A Thermodynamic Model for Silica and Aluminum in Alkaline Solutions with High Ionic Strength at Elevated Temperatures up to 100 °C: Applications to Zeolites, Revision 1

Yongliang Xiong*

Sandia National Laboratories
Carlsbad Programs Group
4100 National Parks Highway
Carlsbad, New Mexico 88220, USA

* Corresponding author, e-mail: yxiong@sandia.gov
Abstract

In this study, a thermodynamic model for silica and aluminum in high ionic strength solutions at elevated temperatures up to 100 °C is constructed. Pitzer equations are utilized for the thermodynamic model construction. This model is valid up to ionic strengths of ~24 m in NaOH solutions with silicate concentrations up to ~1.5 m. The speciation of silica (including monomers and polymers) and aluminum at elevated temperatures is taken into account. Also, the equilibrium constants for silicic acid and its polymer species (H₄SiO₄, H₅Si₂O₇⁻, H₄Si₂O₇²⁻, and H₅Si₃O₇₃⁻) at elevated temperatures up to 100 °C, are obtained based on theoretical calculations. Using this thermodynamic model, thermodynamic properties, including equilibrium constants, and respective reaction enthalpies are obtained for sodium silicates, zeolite A, and the amorphous form of zeolite A, based on solubility experiments at elevated temperatures. The equilibrium constants for zeolite A and amorphous precursor of zeolite A regarding the following reactions up to 100 °C,

NaAlSiO₄•2.25H₂O(cr) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O(l) (I)

and

NaAlSiO₄•2.25H₂O(am) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O(l) (II)

can be expressed as follows

\[ \log K_f = \frac{7963 \pm 327}{T} - 16.46 \pm 0.96 \] (III)

and
where $T$ is temperature in Kevin.

The enthalpy of formation from elements, Gibbs free energy of formation from elements, and standard entropy derived for zeolite A and the amorphous form of zeolite A with the chemical formulas mentioned above at 25 °C and 1 bar are $-2738 \pm 5$ kJ mol$^{-1}$, $-2541 \pm 2$ kJ mol$^{-1}$, $373 \pm 10$ J K$^{-1}$ mol$^{-1}$; and $-2642 \pm 3$ kJ mol$^{-1}$, $-2527 \pm 2$ kJ mol$^{-1}$, and $648 \pm 10$ J K$^{-1}$ mol$^{-1}$, respectively. The enthalpy of formation from elements for zeolite A derived in this study based on solubility experiments in hydrothermal solutions agrees well with those obtained by calorimetric measurements and by theoretical calculations.
1. Introduction

Zeolites have numerous applications ranging from their usage as absorbents and catalysts to wasteforms for radioactive iodine (\(^{129}\text{I}\)) and for contaminated electrolytes from electrorefinery of used nuclear fuel (e.g., Sheppard et al., 2006). The formation of zeolites occurs under alkaline conditions in solutions typically with high ionic strengths. For instance, in the hydrothermal synthesis of zeolite A, zeolite X and mordenite, NaOH solutions ranging from 0.6 m to 2 m were frequently used (e.g., Čižmek et al., 1991a; 1991b; 1992; Šefčík and McCormick, 1997a). It has also been noted that in high-level nuclear waste processing, concentrated waste liquor is produced by evaporation at elevated temperatures up to 140 °C, and the resulting concentrated waste liquor can have NaOH concentrations up to \(~7\) m (Addai-Mensah et al., 2004). In laboratory experiments designed to simulate the concentrated waste liquor from high-level nuclear waste processing, the precipitation of zeolite A and the amorphous form zeolite A from the simulated concentrated waste liquor has been observed (Addai-Mensah et al., 2004).

Finally, at the Hanford site in Washington State, the leakage of the alkaline solutions with high ionic strengths contained in the waste tanks into sediments resulted in formation of cancrinite and sodalite, which belong to the zeolite family, in the sediment below a waste tank (Chorover et al., 2003; Mashal et al., 2004; Serne et al., 2002; Zhao et al., 2004).

While there are experimental data on zeolites (see Table 1), the formation of zeolites at elevated temperatures is not well understood owing to the lack of reliable thermodynamic model(s). Furthermore, there is the current inability to address simultaneously the behavior of both silica and aluminum, the two most rock-forming
elements, in high ionic strength solutions at elevated temperatures. The high ionic
strength nature of solutions and high concentrations of silica are the factors contributing
to the difficulties associated with the development of a thermodynamic model at elevated
temperatures for zeolites. For instance, when zeolite A is formed, the concentration of
silica can reach up to 0.12 m (Šefčík and McCormick, 1997). The nature of high ionic
strength solutions requires that the activity coefficient model be valid to high ionic
strengths. The high concentrations of silica require that polymers of silica species be
considered, as polymerized silica species might become important when silica
concentrations are higher than 0.01 m (e.g., Felmy et al., 2001).

A thermodynamic model applicable to concentrated solutions at elevated
temperatures is therefore needed in order to predict the formation of zeolites including
zeolite A, amorphous precursor of zeolite A, cancrinite and sodalite, at elevated
temperatures. In addition, such a model would provide valuable guidance in the
synthesis of zeolites. A thermodynamic model would also provide a better understanding
of interactions of waste solutions with the sediments, like the interactions observed in the
Hanford studies mentioned above. Consequently, the objective of this study is to develop
a thermodynamic model for aluminum and silica species at elevated temperatures valid to
high ionic strengths. Using this model, equilibrium constants for sodium silicates,
zeolite A, and amorphous precursor of zeolite A can be retrieved from hydrothermal
experiments. Applications of the model to other zeolite species will be presented in the
future.
2. The Thermodynamic Model

In this study, the standard state for a solid phase is defined as its pure end-member with unit activity at temperatures and pressures of interest. The standard state of the solvent in aqueous solutions is pure solvent at temperatures and pressures of interest. The standard state for an aqueous solute is a hypothetical 1 molal (m) solution referred to infinite dilution at temperatures and pressures of interest. Gibbs free energies and enthalpies of formation reported in this study correspond to the formation from chemical elements at their reference states.

The Pitzer model is adopted in this study for calculations of activity coefficients of aqueous species. The detailed descriptions about Pitzer equations are provided in Pitzer (1991). In the following, an equation for calculation of the activity coefficient of $\text{Al(OH)}_4^-$ in NaOH medium without consideration of triple interactions in the Pitzer model is provided as an example,

\[
\ln \gamma = -A \phi \left[ \frac{\sqrt{I_m}}{1 + 1.2 \sqrt{I_m}} + \frac{2}{1.2} \ln (1 + 1.2 \sqrt{I_m}) \right] + m \left[ 2 \beta^{(0)} + \frac{2 \beta^{(1)}}{\alpha^2 \times I_m} \left[ 1 - (1 + \alpha \sqrt{I_m} - \frac{\alpha^2 I_m}{2} e^{-\alpha \sqrt{I_m}}) \right] \right] + \frac{3m^2}{2} C^\phi
\]

where $\gamma$ is activity coefficient, $A \phi$ is Deby-Huckel slope for osmotic coefficient; $I_m$ is ionic strength on molality scale; $m$ is molality of $\text{Na}^+$; $\alpha$ is equal to 2; and $\beta^{(0)}$, $\beta^{(1)}$ and $C^\phi$ are Pitzer binary interaction coefficients between $\text{Na}^+$ and $\text{Al(OH)}_4^-$. The uncertainties reported in this study are two standard deviations ($2\sigma$). Error propagations are calculated based on uncertainties associated with regressions and
equilibrium constants in the model. In some cases, especially for $\Delta_f G$, $\Delta_f H$, and $S^0$, uncertainties could be underestimated, as uncertainties for auxiliary data are not available, and therefore not included.

2.1 Thermodynamic Constants and Pitzer Interaction Parameters of Si and Al Aqueous Species Chosen and Extrapolated from the Literature

The following monomer, dimer and trimer silica species are considered in this thermodynamic model:

1. $H_4SiO_4 = H^+ + H_3SiO_4^-$ (2)
2. $H_3SiO_4^- = H^+ + H_2SiO_4^{2-}$ (3)
3. $2H_4SiO_4 = H^+ + H_5Si_2O_7^- + H_2O$ (4)
4. $H_5Si_2O_7^- = H^+ + H_4Si_2O_7^{2-}$ (5)
5. $3H_4SiO_4 = 3H^+ + H_5Si_3O_{10}^{3-} + 2H_2O$ (6)

The equilibrium constants for Reaction 2 are obtained from Fleming and Crerar (1982) (Table 2). The first dissociation constants of monomeric silicic acid from Fleming and Crerar (1982) are selected, as these authors regressed a number of experimentally determined first dissociation constants for monomeric silicic acid from previous studies, including the most reliable values determined via the solubility and potentiometric
methods (e.g., Seward, 1974; Busey and Mesmer, 1977). However, uncertainties associated with their dissociation constants were not given in Fleming and Crerar (1982). In this study, uncertainties are assigned based on the respective benchmark values of Busey and Mesmer (1977). The second dissociation constants (log $K_2$) of monomeric silicic acid at elevated temperatures (Reaction 3) are estimated by using the one-term isocoulombic approach with phosphoric acid as the model substance as employed before (e.g., Xiong, 2003; 2007) (Table 2), based on the log $K_2$ at 25 °C from Hershey and Millero (1986). The equilibrium constants at 25 °C for dimer and trimer silica species (Reactions 4–6) are from Felmy et al. (2001). The log $K$ at elevated temperatures for those dimer and trimer silica species are also predicted in this study in a manner similar to that described above for the log $K_2$ of monomer silicic acid. They are tabulated in Table 2. The uncertainty for log $K$’s of well balanced isocoulombic reactions is usually within ± 0.50 up to 200 °C (Gu et al., 1994). Therefore, an uncertainty of ± 0.25 is assigned to all predicted values at elevated temperatures up to 100 °C.

An example of a well-balanced isocoulombic reaction is given in the following,

$$H_4SiO_4 + H_2PO_4^- = H_3PO_4 + H_3SiO_4^- \quad (7)$$

In the above reaction, $H_4SiO_4$ with zero charge on the left side is balanced by $H_3PO_4$ with zero charge on the right side. Similarly, $H_2PO_4^-$ with one negative charge on the left side is balanced by $H_3SiO_4^-$ with one negative charge on the right side.

It should be noted that regarding polymers of silica species, only dimer and trimer are considered in this study, and tetramer and hexamer are excluded. This consideration
is primarily based on numerous studies which have indicated that monomer, dimer, and
trimer are adequate in descriptions of silica solutions with high concentrations (e.g.,
Čižmek et al., 1992; Sefčík and McCormic, 1997b; Hunt et al., 2011). For instance, in a
Raman spectroscopic study on silica speciation in concentrated silica solutions up to
5.0 m with KOH as a supporting solution ranging from 0.08 m to 8.0 m conducted by
Hunt et al. (2011), the authors demonstrate that monomer, dimer and trimer species are
sufficient to describe the silica species in highly concentrated silicate solutions with total
silica concentrations up to 5.0 m.

The following species are incorporated into the aluminum thermodynamic model:

\[
\text{Al(OH)}_4^- = \text{Al}^{3+} + 4\text{OH}^- \quad (8)
\]

Equilibrium constants for Reaction 8 at various temperatures are tabulated in Table 3.

The equilibrium constants for Reaction 8 are obtained from Wesolowski (1992)
based on the following reaction:

\[
\text{Al(OH)}_4^- = \text{Al(OH)}_3(\text{cr, gibbsite}) + \text{OH}^- \quad (9)
\]

in combination with the equilibrium constants for the following reaction calculated from
SUPCRT (Johnson et al., 1992):

\[
\text{Al(OH)}_3(\text{cr, gibbsite}) = \text{Al}^{3+} + 3\text{OH}^- \quad (10)
\]
In this study, the hydrolysis constants of Al(OH)$_4^-$ described above determined by Wesolowski (1992) are selected because they are consistent with the respective Pitzer parameters adopted in this study. Although Tagirov and Schott (2001) provided the revised hydrolysis constants for Al(OH)$_4^-$, those values are not adopted in this study, as the activity coefficient model employed by Tagirov and Schott (2001) is different from the Pitzer model.

In the present model, regarding aluminum species in alkaline solutions, only Al(OH)$_4$ is included, and no aluminum polymeric species are considered. This consideration is based on the fact that aqueous aluminum species exists as Al(OH)$_4$ in neutral and basic solutions when $\Sigma Al$ is less than 1.5 m (Moolenaar et al., 1970; Baes and Mesmer, 1976; Castet et al., 1993). The formation of aluminum polymeric species requires $\Sigma Al \geq 1.5$ m (Moolenaar et al., 1970). In the presence of silica, aluminum concentrations are much lower than 1.5 m. As solutions in which zeolites are formed contain both silica and aluminum, aluminum polymeric species will not be important.

It should be noted that although Zhou et al. (2003) also determined the Pitzer interaction parameters for the interaction between Na$^+$ and Al(OH)$_4^-$, the Pitzer parameters of Wesolowski (1992) are adopted in the current model as they are consistent with the hydrolysis constants of Al(OH)$_4^-$ used in this study.

In addition, there are some studies suggesting that the aqueous Al-Si complex(es) such as Al(OH)$_3$H$_3$SiO$_4^-$ and SiAlO$_3$(OH)$_4^-$ might be present in aqueous solutions (e.g., Pokrovski et al., 1998; Salvi et al., 1998; Gout et al., 2000). Salvi et al. (1998) suggested that Al(OH)$_3$H$_3$SiO$_4^-$ with log $\beta_1 = 2.32$ could be present at 300 °C, corresponding to the following reaction,
\[
\text{Al(OH)}_4^- + \text{H}_4\text{SiO}_4(\text{aq}) = \text{Al(OH)}_3\text{H}_3\text{SiO}_4^- + \text{H}_2\text{O}(l) \quad (11)
\]

Gout et al. (2000) mentioned that a weak complex, SiAlO$_3$(OH)$_4^{3-}$ (also formulated as Al(OH)$_3$HSiO$_4^{3-}$) with log $\beta_i = 0.53$ (I = 1.2 m), could be present at 20 °C, with reference to the following reaction,

\[
\text{Al(OH)}_4^- + \text{H}_2\text{SiO}_4^{2-} = \text{Al(OH)}_3\text{HSiO}_4^{3-} + \text{H}_2\text{O}(l) \quad (12)
\]

The $\beta_i$ at infinite dilution can be expressed as,

\[
\beta_i = \frac{(m_{\text{Al(OH)}_3\text{HSiO}_4^{3-}})(\gamma_{\text{Al(OH)}_3\text{HSiO}_4^{3-}})(a_{\text{H}_2\text{O}})}{(m_{\text{Al(OH)}_4^-})(m_{\text{H}_2\text{SiO}_4^{2-}})(\gamma_{\text{H}_2\text{SiO}_4^{2-}})} = \beta_i' \times \frac{(\gamma_{\text{Al(OH)}_3\text{HSiO}_4^{3-}})(a_{\text{H}_2\text{O}})}{(\gamma_{\text{Al(OH)}_3\text{HSiO}_4^{3-}})(\gamma_{\text{H}_2\text{SiO}_4^{2-}})}
\]

Other studies (e.g., Yokoyama et al., 1988) suggested that the formation of aqueous complex(es) of aluminum with silica is likely in dilute NaOH solutions at room temperature, but such a complex is absent in high alkaline solutions such as 1.0 m NaOH, typical of solutions in which zeolites are stable. Yokoyama et al. (1988) also demonstrated that higher temperatures destabilize the complex of aluminum and silicate that formed in dilute NaOH solutions.
However, since the presence or absence of Al-Si complexes is an important issue, as an independent constrain on the existence of Al-Si complexes in alkaline solutions, the $\text{Al(OH)}_3\text{HSiO}_4^{3-}$ complex is tested to see whether it can improve the modeling at 30 °C and 50 °C (see following sections), as this complex was proposed to be present at 20 °C. In doing this, as there is only one apparent formation constant at one ionic strength (i.e., 1.2 m) in NaCl, the $\log \beta_i$ at infinite dilution at 20 °C is first estimated by using the Brønsted-Guggenheim-Scatchard SIT model (Grenthe et al., 1992). According to the SIT model regarding Eqs. 12-13, we have,

$$\log \beta_i = \log \beta_i^I - 4D + \log a_{H_2O} + \Delta \varepsilon(Eq.11) \times I_m$$

$$\Delta \varepsilon(Eq.11) = \varepsilon(\text{Na}^+, \text{Al(OH)}_3\text{HSiO}_4^{3+}) - \varepsilon(\text{Na}^+, \text{Al(OH)}_3^-) - \varepsilon(\text{Na}^+, \text{H}_2\text{SiO}_4^{2-})$$

where $\log \beta_i^I$ is the formation constant regarding Reaction 12 at a certain ionic strength defined in Eq. (13); $I_m$ ionic strength on molal scale; $\varepsilon$’s are the SIT interaction coefficients; $a_{H_2O}$ is activity of water; and D is the Debye-Hückel term given below,

$$D = \frac{A_\gamma \sqrt{I_m}}{1 + 1.5 \sqrt{I_m}}$$

where $A_\gamma$ the Debye-Hückel slope for activity coefficient.
Substituting $\log \beta_1^{\gamma} = 0.53$ at 20 °C from Gout et al. (2000), $A_{\gamma} = 0.5059$ at 20 °C from Helgeson and Kirkham (1974), $a_{H_2O} = 0.9600$ for 1.2 m NaCl calculated using

EQ3/6, $\varepsilon(Na^+, Al(OH)_3 HSiO_4^{--}) = \varepsilon(Na^+, Si_2O_6(OH)_3^{--}) = -0.25 \pm 0.03$, 

$\varepsilon(Na^+, H_2SiO_4^{2--}) = \varepsilon(Na^+, SiO_2(OH)_2^{2--}) = -0.10 \pm 0.07$, and 

$\varepsilon(Na^+, Al(OH)_4^{--}) = \varepsilon(Na^+, B(OH)_4^{--}) = -0.07 \pm 0.05$, all from Grenthe et al. (1992), and

$I_m = 1.2$ m into Eqs. 14–16, $\log \beta_1$ is estimated as $-0.42 \pm 0.10$ at 20 °C. As Reaction 12 is in a semi-isocoulombic form, the $\log \beta_1$ at 20 °C is directly extrapolated to 25 °C and other temperatures using the one-term isocoulombic principle (Table 2).

2.2 Pitzer Interaction Parameters in the Na-Si-OH Systems Obtained in This Study

In this study, the Pitzer binary interaction parameters for $Na^+–H_5Si_2O_7^{--}$, $Na^+–H_4Si_2O_7^{2--}$, and $Na^+–H_5Si_3O_{10}^{3--}$ are calculated according to the method of Plyasumov et al. (1998) based on the respective SIT coefficients for these interactions (Table 3).

The high order Pitzer parameters, $\theta_{ij}$ for $OH^{--}–H_2SiO_4^{2--}$ interaction, and $\Psi_{ijk}$ for $OH^{--}–H_2SiO_4^{2--}–Na^+$ interaction, are evaluated from solubility data of the sodium silicate, $3Na_2O\cdot2SiO_2\cdot11H_2O$, in NaOH solutions up to ~19 m from Sprauer and Pearce (1940).

In addition, $(\frac{\partial \theta_i}{\partial T})_p$ for $OH^{--}–H_2SiO_4^{2--}$ interaction is evaluated from solubility data of $Na_3HSiO_4\cdot5H_2O$ and $Na_3HSiO_4\cdot2H_2O$ at elevated temperatures from Baker et al. (1950) (Table 3).
For the testing purpose, the Pitzer binary interaction parameters for Na\(^+\)-Al(OH)\(_3\)HSiO\(_4\)\(^{3-}\) are assumed to be the same as those for Na\(^+\)–H\(_5\)Si\(_3\)O\(_{10}\)\(^{3-}\) (Table 3).

Sprauer and Pearce (1940) experimentally determined solubilities of 3Na\(_2\)O•2SiO\(_2\)•11H\(_2\)O in NaOH solutions at 25 °C. Their experiments approached equilibrium from supersaturation in about one month (Table 1). Their experiments were in NaOH solutions with very high concentrations up to ~19.0 m. The solubility reaction for 3Na\(_2\)O•2SiO\(_2\)•11H\(_2\)O can be expressed as follows, using H\(_4\)SiO\(_4\)(aq) as a basis species for silica,

\[
3\text{Na}_2\text{O} \cdot 2\text{SiO}_2 \cdot 11\text{H}_2\text{O} + 6\text{H}^+ = 6\text{Na}^+ + 2\text{H}_4\text{SiO}_4(aq) + 10\text{H}_2\text{O}(l)
\] (17)

The log \( K \) for Reaction 17 is evaluated as 83.83 ± 0.62 (Table 3) along with \( \theta_{ij} \) for OH\(^-\)-H\(_2\)SiO\(_4\)\(^{2-}\) interaction, and \( \Psi_{ijk} \) for OH\(^-\)-H\(_2\)SiO\(_4\)\(^{2-}\)–Na\(^+\) interaction (Table 4).

In Figure 1, predicted solubilities of 3Na\(_2\)O•2SiO\(_2\)•11H\(_2\)O at 25 °C are compared with experimental data from Sprauer and Pearce (1940). From the figure, it is clear that the model reproduces experimental data generally within a factor of ~3.

Baker et al. (1950) investigated solubilities of Na\(_3\)HSiO\(_4\)•2H\(_2\)O and Na\(_3\)HSiO\(_4\)•5H\(_2\)O in NaOH solutions up to ~24 m at elevated temperatures to 90 °C. Their experiments approached equilibrium from both undersaturation and supersaturation in a few of weeks. The dissolution of Na\(_3\)HSiO\(_4\)•5H\(_2\)O can be expressed as:

\[
\text{Na}_3\text{HSiO}_4 \cdot 5\text{H}_2\text{O} + 3\text{H}^+ = 3\text{Na}^+ + \text{H}_4\text{SiO}_4(aq) + 5\text{H}_2\text{O}(l)
\] (18)
Based on solubility data from Baker et al. (1950) on Na$_3$HSiO$_4$•5H$_2$O, the equilibrium constant for Na$_3$HSiO$_4$•5H$_2$O at 50 °C is obtained as 41.50 ± 0.35 (Table 4).

Similarly, the dissolution of Na$_3$HSiO$_4$•2H$_2$O can be expressed as:

$$\text{Na}_3\text{HSiO}_4 \cdot 2\text{H}_2\text{O} + 3\text{H}^+ = 3\text{Na}^+ + \text{H}_4\text{SiO}_4(\text{aq}) + 2\text{H}_2\text{O}(\text{l})$$ (19)
In this study, it is assumed that interaction parameters are constant over the temperature range from 25 °C to 100 °C, except for $\theta_{ij}$ for OH$^-$–H$_2$SiO$_4$$^{2-}$ interaction, which is necessary to modeling solubility of sodium silicates in highly concentrated NaOH solutions. This assumption is based on the observation that the Pitzer interaction parameters do not change significantly over a narrow range of temperature. For instance, for the interaction of Na$^+$ with Al(OH)$_4^-$, the temperature derivatives of Pitzer interaction parameters are very small over this temperature range, i.e., $(\frac{\partial \beta^{(0)}}{\partial T})_p = 8.0 \times 10^{-5}$, $(\frac{\partial \beta^{(1)}}{\partial T})_p = 2.7 \times 10^{-4}$, and $(\frac{\partial \phi^{(0)}}{\partial T})_p = 6.9 \times 10^{-5}$, based on the respective interaction parameters as a function of temperature from Wesolowski (1992). In addition, in the validation test (see the following section), model-predicted values are in satisfactory agreement with independent experimental values.

The data0.PIT database in the EQ3/6 code can be used to compute activity coefficients by using the Pitzer equations up to 100 °C (Wolery, 1992). The original data0.PIT database does not contain the following species: H$_4$SiO$_4$(aq), H$_3$SiO$_4^-$, H$_2$SiO$_4$$^{2-}$, H$_2$Si$_2$O$_7^-$, H$_4$Si$_2$O$_7$$^{2-}$, H$_5$Si$_3$O$_{10}^-$, and Al(OH)$_4^-$. By incorporating the equilibrium constants for the respective reaction detailed in Table 2, and the relevant Pitzer interaction parameters from Hershey and Millero (1986), Wesolowski (1992), Azarousal et al. (1997), and this study tabulated in Table 3 into the data0.PIT database, the PIT database is modified to be able to model high ionic strength solutions with high concentrations of Si and Al up to 100 °C.
2.3 Model Validation

A validation test is performed for the Na-Si-OH model developed in this study. For this purpose, predicted solubilities of amorphous silica in alkaline solution with a wide range ionic strengths are compared with experimental data, which are independent from the model development. The dissolution of amorphous silica can be expressed as,

\[
\text{SiO}_2(\text{am}) + 2\text{H}_2\text{O} = 2\text{H}^+ + \text{H}_2\text{SiO}_4^{2-}
\]  

(20)

The log $K$ for Reaction 20, is -25.81 from Weber and Hunt (2003). In Figure 4, predicted solubilities of amorphous silica are compared with experimental data in NaOH solutions at 25 °C from Alexander et al. (1954), and in NaOH + NaNO$_3$ mixtures at 25 °C from Weber and Hunt (2003). From Figure 4, it is obvious that solubilities predicted by the model are in good agreement with model-independent experimental data over the entire ionic strength range from very dilute to ~4.5 m in alkaline solutions.

3. Model Applications

In the following, the model developed above is applied to calculations of speciation of silica species as a function of pH at temperatures of 25 °C and 100 °C, and to retrieval of thermodynamic data from hydrothermal solubility data on zeolite A and amorphous precursor of zeolite A.
3.1 Speciation of Silica Species

In Figure 5A-C, speciation of silica species as a function of pH at 25 °C at different total silica concentrations is displayed. At $\Sigma Si = 0.01$ m, monomer species are the dominant species (Figure 5A). In the pH range from ~7 to ~9, the polymeric species $H_3Si_2O_7^-$ can account for up to 2% of the total dissolved silica (Figure 5A). At $\Sigma Si = 0.1$ m, although monomer species are still the dominant species, contributions from polymeric species become significant (Figure 5B). In the pH range from ~6 to ~10, $H_3Si_2O_7^-$ can account for up to ~20% of the total dissolved silica. In the pH range from ~9 to ~13, both $H_4Si_2O_7^{2-}$ and $H_5Si_3O_{10}^{3-}$ can account for up to ~5% of the total dissolved silica (Figure 5B). At $\Sigma Si = 1$ m, contributions from polymeric species to the total dissolved silica become important (Figure 5C). In the pH range up to 9, $H_5Si_2O_7^-$ can account for up to ~50% of the total dissolved silica. In the pH range from 8 to 14, $H_5Si_3O_{10}^{3-}$ can account for up to ~40% of the total dissolved silica, and $H_4Si_2O_7^{2-}$ can account for up to ~10% of the total dissolved silica. However, the trend indicates that $H_2SiO_4^{2-}$ will be the dominant species above pH 14 even at $\Sigma Si = 1$ m.

Similarly, speciation of silica at different total concentrations of silica as a function of pH at 100 °C is displayed in Figure 6A-C. The polymeric silica species are important when $\Sigma Si = 1$ m (Figure 6C). However, when pH higher than 11.5, the contributions from polymeric silica species diminish, and above pH 13, the contributions from polymeric silica species become insignificant (Figure 6C).

It should be mentioned that under highly basic conditions where zeolites are formed as discussed below, $H_2SiO_4^{2-}$ is the dominant species.
3.2 Calculation of Equilibrium Constants of Zeolites

Using the above thermodynamic model, the thermodynamic equilibrium constants of zeolites can be retrieved from solubility experiments at elevated temperatures. The main criterion of selection of solubility data is that equilibrium state must be attained. For precipitation and dissolution of zeolites at elevated temperatures, detailed kinetic studies have demonstrated that equilibrium state is rapidly attained. Addai-Mensah et al. (2004) performed detailed experiments from both undersaturation and supersaturation on zeolite A and amorphous precursor of zeolite A in NaOH solutions at 65 °C. Their experimental results demonstrate that the reversal was attained in about one minute (~50 seconds), indicating fast kinetics to reach equilibrium. Antonic et al. (1993) also indicated that the equilibrium was attained at about 20 minutes for dissolution of zeolite A at 80 °C. For experimental results in which the equilibrium state was not explicitly mentioned, steady state concentrations after 100 minutes are treated as equilibrium concentrations. In this study, as sufficient data for zeolite A and amorphous precursor of zeolite A have been located, equilibrium constants for these two phases are computed. For zeolite X (NaAlSi_{1.23}O_{4.46}•3.07H_2O), Čižmek et al. (1991b) conducted solubility experiments on zeolite X in 2.06 m NaOH solution from 65 °C to 80 °C at 5 °C increment, and their data set has only one data point at each temperature. Roozeboom et al. (1983) had one single data point for solubility of zeolite X in ~1.02 m NaOH at 98 °C. Therefore, sufficient data have not been located for zeolite X. Consequently, no attempt has been made to compute equilibrium constants for zeolite X at this time.
Equilibrium constants are obtained according to the computer modeling. The computer modeling is performed by using EQ3/6 version 8.0a (Wolery et al., 2010; Xiong, 2011). The essence of the modeling is to minimize the difference between experimental and model predicted values, as detailed in the previous publication (Nemer et al., 2011). The retrieval of the equilibrium constant for amorphous precursor of zeolite A at 30 °C can be served as an example. First, solubility data from Ejaz and Graham (1999) and Addai-Mensah et al. (2004) were used to generate EQ3NR input files. Second, an initial guess for the log $K$ was made. Third, by changing the log $K$ into different values, a series of sums of squares of residuals between experimental solubilities and predicted solubilities were obtained. The final log $K$ corresponds to the minimized sum of squares of residuals.

3.2.1 Calculation of Equilibrium Constants of Zeolite A

A number of researchers have investigated solubilities of zeolite A in NaOH solutions at elevated temperatures. The sources of solubility data used for obtaining equilibrium constants in this study and their respective experimental conditions are listed in Table 1. Based on those experimental data, the equilibrium constants for the following reaction using $\text{Al}^{3+}$ and $\text{H}_4\text{SiO}_4(\text{aq})$ as basis species of aluminum and silica, respectively,

$$\text{NaAlSiO}_4\cdot2.25\text{H}_2\text{O(\text{cr})} + 4\text{H}^+ = \text{Na}^+ + \text{Al}^{3+} + \text{H}_4\text{SiO}_4(\text{aq}) + 2.25\text{H}_2\text{O(\text{l})} \quad (21)$$

are obtained (Table 3). In computation of equilibrium constants, all concentrations on molar scale are converted to molal scale according to the following equation based on
densities of supporting solutions used in experiments at respective temperatures, which
are calculated from density equations of Söhnel and Novotný (1985),

\[ m_i = \frac{1000 \times M_i}{1000 \rho - \sum_i M_i E_i} \]  (22)

where \( m_i \) is concentration of \( i \) species on molality scale, \( \rho \) density of solution, \( M_i \) is
ccentration of \( i \) species on molarity scale, and \( E_i \) the molecular weight of \( i \) species.

In Figure 7, predicted solubilities of zeolite A at temperatures of 30 °C, 65 °C, 70 °C, 80 °C, 90 °C, and 100 °C, are compared with experimental data at the respective
temperatures. At 30 °C, there is a scatter in experimental data from various researchers
with a difference of one order of magnitude, and the model seems to fit experimental data
within a factor of ~5 except for that it underpredicts at \( I \approx 6 \text{ m} \) by a factor of ~7
(Figure 7A). At 50 °C, the values produced by the model agree with experimental values
within a factor of ~1.3 (Figure 7A). At 65 °C, the model reproduces experimental data
within a factor of ~2 (Figure 7A). At 70 °C, the values predicted by the model agree with
experimental values within a factor of ~1.4 (Figure 7B). At 80 °C, the model reproduces
solubilities in the ionic strength range from ~3 m to ~5 m within a factor of ~3.5 in
comparison with the experimental data (Figure 7B), whereas solubilities predicted by the
model agree with experimental solubilities within a factor of ~1.1 in the ionic strength
range from ~0.5 m to ~2 m. At 90 °C, solubilities calculated by the model generally
agree with experimental values within a factor of ~2 (Figure 7C). At 100 °C, solubilities
computed by the model are in agreement with experimental values within a factor of ~1.2
(Figure 7C). In all above descriptions, the Al-Si complex, Al(OH)₃HSiO₄³⁻, was not included in calculations.

At 30 °C, the inclusion of Al(OH)₃HSiO₄³⁻ yields a log K of 9.96 ± 0.30 for Reaction 21, in comparison with a value of 10.23 ± 0.31 for the log K produced by the model without the above Al-Si complex. These two values are statistically indistinguishable. However, as indicated by Figure 7A, the introduction of the above Al-Si complex improves the fitting at 30 °C; the model reproduces the solubility at I ≈ 6 m within a factor of ~2 in comparison with the experimental data point. Notice that as described before, the model without Al(OH)₃HSiO₄³⁻ reproduces the solubility at that ionic strength within a factor of ~7.

On the other hand, at 50 °C, the model with the Al-Si complex produces a log K of 7.41 ± 0.40 for Reaction 21, whereas the model without the Al-Si complex results in a log K of 7.95 ± 0.30. Figure 7A shows that the model without the Al-Si complex performs better. The reason for the poor performance of the model with Al-Si complex is not clear. One possibility may be that the existence of the Al-Si complex at 50 °C is uncertain, as Gout et al. (2000) only mentioned its presence at 20 °C. The model with the Al-Si complex at 30 °C is close to 20 °C, and the presence of the Al-Si complex at 30 °C is of high certainty, if there is such a complex, explaining the better performance of the model with the Al-Si complex at 30 °C. Another possibility is that the Pitzer parameters for Na⁺-Al(OH)₃HSiO₄³⁻ might be problematic. This seems unlikely, as the same set of parameters is also used at 30 °C, which results in a desirable performance. However, the final resolution of this issue requires independent evaluation of the Pitzer parameters and to see if they can improve the model with the Al-Si complex.
Judging from the performance of the models with and without the Al-Si complex at 30 °C and 50 °C, it can be concluded that the model without the Al-Si complex is adequate at temperatures equal to or higher than 50 °C. Therefore, no further testing was performed at higher temperatures. Anyhow, the testing at 30 °C seems to provide independent support for the existence of Al(OH)$_3$HSiO$_4^{3-}$ around 20 °C.

Based on the linear regression of temperature dependence of equilibrium constants, $\Delta_r H$ for Reaction 21 is derived as -154 ± 2 kJ mol$^{-1}$ (Table 5). According to Figure 8, the log $K$ at 25 °C is extrapolated as 10.24 ± 0.31. Using the above thermodynamic properties for Reaction 21, the thermodynamic properties of zeolite A at 25 °C and 1 bar are derived (Table 6). In derivation of these thermodynamic properties, the auxiliary thermodynamic data for Na$^+$ and Al$^{3+}$ are from the DATA0.PIT database. The auxiliary thermodynamic data for H$_4$SiO$_4$(aq), which are not present in the DATA0.PIT database, are from the NBS Thermodynamic Table (Wagman et al., 1982). The latter database is in principle consistent with the DATA0.PIT database. However, as the thermodynamic properties obtained in this study are log $K$’s and $\Delta_r H$, Gibbs free energies of formation, enthalpies of formation, and standard entropies can be re-derived to be consistent with other database(s) of interest, if needed.

3.2.2 Calculation of Equilibrium Constants of Amorphous Precursor of Zeolite A

In the synthesis of zeolite A, its amorphous precursor is usually formed first. Therefore, it is also important to know the thermodynamic properties of the amorphous precursor of zeolite A to optimize synthesis. Similar to Reaction 21, the dissolution reaction of the amorphous precursor of zeolite A is:
Ejaz and Graham (1999) conducted systematic solubility studies on the amorphous precursor of zeolite A in NaOH solutions from 3.0 m to 4.5 m at temperatures of 30 °C, 50 °C, 65 °C and 80 °C (Table 1). In addition, Addai-Mensal et al. (2004) also conducted solubility experiments on amorphous precursor of zeolite A in 3.0 m and 6.3 m NaOH solutions at 30 °C and 65 °C (Table 1). Therefore, based on solubility data from both Ejaz and Graham (1999) and Addai-Mensal et al. (2004), the equilibrium constants for the amorphous precursor of zeolite A are calculated (Table 4).

In Figure 9, predicted solubilities of amorphous precursor of zeolite A are compared with experimental data. At 30 °C and 50 °C, the solubility curves for these two temperatures are very close, and the solubilities predicted by the model match experimental values within a factor of ~1.6 (Figure 9A). At 65 °C, the solubilities predicted by the model are in agreement with experimental values within a factor of ~2.5 for the majority of the data points, and within a factor of ~4.5 for the data point at I≈6 m (Figure 9B). At 80 °C, the model matches the experimental solubilities in high ionic strength range within a factor of ~2, but within a factor of ~5 to ~10 at a low ionic strength (~0.2 m) (Figure 9B).

According to the linear regression of temperature dependence of equilibrium constants (Figure 8), ΔrH for Reaction 23 is obtained as −248 ± 3 kJ mol⁻¹ (Table 5). Based on Figure 8, the log K for Reaction 23 at 25 °C is extrapolated as 12.68 ± 0.35. In accordance with the above thermodynamic properties for Reaction 23, the...
thermodynamic properties of amorphous precursor of zeolite A at 25 °C and 1 bar are derived (Table 6).

3.2.4 Validation of Calculations, and Discussions

The enthalpy of formation for zeolite A derived from solubility studies as a function of temperature in this study is compared with the value from calorimetric measurements. The enthalpy of formation from elements for zeolite A obtained in this study, -2738 ± 5 kJ mol⁻¹, is in excellent agreement with the value obtained by calorimetric measurements (-2731.3 ± 1.8 kJ mol⁻¹) (Turner et al., 2008), and compares favorably with the value obtained by theoretical calculations (-2739 kJ mol⁻¹) (Mathieu and Viellard, 2010). In these two studies, zeolite A is formulated as Na₀.5₀₆₇A₁₀.₅₀₁S₁₀.₄₉₇₄O₁.₉₉₉₆₅•₁.₀₉₀₆H₂O. For comparison purpose, their values are scaled relative to four oxygen atoms for the stoichiometry of zeolite A adopted in this study, i.e., NaAlSiO₄•2.25H₂O. This favorable comparison also validates the model developed in this study.

The Gibbs free energy of formation for zeolite A at 25 °C obtained in this study can also be compared with that of Caullet et al. (1980). Caullet et al. (1980) determined the log $K$ for zeolite A with a formula of NaAlSi₁₀.₀₆O₄.₁₂•2.₄H₂O for the following reaction from solubility experiments in 0.02 m, 0.1 m and 0.₅ m NaOH solutions at 25 °C,

$$\text{NaAlSi}_{10.06}\text{O}_{4.12}\cdot2.4\text{H}_2\text{O} + 1.72\text{H}_2\text{O}(l) = \text{Al(OH)}_4^- + 1.06\text{H}_4\text{SiO}_4(aq) + \text{Na}^+ \quad (24)$$
The average thermodynamic equilibrium constant, log $K$ at 25 °C, for Reaction 24 obtained by Caullet et al. (1980) is $-11.20 \pm 0.16$. Based on the log $K$ value for Reaction 24, the Gibbs free energy change for Reaction 24 is calculated as $63.9 \pm 0.9$ kJ mol$^{-1}$. According to the auxiliary thermodynamic data for species in Reaction 24 from the NBS Thermodynamic Table, the Gibbs free energy of formation, $\Delta G$, for NaAlSi$_{1.06}$O$_{4.12}$•2.4H$_2$O, is derived as $-2619 \pm 2$ kJ mol$^{-1}$. When it is scaled to four oxygen atoms for the stoichiometry of zeolite A adopted in this study, in the same procedure applied in the above validation of calculation of enthalpy of formation for zeolite A, the Gibbs free energy of formation becomes $-2543 \pm 2$ kJ mol$^{-1}$ for the formula of Na$_{0.97}$Al$_{0.97}$Si$_{1.03}$O$_4$•2.33H$_2$O. This value is in good agreement with the value of $-2541 \pm 2$ kJ mol$^{-1}$ obtained in this study. The corresponding log $K$ at 25 °C for the following reaction,

$$\text{Na}_{0.97}\text{Al}_{0.97}\text{Si}_{1.03}\text{O}_4\text{•2.33H}_2\text{O(cr)} + 3.88\text{H}^+ = 0.97\text{Na}^+ + 0.97\text{Al}^{3+} + 1.03\text{H}_4\text{SiO}_4\text{(aq)} + 2.21\text{H}_2\text{O(l)}$$

would be $11.33 \pm 0.20$.

Qiu et al. (2000) measured the entropy of dehydrated zeolite A with a formula of Na$_{96}$Al$_{96}$Si$_{96}$O$_{394}$ as 13030 J K$^{-1}$ mol$^{-1}$ at 25 °C. When the above formula is scaled to four oxygen atom, the corresponding entropy becomes 135.73 J K$^{-1}$ mol$^{-1}$ for dehydrated zeolite A with a formula of NaAlSiO$_4$. Viellard (2010) uses a value of 52.0 J K$^{-1}$ mol$^{-1}$ for the entropy of zeolitic water for predictions of entropies of zeolites. Accordingly, using the above values from Qiu et al. (2000) and Viellard (2010), the entropy for
zeolite A with a formula of NaAlSiO$_4$$\cdot$2.25H$_2$O would be expected to be 253 J K$^{-1}$ mol$^{-1}$. This value would differ from the entropy obtained in this study by about 100 J K$^{-1}$ mol$^{-1}$, which was consistently derived from $\Delta_r H$ and $\Delta_r G$ from equilibrium observations.

As the enthalpy of formation for zeolite A derived in this study is consistent with the value obtained by the calorimetric method as mentioned above, we might use $\Delta_f H = 2738$ kJ mol$^{-1}$ and $S^o = 253$ J K$^{-1}$ mol$^{-1}$ for zeolite A to calculate the log $K$ at 25 °C for Reaction 21. That would result in a value for the log $K$ at 25 °C for Reaction 21 to be 16.50. This value is severely discordant with the log $K$ (10.23 ± 0.31) for Reaction 21 obtained in this study, and the similar log $K$ (11.33 ± 0.20) for Reaction 25 obtained by Caullet et al. (1980), both at 25 °C, and it would predict solubilities of zeolite A by at least five orders of magnitude higher than the observed solubilities. Therefore, it seems that further studies on entropies of both zeolite A and dehydrated zeolite A by calorimetric measurements are required, as the existing value does not agree with equilibrium observations.

4 Conclusions

In this study, a thermodynamic model for silica and aluminum is developed, valid to high ionic strength at elevated temperatures up to 100 °C. This model is useful for understanding the geochemical behaviors of Si and Al in concentrated hydrothermal solutions, and for guiding hydrothermal synthesis of zeolites. This model enable us to calculate equilibrium constants of sodium silicates, zeolite A, and amorphous precursor of zeolite A, from hydrothermal solubility experiments. In the near future, equilibrium constants for other zeolite species such as sodalite and cancrinite will be evaluated from
Because of the discordance of one single existing entropy for dehydrated zeolite from calorimetric measurements with equilibrium observations, it is suggested that further calorimetric studies on both zeolite A and dehydrated zeolite A are needed to resolve such a discordance.

Acknowledgements

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References


Table 1. Sources and experimental conditions for solubility experiments on sodium silicates, zeolite A and amorphous precursor of zeolite A from which solubility data are used for computation of equilibrium constants

<table>
<thead>
<tr>
<th>Solubility-controlling phase</th>
<th>Authors</th>
<th>T°C</th>
<th>Aqueous Solution, NaOH, Molal, m</th>
<th>$\Sigma$Si, Molal</th>
<th>$\Sigma$Al, Molal, m</th>
<th>Remarks</th>
</tr>
</thead>
<tbody>
<tr>
<td>$3Na_2O\cdot2SiO_2\cdot11H_2O$</td>
<td>Sprauer and Pearce (1940)</td>
<td>25</td>
<td>12.3–18.8</td>
<td>7.93×10⁻² to 3.62×10⁻¹</td>
<td>None</td>
<td>Approaching equilibrium: from supersaturation. Experimental duration: one month. Usage of data: calculation of Pitzer parameter.</td>
</tr>
<tr>
<td>Na$_3$HSiO$_4$•5H$_2$O</td>
<td>Baker et al. (1950)</td>
<td>50</td>
<td>12.6–19.6</td>
<td>2.78×10⁻¹ to 1.47</td>
<td>None</td>
<td>Approaching equilibrium: from both under- and supersaturation. Experimental duration: a few weeks. Usage of data: calculation of Pitzer parameter.</td>
</tr>
<tr>
<td>Na$_3$HSiO$_4$•2H$_2$O</td>
<td>Baker et al. (1950)</td>
<td>50, 70, 90</td>
<td>15.9-23.2</td>
<td>7.43×10⁻² to 2.14</td>
<td>None</td>
<td>Approaching equilibrium: from both under- and supersaturation. Experimental duration: a few weeks. Usage of data: calculation of Pitzer parameter.</td>
</tr>
<tr>
<td>Zeolite A</td>
<td>Ciric (1968)</td>
<td>100</td>
<td>0.459–1.72</td>
<td>6.26×10⁻³ to 2.76×10⁻²</td>
<td>1.31×10⁻¹ to 1.78×10⁻¹</td>
<td>Approaching equilibrium: from supersaturation. Experimental duration: longer than 100 min. Usage of data: calculation of Pitzer parameter.</td>
</tr>
<tr>
<td>Authors</td>
<td>Temperature</td>
<td>$P$ range</td>
<td>$K$ range</td>
<td>$K$ range</td>
<td>Notes</td>
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</tr>
<tr>
<td>Zhdanov (1971)</td>
<td>90</td>
<td>1.63–4.27</td>
<td>2.11×10⁻³</td>
<td>2.69×10⁻² to 1.29</td>
<td>Approaching equilibrium: from undersaturation. Experimental duration: longer than 100 min. Usage of data: calculation of $\log K$ for zeolite A.</td>
<td></td>
</tr>
<tr>
<td>Wieker and Fahlke (1985)</td>
<td>80</td>
<td>2.17</td>
<td>2.37×10⁻²</td>
<td>1.96×10⁻²</td>
<td>Approaching equilibrium: from undersaturation. Experimental duration longer: than 100 min. Usage of data: calculation of $\log K$ for zeolite A.</td>
<td></td>
</tr>
<tr>
<td>Čižmek et al. (1991a)</td>
<td>65, 70, 80</td>
<td>1.02–2.06</td>
<td>1.17×10⁻² 1.11×10⁻² to 1.83×10⁻²</td>
<td>1.83×10⁻²</td>
<td>Approaching equilibrium: from undersaturation. Experimental duration longer than 100 min. Usage of data: calculation of $\log K$ for zeolite A.</td>
<td></td>
</tr>
<tr>
<td>Myatt (1994)</td>
<td>90</td>
<td>2.49–3.34</td>
<td>8.47×10⁻³ 1.04×10⁻² to 1.25×10⁻¹</td>
<td>6.13×10⁻² to 1.25×10⁻¹</td>
<td>Approaching equilibrium from supersaturation. Experimental duration: longer than 100 min. Usage of data: calculation of $\log K$ for zeolite A.</td>
<td></td>
</tr>
<tr>
<td>Ejaz and Graham (1999)</td>
<td>30, 50, 65, 80</td>
<td>3.04–4.46</td>
<td>3.28×10⁻³ 4.32×10⁻³ to 1.55×10⁻²</td>
<td>1.69×10⁻² to 1.69×10⁻²</td>
<td>Approaching equilibrium: from undersaturation. Experimental duration: longer</td>
<td></td>
</tr>
<tr>
<td>Study</td>
<td>Temperature (°C)</td>
<td>pH 70</td>
<td>pK (_70)</td>
<td>pK (_{2.05})</td>
<td>pK (<em>{1.51×10^{-2}}) to (</em>{1.81×10^{-2}})</td>
<td>pK (<em>{1.46×10^{-2}}) to (</em>{1.71×10^{-2}})</td>
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<td>---------------------------</td>
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<tr>
<td>Kosanović et al. (2000)</td>
<td>70</td>
<td>2.05</td>
<td>1.51×10(^{-2}) to 1.81×10(^{-2})</td>
<td>1.46×10(^{-2}) to 1.71×10(^{-2})</td>
<td>Approaching equilibrium: from undersaturation.</td>
<td>Experimental duration: longer than 100 min.</td>
</tr>
<tr>
<td>Addai-Mensal et al. (2004)</td>
<td>30, 65</td>
<td>3.02–6.31</td>
<td>1.19×10(^{-2}) to 4.63×10(^{-2})</td>
<td>1.42×10(^{-2}) to 5.43×10(^{-2})</td>
<td>Approaching equilibrium: from both undersaturation and supersaturation.</td>
<td>Experimental duration: longer than 50 min.</td>
</tr>
<tr>
<td>Bosnar et al. (2005)</td>
<td>80</td>
<td>1.49–1.58</td>
<td>8.63×10(^{-3}) to 6.15×10(^{-2})</td>
<td>2.06×10(^{-3}) to 1.03×10(^{-2})</td>
<td>Approaching equilibrium: from supersaturation.</td>
<td>Experimental duration longer than 100 min.</td>
</tr>
<tr>
<td>Amorphous precursor of zeolite A</td>
<td>30, 50, 65, 80</td>
<td>3.04–4.46</td>
<td>4.70×10(^{-2}) to 7.57×10(^{-2})</td>
<td>3.48×10(^{-2}) to 5.54×10(^{-2})</td>
<td>Approaching equilibrium: from undersaturation.</td>
<td>Experimental duration: longer than 100 min.</td>
</tr>
<tr>
<td>Zeolite A.</td>
<td>Addai-Mensal et al. (2004)</td>
<td>30, 65</td>
<td>3.02–6.31</td>
<td>$7.09 \times 10^{-2}$ to $2.20 \times 10^{-1}$</td>
<td>$7.99 \times 10^{-2}$ to $2.19 \times 10^{-1}$</td>
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<tr>
<td>Approaching equilibrium: from both under-saturation and supersaturation.</td>
<td></td>
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<td></td>
<td></td>
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<tr>
<td>Experimental duration: longer than 50 min.</td>
<td></td>
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</tr>
<tr>
<td>Usage of data: calculation of log $K$ for amorphous zeolite A.</td>
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</tr>
</tbody>
</table>
Table 2. Equilibrium constants for silica and aluminum species considered in this study up to 100 °C

<table>
<thead>
<tr>
<th>Reaction</th>
<th>T, °C</th>
<th>log K</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\text{H}_4\text{SiO}_4 = \text{H}^+ + \text{H}_3\text{SiO}_4^-$</td>
<td>25</td>
<td>$-9.68 \pm 0.14^A$</td>
<td>Fleming and Crerar, 1982</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-9.34 \pm 0.16^A$</td>
<td>Fleming and Crerar, 1982</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-9.10 \pm 0.17^A$</td>
<td>Fleming and Crerar, 1982</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-8.94 \pm 0.16^A$</td>
<td>Fleming and Crerar, 1982</td>
</tr>
<tr>
<td>$\text{H}_3\text{SiO}_4^- = \text{H}^+ + \text{H}_2\text{SiO}_4^{2-}$</td>
<td>25</td>
<td>$-13.45 \pm 0.07$</td>
<td>Hershey and Millero, 1986</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-12.95 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-12.56 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-12.28 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td>$2\text{H}_4\text{SiO}_4 = \text{H}^+ + \text{H}_5\text{Si}_2\text{O}_7^- + \text{H}_2\text{O}$</td>
<td>25</td>
<td>$-8.50^C$</td>
<td>Felmy et al., 2001</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-8.14 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-7.86 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-7.65 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td>$\text{H}_5\text{Si}_2\text{O}_7^- = \text{H}^+ + \text{H}_4\text{Si}_2\text{O}_7^{2-}$</td>
<td>25</td>
<td>$-10.90^C$</td>
<td>Felmy et al., 2001</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-10.59 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-10.38 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-10.24 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td>$3\text{H}_4\text{SiO}_4 = 3\text{H}^+ + \text{H}_5\text{Si}<em>3\text{O}</em>{10}^{3-} + 2\text{H}_2\text{O}$</td>
<td>25</td>
<td>$-29.40^C$</td>
<td>Felmy et al., 2001</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-28.75 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-28.34 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-28.11 \pm 0.25^B$</td>
<td>This study</td>
</tr>
<tr>
<td>$\text{Al(OH)}_4^- = \text{Al}^{3+} + 4\text{OH}^-$</td>
<td>25</td>
<td>$-34.05 \pm 0.05$</td>
<td>Derived from Wesolowski, 1992</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-33.44 \pm 0.05$</td>
<td>Derived from Wesolowski, 1992</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-33.11 \pm 0.05$</td>
<td>Derived from Wesolowski, 1992</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-32.99 \pm 0.05$</td>
<td>Derived from Wesolowski, 1992</td>
</tr>
<tr>
<td>$\text{Al(OH)}_4^- + \text{H}_2\text{SiO}_4^{2-} = \text{Al(OH)}_3\text{HSiO}_4^{3-} + \text{H}_2\text{O(l)}$</td>
<td>20</td>
<td>$-0.42 \pm 0.10$</td>
<td>This study$^D$, estimated from log $\beta_i^G$ (Gout et al., 2000) with the SIT model</td>
</tr>
<tr>
<td></td>
<td>25</td>
<td>$-0.42 \pm 0.15$</td>
<td>This study$^D$</td>
</tr>
<tr>
<td></td>
<td>50</td>
<td>$-0.38 \pm 0.25^B$</td>
<td>This study$^D$</td>
</tr>
<tr>
<td></td>
<td>75</td>
<td>$-0.35 \pm 0.25^B$</td>
<td>This study$^D$</td>
</tr>
<tr>
<td></td>
<td>100</td>
<td>$-0.33 \pm 0.25^B$</td>
<td>This study$^D$</td>
</tr>
</tbody>
</table>

$^A$ Uncertainties were not given in Fleming and Crerar (1982). Uncertainties are assigned based on the respective benchmark values of Busey and Mesmer (1977).

$^B$ An uncertainty of ± 0.25 is assigned to all predicted values at elevated temperatures up to 100 °C. See text for details.

$^C$ Uncertainties were not given in Felmy et al. (2001).

$^D$ The complex, Al(OH)$_3$HSiO$_4^{3-}$, is used to test whether it can improve the modeling at 30 °C and 50 °C only.
Table 3. Pitzer interaction parameters employed in this study

<table>
<thead>
<tr>
<th>Interaction Pair</th>
<th>$\beta^{(0)}$</th>
<th>$\beta^{(1)}$</th>
<th>$C^0$</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Na$^+$–H$_3$SiO$_4^-$</td>
<td>0.043±0.019</td>
<td>0.24±0.11</td>
<td></td>
<td>Hershey and Millero (1986)</td>
</tr>
<tr>
<td>Na$^+$–H$_2$SiO$_4^{2-}$</td>
<td>0.32±0.08</td>
<td>0.13±0.50</td>
<td></td>
<td>Hershey and Millero (1986)</td>
</tr>
<tr>
<td>Na$^+$–H$_5$Si$_2$O$_7^{-}$</td>
<td>−0.0571±0.04</td>
<td>0.34±0.13</td>
<td></td>
<td>This Study $^A$</td>
</tr>
<tr>
<td>Na$^+$–H$_4$Si$_2$O$_7^{2-}$</td>
<td>−0.0227±0.06</td>
<td>1.56±0.40</td>
<td></td>
<td>This Study $^B$</td>
</tr>
<tr>
<td>Na$^+$–H$_5$Si$<em>3$O$</em>{10}^{3-}$</td>
<td>0.078±0.03</td>
<td>4.29±0.80</td>
<td></td>
<td>This Study $^C$</td>
</tr>
<tr>
<td>Na$^+$–Al(OH)$_3$HSiO$_4^{3-}$</td>
<td>0.078±0.03</td>
<td>4.29±0.80</td>
<td></td>
<td>This Study $^D$</td>
</tr>
<tr>
<td>Na$^+$–Al(OH)$_4^{-}$</td>
<td>0.051</td>
<td>0.25</td>
<td>−0.00090</td>
<td>Wesolowski (1992) $^E$</td>
</tr>
</tbody>
</table>

Interaction Involving Neutral Species and Mixing Parameters

<table>
<thead>
<tr>
<th>Interaction Pair</th>
<th>$\lambda_{ij}$</th>
<th>$\theta_{ij}$</th>
<th>$\Psi_{ijk}$</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Na$^+$–H$_3$SiO$_4^{0}$</td>
<td>0.0925</td>
<td></td>
<td></td>
<td>Azaroual et al. (1997) $^E$</td>
</tr>
<tr>
<td>OH$^-$–H$_2$SiO$_4^{2-}$</td>
<td></td>
<td>−0.0812±0.003</td>
<td></td>
<td>This Study $^F$</td>
</tr>
<tr>
<td>OH$^-$–H$_2$SiO$_4^{2-}$–Na$^+$</td>
<td></td>
<td></td>
<td>$\left(\frac{\partial \theta_{ij}}{\partial T}\right)_p = -9.35±0.44\times10^{-5}$</td>
<td>This Study $^F$</td>
</tr>
</tbody>
</table>

$^A$ Calculated from the estimation method of Plyasunov et al. (1998) for 1:1 interaction, based on $\varepsilon$(Na$^+$, Si$_2$O$_2$(OH)$_5^{-}$) of −0.08 ± 0.04, which is from Grenthe et al. (1992). In the method of Plyasunov et al. (1998), uncertainty was not given to $\beta^{(1)}$. The uncertainty assigned here is two standard deviations from the average $\beta^{(1)}$ for 1:1 interaction computed in Plyasunov et al. (1998).

$^B$ Calculated from the estimation method of Plyasunov et al. (1998) for 1:2 interaction, based on $\varepsilon$(Na$^+$, Si$_2$O$_3$(OH)$_4^{2-}$) of −0.15 ± 0.06 from Grenthe et al. (1992). In the method of Plyasunov et al. (1998), uncertainty was not given to $\beta^{(1)}$. The uncertainty assigned here is two standard deviations from the average $\beta^{(1)}$ for 1:2 interaction computed in Plyasunov et al. (1998).

$^C$ Calculated from the estimation method of Plyasunov et al. (1998) for 1:3 interaction, based on $\varepsilon$(Na$^+$, Si$_3$O$_6$(OH)$_3^{3-}$) of −0.25 ± 0.03 from Grenthe et al. (1992). In the method of Plyasunov et al. (1998), uncertainty was not given to $\beta^{(1)}$. The uncertainty assigned here is two standard deviations from the average $\beta^{(1)}$ for 1:3 interaction computed in Plyasunov et al. (1998).

$^D$ Interaction parameters are assigned to be the same as those for Na$^+$–H$_2$Si$_3$O$_{10}^{3-}$.

$^E$ Uncertainties were not given in Wesolowski (1992) and Azaroual et al. (1997), respectively.

$^F$ Evaluated from solubility data of 3Na$_2$O•2SiO$_2$•11H$_2$O in NaOH solutions up to ~19 m from Sprauer and Pearce (1940). The temperature dependence of $\theta_{ij}$, $\left(\frac{\partial \theta_{ij}}{\partial T}\right)_p$, is evaluated from solubility data of Na$_3$HSiO$_4$•5H$_2$O at 50°C, and of Na$_3$HSiO$_4$•2H$_2$O at 50°C, 70°C, and 90°C in NaOH solutions up to ~24 m from Baker et al. (1950).
Table 4. Equilibrium constants of sodium silicates and zeolites retrieved from solubility experiments in this study*

<table>
<thead>
<tr>
<th>T, °C</th>
<th>log K ± 2σ</th>
<th>Reaction</th>
</tr>
</thead>
<tbody>
<tr>
<td>25</td>
<td>83.83 ± 0.62</td>
<td>3Na₂O•2SiO₂•11H₂O + 6H⁺ = 6Na⁺ + 2H₄SiO₄(aq) + 10H₂O (l)</td>
</tr>
<tr>
<td>50</td>
<td>47.03 ± 0.35</td>
<td>Na₃HSiO₄•2H₂O + 3H⁺ = 3Na⁺ + H₄SiO₄(aq) + 2H₂O (l)</td>
</tr>
<tr>
<td>70</td>
<td>45.69 ± 0.31</td>
<td>Na₃HSiO₄•5H₂O + 3H⁺ = 3Na⁺ + H₄SiO₄(aq) + 5H₂O (l)</td>
</tr>
<tr>
<td>90</td>
<td>46.03 ± 0.30</td>
<td>NaAlSiO₄•2.25H₂O(cr, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>50</td>
<td>41.50 ± 0.35</td>
<td>Na₃HSiO₄•5H₂O + 3H⁺ = 3Na⁺ + H₄SiO₄(aq) + 5H₂O (l)</td>
</tr>
<tr>
<td>25</td>
<td>10.24 ± 0.31</td>
<td>NaAlSiO₄•2.25H₂O(am, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>30</td>
<td>10.23 ± 0.31</td>
<td>NaAlSiO₄•2.25H₂O(cr, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>50</td>
<td>7.95 ± 0.30</td>
<td>NaAlSiO₄•2.25H₂O(am, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>65</td>
<td>6.70 ± 0.33</td>
<td>NaAlSiO₄•2.25H₂O(cr, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>70</td>
<td>6.54 ± 0.30</td>
<td>NaAlSiO₄•2.25H₂O(am, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>80</td>
<td>5.74 ± 0.45</td>
<td>NaAlSiO₄•2.25H₂O(cr, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>90</td>
<td>5.54 ± 0.36</td>
<td>NaAlSiO₄•2.25H₂O(am, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
<tr>
<td>100</td>
<td>5.34 ± 0.30</td>
<td>NaAlSiO₄•2.25H₂O(cr, zeolite A) + 4H⁺ = Na⁺ + Al³⁺ + H₄SiO₄(aq) + 2.25H₂O</td>
</tr>
</tbody>
</table>

* Experimental conditions for hydrothermal experiments from which solubility data are used for computation of equilibrium constants are detailed in text and Table 1.

A Extrapolated to the reference temperature, 25 °C, based on the linear relation between log K and reciprocal temperature in Kevin for zeolite A.

B Extrapolated to the reference temperature, 25 °C, based on the linear relation between log K and reciprocal temperature in Kevin for amorphous precursor of zeolite A.
Table 5. Enthalpy changes for reactions involving zeolites derived in this study

<table>
<thead>
<tr>
<th>Reactions</th>
<th>$\Delta H$, kJ mol$^{-1}$*</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\text{Na}_3\text{HSiO}_4\cdot\text{2H}_2\text{O} + 3\text{H}^+ = 3\text{Na}^+ + \text{H}_4\text{SiO}_4(\text{aq}) + 2\text{H}_2\text{O}(\text{l})$</td>
<td>-58 ± 45 (2σ)</td>
</tr>
<tr>
<td>$\text{NaAlSiO}_4\cdot\text{2.25H}_2\text{O(cr, zeolite A)} + 4\text{H}^+ = \text{Na}^+ + \text{Al}^{3+} + \text{H}_4\text{SiO}_4(\text{aq}) + 2.25\text{H}_2\text{O}(\text{l})$</td>
<td>-152 ± 5 (2σ)</td>
</tr>
<tr>
<td>$\text{NaAlSiO}_4\cdot\text{2.25H}_2\text{O(am, zeolite A)} + 4\text{H}^+ = \text{Na}^+ + \text{Al}^{3+} + \text{H}_4\text{SiO}_4(\text{aq}) + 2.25\text{H}_2\text{O}(\text{l})$</td>
<td>-248 ± 3 (2σ)</td>
</tr>
</tbody>
</table>

* Uncertainties account for the errors from regressions only. The overall uncertainties could be higher than those provided here.
Table 6. Thermodynamic properties of zeolite A and the amorphous form of zeolite A at 25 °C and 1 bar derived in this study*

<table>
<thead>
<tr>
<th>Properties</th>
<th>Values (±2σ)</th>
<th>Remarks</th>
</tr>
</thead>
<tbody>
<tr>
<td>ΔfH°, Zeolite A, NaAlSiO₄•2.25H₂O(cr)</td>
<td>−2738 ± 5 kJ mol⁻¹</td>
<td>Based on ΔfH derived from temperature dependence of equilibrium constant</td>
</tr>
<tr>
<td>ΔfG°, Zeolite A, NaAlSiO₄•2.25H₂O(cr)</td>
<td>−2541 ± 2 kJ mol⁻¹</td>
<td>Based on ΔfG derived from log K extrapolated to 25 °C.</td>
</tr>
<tr>
<td>S°, Zeolite A, NaAlSiO₄•2.25H₂O(cr)</td>
<td>373 ± 10 J K⁻¹ mol⁻¹</td>
<td>Based on ΔfS calculated from the Gibbs-Helmholtz equation</td>
</tr>
<tr>
<td>ΔfH°, Amorphous precursor of zeolite A, NaAlSiO₄•2.25H₂O(am)</td>
<td>−2642 ± 3 kJ mol⁻¹</td>
<td>Based on ΔfH derived from temperature dependence of equilibrium constant</td>
</tr>
<tr>
<td>ΔfG°, Amorphous precursor of zeolite A, NaAlSiO₄•2.25H₂O(am)</td>
<td>−2527 ± 2 kJ mol⁻¹</td>
<td>Based on ΔfG derived from log K extrapolated to 25 °C.</td>
</tr>
<tr>
<td>S°, Amorphous precursor of zeolite A, NaAlSiO₄•2.25H₂O(am)</td>
<td>648 ± 10 J K⁻¹ mol⁻¹</td>
<td>Based on ΔfS calculated from the Gibbs-Helmholtz equation</td>
</tr>
</tbody>
</table>

* All properties refer to formation from elements.
Figure Captions

Figure 1. Comparison of model predicted solubilities of 3Na₂O•2SiO₂•11H₂O at 25 °C, Na₃HSiO₄•5H₂O, and Na₃HSiO₄•2H₂O at 50 °C, with experimental values. The size of error bars is equal to or smaller than the symbol size.

Figure 2. Comparison of model predicted solubilities of Na₃HSiO₄•2H₂O at 70 °C and 90 °C with experimental values. The size of error bars is equal to or smaller than the symbol size.

Figure 3. A plot showing equilibrium constants of Na₃HSiO₄•2H₂O as a function of reciprocal temperatures in Kevin.

Figure 4. Comparison of model predicted solubilities of amorphous silica with experimental values, which are independent from the model development. The size of error bars is equal to or smaller than the symbol size.

Figure 5. Speciation of silica at different total concentrations of silica as a function of pH at 25 °C. (A) total silica concentration is 0.01 m; (B) total silica concentration is 0.1 m; and (C) total silica concentration is 1 m. Notice that in acidic pH range, the solution is supersaturated with amorphous silica.

Figure 6. Speciation of silica at different total concentrations of silica as a function of pH at 100 °C. (A) total silica concentration is 0.01 m; (B) total silica concentration is 0.1 m; and (C) total silica concentration is 1 m. Notice that in acidic pH range, the solution is supersaturated with amorphous silica/quartz.

Figure 7. Comparison of model predicted solubilities of zeolite A with experimental values. (A) at 30 °C, 50 °C and 65 °C; (B) at 70 °C and 80 °C; and (C) at 90 °C and 100 °C. The size of error bars is equal to or smaller than the symbol size.

Figure 8. A plot showing equilibrium constants of zeolite A and amorphous precursor of zeolite A as a function of reciprocal temperatures in Kevin. The size of error bars is equal to or smaller than the symbol size.

Figure 9. Comparison of model predicted solubilities of amorphous precursor of zeolite A with experimental values. (A) at 30 °C and 50 °C; and (B) at 65 °C and 80 °C. The size of error bars is equal to or smaller than the symbol size.
3Na₂O•2SiO₂•11H₂O, 25°C, experiment (Sprauer and Pearce, 1940)

Na₃HSiO₄•5H₂O, 50°C, experiment (Baker et al., 1950)

Na₃HSiO₄•2H₂O, 50°C, experiment (Baker et al., 1950)

3Na₂O•2SiO₂•11H₂O, 25°C, model, log K = 83.83

Na₃HSiO₄•5H₂O, 50°C, model, log K = 41.50

Na₃HSiO₄•2H₂O, 50°C, model, log K = 47.03

Figure 1
Figure 2

Na$_3$HSiO$_4$$\cdot$2H$_2$O + 3H$^+$ = 3Na$^+$ + H$_4$SiO$_4$(aq) + 2H$_2$O(l)
Figure 3

The graph shows a linear relationship between 1/T (in K) and log K, with the equation:

\[ y = 3035 \pm 2600x + 37 \pm 8 \]

and

\[ R^2 = 0.5516 \]
Figure 4
Figure 5A
\[ \Sigma Si = 0.1 \; \text{m} \]

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Figure 5C
ΣSi = 0.01 m

100 °C

% of Total Silica

pH

Figure 6A
Figure 6B

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Figure 6C

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Figure 7A

\[
\text{NaAlSiO}_4 \cdot 2.25\text{H}_2\text{O} (\text{cr}) + 4\text{H}^+ = \text{Na}^+ + \text{Al}^{3+} + \text{H}_4\text{SiO}_4 (\text{aq}) + 2.25\text{H}_2\text{O} (\text{l})
\]
Figure 7B

NaAlSiO$_4$$\cdot$2.25H$_2$O(cr) + 4H$^+$ = Na$^+$ + Al$^{3+}$ + H$_2$SiO$_4$(aq) + 2.25H$_2$O(l)

Zeolite A, 70ºC, experiment (Cizmek et al., 1991a)
Zeolite A, 70ºC, experiment (Kosanovic et al., 2008)
Zeolite A, 80ºC, experiment (Wieker and Fahlke, 1985)
Zeolite A, 80ºC, experiment (Cizmek et al., 1991a)
Zeolite A, 80ºC, experiment (Ejaz and Graham, 1999)
Zeolite A, 80ºC, experiment (Bosnar et al., 2005)
Zeolite A, 70ºC, model, log K = 6.54
Zeolite A, 80ºC, model, log K = 5.74

Molality of NaOH, m$_{NaOH}$

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Figure 7C

NaAlSiO$_4$·2.25H$_2$O(cr) + 4H$^+$ = Na$^+$ + Al$^{3+}$ + H$_2$SiO$_4$(aq) + 2.25H$_2$O(l)

| Zeolite A, 90ºC, experiment (Zhdanov, 1971) |
| Zeolite A, 90ºC, experiment (Myatt, 1994) |
| Zeolite A, 100ºC, experiment (Ciric, 1968) |
| Zeolite A, 90ºC, model, log K = 5.54 |
| Zeolite A, 100ºC, model, log K = 5.34 |
y = 7963±327x - 16.46±0.96
R² = 0.9597, Zeolite A

y = 12971±160x - 30.8±0.5
R² = 0.9979, Amorphous Precursor of Zeolite A

Figure 8
Amorphous precursor of zeolite A, 30°C, experiment (Ejaz and Graham, 1999)
Amorphous precursor of zeolite A, 30°C, experiment (Addai-Mensah et al., 2008)
Amorphous precursor of zeolite A, 50°C, experiment (Ejaz and Graham, 1999)
Amorphous precursor of zeolite A, 30°C, model, log K = 11.94
Amorphous precursor of zeolite A, 50°C, model, log K = 9.30

NaAlSiO_4·2.25H_2O(am) + 4H^+ = Na^+ + Al^{3+} + H_4SiO_4(aq) + 2.25H_2O
\[ \text{NaAlSiO}_4\cdot2.25\text{H}_2\text{O}(\text{am}) + 4\text{H}^+ = \text{Na}^+ + \text{Al}^{3+} + \text{H}_4\text{SiO}_4(\text{aq}) + 2.25\text{H}_2\text{O} \]

Figure 9B