A Thermodynamic Model for Silica and Aluminum in **Alkaline Solutions with High Ionic Strength at Elevated** Temperatures up to 100 °C : Applications to Zeolites, **Revision 1** Yongliang Xiong^{*} Sandia National Laboratories Carlsbad Programs Group 4100 National Parks Highway Carlsbad, New Mexico 88220, USA

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37 Abstract

38 In this study, a thermodynamic model for silica and aluminum in high ionic strength solutions at elevated temperatures up to 100 °C is constructed. Pitzer equations 39 40 are utilized for the thermodynamic model construction. This model is valid up to ionic 41 strengths of ~24 m in NaOH solutions with silicate concentrations up to ~1.5 m. The 42 speciation of silica (including monomers and polymers) and aluminum at elevated 43 temperatures is taken into account. Also, the equilibrium constants for silicic acid and its polymer species (H₄SiO₄, H₅Si₂O₇⁻, H₄Si₂O₇²⁻, and H₅Si₃O₇³⁻) at elevated temperatures 44 up to 100 °C, are obtained based on theoretical calculations. Using this thermodynamic 45 46 model, thermodynamic properties, including equilibrium constants, and respective 47 reaction enthalpies are obtained for sodium silicates, zeolite A, and the amorphous form 48 of zeolite A, based on solubility experiments at elevated temperatures. The equilibrium 49 constants for zeolite A and amorphous precursor of zeolite A regarding the following 50 reactions up to 100 °C,

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53 and

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56 can be expressed as follows

57

- $\log K_I = \frac{7963 \pm 327}{T} 16.46 \pm 0.96$ (III)
- 59 and

60
$$\log K_{II} = \frac{12971 \pm 160}{T} - 30.80 \pm 0.50$$
 (IV)

61

62 where *T* is temperature in Kevin.

63 The enthalpy of formation from elements, Gibbs free energy of formation from 64 elements, and standard entropy derived for zeolite A and the amorphous form of zeolite A with the chemical formulas mentioned above at 25 $^{\circ}$ C and 1 bar are -2738 ± 5 65 kJ mol⁻¹, -2541 ± 2 kJ mol⁻¹, 373 ± 10 J K⁻¹ mol⁻¹; and -2642 ± 3 kJ mol⁻¹, -2527 ± 2 66 kJ mol⁻¹, and 648 \pm 10 J K⁻¹ mol⁻¹, respectively. The enthalpy of formation from 67 68 elements for zeolite A derived in this study based on solubility experiments in 69 hydrothermal solutions agrees well with those obtained by calorimetric measurements 70 and by theoretical calculations.

71

73 **1. Introduction**

74

75 Zeolites have numerous applications ranging from their usage as absorbents and catalysts to wasteforms for radioactive iodine (¹²⁹I) and for contaminated electrolytes 76 77 from electrorefinery of used nuclear fuel (e.g., Sheppard et al., 2006). The formation of 78 zeolites occurs under alkaline conditions in solutions typically with high ionic strengths. 79 For instance, in the hydrothermal synthesis of zeolite A, zeolite X and mordenite, NaOH 80 solutions ranging from 0.6 m to 2 m were frequently used (e.g., Čižmek et al., 1991a; 81 1991b; 1992; Šefčik and McCormick, 1997a). It has also been noted that in high-level nuclear waste processing, concentrated waste liquor is produced by evaporation at 82 83 elevated temperatures up to 140 °C, and the resulting concentrated waste liquor can have 84 NaOH concentrations up to $\sim 7 \text{ m}$ (Addai-Mensah et al., 2004). In laboratory experiments 85 designed to simulate the concentrated waste liquor from high-level nuclear waste processing, the precipitation of zeolite A and the amorphous form zeolite A from the 86 87 simulated concentrated waste liquor has been observed (Addai-Mensah et al., 2004). 88 Finally, at the Hanford site in Washington State, the leakage of the alkaline solutions with 89 high ionic strengths contained in the waste tanks into sediments resulted in formation of 90 cancrinite and sodalite, which belong to the zeolite family, in the sediment below a waste 91 tank (Chorover et al., 2003; Mashal et al., 2004; Serne et al., 2002; Zhao et al., 2004). 92 While there are experimental data on zeolites (see Table 1), the formation of

93 zeolites at elevated temperatures is not well understood owing to the lack of reliable 94 thermodynamic model(s).

95

simultaneously the behavior of both silica and aluminum, the two most rock-forming

Furthermore, there is the current inability to address

96 elements, in high ionic strength solutions at elevated temperatures. The high ionic 97 strength nature of solutions and high concentrations of silica are the factors contributing 98 to the difficulties associated with the development of a thermodynamic model at elevated 99 temperatures for zeolites. For instance, when zeolite A is formed, the concentration of 100 silica can reach up to 0.12 m (Sefčik and McCormick, 1997). The nature of high ionic 101 strength solutions requires that the activity coefficient model be valid to high ionic 102 strengths. The high concentrations of silica require that polymers of silica species be 103 considered, as polymerized silica species might become important when silica 104 concentrations are higher than 0.01 m (e.g., Felmy et al., 2001).

105 A thermodynamic model applicable to concentrated solutions at elevated 106 temperatures is therefore needed in order to predict the formation of zeolites including 107 zeolite A, amorphous precursor of zeolite A, cancrinite and sodalite, at elevated 108 In addition, such a model would provide valuable guidance in the temperatures. 109 synthesis of zeolites. A thermodynamic model would also provide a better understanding 110 of interactions of waste solutions with the sediments, like the interactions observed in the 111 Hanford studies mentioned above. Consequently, the objective of this study is to develop 112 a thermodynamic model for aluminum and silica species at elevated temperatures valid to 113 high ionic strengths. Using this model, equilibrium constants for sodium silicates, 114 zeolite A, and amorphous precursor of zeolite A can be retrieved from hydrothermal 115 experiments. Applications of the model to other zeolite species will be presented in the 116 future.

117

119 2. The Thermodynamic Model

120

In this study, the standard state for a solid phase is defined as its pure end-member with unit activity at temperatures and pressures of interest. The standard state of the solvent in aqueous solutions is pure solvent at temperatures and pressures of interest. The standard state for an aqueous solute is a hypothetical 1 molal (m) solution referred to infinite dilution at temperatures and pressures of interest. Gibbs free energies and enthalpies of formation reported in this study correspond to the formation from chemical elements at their reference states.

The Pitzer model is adopted in this study for calculations of activity coefficients of aqueous species. The detailed descriptions about Pitzer equations are provided in Pitzer (1991). In the following, an equation for calculation of the activity coefficient of $Al(OH)_4^-$ in NaOH medium without consideration of triple interactions in the Pitzer model is provided as an example,

133

134
$$\ln \gamma = -A_{\phi} \left[\frac{\sqrt{I_m}}{1 + 1.2\sqrt{I_m}} + \frac{2}{1.2} \ln(1 + 1.2\sqrt{I_m}) \right] + m \left\{ 2\beta^{(0)} + \frac{2\beta^{(1)}}{\alpha^2 \times I_m} \left[1 - (1 + \alpha\sqrt{I_m} - \frac{\alpha^2 I_m}{2})e^{-\alpha\sqrt{I_m}} \right] \right\} + \frac{3m^2}{2}C^{\phi}$$
135 (1)

136

137 where γ is activity coefficient, A_{ϕ} is Deby-Huckel slope for osmotic coefficient; I_m is 138 ionic strength on molality scale; *m* is molality of Na⁺; α is equal to 2; and $\beta^{(0)}$, $\beta^{(1)}$ and C^{ϕ} 139 are Pitzer binary interaction coefficients between Na⁺ and Al(OH)₄⁻.

140 The uncertainties reported in this study are two standard deviations (2σ) . Error 141 propagations are calculated based on uncertainties associated with regressions and

142	equilibrium constants in the model. In some cases, especially for $\Delta_f G$, $\Delta_f H$, and S ^o ,								
143	uncertainties could be underestimated, as uncertainties for auxiliary data are not									
144	available, and therefore not included.									
145										
146	2.1 Thermodynamic Constants and Pitzer Interaction Parameters of Si and Al	Aqueous								
147	Species Chosen and Extrapolated from the Literature									
148										
149	The following monomer, dimer and trimer silica species are considered	ed in this								
150	thermodynamic model:									
151										
152	$H_4SiO_4 = H^+ + H_3SiO_4^-$	(2)								
153										
154	$H_3SiO_4^- = H^+ + H_2SiO_4^{2-}$	(3)								
155										
156	$2H_4SiO_4 = H^+ + H_5Si_2O_7^- + H_2O$	(4)								
157										
158	$H_5Si_2O_7^- = H^+ + H_4Si_2O_7^{2-}$	(5)								
159										
160	$3H_4SiO_4 = 3H^+ + H_5Si_3O_{10}^{3-} + 2H_2O$	(6)								
161										
162										
163	The equilibrium constants for Reaction 2 are obtained from Fleming and Crer	ar (1982)								
164	(Table 2). The first dissociation constants of monomeric silicic acid from Fle	ming and								
165	Crerar (1982) are selected, as these authors regressed a number of exper	rimentally								
166	determined first dissociation constants for monomeric silicic acid from previou	is studies,								
167	including the most reliable values determined via the solubility and poter	ntiometric								

168 methods (e.g., Seward, 1974; Busey and Mesmer, 1977). However, uncertainties 169 associated with their dissociation constants were not given in Fleming and Crerar (1982). 170 In this study, uncertainties are assigned based on the respective benchmark values of 171 Busey and Mesmer (1977). The second dissociation constants (log K_2) of monomeric 172 silicic acid at elevated temperatures (Reaction 3) are estimated by using the one-term 173 isocoulombic approach with phosphoric acid as the model substance as employed before 174 (e.g., Xiong, 2003; 2007) (Table 2), based on the log K_2 at 25 °C from Hershey and 175 Millero (1986). The equilibrium constants at 25 °C for dimer and trimer silica species 176 (Reactions 4–6) are from Felmy et al. (2001). The log K at elevated temperatures for 177 those dimer and trimer silica species are also predicted in this study in a manner similar 178 to that described above for the log K_2 of monomer silicic acid. They are tabulated in 179 Table 2. The uncertainty for log K's of well balanced isocoulombic reactions is usually 180 within ± 0.50 up to 200 °C (Gu et al., 1994). Therefore, an uncertainty of ± 0.25 is 181 assigned to all predicted values at elevated temperatures up to 100 °C.

182 An example of a well-balanced isocoulombic reaction is given in the following,

183

184
$$H_4SiO_4 + H_2PO_4^- = H_3PO_4 + H_3SiO_4^-$$
 (7)

185

186 In the above reaction, H_4SiO_4 with zero charge on the left side is balanced by H_3PO_4 with 187 zero charge on the right side. Similarly, $H_2PO_4^-$ with one negative charge on the left side 188 is balanced by $H_3SiO_4^-$ with one negative charge on the right side.

189 It should be noted that regarding polymers of silica species, only dimer and trimer 190 are considered in this study, and tetramer and hexamer are excluded. This consideration

191	is primarily based on numerous studies which have indicated that monomer, dimer, and								
192	trimer are adequate in descriptions of silica solutions with high concentrations (e.g.,								
193	Čižmek et al., 1992; Sefćik and McCormic, 1997b; Hunt et al., 2011). For instance, in a								
194	Raman spectroscopic study on silica speciation in concentrated silica solutions up to								
195	5.0 m with KOH as a supporting solution ranging from 0.08 m to 8.0 m conducted by								
196	Hunt et al. (2011), the authors demonstrate that monomer, dimer and trimer species are								
197	sufficient to describe the silica species in highly concentrated silicate solutions with total								
198	silica concentrations up to 5.0 m.								
199	The following species are incorporated into the aluminum thermodynamic model:								
200									
201	$Al(OH)_4^- = Al^{3+} + 4OH^-$ (8)								
202									
203	Equilibrium constants for Reaction 8 at various temperatures are tabulated in Table 3.								
204	The equilibrium constants for Reaction 8 are obtained from Wesolowski (1992)								
205	based on the following reaction:								
206									
207	$Al(OH)_4^- = Al(OH)_3(cr, gibbsite) + OH^-$ (9)								
208									
209	in combination with the equilibrium constants for the following reaction calculated from								
210	SUPCRT (Johnson et al., 1992):								
211									
212	$Al(OH)_{3}(cr, gibbsite) = Al^{3+} + 3OH^{-} $ (10)								
213									

In this study, the hydrolysis constants of $Al(OH)_4^-$ described above determined by Wesolowski (1992) are selected because they are consistent with the respective Pitzer parameters adopted in this study. Although Tagirov and Schott (2001) provided the revised hydrolysis constants for $Al(OH)_4^-$, those values are not adopted in this study, as the activity coefficient model employed by Tagirov and Schott (2001) is different from the Pitzer model.

220 In the present model, regarding aluminum species in alkaline solutions, only 221 $Al(OH)_4^{-}$ is included, and no aluminum polymeric species are considered. This 222 consideration is based on the fact that aqueous aluminum species exists as $Al(OH)_4$ in neutral and basic solutions when $m_{\Sigma A1}$ is less than 1.5 m (Moolenaar et al., 1970; Baes 223 224 and Mesmer, 1976; Castet et al., 1993). The formation of aluminum polymeric species requires $m_{\Sigma A1} \ge 1.5$ m (Moolenaar et al., 1970). In the presence of silica, aluminum 225 226 concentrations are much lower than 1.5 m. As solutions in which zeolites are formed 227 contain both silica and aluminum, aluminum polymeric species will not be important.

It should be noted that although Zhou et al. (2003) also determined the Pitzer interaction parameters for the interaction between Na^+ and $Al(OH)_4^-$, the Pitzer parameters of Wesolowski (1992) are adopted in the current model as they are consistent with the hydrolysis constants of $Al(OH)_4^-$ used in this study.

In addition, there are some studies suggesting that the aqueous Al-Si complex(es) such as Al(OH)₃H₃SiO₄⁻ and SiAlO₃(OH)₄³⁻ might be present in aqueous solutions (e.g., Pokrovski et al., 1998; Salvi et al., 1998; Gout et al., 2000). Salvi et al. (1998) suggested that Al(OH)₃H₃SiO₄⁻ with log β_I = 2.32 could be present at 300 °C, corresponding to the following reaction,

8/29

(13)

237

238
$$Al(OH)_4^- + H_4SiO_4(aq) = Al(OH)_3H_3SiO_4^- + H_2O(1)$$
 (11)

239

Gout et al. (2000) mentioned that a weak complex, $SiAlO_3(OH)_4^{3-}$ (also formulated as $Al(OH)_3HSiO_4^{3-}$) with log $\beta_I^I = 0.53$ (I = 1.2 m), could be present at 20 °C, with reference to the following reaction,

243

244
$$Al(OH)_4^- + H_2SiO_4^{2-} = Al(OH)_3HSiO_4^{3-} + H_2O(1)$$
 (12)

245

246 The β_l at infinite dilution can be expressed as,

248
$$\beta_{1} = \frac{(m_{Al(OH)_{3}HSiO_{4}^{3-}})(\gamma_{Al(OH)_{3}HSiO_{4}^{3-}})(a_{H_{2}O})}{(m_{Al(OH)_{4}^{-}})(\gamma_{Al(OH)_{4}^{-}})(m_{H_{2}SiO_{4}^{2-}})(\gamma_{H_{2}SiO_{4}^{2-}})} = \frac{(m_{Al(OH)_{3}HSiO_{4}^{3-}})}{(m_{Al(OH)_{4}^{-}})(m_{H_{2}SiO_{4}^{2-}})} \times \frac{(\gamma_{Al(OH)_{3}HSiO_{4}^{3-}})(a_{H_{2}O})}{(\gamma_{Al(OH)_{4}^{-}})(\gamma_{H_{2}SiO_{4}^{2-}})}$$
$$= \beta_{1}^{I} \times \frac{(\gamma_{Al(OH)_{3}HSiO_{4}^{3-}})(a_{H_{2}O})}{(\gamma_{Al(OH)_{4}^{-}})(\gamma_{H_{2}SiO_{4}^{2-}})}$$

249

250

Other studies (e.g., Yokoyama et al., 1988) suggested that the formation of aqueous complex(es) of aluminum with silica is likely in dilute NaOH solutions at room temperature, but such a complex is absent in high alkaline solutions such as 1.0 m NaOH, typical of solutions in which zeolites are stable. Yokoyama et al. (1988) also demonstrated that higher temperatures destabilize the complex of aluminum and silicate that formed in dilute NaOH solutions.

as an independent constrain on the existence of Al-Si complexes in alkaline solutions, the Al(OH)₃HSiO₄³⁻ complex is tested to see whether it can improve the modeling at 30 °C and 50 °C (see following sections), as this complex was proposed to be present at 20 °C. In doing this, as there is only one apparent formation constant at one ionic strength (i.e., 1.2 m) in NaCl, the log β_I at infinite dilution at 20 °C is first estimated by using the Brønsted-Guggenheim-Scatchard SIT model (Grenthe et al., 1992). According to the SIT model regarding Eqs. 12-13, we have,

265

257

266
$$\log \beta_1 = \log \beta_1^I - 4D + \log a_{H,O} + \Delta \varepsilon (Eq.11) \times I_m$$
(14)

267

$$268 \qquad \Delta \varepsilon(Eq.11) = \varepsilon(Na^+, Al(OH)_3 HSiO_4^{3-}) - \varepsilon(Na^+, Al(OH)_4^-) - \varepsilon(Na^+, H_2SiO_4^{2-})$$
(15)

269

where $\log \beta_1^{I}$ is the formation constant regarding Reaction 12 at a certain ionic strength defined in Eq. (13); I_m ionic strength on molal scale; ε 's are the SIT interaction coefficients; a_{H_2O} is activity of water; and D is the Debye-Hückel term given below,

273

274
$$D = \frac{A_{\gamma}\sqrt{I_m}}{1+1.5\sqrt{I_m}}$$
(16)

275

276 where A_{γ} the Debye-Hückel slope for activity coefficient.

Substituting
$$\log \beta_1^I = 0.53$$
 at 20 °C from Gout et al. (2000), $A_{\gamma} = 0.5059$ at 20 °C
from Helgeson and Kirkham (1974), $a_{H_2O} = 0.9600$ for 1.2 m NaCl calculated using
EQ3/6, $\varepsilon(Na^+, Al(OH)_3 HSiO_4^{3-}) \approx \varepsilon(Na^+, Si_3O_6(OH)_3^{3-}) = -0.25 \pm 0.03$,
 $\varepsilon(Na^+, H_2SiO_4^{2-}) = \varepsilon(Na^+, SiO_2(OH)_2^{2-}) = -0.10 \pm 0.07$, and
 $\varepsilon(Na^+, Al(OH)_4^-) \approx \varepsilon(Na^+, B(OH)_4^-) = -0.07 \pm 0.05$, all from Grenthe et al. (1992), and
 $I_m = 1.2$ m into Eqs. 14–16, $\log \beta_I$ is estimated as -0.42 ± 0.10 at 20 °C. As Reaction 12
is in a semi-isocoulombic form, the $\log \beta_I$ at 20 °C is directly extrapolated to 25 °C and
other temperatures using the one-term isocoulombic principle (Table 2).
2.2 *Pitzer Interaction Parameters in the Na-Si-OH Systems Obtained in This Study*
In this study, the Pitzer binary interaction parameters for Na⁺-H₅Si₂O₇⁻,

In this study, the Pitzer binary interaction parameters for Na⁻-H₅Si₂O₇, Na⁺-H₄Si₂O₇²⁻, and Na⁺-H₅Si₃O₁₀³⁻ are calculated according to the method of Plyasumov et al. (1998) based on the respective SIT coefficients for these interactions (Table 3).

The high order Pitzer parameters, θ_{ij} for OH⁻-H₂SiO₄²⁻ interaction, and Ψ_{ijk} for OH⁻-H₂SiO₄²⁻-Na⁺ interaction, are evaluated from solubility data of the sodium silicate, 3Na₂O•2SiO₂•11H₂O, in NaOH solutions up to ~19 m from Sprauer and Pearce (1940). In addition, $(\frac{\partial \theta_{ij}}{\partial T})_p$ for OH⁻-H₂SiO₄²⁻ interaction is evaluated from solubility data of Na₃HSiO₄•5H₂O and Na₃HSiO₄•2H₂O at elevated temperatures from Baker et al. (1950) (Table 3).

298	For the testing purpose, the Pitzer binary interaction parameters for
299	$Na^{+}-Al(OH)_{3}HSiO_{4}^{3-}$ are assumed to be the same as those for $Na^{+}-H_{5}Si_{3}O_{10}^{3-}$ (Table 3).
300	Sprauer and Pearce (1940) experimentally determined solubilities of
301	3Na ₂ O•2SiO ₂ •11H ₂ O in NaOH solutions at 25 °C. Their experiments approached
302	equilibrium from supersaturation in about one month (Table 1). Their experiments were
303	in NaOH solutions with very high concentrations up to \sim 19.0 m. The solubility reaction
304	for $3Na_2O \cdot 2SiO_2 \cdot 11H_2O$ can be expressed as follows, using $H_4SiO_4(aq)$ as a basis species
305	for silica,
306	
307	$3Na_2O \cdot 2SiO_2 \cdot 11H_2O + 6H^+ = 6Na^+ + 2H_4SiO_4(aq) + 10H_2O(l)$ (17)
308	
309	The log K for Reaction 17 is evaluated as 83.83 \pm 0.62 (Table 3) along with θ_{ij} for
310	$OH^{-}H_{2}SiO_{4}^{2-}$ interaction, and Ψ_{ijk} for $OH^{-}H_{2}SiO_{4}^{2-}-Na^{+}$ interaction (Table 4).
311	In Figure 1, predicted solubilities of 3Na ₂ O•2SiO ₂ •11H ₂ O at 25 °C are compared
312	with experimental data from Sprauer and Pearce (1940). From the figure, it is clear that
313	the model reproduces experimental data generally within a factor of \sim 3.
314	Baker et al. (1950) investigated solubilities of Na ₃ HSiO ₄ •2H ₂ O and
315	$Na_3HSiO_4 \bullet 5H_2O$ in NaOH solutions up to ~24 m at elevated temperatures to 90 °C. Their
316	experiments approached equilibrium from both undersaturation and supersaturation in a
317	few of weeks. The dissolution of $Na_3HSiO_4 \cdot 5H_2O$ can be expressed as:
318	
319	$Na_{3}HSiO_{4} \bullet 5H_{2}O + 3H^{+} = 3Na^{+} + H_{4}SiO_{4}(aq) + 5H_{2}O(l) $ (18)
320	

321	Based on solubility data from Baker et al. (1950) on Na ₃ HSiO ₄ •5H ₂ O, the equilibrium
322	constant for Na ₃ HSiO ₄ •5H ₂ O at 50 °C is obtained as 41.50 ± 0.35 (Table 4).
323	Similarly, the dissolution of Na ₃ HSiO ₄ •2H ₂ O can be expressed as:
324	
325	$Na_{3}HSiO_{4} \bullet 2H_{2}O + 3H^{+} = 3Na^{+} + H_{4}SiO_{4}(aq) + 2H_{2}O(l) $ (19)
326	
327	Based on solubility data from Baker et al. (1950) on Na ₃ HSiO ₄ •2H ₂ O, the equilibrium
328	constants for Na ₃ HSiO ₄ •2H ₂ O at 50 °C, 70 °C and 90 °C are obtained as 47.03 \pm 0.35,
329	45.69 ± 0.31 , and 46.03 ± 0.30 , respectively (Table 4). In combination with experimental
330	data for Na ₃ HSiO ₄ •5H ₂ O at 50 °C, the temperature dependence of $\theta_{OH^-, H_2SiO_4^{2-}}$, $(\frac{\partial \theta_{ij}}{\partial T})_p$,
331	is also evaluated (Table 3).
332	In Figure 1, predicted solubilities of Na ₃ HSiO ₄ •5H ₂ O and Na ₃ HSiO ₄ •2H ₂ O at
333	50 °C are also compared with experimental data from Baker et al. (1950). For
334	Na ₃ HSiO ₄ •5H ₂ O, the model reasonably reproduces experimental data within a factor of
335	~1.5, but at I \approx 19.5 m, it is within a factor of ~5.5. For Na ₃ HSiO ₄ •2H ₂ O, the model
336	predicts solubilities at I \approx 20 m within a factor of \sim 3, but it reproduces experimental data
337	at higher ionic strengths within a factor of ~2.
338	In Figure 2, predicted solubilities of Na ₃ HSiO ₄ •2H ₂ O at 70 °C and 90 °C are
339	compared with experimental data. Figure 2 demonstrates that the model reproduces
340	experimental data within a factor of ~ 2 for the majority of the data points.
341	Based on the linear regression of temperature dependence of equilibrium
342	constants, $\Delta_r H$ for Reaction 19 is derived as -58 ± 45 (2 σ) kJ mol ⁻¹ (Figure 3, and
343	Table 5).

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In this study, it is assumed that interaction parameters are constant over the temperature range from 25 °C to 100 °C, except for θ_{ij} for OH⁻-H₂SiO₄²⁻ interaction, which is necessary to modeling solubility of sodium silicates in highly concentrated NaOH solutions. This assumption is based on the observation that the Pitzer interaction parameters do not change significantly over a narrow range of temperature. For instance, for the interaction of Na⁺ with Al(OH)₄⁻, the temperature derivatives of Pitzer interaction

350 parameters are very small over this temperature range, i.e., $(\frac{\partial \beta^{(0)}}{\partial T})_p = 8.0 \times 10^{-5}$,

351
$$\left(\frac{\partial \beta^{(1)}}{\partial T}\right)_P = 2.7 \times 10^{-4}$$
, and $\left(\frac{\partial C^{(\phi)}}{\partial T}\right)_P = 6.9 \times 10^{-5}$, based on the respective interaction

parameters as a function of temperature from Wesolowski (1992). In addition, in the
validation test (see the following section), model-predicted values are in satisfactory
agreement with independent experimental values.

355 The data0.PIT database in the EQ3/6 code can be used to compute activity 356 coefficients by using the Pitzer equations up to 100 °C (Wolery, 1992). The original 357 data0.PIT database does not contain the following species: H₄SiO₄(aq), H₃SiO₄, $H_2SiO_4^{2-}$, $H_5Si_2O_7^{-}$, $H_4Si_2O_7^{2-}$, $H_5Si_3O_{10}^{3-}$, and $Al(OH)_4^{-}$. By incorporating the 358 359 equilibrium constants for the respective reaction detailed in Table 2, and the relevant 360 Pitzer interaction parameters from Hershey and Millero (1986), Wesolowski (1992), 361 Azarousal et al. (1997), and this study tabulated in Table 3 into the data0.PIT database, 362 the PIT database is modified to be able to model high ionic strength solutions with high concentrations of Si and Al up to 100 °C. 363

364

366 2.3 Model Validation

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A validation test is performed for the Na-Si-OH model developed in this study. For this purpose, predicted solubilities of amorphous silica in alkaline solution with a wide range ionic strengths are compared with experimental data, which are independent from the model development. The dissolution of amorphous silica can be expressed as, SiO₂(am) + 2H₂O = 2H⁺ + H₂SiO₄²⁻ (20) SiO₂(am) + 2H₂O = 2H⁺ + H₂SiO₄²⁻ (20) The log *K* for Reaction 20, is -25.81 from Weber and Hunt (2003). In Figure 4, predicted

solubilities of amorphous silica are compared with experimental data in NaOH solutions at 25 °C from Alexander et al. (1954), and in NaOH + NaNO₃ mixtures at 25 °C from Weber and Hunt (2003). From Figure 4, it is obvious that solubilities predicted by the model are in good agreement with model-independent experimental data over the entire ionic strength range from very dilute to ~4.5 m in alkaline solutions.

381

382 3. Model Applications

In the following, the model developed above is applied to calculations of speciation of silica species as a function of pH at temperatures of 25 °C and 100 °C, and to retrieval of thermodynamic data from hydrothermal solubility data on zeolite A and amorphous precursor of zeolite A.

387

389 3.1 Speciation of Silica Species

390

In Figure 5A-C, speciation of silica species as a function of pH at 25 °C at 391 392 different total silica concentrations is displayed. At $\Sigma Si = 0.01$ m, monomer species are the dominant species (Figure 5A). In the pH range from \sim 7 to \sim 9, the polymeric species 393 394 $H_5Si_2O_7^-$ can account for up to 2% of the total dissolved silica (Figure 5A). At $\Sigma Si =$ 395 0.1 m, although monomer species are still the dominant species, contributions from polymeric species become significant (Figure 5B). In the pH range from ~6 to ~10, 396 397 $H_5Si_2O_7$ can account for up to ~20 % of the total dissolved silica. In the pH range from ~9 to ~13, both $H_4Si_2O_7^{2-}$ and $H_5Si_3O_{10}^{3-}$ can account for up to ~5% of the total dissolved 398 silica (Figure 5B). At Σ Si = 1 m, contributions from polymeric species to the total 399 400 dissolved silica become important (Figure 5C). In the pH range up to 9, $H_5Si_2O_7^-$ can account for up to ~50% of the total dissolved silica. In the pH range from 8 to 14, 401 $H_5Si_3O_{10}^{3-}$ can account for up to ~40% of the total dissolved silica, and $H_4Si_2O_7^{2-}$ can 402 403 account for up to $\sim 10\%$ of the total dissolved silica. However, the trend indicates that $H_2SiO_4^{2-}$ will be the dominant species above pH 14 even at $\Sigma Si = 1$ m. 404

Similarly, speciation of silica at different total concentrations of silica as a function of pH at 100 °C is displayed in Figure 6A-C. The polymeric silica species are important when $\Sigma Si = 1$ m (Figure 6C). However, when pH higher than 11.5, the contributions from polymeric silica species diminish, and above pH 13, the contributions from polymeric silica species become insignificant (Figure 6C).

410 It should be mentioned that under highly basic conditions where zeolites are 411 formed as discussed blow, $H_2SiO_4^{2-}$ is the dominant species.

413 3.2 Calculation of Equilibrium Constants of Zeolites

414 Using the above thermodynamic model, the thermodynamic equilibrium constants 415 of zeolites can be retrieved from solubility experiments at elevated temperatures. The 416 main criterion of selection of solubility data is that equilibrium state must be attained. 417 For precipitation and dissolution of zeolites at elevated temperatures, detailed kinetic 418 studies equilibrium have demonstrated that rapidly attained. state is 419 Addai-Mensah et al. (2004) performed detailed experiments from both undersaturation 420 and supersaturation on zeolite A and amorphous precursor of zeolite A in NaOH 421 solutions at 65 °C. Their experimental results demonstrate that the reversal was attained 422 in about one minute (~50 seconds), indicating fast kinetics to reach equilibrium. 423 Antonic et al. (1993) also indicated that the equilibrium was attained at about 20 minutes 424 for dissolution of zeolite A at 80 °C. For experimental results in which the equilibrium 425 state was not explicitly mentioned, steady state concentrations after 100 minutes are 426 treated as equilibrium concentrations. In this study, as sufficient data for zeolite A and 427 amorphous precursor of zeolite A have been located, equilibrium constants for these two phases are computed. For zeolite X (NaAlSi_{1.23}O_{4.46}•3.07H₂O), Čižmek et al. (1991b) 428 429 conducted solubility experiments on zeolite X in 2.06 m NaOH solution from 65 °C to 80 °C at 5 °C increment, and their data set has only one data point at each temperature. 430 431 Roozeboom et al. (1983) had one single data point for solubility of zeolite X in \sim 1.02 m 432 NaOH at 98 °C. Therefore, sufficient data have not been located for zeolite X. 433 Consequently, no attempt has been made to compute equilibrium constants for zeolite X 434 at this time.

435 Equilibrium constants are obtained according to the computer modeling. The 436 computer modeling is performed by using EQ3/6 version 8.0a (Wolery et al., 2010; 437 Xiong, 2011). The essence of the modeling is to minimize the difference between 438 experimental and model predicted values, as detailed in the previous publication 439 (Nemer et al., 2011). The retrieval of the equilibrium constant for amorphous precursor 440 of zeolite A at 30 °C can be served as an example. First, solubility data from Ejaz and 441 Graham (1999) and Addai-Mensah et al. (2004) were used to generate EQ3NR input files. 442 Second, an initial guess for the log K was made. Third, by changing the log K into 443 different values, a series of sums of squares of residuals between experimental 444 solubilities and predicted solubilities were obtained. The final log K corresponds to the 445 minimized sum of squares of residuals.

446

447 3.2.1 Calculation of Equilibrium Constants of Zeolite A

448

449 A number of researchers have investigated solubilities of zeolite A in NaOH 450 solutions at elevated temperatures. The sources of solubility data used for obtaining 451 equilibrium constants in this study and their respective experimental conditions are listed 452 in Table 1. Based on those experimental data, the equilibrium constants for the following 453 reaction using Al^{3+} and $H_4SiO_4(aq)$ as basis species of aluminum and silica, respectively,

454

456

457 are obtained (Table 3). In computation of equilibrium constants, all concentrations on458 molar scale are converted to molal scale according to the following equation based on

densities of supporting solutions used in experiments at respective temperatures, which
are calculated from density equations of Sőhnel and Novotný (1985),

461

462
$$m_i = \frac{1000 \times M_i}{1000\rho - \sum_{i}^{j} M_i E_i}$$
(22)

463

464 where m_i is concentration of *i* species on molality scale, ρ density of solution, M_i is 465 concentration of *i* species on molarity scale, and E_i the molecular weight of *i* species.

In Figure 7, predicted solubilities of zeolite A at temperatures of 30 °C, 65 °C, 466 70 °C, 80 °C, 90 °C, and 100 °C, are compared with experimental data at the respective 467 468 temperatures. At 30 °C, there is a scatter in experimental data from various researchers 469 with a difference of one order of magnitude, and the model seems to fit experimental data within a factor of ~ 5 except for that it underpredicts at I ≈ 6 m by a factor of ~ 7 470 471 (Figure 7A). At 50 °C, the values produced by the model agree with experimental values within a factor of ~1.3 (Figure 7A). At 65 °C, the model reproduces experimental data 472 473 within a factor of ~ 2 (Figure 7A). At 70 °C, the values predicted by the model agree with 474 experimental values within a factor of ~1.4 (Figure 7B). At 80 °C, the model reproduces 475 solubilities in the ionic strength range from ~ 3 m to ~ 5 m within a factor of ~ 3.5 in 476 comparison with the experimental data (Figure 7B), whereas solubilities predicted by the 477 model agree with experimental solubilities within a factor of ~ 1.1 in the ionic strength 478 range from ~ 0.5 m to ~ 2 m. At 90 °C, solubilities calculated by the model generally 479 agree with experimental values within a factor of ~2 (Figure 7C). At 100 °C, solubilities 480 computed by the model are in agreement with experimental values within a factor of ~ 1.2 481 (Figure 7C). In all above descriptions, the Al-Si complex, $Al(OH)_3HSiO_4^{3-}$, was not 482 included in calculations.

At 30 °C, the inclusion of Al(OH)₃HSiO₄³⁻ yields a log K of 9.96 ± 0.30 for 483 Reaction 21, in comparison with a value of 10.23 ± 0.31 for the log K produced by the 484 485 model without the above Al-Si complex. These two values are statistically 486 indistinguishable. However, as indicated by Figure 7A, the introduction of the above Al-Si complex improves the fitting at 30 °C; the model reproduces the solubility at I ≈ 6 487 m within a factor of ~ 2 in comparison with the experimental data point. Notice that as 488 described before, the model without $Al(OH)_3HSiO_4^{3-}$ reproduces the solubility at that 489 490 ionic strength within a factor of \sim 7.

491 On the other hand, at 50 °C, the model with the Al-Si complex produces a log K 492 of 7.41 ± 0.40 for Reaction 21, whereas the model without the Al-Si complex results in a log K of 7.95 \pm 0.30. Figure 7A shows that the model without the Al-Si complex 493 494 performs better. The reason for the poor performance of the model with Al-Si complex is not clear. One possibility may be that the existence of the Al-Si complex at 50 °C is 495 uncertain, as Gout et al. (2000) only mentioned its presence at 20 °C. The model with the 496 497 Al-Si complex at 30 °C is close to 20 °C, and the presence of the Al-Si complex at 30 °C 498 is of high certainty, if there is such a complex, explaining the better performance of the 499 model with the Al-Si complex at 30 °C. Another possibility is that the Pitzer parameters for Na^+ -Al(OH)₃HSiO₄³⁻ might be problematic. This seems unlikely, as the same set of 500 501 parameters is also used at 30 °C, which results in a desirable performance. However, the 502 final resolution of this issue requires independent evaluation of the Pitzer parameters and 503 to see if they can improve the model with the Al-Si complex.

Judging from the performance of the models with and without the Al-Si complex at 30 $^{\circ}$ C and 50 $^{\circ}$ C, it can be concluded that the model without the Al-Si complex is adequate at temperatures equal to or higher than 50 $^{\circ}$ C. Therefore, no further testing was performed at higher temperatures. Anyhow, the testing at 30 $^{\circ}$ C seems to provide independent support for the existence of Al(OH)₃HSiO₄^{3–} around 20 $^{\circ}$ C.

509 Based on the linear regression of temperature dependence of equilibrium constants, Δ_r H for Reaction 21 is derived as -154 ± 2 kJ mol⁻¹ (Table 5). According to 510 Figure 8, the log K at 25 °C is extrapolated as 10.24 ± 0.31 . Using the above 511 512 thermodynamic properties for Reaction 21, the thermodynamic properties of zeolite A at 25 °C and 1 bar are derived (Table 6). In derivation of these thermodynamic properties, 513 the auxiliary thermodynamic data for Na⁺ and Al³⁺ are from the DATA0.PIT database. 514 515 The auxiliary thermodynamic data for $H_4SiO_4(aq)$, which are not present in the 516 DATA0.PIT database, are from the NBS Thermodynamic Table (Wagman et al., 1982). 517 The latter database is in principle consistent with the DATA0.PIT database. However, as 518 the thermodynamic properties obtained in this study are log K's and Δ_r H, Gibbs free 519 energies of formation, enthalpies of formation, and standard entropies can be re-derived 520 to be consistent with other database(s) of interest, if needed.

521

522 3.2.2 Calculation of Equilibrium Constants of Amorphous Precursor of Zeolite A

523

In the synthesis of zeolite A, its amorphous precursor is usually formed first. Therefore, it is also important to know the thermodynamic properties of the amorphous precursor of zeolite A to optimize synthesis. Similar to Reaction 21, the dissolution reaction of the amorphous precursor of zeolite A is:

530

Ejaz and Graham (1999) conducted systematic solubility studies on the amorphous precursor of zeolite A in NaOH solutions from 3.0 m to 4.5 m at temperatures of 30 °C, 50 °C, 65 °C and 80 °C (Table 1). In addition, Addai-Mensal et al. (2004) also conducted solubility experiments on amorphous precursor of zeolite A in 3.0 m and 6.3 m NaOH solutions at 30 °C and 65 °C (Table 1). Therefore, based on solubility data from both Ejaz and Graham (1999) and Addai-Mensal et al. (2004), the equilibrium constants for the amorphous precursor of zeolite A are calculated (Table 4).

538 In Figure 9, predicted solubilities of amorphous precursor of zeolite A are 539 compared with experimental data. At 30 °C and 50 °C, the solubility curves for these two 540 temperatures are very close, and the solubilities predicted by the model match experimental values within a factor of ~1.6 (Figure 9A). At 65 °C, the solubilities 541 542 predicted by the model are in agreement with experimental values within a factor of ~ 2.5 543 for the majority of the data points, and within a factor of ~4.5 for the data point at I≈6 m (Figure 9B). At 80 °C, the model matches the experimental solubilities in high 544 545 ionic strength range within a factor of ~ 2 , but within a factor of ~ 5 to ~ 10 at a low ionic 546 strength (~ 0.2 m) (Figure 9B).

547 According to the linear regression of temperature dependence of equilibrium 548 constants (Figure 8), Δ_r H for Reaction 23 is obtained as -248 ± 3 kJ mol⁻¹ (Table 5). 549 Based on Figure 8, the log *K* for Reaction 23 at 25 °C is extrapolated as 12.68 ± 0.35. In 550 accordance with the above thermodynamic properties for Reaction 23, the 551 thermodynamic properties of amorphous precursor of zeolite A at 25 °C and 1 bar are 552 derived (Table 6).

553

- 554 3.2.4 Validation of Calculations, and Discussions
- 555

556 The enthalpy of formation for zeolite A derived from solubility studies as a 557 function of temperature in this study is compared with the value from calorimetric 558 measurements. The enthalpy of formation from elements for zeolite A obtained in this study, -2738 ± 5 kJ mol⁻¹, is in excellent agreement with the value obtained by 559 calorimetric measurements $(-2731.3 \pm 1.8 \text{ kJ mol}^{-1})$ (Turner et al., 2008), and compares 560 561 favorably with the value obtained by theoretical calculations (-2739 kJ mol⁻¹) (Mathieu 562 Viellard, 2010). In these two studies, zeolite A is formulated and as $Na_{0.5067}Al_{0.501}Si_{0.4974}O_{1.99965} \cdot 1.0906H_2O$. For comparison purpose, their values are scaled 563 564 relative to four oxygen atoms for the stoichiometry of zeolite A adopted in this study, i.e., 565 NaAlSiO₄ \cdot 2.25H₂O. This favorable comparison also validates the model developed in 566 this study.

The Gibbs free energy of formation for zeolite A at 25 °C obtained in this study can also be compared with that of Caullet et al. (1980). Caullet et al. (1980) determined the log *K* for zeolite A with a formula of NaAlSi_{1.06}O_{4.12}•2.4H₂O for the following reaction from solubility experiments in 0.02 m, 0.1 m and 0.5 m NaOH solutions at $25 ^{\circ}$ C,

The average thermodynamic equilibrium constant, log K at 25 °C, for Reaction 24 575 obtained by Caullet et al. (1980) is -11.20 ± 0.16 . Based on the log K value for 576 Reaction 24, the Gibbs free energy change for Reaction 24 is calculated as 63.9 ± 0.9 577 kJ mol⁻¹. According to the auxiliary thermodynamic data for species in Reaction 24 from 578 579 the NBS Thermodynamic Table, the Gibbs free energy of formation, ΔG , for NaAlSi_{1.06}O_{4.12}•2.4H₂O, is derived as -2619 ± 2 kJ mol⁻¹. When it is scaled to four 580 581 oxygen atoms for the stoichiometry of zeolite A adopted in this study, in the same 582 procedure applied in the above validation of calculation of enthalpy of formation for zeolite A, the Gibbs free energy of formation becomes $-2543 \pm 2 \text{ kJ mol}^{-1}$ for the formula 583 584 of Na_{0.97}Al_{0.97}Si_{1.03}O₄•2.33H₂O. This value is in good agreement with the value of -2541 ± 2 kJ mol⁻¹ obtained in this study. The corresponding log K at 25 °C for the 585 586 following reaction,

588

$$Na_{0.97}Al_{0.97}Si_{1.03}O_4 \bullet 2.33H_2O(cr) + 3.88H^+ =$$

0.97Na^+ + 0.97Al^{3+} + 1.03H_4SiO_4(aq) + 2.21H_2O(l)

589

590 would be 11.33 ± 0.20 .

Qiu et al. (2000) measured the entropy of dehydrated zeolite A with a formula of Na₉₆Al₉₆Si₉₆O₃₉₄ as 13030 J K⁻¹ mol⁻¹ at 25 °C. When the above formula is scaled to four oxygen atom, the corresponding entropy becomes 135.73 J K⁻¹ mol⁻¹ for dehydrated zeolite A with a formula of NaAlSiO₄. Viellard (2010) uses a value of 52.0 J K⁻¹ mol⁻¹ for the entropy of zeolitic water for predictions of entropies of zeolites. Accordingly, using the above values from Qiu et al. (2000) and Viellard (2010), the entropy for

(25)

zeolite A with a formula of NaAlSiO₄•2.25H₂O would be expected to be 253 J K⁻¹ mol⁻¹. 597 This value would differ from the entropy obtained in this study by about 100 J K^{-1} mol⁻¹. 598 599 which was consistently derived from $\Delta_r H$ and $\Delta_r G$ from equilibrium observations. 600 As the enthalpy of formation for zeolite A derived in this study is consistent with 601 the value obtained by the calorimetric method as mentioned above, we might use $\Delta_t H =$ 2738 kJ mol⁻¹ and S^o = 253 J K⁻¹ mol⁻¹ for zeolite A to calculate the log K at 25 °C for 602 603 Reaction 21. That would result in a value for the log K at 25 °C for Reaction 21 to be 604 16.50. This value is severely discordant with the log K (10.23 \pm 0.31) for Reaction 21 605 obtained in this study, and the similar log K (11.33 \pm 0.20) for Reaction 25 obtained by Caullet et al. (1980), both at 25 °C, and it would predict solubilities of zeolite A by at 606 607 least five orders of magnitude higher than the observed solubilities. Therefore, it seems 608 that further studies on entropies of both zeolite A and dehydrated zeolite A by 609 calorimetric measurements are required, as the existing value does not agree with 610 equilibrium observations.

611

612 4 Conclusions

613

In this study, a thermodynamic model for silica and aluminum is developed, valid to high ionic strength at elevated temperatures up to 100 °C. This model is useful for understanding the geochemical behaviors of Si and Al in concentrated hydrothermal solutions, and for guiding hydrothermal synthesis of zeolites. This model enable us to calculate equilibrium constants of sodium silicates, zeolite A, and amorphous precursor of zeolite A, from hydrothermal solubility experiments. In the near future, equilibrium constants for other zeolite species such as sodalite and cancrinite will be evaluated from

621	hydrothermal experiments. Because of the discordance of one single existing entropy for
622	dehydrated zeolite from calorimetric measurements with equilibrium observations, it is
623	suggested that further calorimetric studies on both zeolite A and dehydrated zeolite A are
624	needed to resolve such a discordance.
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637 **References**

638

Addai-Mensah, J., Li, J., Rosencrance, S., Wilmarth, W. (2004) Solubility of amorphous
sodium aluminosilicate and zeolite A crystals in caustic and nitrate/nitrite-rich caustic
aluminate liquors. Journal of Chemical and Engineering Data, 49, 1682–1687.

642

Alexander, G.B., Heston, W.M., and Iler, R.K. (1954) The solubility of amorphous silica
in water. Journal of American Chemical Society, 58, 453–455.

645

Antonić, T., Čižmek, A., Kosanović, C., and Subotić, B. (1993) Dissolution of
amorphous aluminosilicate zeolite precursors in alkaline solutions. Journal of Chemical
Society Faraday Transactions, 89, 1817–1822.

649

Azaroual, M., Fouillac, C., and Matray, J.M. (1997) Solubility of silica polymorphs in
electrolyte solutions, I. Activity coefficient of aqueous silica from 25 °C to 250 °C,
Pitzer's parameterization. Chemical Geology, 140, 155–165.

653

Baes, C.F., and Mesmer, R.E. (1976) The Hydrolysis of Cations. Wiley, New York. 655

Baker, C.L., Jue, L.R., and Wills, J.H. (1950) The system $Na_2O-SiO_2-H_2O$ at 50, 70 and 90 °C. Journal of Physical Chemistry, 72, 5369–5382.

658

Bosnar, S., Bronić, J., Krznarić, I., and Subotić, B., (2005) Influence of the
concentrations of aluminium and silicon in the liquid phase on the growth kinetics of
zeolite A and X microcrystals. Croatica Chemica Acta, 78, 1–8.

662

Busey, R.H., and Mesmer, R.E. (1977) Ionization of silicic acid and polysilicate
formation in aqueous chloride solutions to 300 °C.. Inorganic Chemistry, 16, 2444–2453.

666 Castet, S., Dandurand, J., Schott, J., and Gout, R. (1993) Boehmite solubility and aqueous 667 aluminum speciation in hydrothermal solutions (90-350 °C): Experimental study and 668 modeling. Geochimica et Cosmochimica Acta, 57, 4869–4884.

669

Caullet, P., Guth, J.-L., and Wey, R. (1980) Solubilité et grandeurs thermodynamiques de
dissolution des zéolites 4A et 13X dans des solutions aqueuses basiques. Bulletin de
Minéralogie, 103, 330-335.

673

674 Chorover, J., Choi, S., Amistadi, M. K., Karthikeyan, K. G., Crosson, G. and Mueller, K.
675 T. (2003) Linking cesium and strontium uptake to kaolinite weathering in simulated tank
676 waste leachate. Environ. Sci. Technol., 37, 2200–2208.

677

678 Ciric, J., (1968). Kinetics of zeolite A crystallization. Journal of Colloid and Interface 679 Science, 28, 315–324.

Čižmek, A., Komunjer, L., Subotić, B., Široki, M., and Rončević, S. (1991b). Kinetics of
zeolite dissolution: Part 2. Dissolution of zeolite X in hot sodium hydroxide. Zeolites,
11, 810–815.

688

Čižmek, A., Komunjer, L., Subotić, B., Široki, M., and Rončević, S. (1992) Kinetics of
zeolite dissolution: Part 3. Dissolution of synthetic mordenite in hot sodium hydroxide.
Zeolites, 12, 190–196.

692

Ejaz, T., Graham, A.G.J. (1999) Solubility of zeolite A and its amorphous precursor
under synthesis conditions. Journal of Chemical and Engineering Data, 44, 574–576.

Felmy, A.R., Cho, H., Rustad, J.R., Mason, M.J. (2001) An aqueous thermodynamic
model for polymerized silica species to high ionic strength. Journal of Solution
Chemistry, 30, 509–525.

Fleming, B.A., Crerar, D.A. (1982) Silicic acid ionization and calculation of silica
 solubility at elevated temperature and pH. Application to geothermal fluid processing and
 re-injection. Geothermics, 11, 15–29

703

699

Gout, R., Pokrovski, G.S., Schott, J., Zwick, A. (2000) Raman spectroscopic study of
aluminum silicate complexes at 20°C in basic solutions. Journal of Solution Chemistry,
29, 1173–1186.

707

Grenthe, I., Fuger, J., Konings, R.J.M., Lemire, R.J., Muller, A.B., Nguyen-Trung, C.,
Wanner, H. (1992) Chemical Thermodynamics of Uranium, Wanner, H., Forest, I.,
Nuclear Energy Agency, Organization for Economic Co-operation, Development, Eds.,
Vol. 1, Chemical Thermodynamics, North Holland Elsevier Science Publishers B.V.,
Amsterdam, The Netherlands, 715 p.

713

Helgeson, H.C., and Kirkham, D.H. (1974) Theoretical prediction of the thermodynamic
behavior of aqueous electrolytes at high pressures and temperatures. II. Debye–Hückel
parameters for activity coefficients and relative partial molal properties. American
Journal of Sciences, 274, 1199–1261.

718

Hershey, J.P., and Millero, F.J. (1986) The dependence of the acidity constants of silicis
acid on NaCl concentration using Pitzer's equation. Marine Chemistry, 18, 101–105.

721

722Hunt, J.D., Kavner, A., Schauble, E.A., Snyder, D., and Manning, C.E., 2011.723Polymerization of aqueous silica in H_2O-K_2O solutions at 25–200 °C and 1 bar to 20724kbar. Chemical Geology, 283, 161–170.

Johnson, J.W., Olkers, E.H., Helgeson, H.C. (1992) SUPCRT92: A software package for calculating the standard molal thermodynamic properties of minerals, gases, aqueous species, and reactions from 1 to 5000 bar and 0 to 1000°C. Computers and Geosciences, 18, 899–947.

730

Kosanović, C., Subotić, B., Kaučič, V., and Škreblin, M. (2000) Dissolution of the
zeolites NaA, potassum exchanged zeolite (KA) and the amorphous and crystalline
phases obtained by thermal treatment of zeolite KA in hot alkaline solution. Physical
Chemistry and Chemical Physics, 2, 3447–3451.

735

Mashal, K., Harsh, J. B., Flury, M. and Felmy, A. R. (2004) Colloid formation in
Hanford sediments reacted with simulated tank waste. Environ. Sci. Technol., 38, 5750–
5756.

- 740 Mathieu, R., Vieilard, P. (2010) A predictive model for the enthalpy formation of 741 zeolites. Microporous and Mesoporous Materials, 132, 335–351.
- 742

Moolenaar, R.J., Evans, J.C., and McKeever, L.D. (1970) The structure of the aluminate
ion in solutions at high pH. Journal of Physical Chemistry, 74, 3629–3636.

Myatt, G.J., Budd, P.M., Price, C., Hollway, F., Carr, S.E. (1994) The influence of
surfactants and water-soluble polymers on the crystallization of zeolite NaA. Zeolites,
14, 190–197.

749

Nemer, M., Xiong, Y.-L., Ismail, A.E., and Jang, J.-H. (2011) Solubility of Fe₂(OH)₃Cl (pure-iron end-member of hibbingite) in NaCl and Na₂SO₄ brines. Chemical Geology 280, 26–32.

753

Pitzer, K.S. (1991) Ion interaction approach: theory and data correlation. In Pitzer, K.S.,
editor, Activity Coefficients in Electrolyte Solutions, 2nd edition, CRC Press, Boca
Raton, Florida, p. 75-153.

757

Plyasunov, A., Fanghanel, T., and Grenthe, I. (1998) Estimation of the Pitzer equation
parameters for aqueous complexes. A case study for uranium at 298.15 K and 1 atm.
Acta Chemica Scandinavica, 52, 250–260.

761

Pokrovski, G.S., Schott, J., Salvi, G., Gout, R. and, Kubicki, J.D. (1998) Structure and
stability of aluminium-silica complexes in neutral to basic solutions. Experimental study
and molecular orbital calculations. Mineralogical Magazine, 62A, 1194-1195.

765

Qiu, L.-Y., Murashov, V., and White, M.A. (2000) Zeolite 4A: heat capacity and
 thermodynamic properties. Solid State Science, 2, 841–846.

768

Roozeboom, F., Robson, H.E., and Chan, S. (1983) Laser Raman study on the crystallization of zeolites A, X, and Y. Zeolites, 3, 321–328.

- 772 Salvi, S., Pokrovski, G.B., and Schott, J. (1998) Experimental investigation of aluminum-
- silica aqueous complexing at 300°C. Chemical Geology, 151, 51–67
- 774

Šefćik, J., and McCormick, A.V. (1997a) What is the solubility of zeolite A.
Microporous Materials, 10, 173–179.

777

Šefćik, J., and McCormick, A.V. (1997b) Thermochemistry of aqueous silicate solution
 precursors to ceramics. AIChE Journal, 43, 2773–2784.

780

Serne, R. J., Clayton, R. E., Kutnyakov, I. V., Last, G. V., LeGore, V. L., Wilson, T. C.,
Schaef, H. T., O'Hara, M. J., Wagnon, K. B., Lanigan, D. C., Brown, C. F., Williams, B.
A., Lindenmeier, C. W., Orr, R. D., Burke, D. S. and Ainsworth, C. C. (2002)
Characterization of Vadose Zone Sediment: Borehole 41-09-39 in the S-SX Waste
Management Area. Pacific Northwest National Laboratory, US Department of Energy,
PNNL-13757-3, Richland, Washington.

- Seward, T.M. (1974) Determination of the first ionization constant of silicic acid from
 quartz solubility in borate buffer solutions to 350 °C. Geochimica et Cosmochimica
 Acta, 38, 1651–1664.
- 791

- Sheppard, G.P., Hriljac, J.A., Maddrell, E.R., Hyatt, N. (2006) Silver zeolites: Iodide
 occlusion and conversion to sodalite-a potential ¹²⁹I waste form? Materials Research
 Society Symposium Proceeding, Vol. 932, 8 pp.
- Sőhnel, O. and Novotný, P. (1985) Densities of aqueous solutions of inorganic
 substances. Elsevier, New York, 335 p.
- 798
- Sprauer, J.W., and Pearce, D.W. (1940) Equilibrium in the systems Na₂O–SiO₂–H₂O and
 Na₂O–Al₂O₃–H₂O at 25 °C. Journal of American Chemical Society, 44, 909–916.
- 801
- Tagirov, B., and Schott, J. (2001) Aluminum speciation in crustal fluids revisited.
 Geochimica et Cosmochimica Acta, 65, 3965–3992.
- Turner, S., Sieber, J.R., Vetter, T.W., Zeisler, R., Marlow, A.F., Moreno-Ramirez, M.G.,
 Davis, M.E., Kennedy, G.J., Borghard, W.G., Yang, S., Navrotsky, Toby, B.H., Kelly,
 J.F., Fletcher, R.A., Windsor, E.S., Verkouteren, J.R., and Leigh, S.D. (2008).
 Characterization of chemical properties, unit cell parameters and particle size distribution
 of three zeolite reference materials: RM 8850 zeolite Y, RM 8851 zeolite A and RM
 8852 ammonium ZSM-5 zeolite. Microporous and Mesoporous Materials, 107, 252–
 267
- 812
- Viellard, P. (2010) A predictive model for the entropies and heat capacities of zeolites.
 European Journal of Mineralogy, 22, 823–836.
- 815
- 816 Wagman, D.D., Evans, W.H., Parker, V.B., Schumm, R.H., Halow, I., Bailey, S., 817 Churney, K., Nuttall, R.L. (1982) The NBS tables of chemical thermodynamic

818 properties: Selected values for inorganic and C_1 and C_2 organic substances in SI units. 819 Journal of Physical and Chemical Reference Data, Volume 11, Supplement No. 2, 392 820 pp.

- 821
- Weber, C., and Hunt, R.D. (2003) Modeling alkaline silicate solutions at 25 °C.
 Industrial Engineering and Chemical Research, 42:6970–6976.
- Wesolowski, D.J. (1992) Aluminum speciation and equilibrium in aqueous solution: I.
 The solubility of gibbsite in the system Na-K-Cl-OH-Al(OH)₄ from 0 °C to 100 °C.
 Geochimica et Cosmochimica Acta, 56, 1065–1091.
- 828

- Wieker, W., and Fahlke, B. (1985) In Drzaj, B., Hocevar, S., Pejovnik, S., Eds., Zeolites,
 Elsevier, Amsterdam, 161.
- Wolery, T.J. (1992) EQ3/6, A Software Package for Geochemical Modeling of Aqueous
 Systems: UCRL-MA-110662 PT I. Lawrence Livermore National Laboratory,
 Livermore, CA, USA.
- 835
- Wolery, T.W., Xiong, Y.-L., and Long, J. (2010) Verification and Validation
 Plan/Validation Document for EQ3/6 Version 8.0a for Actinide Chemistry, Document
 Version 8.10. Carlsbad, NM: Sandia National laboratories. ERMS 550239.
- 839
- Xiong, Y.-L. (2003) Predicted equilibrium constants for solid and aqueous selenium
 species to 300 °C: Applications to selenium-rich mineral deposits. Ore Geology
 Reviews, 23, 259–276.
- 843
- Xiong, Y.-L. (2007) Hydrothermal thallium mineralization up to 300 °C: a
 thermodynamic approach. Ore Geology Reviews, 32, 291–313.
- 846
- Xiong, Y.-L. (2011) WIPP Verification and Validation Plan/Validation Document for
 EQ3/6 Version 8.0a for Actinide Chemistry, Revision 1, Document Version 8.20.
 Supersedes ERMS 550239. Carlsbad, NM. Sandia National Laboratories. ERMS
 555358.
- Yokoyama, T., Kinoshita, S., Wakita, H. and Tarutani, T. (1988) ²⁷Al NMR Study on the
 Interaction between aluminate and silicate ions in alkaline solution. Bulletin of the
 Chemical Society of Japan, 61, 1002–1004.
- 855
- Zhao, H., Deng, Y., Harsh, J. B., Flury, M. and Boyle, J. (2004) Alteration of kaolinite to
 cancrinite and sodalite by simulated Hanford Tank Wastes and its impact on cesium
 retention. Clays Clay Mineralogy, 52, 1–13.
- 859
- 860 Zhdanov, S.P. (1971) Some Problems of Zeolite Crystallization. In Flanigen, E.M., and
- 861 Sand, L.B., Eds., Molecular Sieve Zeolites-I. American Chemical Society, Washington,
- 862 D.C., pp. 20-43.

- Zhou, J., Chen, Q.-Y., Li, J., Yin, Z.-L., Zhou, X., Zhang, P.-M. (2003). Isopiestic
 measurement of the osmotic and activity coefficients for the NaOH–NaAl(OH)₄–H₂O
- 866 system at 313.2 K. Geochimica et Cosmochimica Acta, 67, 3459–3472.

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Table 1. Sources and experimental conditions for solubility experiments on sodium

- silicates, zeolite A and amorphous precursor of zeolite A from which solubility data are
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used for computation of equilibrium constants

Solubility- controlling phase	Authors	T°C	Aqueous Solution, NaOH, Molal, m	ΣSi, Molal	ΣAl, Molal, m	Remarks
3Na ₂ O•2SiO ₂ •11H ₂ O	Sprauer and Pearce (1940)	25	12.3– 18.8	7.93×10 ⁻² to 3.62×10 ⁻¹	None	Approaching equilibrium: from supersaturation. Experimental duration: one month. Usage of data: calculation of Pitzer parameter
Na ₃ HSiO ₄ •5H ₂ O	Baker et al. (1950)	50	12.6– 19.6	2.78×10 ⁻¹ to 1.47	None	Approaching equilibrium: from both under- saturation and supersaturation. Experimental duration: a few weeks. Usage of data: calculation of Pitzer parameter
Na ₃ HSiO ₄ •2H ₂ O	Baker et al. (1950)	50, 70, 90	15.9-23.2	7.43×10 ⁻² to 2.14	None	Approaching equilibrium: from both under- saturation and supersaturation. Experimental duration: a few weeks. Usage of data: calculation of Pitzer parameter.
Zeolite A	Ciric (1968)	100	0.459– 1.72	6.26×10 ⁻³ to 2.76×10 ⁻²	1.31×10 ⁻¹ to 1.78×10 ⁻¹	Approaching equilibrium: from supersaturation. Experimental duration: longer than 100 min. Usage of data: calculation of

					log <i>K</i> for zeolite A
Zhdanov (1971)	90	1.63– 4.27	2.11×10^{-3} to 2.69×10^{-2}	2.69×10 ⁻² to 1.29	Approaching equilibrium: from undersaturation.
			2.07 10		Experimental duration: longer than 100 min
					Usage of data:
					log K for zeolite A.
Wieker and Fahlke (1985)	80	2.17	2.37×10 ⁻²	1.96×10 ⁻²	Approaching equilibrium: from undersaturation.
					Experimental duration longer: than 100 min.
					Usage of data: calculation of log <i>K</i> for zeolite A.
Čižmek et al. (1991a)	65, 70, 80	1.02– 2.06	1.17×10^{-2} to 1.83×10^{-2}	1.11×10^{-2} to 1.83×10^{-2}	Approaching equilibrium: from undersaturation.
	00		1.05*10	1.05*10	Experimental duration longer than 100 min.
					Usage of data: calculation of log <i>K</i> for zeolite A.
Myatt (1004)	90	2.49-	8.47×10 ⁻³	1.04×10^{-2}	Approaching equilibrium from
(1994)		3.34	6.13×10^{-2}	1.25×10^{-1}	supersaturation.
					Experimental duration: longer than 100 min.
					Usage of data: calculation of log <i>K</i> for zeolite A.
Ejaz and Graham	30, 50	3.04-	3.28×10^{-3}	4.32×10^{-3}	Approaching equilibrium: from
(1999)	65,	7.40	1.69×10 ⁻²	1.55×10^{-2}	undersaturation.
	80				Experimental duration: longer

						than 100 min.
						Usage of data: calculation of log K for zeolite A
	Kosanović	70	2 05	1.51×10^{-2}	1.46×10^{-2}	Approaching
	et al. (2000)	, 0	2.00	to 1.81×10 ⁻²	to 1.71×10 ⁻²	equilibrium: from undersaturation.
						Experimental duration: longer than 100 min.
						Usage of data: calculation of log K for zeolite A.
	Addai- Mensal et al. (2004)	30, 65	3.02– 6.31	1.19×10^{-2} to 4.63×10^{-2}	1.42×10^{-2} to 5.43×10^{-2}	Approaching equilibrium: from both under- saturation and supersaturation.
						Experimental duration: longer than 50 min.
						Usage of data: calculation of log <i>K</i> for zeolite A.
	Bosnar et al. (2005)	80	1.49– 1.58	$\begin{array}{c} 8.63 \times 10^{-3} \\ \text{to} \\ 6.15 \times 10^{-2} \end{array}$	2.06×10^{-3} to 1.03×10 ⁻²	Approaching equilibrium: from supersaturation.
						Experimental duration longer than 100 min.
						Usage of data: calculation of log <i>K</i> for zeolite A.
Amorphous precursor of zeolite A	Ejaz and Graham (1999)	30, 50, 65,	3.04– 4.46	4.70×10 ⁻² to 7.57×10 ⁻²	3.48×10^{-2} to 5.54×10^{-2}	Approaching equilibrium: from undersaturation.
		80				Experimental duration: longer than 100 min.
						Usage of data: calculation of log K for amorphous

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					zeolite A.
Addai- Mensal et al. (2004)	30, 65	3.02– 6.31	7.09×10^{-2} to 2.20×10 ⁻¹	7.99×10 ⁻² to 2.19×10 ⁻¹	Approaching equilibrium: from both under- saturation and supersaturation.
					Experimental duration: longer than 50 min.
					Usage of data: calculation of log K for amorphous
					zeolite A.

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Table 2. Equilibrium constants for silica and aluminum species considered in this study up to 100 °C

	uptor	00 C	
Reaction	Т, ^о С	log K	Reference
$H_4SiO_4 = H^+ + H_3SiO_4^-$	25	$-9.68 \pm 0.14^{\text{A}}$	Fleming and Crerar, 1982
	50	$-9.34 \pm 0.16^{\text{A}}$	Fleming and Crerar, 1982
	75	$-9.10 \pm 0.17^{\text{A}}$	Fleming and Crerar, 1982
	100	$-8.94 \pm 0.16^{\text{A}}$	Fleming and Crerar, 1982
$H_3SiO_4^- = H^+ + H_2SiO_4^{2-}$	25	-13.45 ± 0.07	Hershey and Millero, 1986
	50	$-12.95 \pm 0.25^{\text{B}}$	This study
	75	$-12.56 \pm 0.25^{\text{B}}$	This study
	100	$-12.28 \pm 0.25^{\text{B}}$	This study
$2H_4SiO_4 = H^+ + H_5Si_2O_7^- + H_2O$	25	-8.50°	Felmy et al., 2001
	50	$-8.14 \pm 0.25^{\text{B}}$	This study
	75	$-7.86 \pm 0.25^{\text{B}}$	This study
	100	$-7.65 \pm 0.25^{\text{B}}$	This study
$H_5Si_2O_7^- = H^+ + H_4Si_2O_7^{2-}$	25	-10.90 [°]	Felmy et al., 2001
	50	$-10.59 \pm 0.25^{\text{B}}$	This study
	75	$-10.38 \pm 0.25^{\text{B}}$	This study
	100	$-10.24 \pm 0.25^{\text{B}}$	This study
$3H_4SiO_4 = 3H^+ + H_5Si_3O_{10}^{3-} + 2H_2O$	25	-29.40 [°]	Felmy et al., 2001
	50	$-28.75 \pm 0.25^{\text{B}}$	This study
	75	$-28.34 \pm 0.25^{\text{B}}$	This study
	100	$-28.11 \pm 0.25^{\text{B}}$	This study
$Al(OH)_4^- = Al^{3+} + 4OH^-$	25	-34.05 ± 0.05	Derived from Wesolowski, 1992
	50	-33.44 ± 0.05	Derived from Wesolowski, 1992
	75	-33.11 ± 0.05	Derived from Wesolowski, 1992
	100	-32.99 ± 0.05	Derived from Wesolowski, 1992
$Al(OH)_4^- + H_2SiO_4^{2-} =$			This study ^D , estimated
$Al(OH)_3HSiO_4^{3-} + H_2O(l)$			from $\log \beta_1^I$ (Gout
			et al., 2000) with the SIT
	20	-0.42 ± 0.10	model
	25	-0.42 ± 0.15	This study ^D
	50	$-0.38 \pm 0.25^{\text{B}}$	This study ^D
	75	$-0.35 \pm 0.25^{\text{B}}$	This study ^D
	100	$-0.33 \pm 0.25^{\text{B}}$	This study ^D

^AUncertainties were not given in Fleming and Crerar (1982). Uncertainties are assigned 877

based on the respective benchmark values of Busey and Mesmer (1977). 878

^B An uncertainty of ± 0.25 is assigned to all predicted values at elevated temperatures up 879

to 100 °C. See text for details. 880

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^C Uncertainties were not given in Felmy et al. (2001). ^D The complex, $Al(OH)_3HSiO_4^{3-}$, is used to test whether it can improve the modeling at 882

30 °C and 50 °C only. 883

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Table 3. Pitzer interaction parameters employed in this study

	Binary Interaction Parameters								
	Interaction Pair	β ⁽⁰	$\beta^{(0)}$		C¢	References			
	Na ⁺ -H ₃ SiO ₄ ⁻	0.043±0	0.043±0.019			Hershey and Millero (1986)			
	Na ⁺ -H ₂ SiO ₄ ²⁻	0.32±0	0.08	0.13±0.50		Hershey and Millero (1986)			
	$Na^{+}-H_{5}Si_{2}O_{7}^{-}$	-0.0571	-0.0571±0.04			This Study ^A			
	$Na^{+}-H_{4}Si_{2}O_{7}^{2-}$	-0.0227	-0.0227±0.06			This Study ^B			
	$Na^{+}-H_{5}Si_{3}O_{10}^{3-}$	0.078±	0.078±0.03			This Study ^C			
	Na ⁺ -Al(OH) ₃ HSiO ₄ ³⁻	0.078±	0.078±0.03			This Study ^D			
	Na ⁺ –Al(OH) ₄ ⁻	0.05	0.051		-0.00090	Wesolowski (1992) ^E			
	Interaction Involving N	Neutral Spe	cies and	l Mixing Para	meters				
	Interaction Pair	λ_{ij}	λ_{ij}		Ψ_{ijk}	References			
	$Na^+ - H_4 SiO_4^0$	0.0925				Azaroual et al. (1997) ^E			
	OH ⁻ –H ₂ SiO ₄ ^{2–}		-0.08	812±0.003		This Study ^F			
			į	$\partial \theta_{ii}$					
			(-	$\frac{1}{\partial T})_{P} =$					
			-9.35	$\pm 0.44 \times 10^{-5}$					
	$OH^H_2SiO_4^{2-}-Na^+$				-0.017±0.02	This Study ^F			
887	^A Calculated from the e	estimation 1	nethod	of Plyasunov	et al. (1998) for	1:1 interaction,			
888	based on $\varepsilon(\text{Na}^+, \text{Si}_2\text{O}_2(\text{OH})_5^-)$ of -0.08 ± 0.04 , which is from Grenthe et al. (1992)								
889	In the method of Plyasunov et al. (1998) uncertainty was not given to $R^{(1)}$. The								
890	uncertainty assigned here is two standard deviations from the average $R^{(1)}$ for 1.1								
891	interaction computed in Plyasumov et al. (1998).								
892	^B Calculated from the estimation method of Plyasunov et al. (1998) for 1:2 interaction,								
893	based on $\varepsilon(\text{Na}^+, \text{Si}_2\text{O}_3(\text{OH})_4^{2-})$ of -0.15 ± 0.06 from Grenthe et al. (1992). In the								
894	method of Plyasunov et al. (1998), uncertainty was not given to $\beta^{(1)}$. The uncertainty								
895	assigned here is tw	o standard	deviatio	ons from the a	verage $\beta^{(1)}$ for 1	:2 interaction			
896	computed in Plyasi	umov et al.	(1998).						
897	^C Calculated from the e	estimation r	nethod	of Plyasunov	et al. (1998) for	1:3 interaction,			
898	based on $\varepsilon(Na^+, Si_2)$	₃ O ₆ (OH) ₃ ³⁻) of -0.2	25 ± 0.03 from	n Grenthe et al.	(1992). In the			
899	method of Plyasun	ov et al. (19	998), un	certainty was	not given to $\beta^{(1)}$	⁾ . The uncertainty			
900	assigned here is tw	o standard	deviatio	ons from the a	verage $\beta^{(1)}$ for 1	:3 interaction			
901	computed in Plyas	umov et al.	(1998).						
902	^D Interaction parameter	rs are assig	ned to b	be the same as	those for Na ⁺ –I	$H_5 Si_3 O_{10}^{3-}$.			
903	^E Uncertainties were no	ot given in	Wesolo	wski (1992) a	nd Azaroual et	al. (1997),			
904	respectively.		_						
905	^T Evaluated from solub	oility data o	$f 3Na_2C$	$-2SiO_2-11H_2$	O in NaOH solu	itions up to ~ 19 m			
906	from Sprauer and I	Pearce (194	0). The	e temperature	dependence of 6	$\Theta_{ij}, (\frac{\partial \Theta_{ij}}{\partial T})_P$, is			
907	evaluated from solution	ubility data	of Na ₃ l	HSiO ₄ •5H ₂ O	at 50° C, and of]	Na ₃ HSiO ₄ •2H ₂ O at			
908	50°C, 70°C, and 90	^o C in NaO	H soluti	ons up to ~ 24	m from Baker	et al. (1950).			
909									
910									

912	experiments in this study*					
	T, °C	$\log K \pm 2\sigma$	Reaction			
	25	83.83 ± 0.62	$3Na_2O \cdot 2SiO_2 \cdot 11H_2O + 6H^+ = 6Na^+ + 2H_4SiO_4(aq) + 10H_2O(l)$			
	50	47.03 ± 0.35	$Na_3HSiO_4 \cdot 2H_2O + 3H^+ = 3Na^+ + H_4SiO_4(aq) + 2H_2O(l)$			
	70	45.69 ± 0.31				
	90	46.03 ± 0.30				
	50	41.50 ± 0.35	$Na_{3}HSiO_{4} \cdot 5H_{2}O + 3H^{+} = 3Na^{+} + H_{4}SiO_{4}(aq) + 5H_{2}O(l)$			
	25	10.24 ± 0.31 ^A				
	30	10.23 ± 0.31				
	50	7.95 ± 0.30	NaAlSiO ₄ •2.25H ₂ O(cr, zeolite A) + 4H ⁺ = Na ⁺ + Al ³⁺ + H ₄ SiO ₄ (aq) + 2.25H ₂ O			
	65	6.70 ± 0.33				
	70	6.54 ± 0.30				
	80	5.74 ± 0.45				
	90	5.54 ± 0.36				
	100	5.34 ± 0.30				
	25	12.68 ± 0.35 ^B	NaAlSiO ₄ •2.25H ₂ O(am, zeolite A) + 4H ⁺ = Na ⁺ + Al ³⁺ + H ₄ SiO ₄ (aq) + 2.25H ₂ O			
	30	11.94 ± 0.20				
	50	9.30 ± 0.30				
	65	7.84 ± 0.35				
	80	5.82 ± 0.22				
913	* Experimental conditions for hydrothermal experiments from which solubility data are					
914	used for computation of equilibrium constants are detailed in text and Table 1.					
915	^A Extrapola	^A Extrapolated to the reference temperature, 25 °C, based on the linear relation between				
916	log K and reciprocal temperature in Kevin for zeolite A.					

911 Table 4. Equilibrium constants of sodium silicates and zeolites retrieved from solubility

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^B Extrapolated to the reference temperature, 25 °C, based on the linear relation between log *K* and reciprocal temperature in Kevin for amorphous precursor of zeolite A.

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Table 5. Enthalpy changes for reactions involving zeolites derived in this study

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Reactions	Δ_r H, kJ mol ⁻¹ *
$Na_{3}HSiO_{4} \cdot 2H_{2}O + 3H^{+} = 3Na^{+} + H_{4}SiO_{4}(aq) + 2H_{2}O(l)$	$-58 \pm 45 (2\sigma)$
$NaAlSiO_4 \cdot 2.25H_2O(cr, zeolite A) + 4H^+ = Na^+ + Al^{3+} + H_4SiO_4(aq) + H_$	
2.25H ₂ O(1)	$-152 \pm 5 (2\sigma)$
$NaAlSiO_4 \cdot 2.25H_2O(am, zeolite A) + 4H^+ = Na^+ + Al^{3+} + H_4SiO_4(aq) + H_4SiO_4(ad) + H_4SiO_4(ad) + H_$	
2.25H ₂ O(1)	$-248 \pm 3 (2\sigma)$

924 * Uncertainties account for the errors from regressions only. The overall uncertainties

925 could be higher than those provided here.

929 Table 6. Thermodynamic properties of zeolite A and the amorphous form of zeolite A at

930

25 °C and 1 bar derived in this study*

Properties	Values $(\pm 2\sigma)$	Remarks
$\Delta_{f} H^{o}$, Zeolite A,	$-2738 \pm 5 \text{ kJ mol}^{-1}$	Based on Δ_r H derived from
$NaAlSiO_4 \cdot 2.25H_2O(cr)$		temperature dependence of
		equilibrium constant
$\Delta_f G^o$, Zeolite A,	$-2541 \pm 2 \text{ kJ mol}^{-1}$	Based on $\Delta_r G$ derived from log K
$NaAlSiO_4 \cdot 2.25H_2O(cr)$		extrapolated to 25 °C.
S ^o , Zeolite A,	$373 \pm 10 \text{ J K}^{-1} \text{ mol}^{-1}$	Based on $\Delta_r S$ calculated from the
$NaAlSiO_4 \cdot 2.25H_2O(cr)$		Gibbs-Helmholtz equation
$\Delta_{f} H^{o}$, Amorphous	$-2642 \pm 3 \text{ kJ mol}^{-1}$	Based on Δ_r H derived from
precursor of zeolite A,		temperature dependence of
NaAlSiO ₄ •2.25H ₂ O(am)		equilibrium constant
$\Delta_f G^{\circ}$, Amorphous	$-2527 \pm 2 \text{ kJ mol}^{-1}$	Based on $\Delta_r G$ derived from log K
precursor of zeolite A,		extrapolated to 25 °C.
$NaAlSiO_4 \cdot 2.25H_2O(am)$		
S°, Amorphous	$648 \pm 10 \text{ J K}^{-1} \text{ mol}^{-1}$	Based on $\Delta_r S$ calculated from the
precursor of zeolite A,		Gibbs-Helmholtz equation
NaAlSiO ₄ •2.25H ₂ O(am)		

931 * All properties refer to formation from elements.

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935 Figure Captions

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Figure 1. Comparison of model predicted solubilities of 3Na₂O•2SiO₂•11H₂O at 25 °C,
Na₃HSiO₄•5H₂O, and Na₃HSiO₄•2H₂O at 50 °C, with experimental values. The size of
error bars is equal to or smaller than the symbol size.

940

Figure 2. Comparison of model predicted solubilities of $Na_3HSiO_4 \cdot 2H_2O$ at 70 °C and 90 °C with experimental values. The size of error bars is equal to or smaller than the 943 symbol size.

944

Figure 3. A plot showing equilibrium constants of Na₃HSiO₄•2H₂O as a function of
reciprocal temperatures in Kevin.

Figure 4. Comparison of model predicted solubilities of amorphous silica with
experimental values, which are independent from the model development. The size of
error bars is equal to or smaller than the symbol size.

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Figure 5. Speciation of silica at different total concentrations of silica as a function of pH
at 25 °C. (A) total silica concentration is 0.01 m; (B) total silica concentration is 0.1 m;
and (C) total silica concentration is 1 m. Notice that in acidic pH range, the solution is
supersaturated with amorphous silica.

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Figure 6. Speciation of silica at different total concentrations of silica as a function of pH
at 100 °C. (A) total silica concentration is 0.01 m; (B) total silica concentration is 0.1 m;
and (C) total silica concentration is 1 m. Notice that in acidic pH range, the solution is
supersaturated with amorphous silica/quartz.

961

Figure 7. Comparison of model predicted solubilities of zeolite A with experimental values. (A) at 30 °C, 50 °C and 65 °C; (B) at 70 °C and 80 °C; and (C) at 90 °C and 100 °C. The size of error bars is equal to or smaller than the symbol size.

965

Figure 8. A plot showing equilibrium constants of zeolite A and amorphous precursor of
zeolite A as a function of reciprocal temperatures in Kevin. The size of error bars is
equal to or smaller than the symbol size.

969

Figure 9. Comparison of model predicted solubilities of amorphous precursor of
zeolite A with experimental values. (A) at 30 °C and 50 °C; and (B) at 65 °C and 80 °C.
The size of error bars is equal to or smaller than the symbol size.

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1010 Figure 5C

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1021 Figure 6B

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