

1 **REVISED 2**

2
3 **Investigation of hydrozincite structure by Infrared and Solid-State NMR**
4 **Spectroscopy**

5
6
7 Roberta Sanna^a, Giovanni De Giudici^{b*}, Andrea Mariano Scorciapino^a, Costantino Floris^a and
8 Mariano Casu^a

9
10 Dipartimento di Scienze Chimiche e Geologiche

11 ^a Complesso Universitario, S.S. 554, I-09042 Monserrato, Cagliari, Italy

12 ^b Via trentino 51, I-09127, Cagliari, Italy

13
14 *corresponding author gbiudic@unica.it

15
16
17 **Abstract**

18
19 To better understand lattice disorder in hydrozincite, natural hydrozincite samples and
20 synthetic analogues were investigated by XRD, FTIR, ¹³C MAS and ¹³C CPMAS NMR. The size
21 of coherent diffraction domains ranges between ~10 nm (Synth1) and ~30 nm (Synth2). FTIR peaks
22 from the antisymmetric CO₃²⁻ stretching ν₃ mode were observed at 1383 cm⁻¹ and 1515 cm⁻¹ in all
23 samples. Peaks due to OH vibrations were observed for all the samples at 3234, 3303 and 3363 cm⁻¹,
24 and were sharp only for the samples having larger crystal domains. The ¹³C MAS and CPMAS
25 spectra showed a main carbon signal at ~164 ppm in the Synth2 sample, while two main signals
26 were observed at ~164 ppm and ~168 ppm in the Synth1 sample. The intensity ratio of the latter
27 signals were found to be independent of contact time, in the investigated range between 0.2 and 30
28 ms.

29 In addition, ¹³C CP MAS dynamics indicates that the Synth1 sample has shorter T_{1ρ} with
30 respect to Synth2. This indicates a more effective process of spin diffusion of proton magnetization
31 in the former due to different structural properties of Synth1 and Synth2 samples. In addition,
32 chemical shift anisotropy analysis was attributed to a structural change in the carbonate group or
33 hydrogen bonding for Synth1 and Synth2. This was interpreted as a deviation from the ideal
34 structure generated by linear and planar lattice defects and / or grain boundaries.

35

36 Keywords: hydrozincite, carbonates, chemical shift anisotropy, nanocrystals

37

38

Introduction

39 Hydrozincite $[\text{Zn}_5(\text{CO}_3)_2(\text{OH})_6]$, is a mineral commonly occurring in the oxidized zones of
40 zinc ores. It is often found as masses or crusts associated with other secondary minerals such as
41 smithsonite, hemimorphite, and aurichalcite (Anthony et al. 2003). Hydrozincite can also precipitate
42 during processes controlled by cyanobacteria (Podda et al. 2000, De Giudici et al. 2009; Medas et
43 al. 2012a, b). More recently, hydrozincite was found to precipitate with hydrocerussite in
44 filamentous fungi (Słaba and Długoński 2011).

45 The crystal structure of hydrozincite was first studied in a highly crystalline sample by Ghose
46 in 1964 using single crystal X-ray diffraction. The structure is monoclinic with $a_0=13.62$, $b_0=6.30$,
47 $c_0=5.42$ Å, $\beta = 95^\circ 50'$. The space group is $C2/m$. Zn occupies two different octahedral sites and
48 one tetrahedral site. The octahedral and tetrahedral sites are in a 3 : 2 ratio. Zinc atoms in
49 tetrahedral coordination occur above holes that are present in the sheets of octahedral zinc atoms.
50 CO_3^{2-} groups bind these sheets parallel to (100). An oxygen atom for each CO_3 group is shared with
51 an octahedrally coordinated Zn, a second oxygen atom is sheared with a tetrahedrally coordinated
52 Zn, and the third oxygen atom is hydrogen bonded to three OH groups.

53 Ghose (1964) found also that synthetic crystals of hydrozincite have low crystallinity, and
54 speculated that replacement of some carbonate groups holding Tetrahedral-Octahedral-Tetrahedral
55 (T-O-T) groups together would result in stacking disorder. Jambor (1966) and Zabinsky (1966)
56 found that infrared spectra from different hydrozincite specimens differed significantly and showed
57 peak broadening. Such a difference was explained invoking contributions of different plane defects.
58 Conversely, Jambor (1964) and, more recently, Hales and Frost (2007) interpreted IR peak
59 broadening of hydrous zinc carbonate as due to the formation of polymorphs. Lattanzi et al. (2007)
60 and De Giudici et al. (2009) investigated the crystallinity of biologically produced hydrozincite in

61 comparison with synthetic and geologic reference standards. As pointed out by Medas et al. (2012
62 b), high precision synchrotron-based X-ray Powder Diffraction (XRPD) patterns of hydrozincites
63 collected in the Rio Naracauli, Sardinia, do not match any known structures of this mineral. In fact,
64 a fairly large difference was recorded for the a_0 cell parameter of the geologic hydrozincite ($13.59 \pm$
65 0.03 \AA) and the biologically produced Naracauli hydrozincite ($13.832 \pm 0.006 \text{ \AA}$).

66 Hydrozincite has been the subject of many studies because of its role in the corrosion of Zn-
67 rich materials (Stoffyn-Egli et al. 1998; Morales and Borges. 2006; Ghosh and Singh 2007), and its
68 involvement in controlling the mobility of zinc both in soils (Uygur and Rimmer 2000) and waters
69 (Mercy et al. 1998; Zuddas and Podda 2005; Wanty et al. 2013). The occurrence of hydrozincite in
70 Zn-contaminated calcareous soils was recently documented by Jacquat et al. (2008). Hydrozincite is
71 a precursor for the preparation of ZnO (Music et al. 2002 and references therein) and it was
72 recently found to be a by-product of Zn-mine-drainage remediation technologies (Perez-Lopez et al.
73 2011).

74 Our interest in hydrozincite originated from the discovery, some years ago, that this mineral
75 directly precipitates from the heavy-metal contaminated waters of Rio Naracauli, Sardinia, in
76 association with a biological photosynthetic community, composed of an algae (*Chlorella sp.*) and a
77 cyanobacteria (*Scytonema sp.*; Podda et al. 2000; Zuddas and Podda 2005). At Naracauli Creek, as
78 already shown in the literature, the precipitation of this biomineral results in the reduction of zinc
79 concentration from 348 to 2 mg/L within a few hundred meters downstream. In addition, Pb
80 concentration in the biominerals attains 6500 mg/kg, Cd concentration is 540 mg/kg, together with
81 many other heavy metals detected in high concentration in the Naracauli hydrozincite (see Table 1
82 in Podda et al. 2000). In a previous study, Transmission Electron Microscopy (TEM) and XRPD
83 analysis indicated a progressive decrease in the size of the particles comparing the biomineral to
84 both the synthetic and geologic samples (De Giudici et al. 2009). Moreover, De Giudici et al.
85 (2009) showed that ^{13}C Magic Angle Spinning (MAS) and Cross Polarization magic angle spinning
86 (CPMAS) NMR spectra have more than one peak for all the investigated samples, despite the fact

87 that carbon atoms occupy a single crystallographic position in the hydrozincite structure. The
88 additional peaks might reflect the presence of lattice defects typical of nanocrystals.

89 The disruption of the periodicity in the stacking order leads to measurable effects in the
90 diffraction pattern of the crystal (see Estevez-Rams et al. 2007, and reference therein), and can
91 affect both shape and chemical shift of IR and NMR peaks (see De Giudici et al. 2009, and
92 reference therein). The information about the stacking arrangement contained in the IR and NMR
93 signal, in many cases is not directly available, and can be further “hidden” by other effects such as
94 other structural imperfections (e.g. grain size, line defects, chemical bonding distortion) that
95 contribute to the broadening of the spectral profiles.

96 The scope of this work is to investigate structure of $[Zn_5(CO_3)_2(OH)_6]$. For this purpose we
97 investigated synthetic hydrozincite samples with different size of single coherent crystal domains by
98 using ^{13}C MAS, CPMAS and Fourier Transform Infrared Spectroscopy (FTIR). These methods are
99 particularly sensitive to small changes in C-O bonding environments, and so represent a suitable
100 method for investigating lattice defects in hydrozincite.

101

102

Materials and methods

103 Samples used in this study

104 Four specimens of hydrozincite were investigated in this study. Sample Synth1 was
105 synthesized according to protocol reported in Garcia-Clavel et al. (1989) slightly modified. Zinc
106 nitrate was added instead of calcium chloride and synthesis was carried out at 373 K by mixing
107 equal amounts of $(NH_4)_2CO_3$ (Aldrich) solution and $Zn(NO_3)_2 \cdot 6H_2O$ (Aldrich) solution. The
108 precipitate was obtained in about 30 min, then was filtered, washed with cold distilled water and
109 dried at room temperature. A second sample was synthesized according to the hydrothermal
110 procedure described by Music et al. (2002). The aqueous solution was made by adding
111 $Zn(NO_3) \cdot 6H_2O$ (Carlo Erba chemicals) with urea (Carlo Erba chemicals). Two autoclaving thermal

112 treatments were applied, first at 60° C for 6 h, and then at 160° C for 20 h. The precipitate obtained
113 was filtered, washed with cold distilled water and dried at room temperature. Hydrothermal
114 hydrozincite is hereafter referred to as Synth2.

115 Besides the two synthetic samples, two museum samples were investigated. One is from the
116 Sa Duchessa Mine (SD), and a second from the Malfidano Mine (MF) respectively, both located in
117 Sardinia, Italy. These samples come from supergene Zn mineralization (calamine), and are more
118 than one million years old (Boni et al. 2003).

119 **X-ray diffraction**

120 XRD was performed with a θ -2 θ conventional diffractometer (Phillips X'PERT MPD) with
121 Cu K α radiation (1.5418 Å). For XRD analysis, 200 mg of each sample was lightly ground in agate
122 mortar and was packed into the sample holder for X-ray diffraction analysis.

123 The particle size of sample was estimated from the full-width at half maximum (FWHM) of
124 the most intense peak according to the Debye-Scherrer formula. The exact FWHM and peak
125 position were obtained by a non linear fitting of the XRD spectra with Log Normal function using
126 the Origin 5.0 program from Microcal Software.

127

128 **Fourier Transformed Infrared Spectroscopy**

129 FTIR spectra were collected in the Mid region from 400 cm⁻¹ to 4000 cm⁻¹ at 4 cm⁻¹ resolution
130 over 64 averaged scans using a Bruker Equinox 55 spectrophotometer. For FTIR analysis of solids,
131 10 mg of solid sample was ground in an agate mortar with 30 mg of KBr and pressed.

132 The overlapping bands observed in the region 2500 – 4000 cm⁻¹ were decomposed using the
133 Origin 7 (Microcal) software package. A sum of Gaussian and Lorentzian functions was used
134 throughout with the minimum number of component bands used for the fitting process. Other
135 functions were also examined but the results were inferior. Line widths, intensities, and frequencies

136 were allowed to vary in the iteration process until reproducible results were obtained with squared
137 correlations (R^2 greater than 0.999).

138

139 **Nuclear Magnetic Resonance**

140 High-resolution NMR spectra were collected using a Varian Unity Inova spectrometer with a
141 9.39 T wide-bore Oxford magnet operating at a proton frequency of 399.952 MHz. ^{13}C MAS
142 experiments for solids were performed by packing ~100 mg of lightly ground (in agate mortar)
143 sample into a 4 mm Si_3N_4 rotor, with 7 kHz spin rate, 1200 s recycle time, 6.3 μs pulse length (90°)
144 and 50 kHz spectral window. ^{13}C CPMAS spectra were collected with contact times of 4 ms, and a
145 recycle time of 10 s. Hexamethylbenzene ($\text{CH}_3 = 17.4$ ppm) was used as external reference for ^{13}C
146 chemical shifts. Variable contact time experiments were performed by arraying 17 different contact
147 times from 0.2 to 30 ms. These values were selected in order to investigate a wide range of contact
148 times allowing us to collect any signal from proton that, in hydrozincite environment, can cross
149 polarize. An acquisition time of 50 ms, ^1H 90° pulse of 4 μs , B1 field strengths of about 36 kHz,
150 4k data points, a recycle delay of 10 s and 2000 transients were used to acquire all the spectra. The
151 acquisition of the ^{13}C MAS NMR spectra of the Synth1 and Synth2 samples was performed with a
152 probe with 7 mm Si_3N_4 rotors at a spinning rate of 6 kHz for sensitivity reasons. MAS experiments
153 were run with a recycle time of 35 min and 80 transients, 45° pulse lengths (4.5 μs) and 50 kHz
154 bandwidth. The ^{13}C spin-lattice relaxation time in the laboratory frame (T_1) was measured by the
155 Torchia (1978) pulse sequence. The ^{13}C Chemical Shift Anisotropies (CSA) (see Feng et al. 2006
156 for more details) were determined from analysis of the spinning side band intensities using the
157 WSOLIDS program (Eichele and Wasylishen, 2000).

158

159

Results and discussion

160 **XRD**

161 Figure 1 shows the XRD patterns of the samples Synth1 and Synth2. The sharp diffraction
162 lines of sample Synth2 can be indexed as a hydrozincite single phase with a monoclinic structure,
163 according to the experimental (PDF Card 19-1458) and calculated (PDF Card 72-1100) reference
164 patterns. XRD patterns of the Synth1 show broader diffraction lines suggesting a decrease in the
165 size of crystal-coherent domains. The average sizes calculated from the main reflection were ~30
166 nm, ~19 nm, ~14 nm and ~10 nm for Synth2, MF, SD and Synth1, respectively. It is worth noting
167 that this decrease in the size of crystal coherent domains might depend on the structural evolution of
168 the mineral and/or the crystallization process (Meldrun and Cölfen 2007; De Giudici et al. 2009).

169

170 **FTIR analysis**

171 The FTIR spectra of the samples MF, SD, Synth1 and Synth2 in the range 400 – 4000 cm^{-1}
172 are shown in Figure 2 and indicate the presence of carbonate groups. The carbonate stretching and
173 bending vibrations were attributed on the basis of the data reported in the literature (Music et al.
174 2002; Stoilova et al. 2002; Hales and Frost 2007). The four peaks at 1520–1390 cm^{-1} are ascribed
175 to the antisymmetric CO_3^{2-} stretching ν_3 mode. The peak at 1050 cm^{-1} is assigned to the ν_1
176 symmetric CO_3^{2-} stretching mode. The strong and sharp peaks at 834 and 705 cm^{-1} are assigned to
177 the ν_2 out-of-plane OCO bending mode and to the ν_4 antisymmetric OCO bending mode,
178 respectively.

179 The ν_1 , ν_2 , ν_3 bands in the Synth1 and SD are broader than observed for MF and Synth2
180 samples. This has been attributed to lattice defects in hydrozincite (Zabinsky 1966; Music et al.
181 2002). Music et al. (2002) reported that the splitting into ν_3 frequency of the two bands at 1512 and
182 1387 cm^{-1} is dependent on $\text{Zn}_5(\text{OH})_6(\text{CO}_3)_2$ synthesis.

183 It should be noted that, spectra of carbonate minerals such as calcite (1421 cm^{-1}), magnesite
184 (1439 cm^{-1}), dolomite (1420 cm^{-1}), and smithsonite (1392 cm^{-1}) are characterized by a single peak,
185 which corresponds to ν_3 degenerate antisymmetric stretching of an undistorted CO_3^{2-} group (White,

186 1974; Nakamoto, 1997). Indeed, distortion of this group eliminates the degeneracy and results in a
187 double peak at ~ 1435 and ~ 1515 cm^{-1} (Sato and Matsuda, 1969; Neumann and Epple 2007). This
188 has been observed in silicate glasses containing dissolved carbon dioxide (Fine and Stolper 1985),
189 ground diopside samples (Kalinkina et al. 2001) and other carbonate minerals such as vaterite (Sato
190 and Matsuda 1969), monohydrocalcite and aragonite (Neumann and Epple 2007).

191 In the range $2500\text{--}4000$ cm^{-1} all the samples exhibit a complex pattern of overlapped bands,
192 which can be ascribed to the stretching modes of the hydroxyl groups and water. As shown in
193 Figure 3 the FTIR spectra of the sample Synth2, MF and SD show three main peaks at 3234 , 3303
194 and 3363 cm^{-1} , that have been attributed to the stretching vibrations of the structural OH groups in
195 hydrozincite (Music et al. 2002). In contrast, the FTIR spectrum of the Synth1 sample shows a
196 broad band, instead of the three peaks with the shoulders as observed in the previous samples. The
197 shoulder at higher wavenumbers can be easily detected in Figure 3. In addition, peaks at 3234 , 3303
198 and 3363 cm^{-1} appear to progressively broaden, going from Synth2 to MF, and to the SD sample.
199 Broadening is even more pronounced in the Synth1 sample, where the three bands are not more
200 identifiable and a unique broad band is observed, shifted to higher wavenumbers. All the spectra
201 show shoulders at $3490\text{--}3520$ cm^{-1} and at $2900\text{--}3100$ cm^{-1} . The complex overlapping of OH
202 stretching bands can be resolved by the application of band component analysis. This approach has
203 been used in the literature for the simulation of infrared spectra of hydrozincite (Hales and Frost
204 2007). In order to fit the Synth1 spectrum, the position of the bands attributed to the stretching
205 vibrations of the structural OH groups found in the Synth2 spectrum were set as initial values before
206 the iteration procedure was started.

207 The FTIR spectra in the $2500\text{--}4000$ cm^{-1} region of Synth1 and Synth2 samples can be fitted
208 with 6 overlapping bands (Fig. 4). Bands at 3375 ± 2 , 3303 ± 1 , 3234 ± 1 cm^{-1} (line width 110 ± 25 ,
209 91 ± 5 and 65 ± 7 cm^{-1} respectively), are found in the Synth2 sample. In the Synth1 sample, $3376 \pm$
210 14 , 3304 ± 7 and 3239 ± 10 cm^{-1} (line width 130 ± 18 , 109 ± 24 and 106 ± 26 cm^{-1} respectively)
211 are found. Results from FTIR fitting does not provide evidence for changes in the band frequency

212 of OH molecular vibrations. However, it reveals a difference between the line width of the 3239
213 cm^{-1} ($65 \pm 7 \text{ cm}^{-1}$) and the 3234 cm^{-1} ($106 \pm 26 \text{ cm}^{-1}$) bands, while the others do not significantly
214 differ.

215 Shoulders observed in the 3510-3542 cm^{-1} region might be due to adsorbed and /or hydrogen
216 bonded water. Adsorbed water, indeed, should show a band at 3550 cm^{-1} . The band at 3473 cm^{-1}
217 was attributed to hydrogen bonded water molecules in hydrotalcites by Frost et al. (2003). Broad
218 bands near 3200-3100 cm^{-1} have been tentatively ascribed to adsorbed water in rosasite group
219 minerals and aurichalcite (Reddy and Frost 2007). Ghose (1964) indicated that hydrozincite always
220 contains some adsorbed water on the surface. Moreover, as suggested by Ghose (1964), part of the
221 water molecules could be structural and/or located in channels running parallel to the (001) plane,
222 thus, between the rows of carbonate groups. Therefore, we can speculate that some of the adsorbed
223 and /or hydrogen bonded water identified in the FTIR spectrum should be attributed to water
224 molecules in these hydrozincite channels.

225

226

227

228 ¹³C NMR analysis

229 The NMR technique is well known to be sensitive to crystal order and to the presence of
230 lattice defects such as stacking faults and/or polytypes. The chemical shifts provide atomic
231 selectivity for well-ordered systems and is influenced by the local environment, extending to only a
232 few spheres of coordination (Mehring 1983; Engelhardt and Michel 1987).

233 The ¹³C MAS spectra of the Synth2, MF and Synth1 hydrozincite samples show overlapping
234 signals in the 162-169 ppm range despite the fact that hydrozincite has only one crystallographically
235 independent site (Fig 5). All the spectra show a main carbon signal at ~164 ppm which is
236 characterized by ¹³C spin lattice relaxation time in the laboratory frame (T_1), $T_1 = 97 \pm 15 \text{ s}$ for
237 Synth1 sample, $T_1 = 5.1 \pm 0.7 \text{ s}$ for MF sample, $T_1 > 150 \text{ s}$ for Synth2 sample. In the Synth1

238 sample a second peak can be observed at ~168 ppm with ^{13}C T_1 (89 ± 27 s) not significantly
239 different from that measured at ~164 ppm.

240 The long ^{13}C T_1 of the carbonate carbon in the Synth1 and Synth2 samples is indicative of a
241 rigid atomic lattice, while the lower value of ^{13}C T_1 recorded for the MF sample is due to the
242 presence of paramagnetic impurities (De Giudici et al. 2009). All the spectra show shoulders that
243 decrease in intensity going from Synth1 to MF to Synth2 samples. Stacking faults or polytypism in
244 the structure of hydrozincite would cause a deviation from the ideal structure and would generate
245 small local distortions resulting in the slight shifts of the carbonate resonance. In addition,
246 nanocrystals such as synthetic hydrozincite have a huge surface area, where crystal truncation
247 produces dangling chemical bonds. In turn, this creates an excess of surface energy that decreases
248 because of aggregation of nanoparticles. For these reasons, De Giudici et al. (2009) ascribed the
249 observed additional peaks to the presence of lattice defects, namely, grain boundaries and stacking
250 modes, in agreement with HRTEM analysis. Figure 5 also shows the ^{13}C MAS spectra obtained
251 with cross polarization (CPMAS) collected at a contact time of 4 ms. No significant differences
252 were observed between MAS and CPMAS spectra.

253 In the ^1H - ^{13}C cross-polarization experiment, magnetization transfer to the carbon nucleus
254 depends upon the contact time period and is related to other structure-dependent parameters, such as
255 the distance between the two nuclei, the number of protons and the proton relaxation times in the
256 rotating frame (Mehring 1983; Slichter 1989). Only signals from carbon atoms having neighboring
257 hydrogen atom can be collected in the cross-polarization spectrum. Moreover, carbon having a
258 different environment and/or a different mobility are expected to show cross-polarization at
259 different contact times. Thus, the total intensity of the ^{13}C spectra should vary with contact time,
260 depending on how the different signal components behave with respect to contact time. However, in
261 a range between 0.2 and 30 ms, the features of the ^{13}C CPMAS spectra of the investigated samples
262 were found to be independent of contact time, even in the Synth1 sample, where a higher
263 concentration of lattice defects have been observed. Then, we argue that all carbon atoms of our

264 hydrozincite samples have neighboring hydrogen atoms. It should be noted that non-H-bonded
265 carbonate groups, if present, would have not been cross polarized at short contact times, but could
266 be cross polarized only at longer contact times.

267 Figures 6 and 7 show the ^{13}C CPMAS spectra of Synth2 and Synth1 samples acquired at two
268 different spinning rates. The spectrum of Synth2 (Fig. 6) shows five bands at 2300 spinning rate ,
269 while ten bands are observed at a 13 kHz spinning rate. The spectrum of the Synth1 (Fig. 7) shows
270 a more intricate pattern, characterized by five bands at a 2.3 kHz spinning rate, and each band
271 shows two separate peaks, corresponding to the two isotropic peaks at 164 and 168 ppm. At a 1.4
272 kHz spinning rate, seven bands are observed (Fig. 7), and again, each band is constituted by two
273 peaks.

274 The sideband patterns were simulated for each peak to obtain the principal values of the
275 chemical shift anisotropy (CSA) tensors in order to obtain information on the local structure and
276 motions of the carbonate group. The CSA for resonances in ^{13}C CPMAS spectra were determined
277 from the analysis of the Spinning Sidebands (SSB) intensities obtained at low spinning rate.

278 Least-square fitting of SSB intensities of the Synth2 sample yields the following CSA values for δ_{iso}
279 = 164 ppm: $\delta_{11} = 218\text{ppm}$, $\delta_{22} = 153\text{ ppm}$, and $\delta_{33} = 120\text{ ppm}$. The uncertainty of the principal
280 values, based on four repeated experiments at different spinning rates, was estimated to be about 2
281 ppm. These values deviate significantly from axial symmetry of carbonate group, in agreement
282 with the differences reported by Ghose (1964) for C-O bond distances in the carbonate group in
283 hydrozincite, C–O1 1.35 Å , C–O2 1.44 Å and for C–O3 1.27 Å.

284 For the Synth1 sample, CSA values were obtained from the SSB intensities of the two signals
285 at 164 and 168 ppm in the CPMAS spectrum recorded at 1.4 kHz. The CSA values of the signal at
286 $\delta_{\text{iso}} = 164\text{ ppm}$ were $\delta_{11} = 216\text{ ppm}$, $\delta_{22} = 154\text{ ppm}$, and $\delta_{33} = 122\text{ pmm}$. The signals at $\delta_{\text{iso}} = 168$
287 ppm has $\delta_{11} = 205\text{ ppm}$, $\delta_{22} = 182\text{ ppm}$, and $\delta_{33} = 115\text{ ppm}$.

288 Significant differences can be noted in the CSA values for the signal at 164 and 168 ppm in
289 the Synth1 spectrum. In particular δ_{22} differs by 28 ppm between the signal at $\delta_{\text{iso}} = 164$ and 168

290 ppm for Synth1. On the contrary, significant differences in CSA values for the signal at 164 ppm in
291 the Synth1 and Synth2 CPMAS spectra were not observed.

292 Previous studies on carbonate and carbonyl compounds indicate that the most shielded
293 principal axis of the CSA, δ_{33} , is usually perpendicular to the carbonate plane, while the less
294 shielded axis, δ_{11} , is placed along the direction defined by the two ether oxygen atoms. This puts the
295 intermediately shielded axis, δ_{22} , along the C=O double bond (Stueber et al. 2002; Robyr et al.
296 1998; Oas et al. 1987).

297 It is peculiar that, in our samples, the most affected parameter is the δ_{22} parallel to the C=O
298 bond, which, as reported by Ghose (1964), is involved in hydrogen bonding with the three OH
299 groups. A relationship between hydrogen bond and the δ_{22} parameter has been observed in previous
300 works on carboxyl groups in aminoacids and peptides (Ando et al. 1988, Gu et al 1994,
301 Gardiennet-Doucet et al. 2006). Gu et al. (1994) observed a strong correlation between the δ_{22}
302 values and the strength of hydrogen bonding. Specifically, they observed that the deshielding effect
303 of deprotonation leads to decreased shielding for the δ_{22} element for the protonated carboxyl groups
304 having strong hydrogen bonding. The reverse occurs for the deprotonated carboxylates.

305 The values of the CSA principal components might be explained by 1) changes in the
306 planarity of the carbonate group, 2) and/or changes in the C-O bond distances induced by deviation
307 from the ideal structure generated by linear and planar lattice defects, grain boundaries and stacking
308 faults, 3) change in H-bonding. In the cases 1 and 2, a variation also of δ_{11} and/or δ_{33} CSA tensors
309 should be expected. On the other hand, significant change in the hydrogen bond is not supported by
310 the analysis of FTIR peaks and the lack of contact-time-dependence observed in ^{13}C CPMAS
311 features of Synth1. Then, a change in H-bonding would be based only on the observed broadening
312 of the FTIR bands in the Synth1 sample. Thus, additional measurement of H position in
313 hydrozincite structure and H bond length are needed to definitely rule out or confirm that changes in
314 H-bonding are responsible for the large difference in δ_{22} CSA tensor element.

315 Additional information can be obtained by ^{13}C CPMAS dynamics spectroscopy (Stejskal and
316 Memory, 1994). In this technique, the intensity data are fitted as a function of the contact time
317 using the standard relationship (Mehring 1983; Slichter 1989):

$$318 \quad I(t) = I_0(1 - T_{\text{CH}}/T_{1\rho})^{-1}[\exp(-t/T_{1\rho}) - \exp(-t/T_{\text{CH}})] \quad (1)$$

319 where I_0 is the peak intensity, T_{CH} is the cross-polarization time between protons and carbons, and
320 $T_{1\rho}$ is the proton spin lattice relaxation time in the rotating frame, allowing the values for these two
321 characteristic times to be obtained. The least squares curve-fitting are shown in Figure 8, while
322 Table 1 reports the values obtained for T_{CH} and $T_{1\rho}$. The curve fitting was carried out using the
323 intensity ($I(t)$) of the peak at 164 ppm for the Synth1, MF and Synth2 samples. The intensity of the
324 168 ppm peak was also analyzed for the sample Synth1.

325 The carbonate signal at 164 ppm of the Synth2 and MF samples exhibits fast build-up of the
326 magnetization and a slow relaxation rate, while the Synth1 signal at 164 ppm is characterized by
327 slower build-up of the magnetization and a faster relaxation rate (Fig . 8). The $T_{1\rho}$ data of Synth2
328 sample show a value >100 ms, larger than in the Synth1 sample, 22.04 ms. The large $T_{1\rho}$ value
329 observed in the Synth2 sample is an indication of the rigid conformation and/or an ineffective spin
330 diffusion. However, the carbonate signal of the Synth2 sample is also characterized by large ^{13}C T_1
331 (>150 s) providing an indication of a highly rigid crystal lattice. This is supported also by the SSB
332 pattern observed for both samples at low spinning rate.

333 The Synth1 sample is also characterized by a large ^{13}C T_1 (97 ± 15 s) and SSB pattern
334 similar to that observed for the Synth2 sample. Thus, we argue that the lower $T_{1\rho}$ data measured in
335 the Synth1 sample must be attributed to a more effective spin diffusion of proton magnetization.

336 The T_{CH} value measured for the carbonate signal at 164 ppm is significantly higher for the
337 Synth1 sample, 1.84 ± 0.15 ms, than for the Synth2 and MF samples, 0.65 ± 0.09 ms. The time
338 constant T_{CH} depends on several parameters such as: the distance between the protons and carbon
339 nuclei, the magnitude of the ^1H - ^{13}C dipolar coupling (it decreases with the number of hydrogen

340 atoms bound to the carbon atom under observation and or the distance between proton and carbon
341 atoms), and the mobility of a structural fragment bearing the particular carbon atom. The efficiency
342 of CP transfer is also inversely dependent on mobility, thus, a decrease in T_{CH} values would reflect
343 an increase in the fragment rigidity. The slower build-up of the magnetization of the signal at 164
344 ppm in the Synth1 sample should be an indication of a decrease in the carbonate rigidity and/or the
345 structural fragment, which would be in agreement with the smaller T_1 ^{13}C value observed in this
346 sample. Moreover, the carbonate signal at 168 ppm in the Synth1 sample shows a $T_{1\rho}$ and T_{CH}
347 value (19.16 ± 0.29 ms and 2.12 ± 0.29 ms, respectively) not significantly different than those
348 measured at 164 ppm (22.04 ± 1.2 ms and 1.84 ± 0.15 ms, respectively). If the $T_{1\rho}$ parameter is
349 sensitive both to molecular motion and to proton spin diffusion, significantly different values would
350 indicate the presence of dynamic processes that are characterized by different rates but are not
351 averaged by proton spin diffusion, which can be possibly attributed to a certain degree of spatial
352 heterogeneity in the sample. The similar values observed suggest that proton spin diffusion in the
353 Synth1 sample is efficient enough to equilibrate dynamical processes, and that the nearest protons
354 are sufficiently close to produce adequate cross polarization of the carbon atoms. The proton spin
355 diffusion might be also mediated by the presence of water molecules in the channels.

356 A ^{13}C NMR study of 28 synthetic and natural carbonates including samples of mineralogical
357 and biogenic origin, showed only a small shift range, from 166.3 ppm for a synthetic magnesian
358 calcite to 169.9 ppm for aragonite (Papenguth et al. 1989). This relatively small chemical shift
359 variation reflects some change in the structure of the CO_3^{2-} group (Papenguth et al. 1989). On the
360 other hand this relatively narrow range is slightly extended to 163.8 ppm in the scapolite series,
361 where the carbonate group is more distorted than in the simple carbonates (Sherriff et al. 1987). A
362 similar chemical shift value is observed in hydrozincite, as shown, supporting the idea that the
363 carbonate group might be similarly distorted. However a tight correlation between carbonate
364 distortion and ^{13}C chemical shift cannot be invoked. In fact, many examples could violate a general

365 rule, for example the carbonate groups in vaterite have higher ^{13}C chemical shifts than observed for
366 calcite, but the carbonate group in this mineral is more distorted than in calcite.

367 Finally, the IR and NMR analysis provided deeper insight on the hydrozincite structure,
368 showing some structural difference in the mineral depending on the synthesis methods. (Ghose
369 1964; Jambor 1964; Zabyński, 1966; Hales and Frost 2007; De Giudici et al. 2009; Medas et al.
370 2012b). Specifically, the elongation of the a_0 lattice parameter by more than 10% (Lattanzi et al.
371 2007) could correspond also to distortion of carbonate groups. In fact, the carbonate groups hold
372 the T-O-T sheet structure together and consequently, a microstructural change of carbonate groups
373 in hydrozincite can be reflected in different layer sequences, in agreement with the observed
374 elongation in the a_0 cell parameter observed by Lattanzi et al. (2007) (see also Medas et al. 2012a).

375 NMR analysis indicates that the Synth1 sample shows shorter $T1\rho$ with respect to the better
376 crystalline Synth2 sample. This seems to be due to spin diffusion to a reservoir of restrictedly
377 mobile H. Such a change in the H spin diffusion processes might be due to different structural
378 properties of Synth1 and Synth2 samples. These are nanometric size, abundance of both line and
379 planar defects, fusion of nanocrystallites during aggregation processes, presence of chemical
380 dangling bonds, elongation of a_0 lattice parameters. While fairly crystalline samples like Synth2
381 show a structure similar to the hydrozincite defined by Ghose (1964), poorly crystalline samples
382 like Synth1 have a significantly different structure, likely a polymorph. Based on the observed
383 NMR features, biologically controlled $[\text{Zn}_5(\text{CO}_3)_2(\text{OH})_6]$ mineralization from Naracauli (De Giudici
384 et al. 2009) has the same structure as the poorly crystalline Synth1 sample.

385

386 **Acknowledgement**

387 This work was supported by funding from EU project UMBRELLA (grant 226870). Comments
388 from the Associate Editor (Brian Phillips) and criticism of three reviewers were appreciated and
389 useful to improve this work.

390

391

392 **References**

393 Ando, S., Ando, I., Shoji, A. and Ozaki, T. (1988) Intermolecular Hydrogen-Bonding Effect on ^{13}C
394 NMR Chemical Shifts of Glycine Residue Carbonyl Carbons of Peptides in the Solid State. Journal
395 of American Chemical Society, 110, 3380-3386.

396

397 Anthony, J. W., Bideaux, R. A., Bladh, K. W. and Nichols, M. C. (2003) Handbook of Mineralogy,
398 pp. 791, Mineral Data Publishing, Tucson, Arizona, USA.

399

400 Boni, M., Gilg, H.A., Aversa, G., and Balassone, G. (2003) The “calamine” of southwest Sardinia:
401 Geology, mineralogy and stable isotope geochemistry of supergene Zn mineralization. Economic
402 Geology, 98, 731–748.

403

404 De Giudici G., Podda F., Sanna R., Musu E., Tombolini R., Cannas C., Musinu A., and Casu M.,
405 (2009) Structural properties of biologically controlled hydrozincite: a HR-TEM and NMR
406 spectroscopic study. American Mineralogist, 94, 1698–1706.

407

408 Eichele, K. and Wasylishen, R.E. (2000) WSOLIDS. Dalhousie University, Halifax, Canada

409

410 Engelhardt, G., and Michel, D. (1987) High-Resolution Solid-State NMR of Silicates and Zeolites,
411 485 p. Wiley, New York.

412

413 Estevez-Rams, E., Penton Madrigal, A., Scardi, P., and Leoni, M., (2007) Powder diffraction
414 characterization of stacking disorder. Zeitschrift fur Kristallographie, Supplement 1 (26), 99-104

415

416 Feng J., Reeder, R.J., and Phillips, B.L., (2006) Observation of bicarbonate in calcite by NMR
417 spectroscopy. American Mineralogist, 91, 957–960.

418

419 Fine, G., and Stolper, E., (1985) Dissolved carbon dioxide in basaltic glasses: concentrations and
420 speciation. Earth Planetary Science Letter, 76, 263–278.

421

- 422 Frost, R. L. , Martens, W., Ding, Z., Klopogge, J. T., and Johnson, T. E. (2003) The role of water
423 in synthesised hydrotalcites of formula $Mg_xZn_{6-x}Cr_2(OH)_{16}(CO_3)_4 \cdot 4H_2O$ and $Ni_xCo_{6-x}Cr_2(OH)_{16}(CO_3)_4 \cdot 4H_2O$ an infrared spectroscopic study. *Spectrochimica Acta Part A*, 59, 291-302.
424
425
- 426 Garcia-Clavel, M., Martinez-Lope, M.J. and Casais-Alvarez, M.T. (1989) Thermoanalytical study
427 of the system $Pb^{2+}-Zn^{2+}$ coprecipitated as binary carbonates. *Thermochimica Acta*, 137, 177–187.
428
- 429 Gardiennet-Doucet, C., Henry, B., and Tekely, P., (2006) Probing the ionisation state of functional
430 groups by chemical shift tensor fingerprints. *Progress in Nuclear Magnetic Resonance Spectroscopy*
431 49, 129–149.
432
- 433 Ghose, S. (1964) The crystal structure of hydrozincite, $Zn_5(OH)_6(CO_3)_2$. *Acta Crystallographica*,
434 17, 1051–1057.
435
- 436 Ghosh, R., and Singh, D.D.N. (2007) Kinetics, mechanism and characterisation of passive film
437 formed on hot dip galvanized coating exposed in simulated concrete pore solution. *Surface and*
438 *Coating Technology*, 201, 7346-7359.
439
- 440 Gu, Z., Zambrano, R. and McDermott, A. (1994) Hydrogen bonding of carboxyl groups in solid-
441 state amino acids and peptides: comparison of carbon chemical shielding, Infrared Frequencies, and
442 Structures, *Journal of American Chemical Society*, 116, 6368-6372.
443
444
- 445 Hales, M.C., and Frost, R.L. (2007) Synthesis and vibrational spectroscopic characterization of
446 synthetic smithsonite and hydrozincite. *Polyhedron*, 26, 4955–4962.
447
- 448 Jacquat, O., Voegelin, A., Villard, A., Marcus, M.A. and Kretzschmar, R. (2008) Formation of Zn-
449 rich phyllosilicate, Zn-layered double hydroxide and hydrozincite in contaminated calcareous soils.
450 *Geochimica et Cosmochimica Acta*, 72, 5037–5054.
451
- 452 Jambor, J.L. (1964) Studies of basic copper and zinc carbonates; Part 1, Synthetic zinc carbonates
453 and their relationship to hydrozincite. *Canadian Mineralogist*, 8, 92–108.
454
- 455 Jambor, J.L. (1966) Natural and synthetic hydrozincites. *Canadian Mineralogist*, 8, 652–653.
456

- 457 Kalinkina, E.V., Kalinkin, A.M., Forsling, W., and Makarov, V.N. (2001) Sorption of atmospheric
458 carbon dioxide and structural changes of Ca and Mg silicate minerals during grinding I. Diopside.
459 International Journal of Mineral Processing, 61, 273–288.
460
- 461 Lattanzi, P, Meneghini, C, De Giudici, G, and Podda, F. (2007) Report of experiment 08-02-636.
462 ESRF, Grenoble.
463
- 464 Medas D., Cidu R., Lattanzi P., Podda F, Wanty R.B. and De Giudici G. (2012a) Hydrozincite
465 seasonal precipitation at Naracauli (Sardinia – Italy): Hydrochemical factors and morphological
466 features of the biomineralization process. Applied Geochemistry, 27, 1814-1820.
467
- 468 Medas D., Cidu R., Lattanzi P., Podda F and De Giudici G. (2012b) Natural Biomineralization in
469 the Contaminated Sediment-Water System at the Ingurtosu Abandoned Mine, in E. Kothe and A.
470 Varma (eds.), Bio-Geo Interactions in Metal-Contaminated Soils, Soil Biology 31, p 113-130,
471 Springer-Verlag Berlin Heidelberg.
472
- 473 Mehring, M. (1983) Principles of High Resolution NMR in Solids, 2nd edition, 342 p. Springer-
474 Verlag, New York.
475
- 476 Meldrum, F., and Cölfen, H. (2007) Controlling mineral morphologies and structures in biological
477 and synthetic systems. Chemical Reviews, 108, 4332–4432.
478
- 479 Mercy, M.A., Rock, P.A., Casey, W.H., and Mokarram, M.M. (1998) Gibbs energies of formation
480 for hydrocerussite $[\text{Pb}(\text{OH})_2 \cdot (\text{PbCO}_3)_2(\text{s})]$ and hydrozincite $\{[\text{Zn}(\text{OH})_2]_3 \cdot (\text{ZnCO}_3)_{2(\text{s})}\}$ at 298 K and
481 1 bar from electrochemical cell measurements. American Mineralogist, 83, 739–745.
482
- 483 Morales, J.F. and Borges, J. (2006) Atmospheric corrosion in subtropical areas: XRD and
484 electrochemical study of zinc atmospheric corrosion products in the province of Santa Cruz de
485 Tenerife (Canary Islands, Spain). Corrosion Science, 48, 361-371.
486
- 487 Music, S., Popovic, S., Maljkovic, M. and Dragcevic, D. (2002) Influence of synthesis procedure
488 on the formation and properties of zinc oxide. Journal of Alloys and Compounds, 347, 324–332.
489

- 490 Nakamoto, K. (1997) Infrared and raman spectra of inorganic and coordination compounds: Part A.
491 theory and applications in inorganic chemistry. Wiley and Sons, pp 387, New York.
492
- 493 Neumann, M. and Epple, M. (2007) Monohydrocalcite and Its Relationship to Hydrated
494 Amorphous Calcium Carbonate in Biominerals. European Journal of Inorganic Chemistry, 2007,
495 1953–1957.
496
- 497 Papenguth, H.W., Kirkpatrick, R.J., Montez, B. and Sandberg, P.A. (1989) ^{13}C MAS NMR
498 spectroscopy of inorganic and biogenic carbonate. American Mineralogist, 74, 1152-1158.
499
- 500 Perez-Lopez, R., Macías, F., Caraballo, M.A., Nieto, J.M., Roman-Ross, G., Tucoulou, R. and
501 Ayora, C. (2011) Mineralogy and geochemistry of Zn-rich mine-drainage precipitates from an
502 MgO passive treatment system by synchrotron-based X-ray analysis. Environmental Science and
503 Technology, 45, 7826–7833.
504
- 505 Podda, F., Zuddas, P., Minacci, A., Pepi, M. and Baldi, F. (2000) Heavy metal coprecipitation with
506 hydrozincite $[\text{Zn}_5(\text{CO}_3)_2(\text{OH})_6]$ from mine waters caused by photosynthetic microorganisms.
507 Applied and Environmental Microbiology, 66, 5092–5098.
508
- 509 Oas, T.G., Hartzell, C.J., McMahon, T.J., Drobny, G.P., and Dahlquist, F.W. (1987) The carbonyl
510 ^{13}C chemical shift tensors of five peptides determined from 5N dipole-coupled chemical shift
511 powder patterns. Journal of American Chemical Society, 109, 5956-5962.
512
513
- 514 Reddy, B.J., and Frost, R.L. (2007) Near infrared spectroscopy of aurichalcite
515 $(\text{Zn,Cu}^{2+})_5(\text{CO}_3)_2(\text{OH})_6$. Journal Near Infrared Spectroscopy, 15, 113-121.
516
- 517 Robyr, P., Utz, M., Gan, Z., Scheurer, C., Tomaselli, M., Suter, U.W. and Ernst R.R. (1998)
518 Orientation of the Chemical Shielding Anisotropy Tensor of the Carbonate Carbon in Diphenyl
519 Carbonate and Its Consequences for NMR Studies on Polycarbonate. Macromolecules 31, 5818-
520 5822.
521
- 522 Sato, M. and Matsuda, S. (1969) Structure of vaterite and infrared spectra, Zeitschrift Fur
523 Kristallographie, 129, S. 405-410.

524

525 Sherriff B.L., Grundy, H.D. and Hartman, J.S. (1987) Occupancy of T sites in the scapolites
526 series: a multinuclear NMR study using Magic-Angle-Spinning. *Canadian Mineralogist*, 25, 717-
527 730.

528

529 Słaba, M. and Długoński, J. (2011) Efficient Zn^{2+} and Pb^{2+} uptake by filamentous fungus
530 *Paecilomyces marquandii* with engagement of metal hydrocarbonates precipitation. *International*
531 *Biodeterioration and Biodegradation*, 65, 964 - 960.

532

533 Slichter, C.P. (1989) *Principles of Magnetic Resonance*, 3rd edition, 655 p. Springer Verlag, New
534 York.

535

536 Stejskal, E.O. and Memory, J.D. (1994) *High-resolution NMR in the solid state*. Oxford University
537 Press, pp 192, New York, NY.

538

539 Stoffyn-Egli, P., Buckley, D. E. and Clyburne, J. A. C. (1998) Corrosion of brass in a marine
540 environment: mineral products and their relationship to variable oxidation and reduction conditions.
541 *Applied Geochemistry*, 13, 643-650.

542

543 Stoilova, D., Koleva, V. and Vassileva, V. (2002) Infrared study of some synthetic phases of
544 malachite ($Cu_2(OH)_2CO_3$)–hydrozincite ($Zn_5(OH)_6(CO_3)_2$) series. *Spectrochimica Acta Part A*, 58,
545 2051–2059.

546

547 Stueber, D., Orendt, A.M., Facelli, J.C., Parry, R.W. and David Grant, D.M. (2002) Carbonates,
548 thiocarbonates, and the corresponding monoalkyl derivatives: III.1 The ^{13}C chemical shift tensors in
549 potassium carbonate, bicarbonate and related monomethyl derivatives. *Solid State Nuclear*
550 *Magnetic Resonance*, 22, 29–49.

551

552 Torchia, D. A. (1978) The measurement of proton-enhanced carbon-13 T_1 values by a method
553 which suppresses artifacts. *Journal of Magnetic Resonance*, 30, 613-616.

554

555 Uygur, V. and Rimmer, D.L. (2000) Reactions of zinc with iron-oxide coated calcite surfaces at
556 alkaline pH. *European Journal of Soil Science*, 51, 511–516.

557 Wanty, R.B., Podda, F., De Giudici, G., Cidu, R. and Lattanzi, P. (2013). Zinc isotope and
558 transition-element dynamics accompanying hydrozincite biomineralization in the Rio Naracauli,
559 Sardinia, Italy. *Chemical Geology*, 337–338 1–10.

560

561

562

563 White, W.B. (1974). The carbonate minerals. In: Farmer, W.C. Ed., *The infrared spectra of*
564 *minerals*. Mineralogical Society, London, pp. 227–284.

565

566 Yang, J., Guan, Y., Verhoeven, M.W.G.M., Santen, R.A. van, Li, C., and Hensen, E.J.M. (2009).
567 Basic metal carbonate supported gold nanoparticles: enhanced performance in aerobic alcohol
568 oxidation. *Green Chemistry*, 11, 322–325.

569

570 Zabinsky, W. (1966) The problem of stacking-order in natural hydrozincite. *Canadian Mineralogist*,
571 8, 649–652.

572

573 Zuddas, P. and Podda, F. (2005) Variations in physico-chemical properties of water associated with
574 bio-precipitation of hydrozincite $[Zn_5(CO_3)_2(OH)_6]$ in the waters of Rio Naracauli, Sardinia (Italy).
575 *Applied Geochemistry*, 20, 507–517.

576

577

578

579

580

581

582

583 Table 1 - Cross polarization parameters for Hydrozincite Synth1, Synth2 and Malfidano samples

584

	Synth1 (168 ppm)	Synth1 (164 ppm)	MF (164 ppm)	Synth2 (164 ppm)
T_{CH} (ms)	2.1±0.3	1.8±0.2	0.6±0.1	0.6±0.2
T_{1ρ} (ms)	19.2±2.9	22.0±1.2	>100	>100

585

586

587

588

589 Captions

590

591 Figure 1 – XRD patterns of synthetic (Synth1 and Synth2) samples.

592

593

594 Figure 2 - FTIR spectra of Synth1, SD, MF and Synth2 samples

595

596

597 Figure 3 - FTIR spectra of Synth1, SD, MF and Synth2 samples in the region 2500 – 4000 cm^{-1}

598

599 Figure 4 – FTIR spectra of Synth1 and Synth2 samples in the region 2500 – 4000 cm^{-1} .

600 Experimental signals are decomposed into individual Gaussians.

601

602 Figure 5 – ^{13}C MAS and $^{13}\text{C}/^1\text{H}$ CPMAS spectra of the Synth1 and Synth2 samples.

603

604 Figure 6. Experimental (a, b) slow spinning ^{13}C CPMAS spectra of carbonate group of Synth2
605 sample and the best-fitting (c). The spinning frequencies were a) 2300 rpm, b) 1400 rpm. The
606 assigned center bands are indicated.

607

608 Figure 7. Experimental (a, b) slow spinning ^{13}C CPMAS spectra of carbonate group of Synth1
609 sample and the best-fitting (c). The spinning frequencies were a) 2300 rpm, b) 1370 rpm. The
610 assigned center bands are indicated.

611

612 Figure 8 - Variable contact time ^{13}C CPMAS and non linear least-square fitting for Synth2 at 164
613 ppm, triangles; Malfidano at 164 ppm, rhombus; Synth1 at 164 ppm, squares; Synth1 at 168 ppm,
614 circles.

615

616

Fig. 1

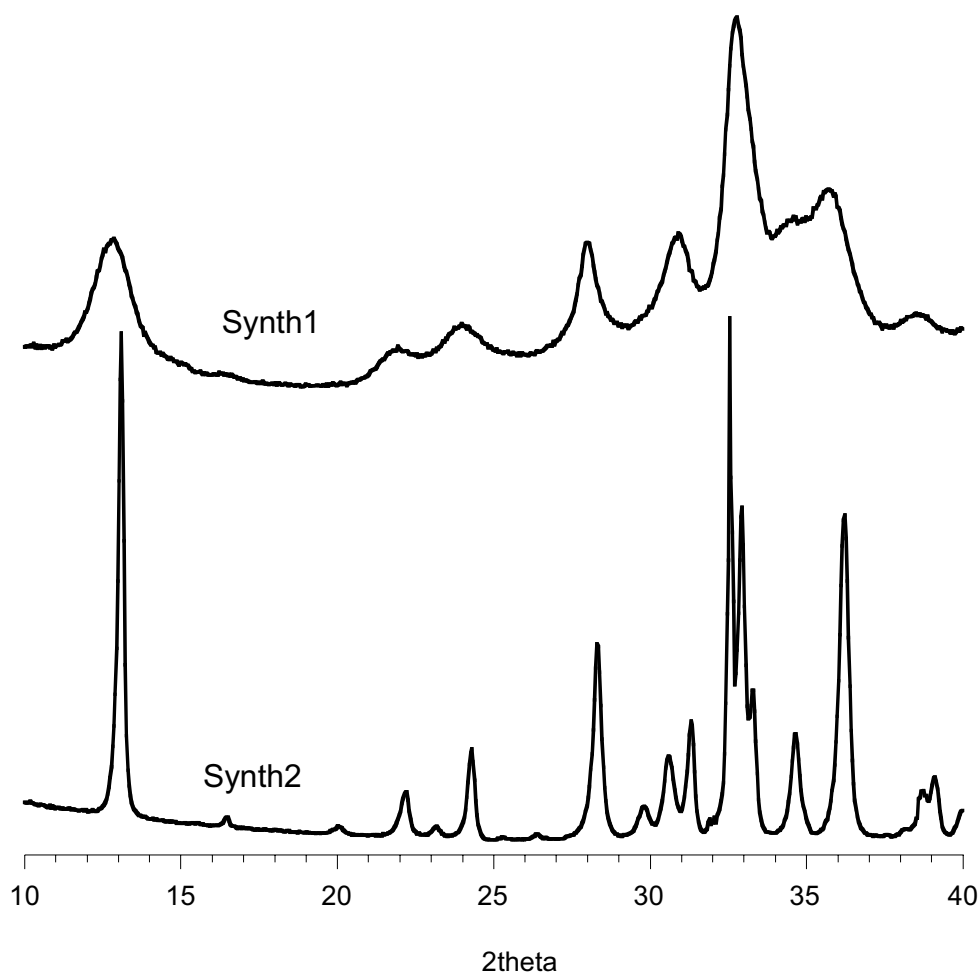


Fig. 2

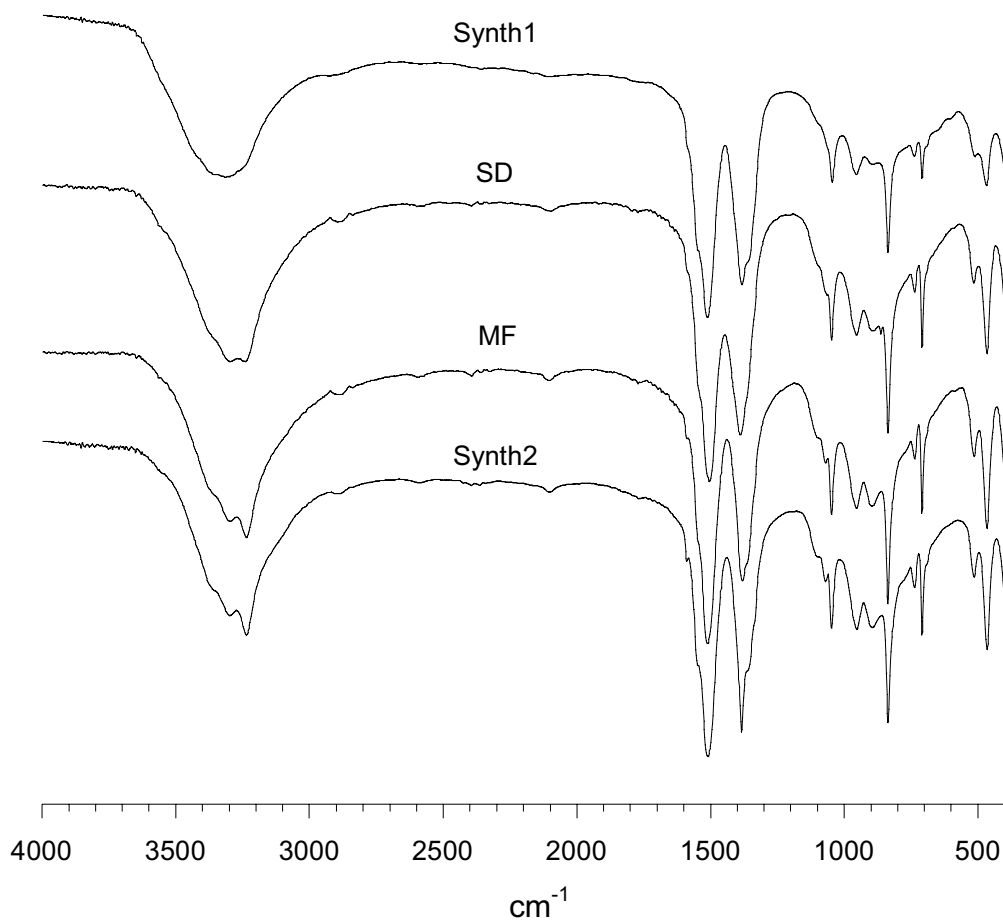


Fig. 3

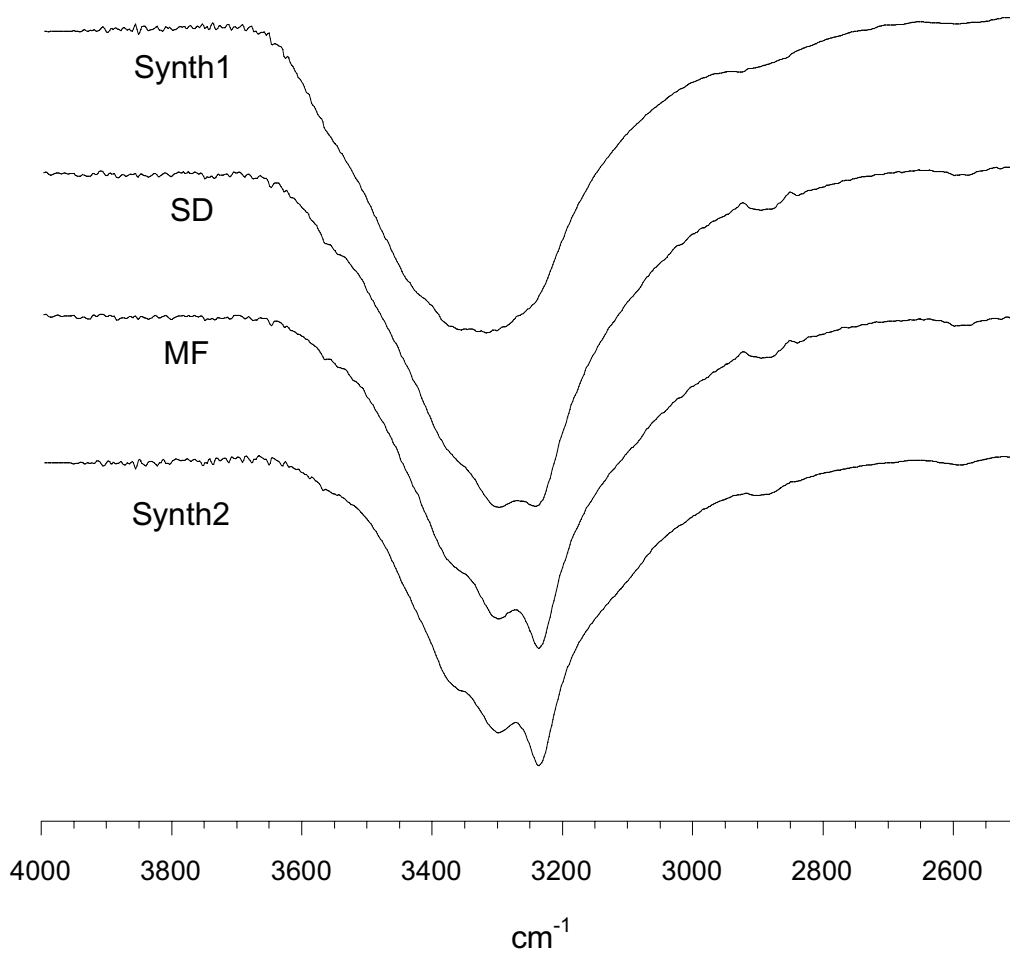


Fig. 4

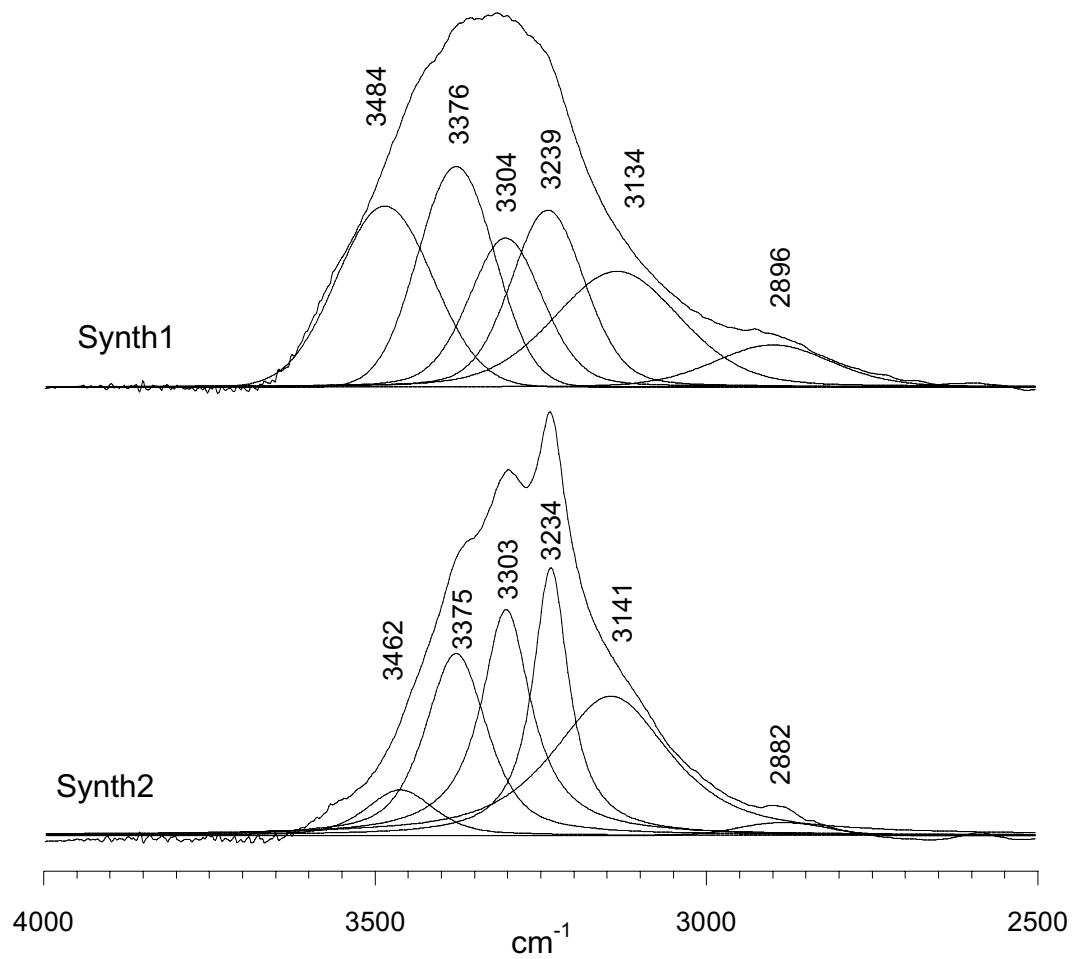


Fig. 5

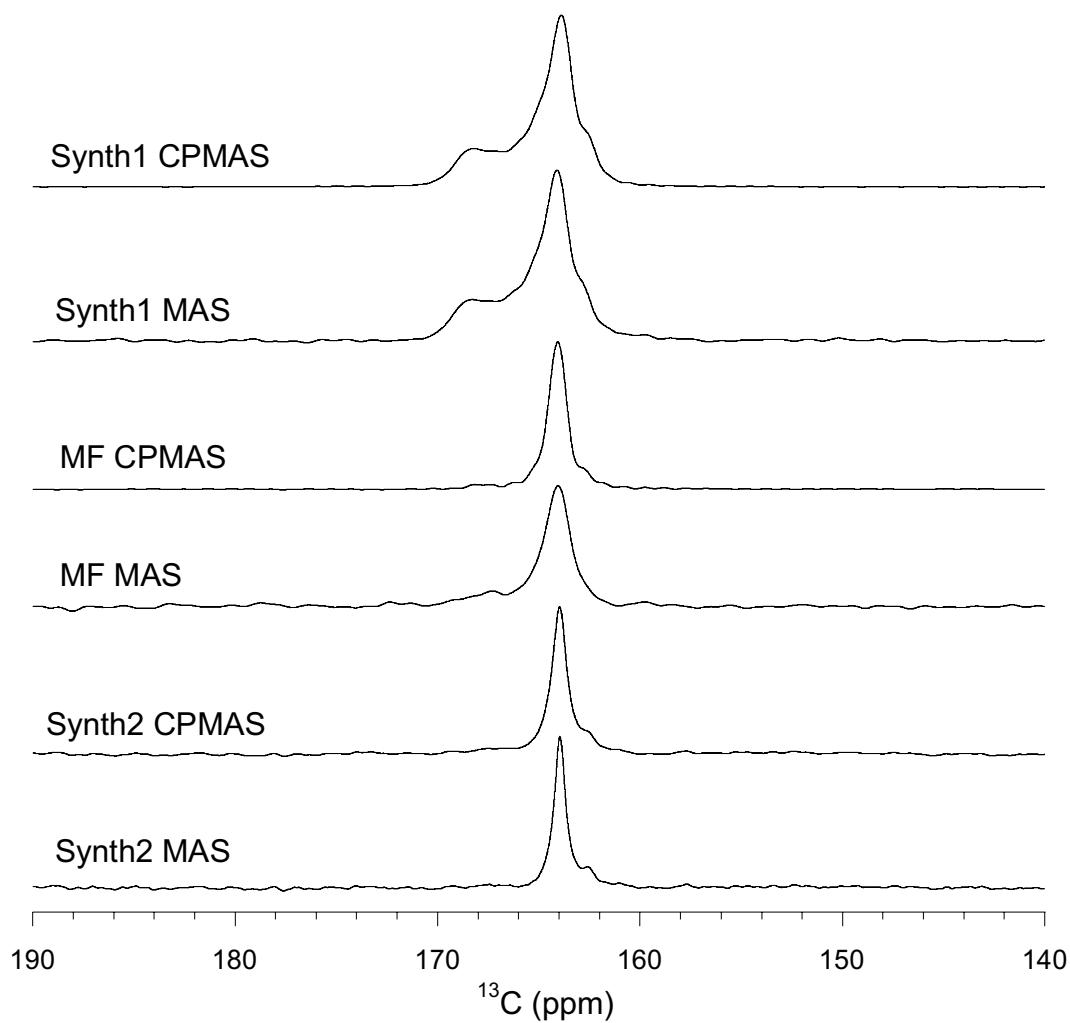


Fig. 6

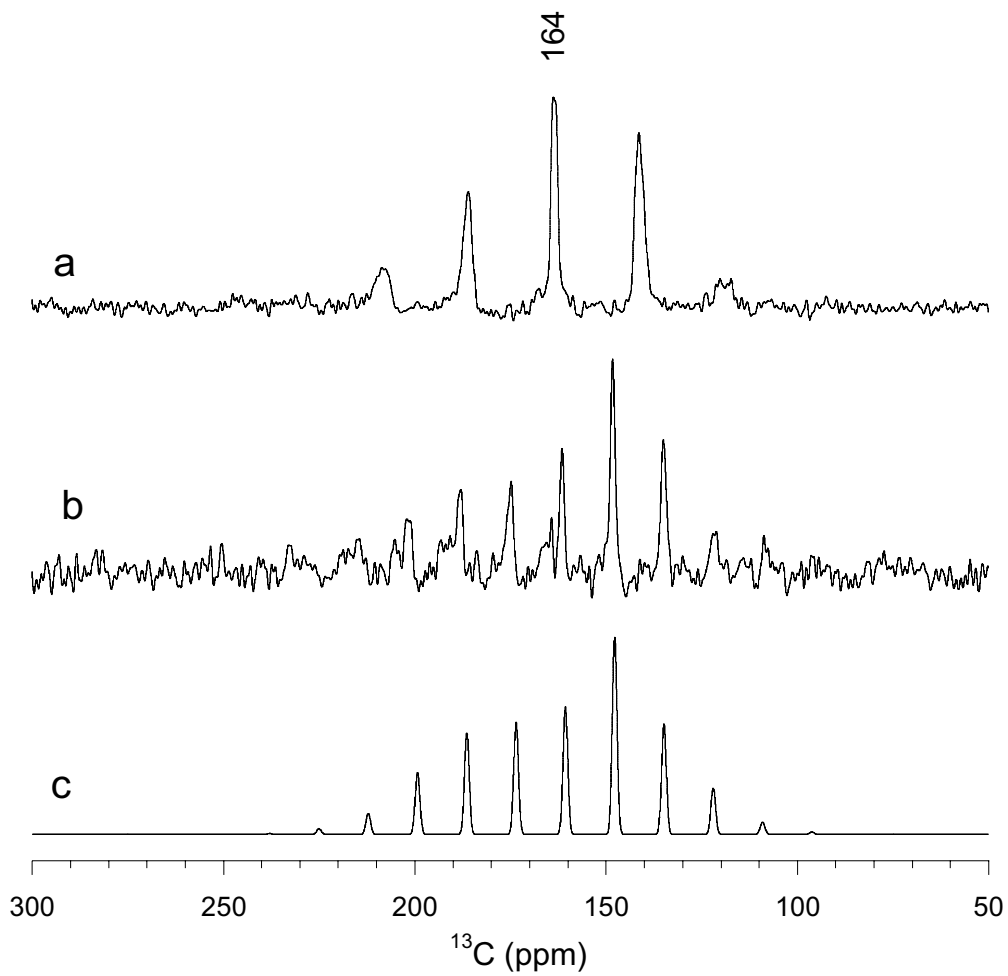


Fig. 7

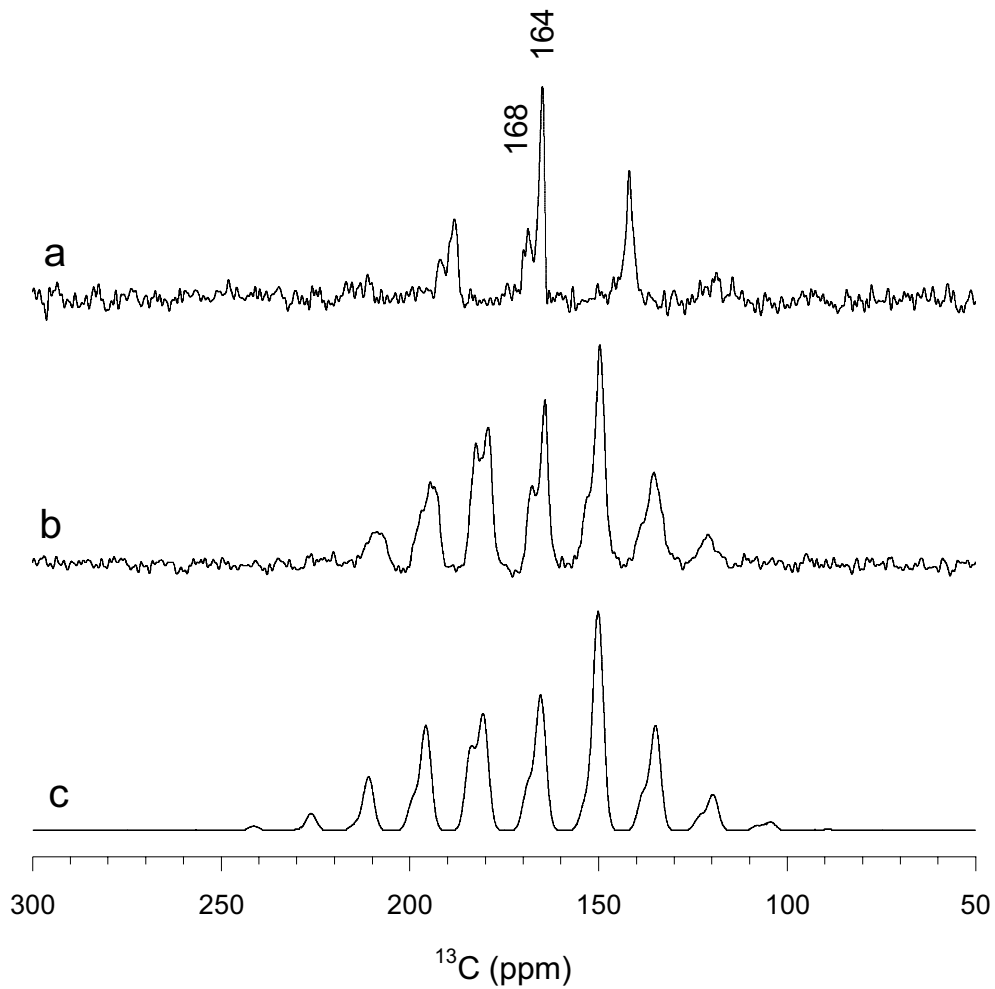


Fig. 8

