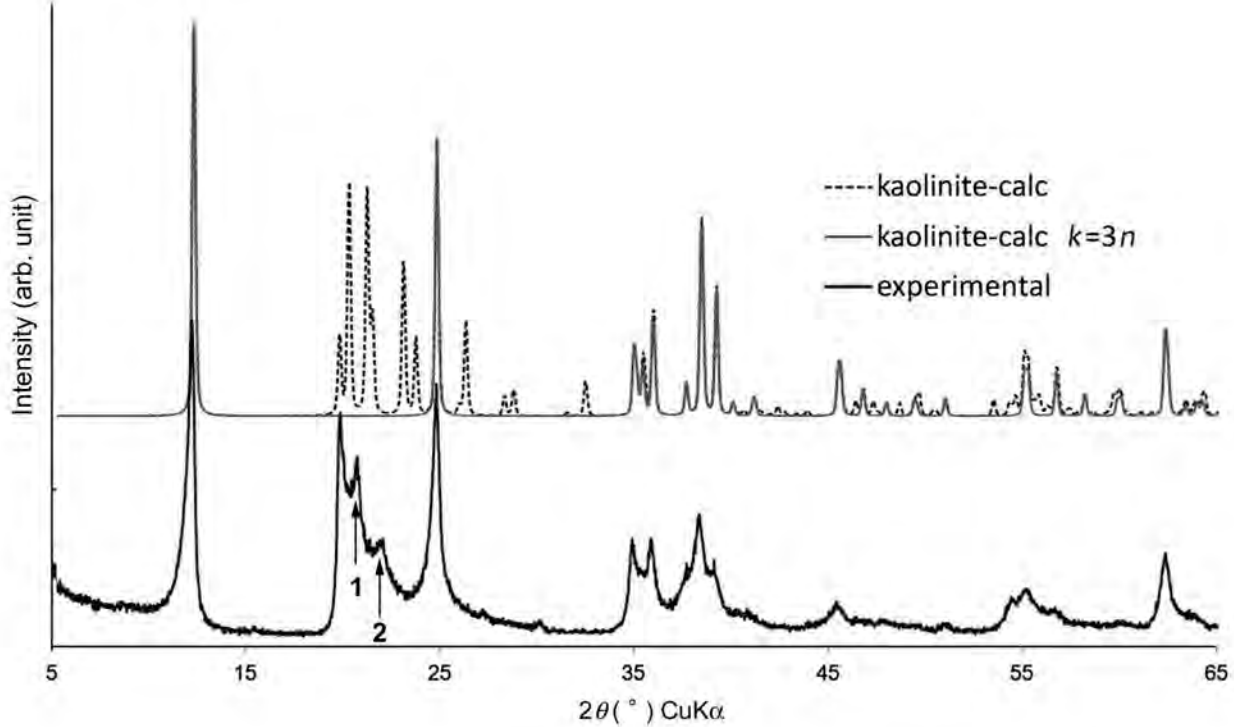
519

520

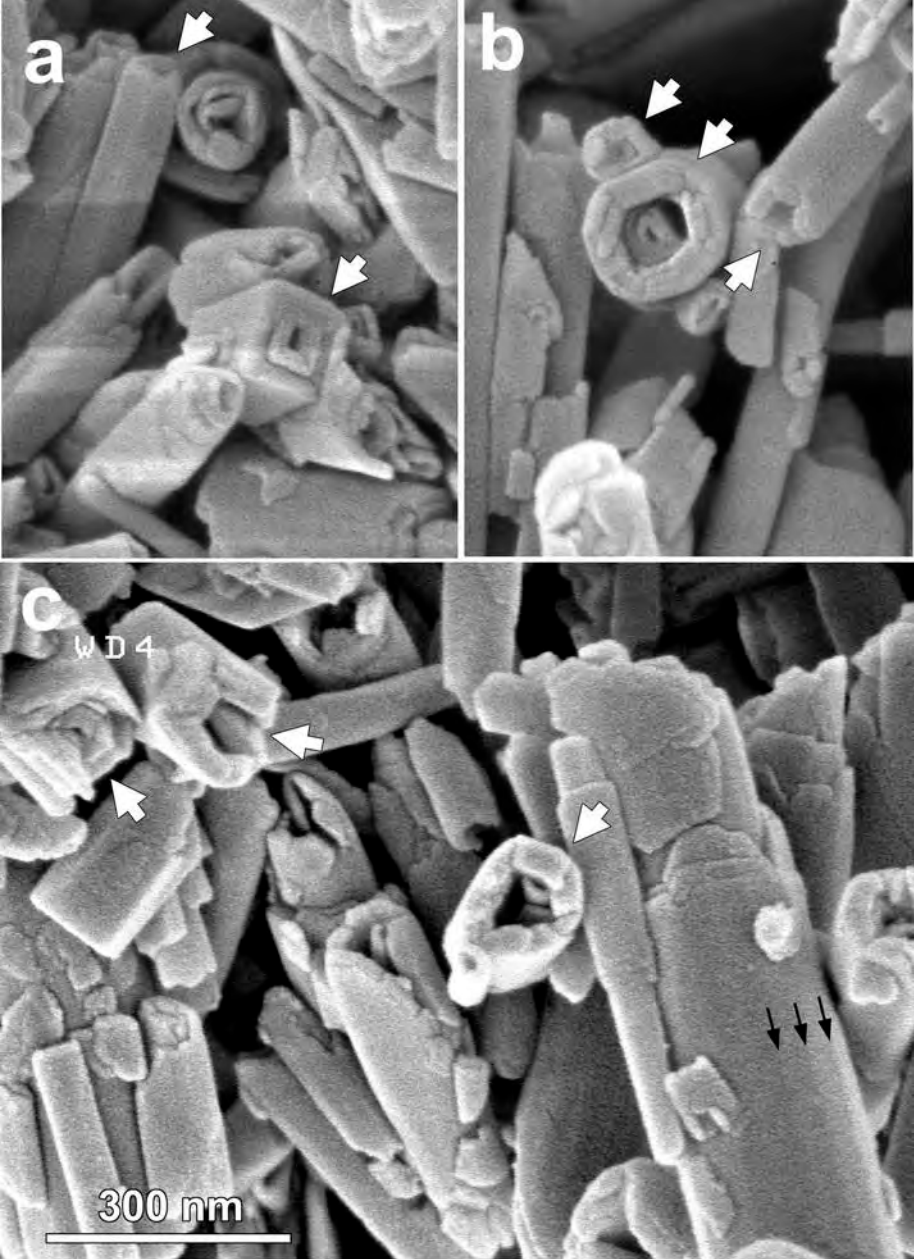
521 **FIGURE 8.** Calculated intensity distribution on the diffraction rows next to  $00l$  in the SAED  
522 pattern in Figure 6c. The unit for the horizontal axis is the value of  $l$  for two-layer periodicity.  
523 The experimental intensity distribution along the diffraction row indicated with the white  
524 bracket in Figure 6c is also shown at the top.

525

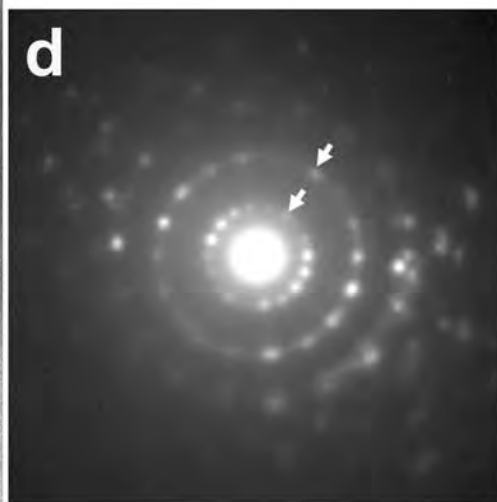
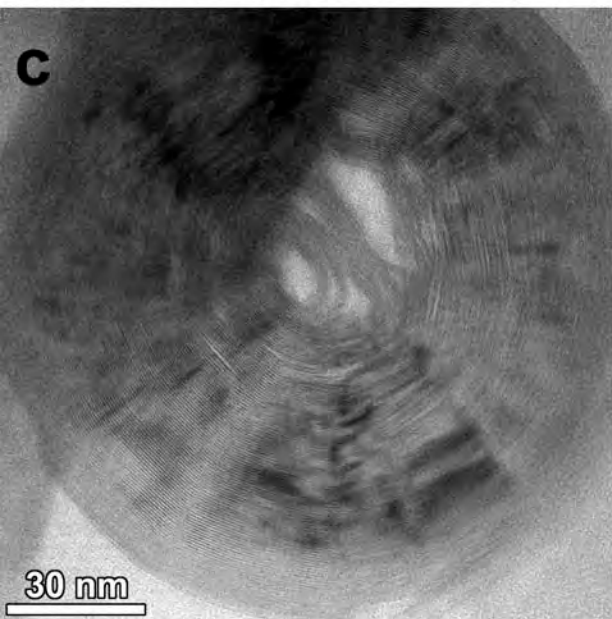
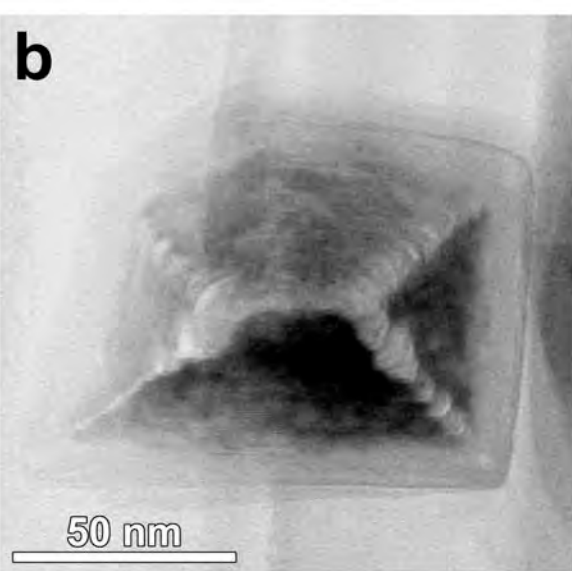
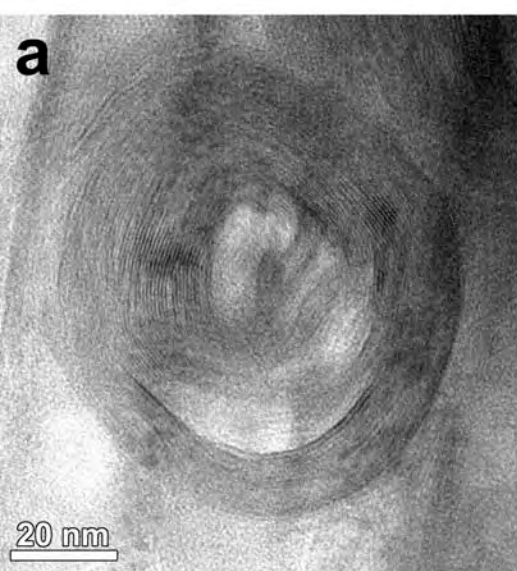
526 **FIGURE 9.** Schematic drawing for the ripening pathway of halloysite. See the text for details.

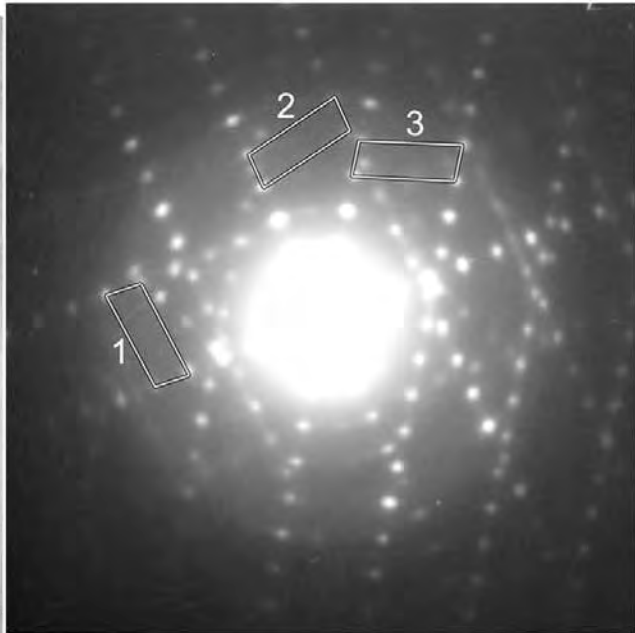


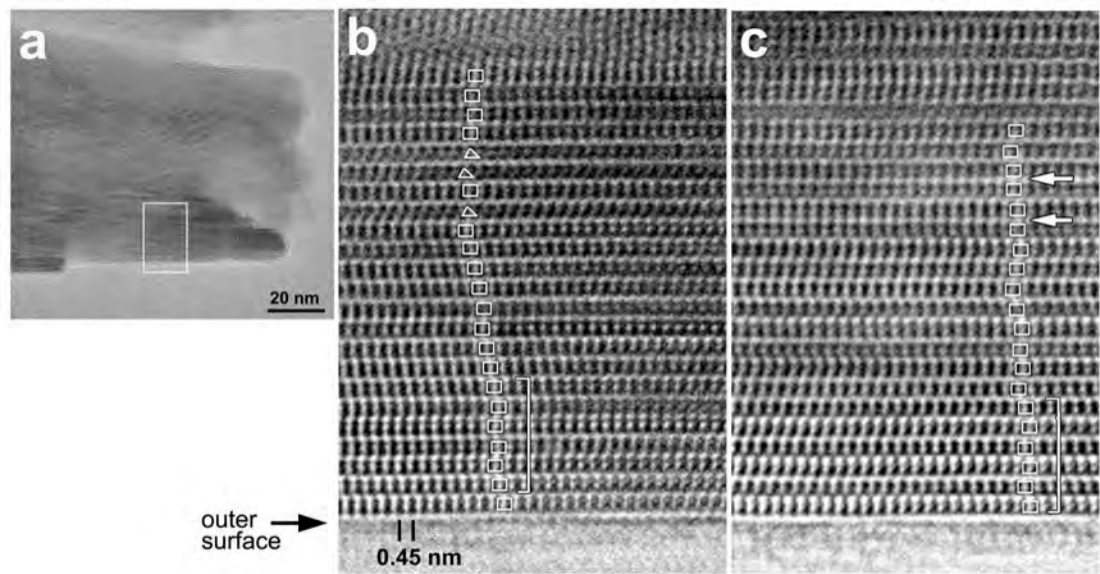




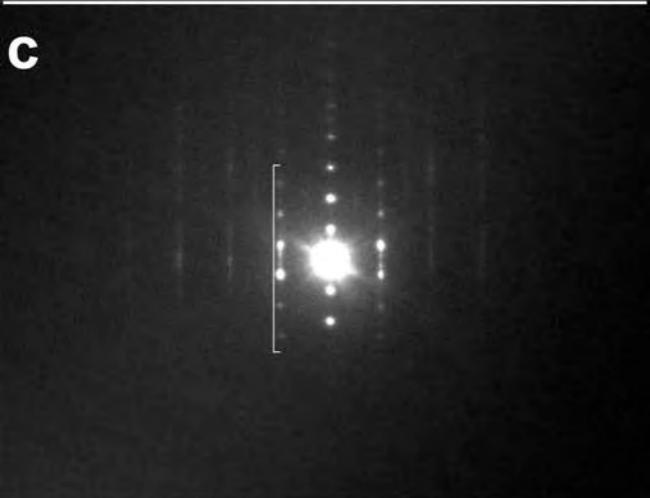
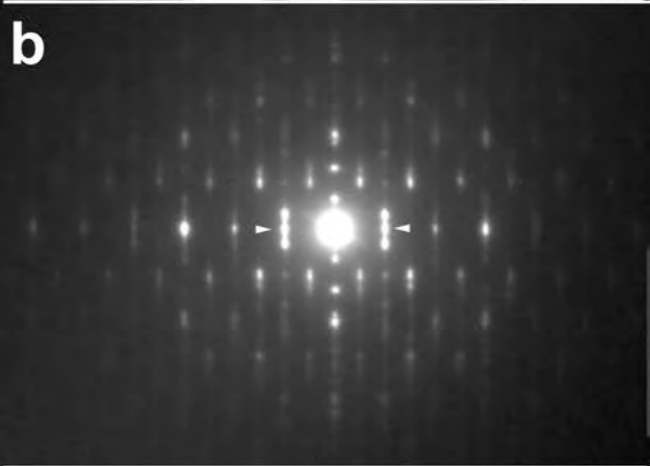
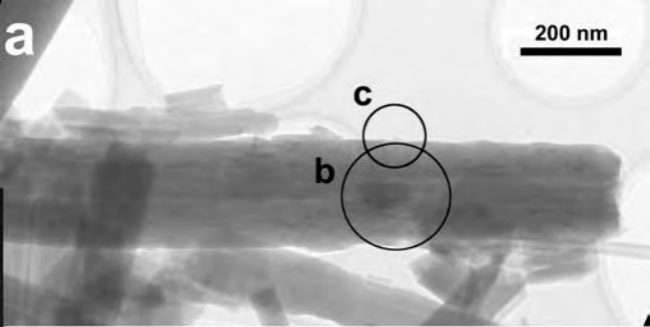
Kogure et al., Figure 2

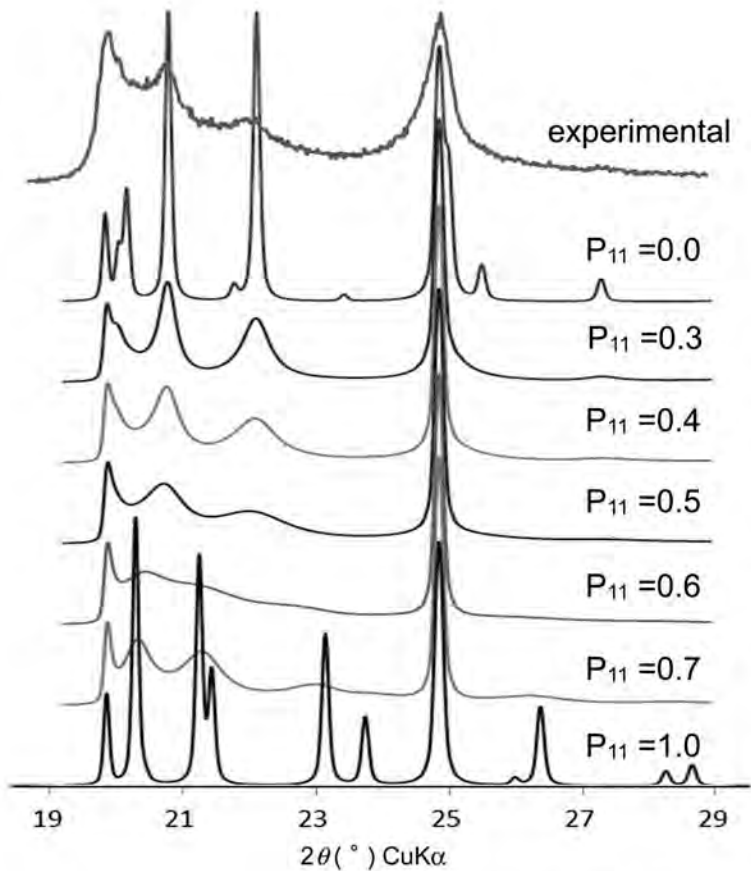




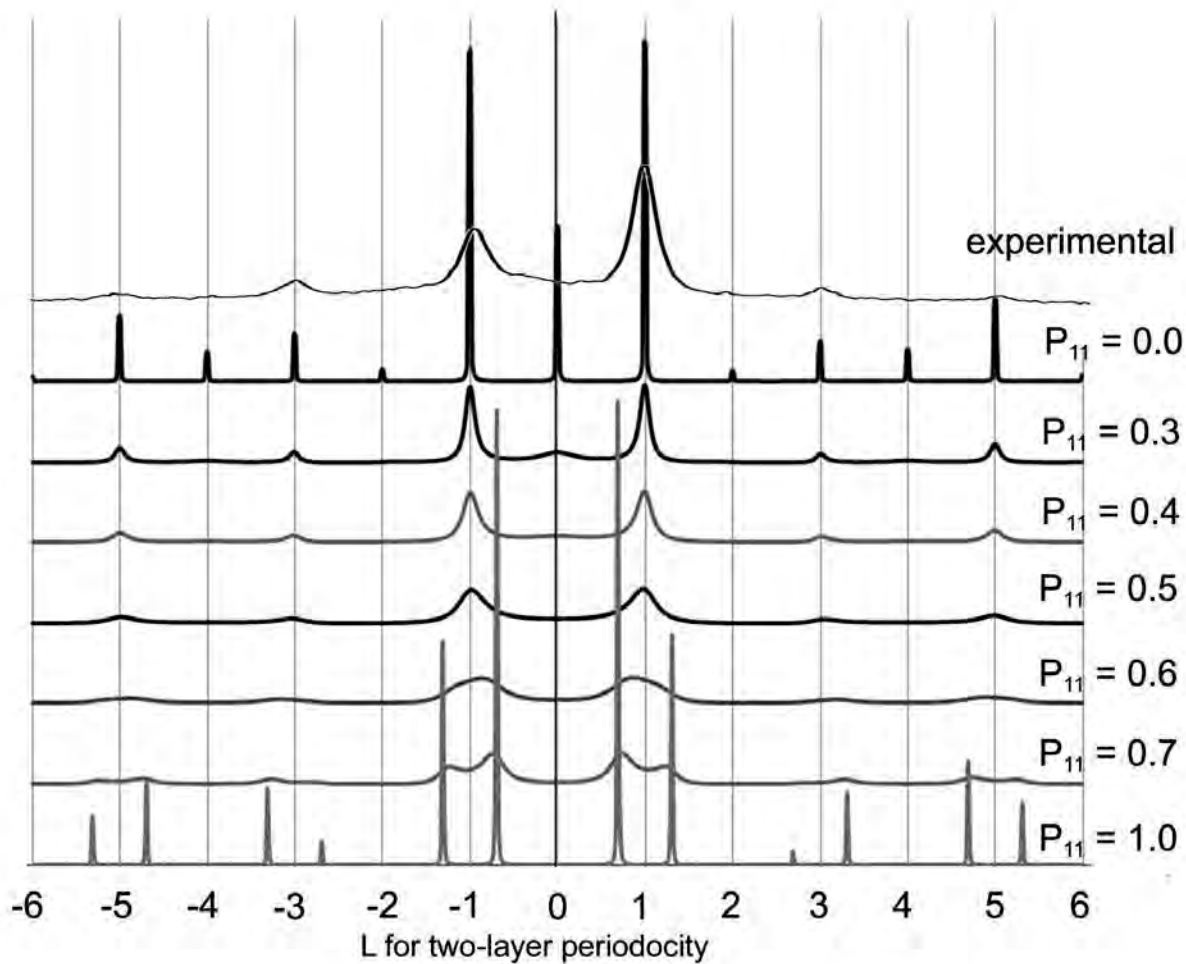


Kogure et al., Figure 5

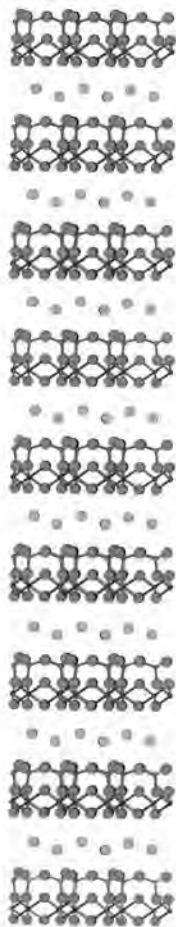
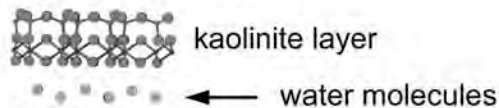




Kogure et al., Figure 7

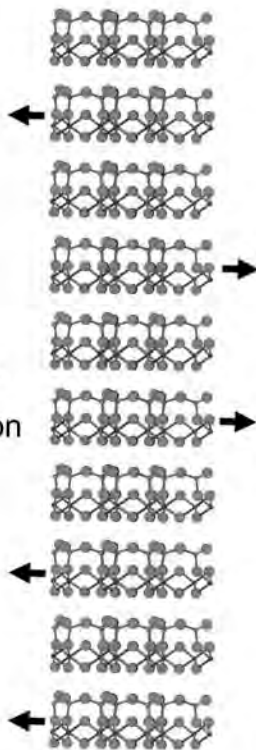


# halloysite (10 Å)



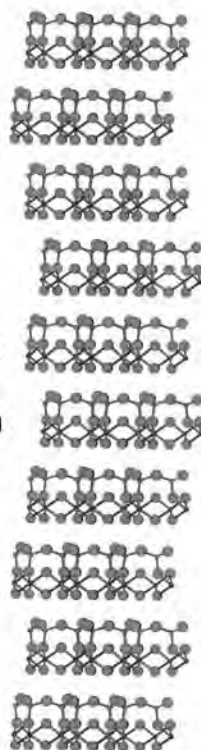
**a**

# halloysite (7 Å)



**b**

# prismatic halloysite (7 Å)



**c**

→ dehydration

flattening  
→  
hydrogen bonding

← tube axis →