

## **revision 1**

### **Crystal chemistry of layered Pb oxychloride minerals with PbO-related structures.**

#### **I. Crystal structure of hereroite, $[\text{Pb}_{32}\text{O}_{20}(\text{O},\square)](\text{AsO}_4)_2((\text{Si},\text{As},\text{V},\text{Mo})\text{O}_4)_2\text{Cl}_{10}$**

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### **Abstract**

The crystal structure of hereroite, a new complex lead oxychloride mineral from the Kombat Mine, Grootfontein, Namibia, has been solved by direct methods and refined to  $R_1 = 0.054$  for 6931 unique observed reflections. The mineral is monoclinic  $C2/c$ ,  $a = 23.139(4)$ ,  $b = 22.684(4)$ ,  $c = 12.389(2)$  Å,  $\beta = 102.090(3)^\circ$ ,  $V = 6358.8(18)$  Å<sup>3</sup>. The structure contains 16 independent Pb sites in strongly asymmetric coordination by O and Cl atoms. There are two tetrahedral sites, from which one (As) is occupied solely by As, whereas the second (T) has the mixed occupancy of  $[\text{Si}_{0.48}\text{As}_{0.29}\text{V}_{0.15}\text{Mo}_{0.09}]$ . There are in total 21 O sites. The O1-O8 sites belong to the  $\text{AsO}_4$  and  $\text{TO}_4$  tetrahedral oxyanions. The other O atoms (O9-O20) are tetrahedrally coordinated by Pb atoms, thus being central for the  $\text{OPb}_4$  oxocentered tetrahedra. The  $\text{OPb}_4$  tetrahedra share edges to form the  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layers that can be described as derivatives of the [OPb] layer from the structure of tetragonal PbO (litharge). The  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layer in hereroite can be obtained from the [OPb] layer by

27 removal of blocks of oxocentered tetrahedra, which results in formation of double-square sevenfold and  
28 square fourfold cavities. The cavities are occupied by the  $\text{AsO}_4$  and  $\text{TO}_4$  tetrahedra, respectively. The  
29 topology of the  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layer is rather complex and can be described as a combination of modules  
30 extracted from the layers of  $\text{OPb}_4$  tetrahedra present in the structures of kombatite and symesite. The  
31 topological functions of tetrahedra within the layer are analysed using the square lattice method, which  
32 shows that each symmetry-independent tetrahedron has its own topological function in the layer  
33 construction. The structure of hereroite belongs to the 2:1 type of layered Pb oxyhalides and consists of  
34 alternating  $\text{PbO}$ -type layers and Cl sheets oriented parallel to the (010) plane.

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37 **Keywords:** hereroite; lead oxyhalides; crystal structure; litharge derivatives; layered structures,  
38 oxocentered tetrahedra, modular structures, method of square lattices.

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## Introduction

Lead oxyhalides occur under a wide range of natural and technological conditions (Post and Buseck, 1985; Welch et al. 1996, 1998, 2000, 2001; Noren et al. 2002; Turner, 2006; Cziczko et al. 2009; Kampf et al. 2010a, b; Turner and Rumsey, 2010). Their formation and precipitation play an important role in the transport of Pb from mines and mill tailings to the biosphere. Lead oxyhalides are also important as inorganic materials with various existing and potential technological applications (Matsumoto et al. 2001; Sigman and Korgel, 2005; Pfitzner and Pohla, 2009; Wan et al. 2011). In nature, lead oxychlorides occur as secondary minerals in oxidation zones of mineral deposits. Their structures may accommodate a wide range of elements, including As, S, V, Mo, P, Si, I, *etc.*, which results in interesting chemical and structural diversity. Of particular interest are minerals with layered PbO-derivative structures, which are structurally related to the Aurivillius phases. Table 1 provides details on chemical composition and crystallographic data for natural PbO-related oxychloride minerals known to date. In this paper, we report on the crystal structure of hereroite,  $[\text{Pb}_{32}\text{O}_{21-x+y}](\text{AsO}_4)_2((\text{Si}_x(\text{As}, \text{V})_{1-x-y}\text{Mo}_y)\text{O}_4)_2\text{Cl}_{10}$ , a new complex lead oxychloride mineral recently described from the Kombat Mine, Grootfontein, Namibia (Turner et al. 2012a). Hereroite was found in close association with asisite, damaraite, kombatite, sahlinitite, vladkrivovichevite, barysilite, quartz, native copper, hausmannite, jacobsonite and manganite. The Kombat mine ore bodies contain primary Pb-Cu-Zn (-Ag) sulfides, which were originally emplaced hydrothermally as fracture fillings. Primary ores were then modified by later epigenetic, hydrothermal and metasomatic replacement events. These events created a range of Pb-Mn-Fe silicate minerals in the Kombat deposit, as well as a variety of late-stage Pb-oxychloride minerals including hereroite (Cairncross, 1997). Hereroite occurs as transparent to translucent intergrown, glassy-looking orange grains up to ~0.5 mm in diameter.

## Background information

66 The crystal structures of PbO-related layered lead oxyhalides are based upon the O-Pb layers  
67 alternating with the sheets  $X$  of  $X^-$  ions ( $X = \text{Cl}, \text{Br}$ ) (Aurivillius 1982, 1983; Krivovichev et al. 2009).

68 The stacking sequence of sheets is either:

69  $\dots X | \text{OPb} | \text{OPb} | X | \text{OPb} | \text{OPb} | \dots$  with the (OPb): $X$  ratio equal to 2:1 ( $X = \text{F}, \text{Cl}, \text{Br}$ ) or

70  $\dots X | \text{OPb} | X | \text{OPb} | X \dots$  with (OPb): $X = 1:1$ .

71 The 2:1 structure has also been observed in synthetic Aurivillius phases, kombatite, sahlinite, asisite  
72 and parkinsonite, whereas the examples of the 1:1 structure are symesite, schwartzembergite, blixite,  
73 and mereheadite.

74 The O-Pb layers can be viewed as derivatives of the [OPb] layer from the structure of litharge, a  
75 tetragonal modification of PbO (Boher et al. 1985) (Fig. 1a). The relationships between the structures  
76 of litharge and its oxychloride derivatives can be conveniently described in terms of structural units  
77 based upon anion-centered tetrahedra. This approach to the description of complex inorganic crystal  
78 structures has many advantages over the traditional cation-centered description where the structure  
79 contains so-called 'additional' O atoms, coordinated by di- and trivalent cations only [see (Krivovichev  
80 and Filatov, 1999a, b; Krivovichev et al. 2004; Siidra et al. 2008c) for reviews and (Chen et al. 2009;  
81 Huvé et al. 2009; Mauck et al. 2010; Li et al. 2011; Kampf et al. 2011; Reshak et al. 2012; Endara et al.  
82 2012) for recent applications]. The structure of litharge is usually described in terms of the  $\text{OPb}_4$   
83 oxocentered tetrahedra, sharing edges to form [OPb] layers (O'Keeffe and Hyde, 1996). Transformation  
84 of the [OPb] layer into one of its derivatives corresponds to the removal of blocks of the  $\text{OPb}_4$   
85 tetrahedra from the former (Fig. 1b,c). The resulting layers have the chemical composition  $[\text{Pb}_m\text{O}_n]^{z+}$ ,  
86 with  $m > n$ . The role of the  $X$  sheets is to compensate the positive charge of the O-Pb layers. However,  
87 a number of additional modifications has also been observed: (1) substitution of the  $\text{OPb}_4$  tetrahedra in  
88 the [OPb] layer by tetrahedral or triangular anions such as  $\text{SO}_4$  (symesite),  $\text{MoO}_4$  (parkinsonite),  $\text{VO}_4$   
89 (kombatite),  $\text{AsO}_4$  (sahlinite),  $\text{SiO}_4$  (asisite),  $\text{CO}_3$  and  $\text{BO}_3$  (mereheadite) etc.; the substitution may be  
90 disordered (as in parkinsonite and asisite) or ordered as in symesite, kombatite and sahlinite; (2)  $\text{O}^{2-} \leftrightarrow$   
91  $\text{OH}^-$  substitution in the [OPb] layers (as observed in blixite and mereheadite); (3) insertion of additional

92 Pb atoms into the  $X$  sheet (observed in mereheadite and synthetic  $\text{Pb}_{31}\text{O}_{22}\text{Br}_{10}\text{Cl}_8$  (Krivovichev et al.  
93 2006)); (4)  $\text{O}^{2-} \leftrightarrow \text{F}^-$  substitution in the [OPb] layers (observed in rumseyite (Turner et al. 2012b)).

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## 95 Experimental

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### 97 Chemical composition

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99 According to Turner et al. (2012a), the empirical formula for hereroite based on the assumption  
100 of  $\text{Pb} = 32$  apfu is  $[\text{Pb}_{32}\text{O}_{20.70}] (\text{AsO}_4)_2 ((\text{Si}_{0.48}\text{As}_{0.29}\text{V}_{0.15}\text{Mo}_{0.09})\text{O}_4)_2\text{Cl}_{9.84}$  which gives the simplified  
101 formula  $[\text{Pb}_{32}\text{O}_{21-x+y}](\text{AsO}_4)_2((\text{Si}_x(\text{As},\text{V})_{1-x-y}\text{Mo}_y)\text{O}_4)_2\text{Cl}_{10}$ , where  $x = 0.48$ ,  $y = 0.08$ . The corresponding  
102 ideal formula can be written as  $[\text{Pb}_{32}\text{O}_{20}(\text{O},\square)](\text{AsO}_4)_2((\text{Si},\text{As},\text{V},\text{Mo})\text{O}_4)_2\text{Cl}_{10}$ .

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### 104 Single crystal X-ray diffraction study

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106 A transparent equant orange crystal was studied using a Bruker Smart Apex II diffractometer at  
107 the Department of Crystallography, St. Petersburg State University, Russia. More than a hemisphere of  
108 X-ray diffraction data ( $\theta_{\text{max}} = 36.15^\circ$ ) with frame widths of  $0.3^\circ$  in  $\omega$ , and with 45 s spent counting for  
109 each frame was collected at room temperature using the  $\text{MoK}\alpha$  radiation. The data were integrated and  
110 corrected for absorption with an empirical ellipsoidal model using the Bruker programs *APEX* and  
111 *XPREP*. The observed systematic absences were consistent with space group  $C2/c$ . The structure was  
112 solved by direct methods and refined to  $R_1 = 0.053$  on the basis of  $F^2$  for all unique data. The *SHELX*  
113 program package was used for all structural calculations (Sheldrick, 2008). The final model included  
114 all atomic positional parameters, anisotropic-displacement parameters for all atoms except oxygen, and  
115 a refinable weighting scheme for the structure factors. The final atomic coordinates and displacement  
116 parameters are given in Table 3, and selected interatomic distances in Table 4. The list of observed and  
117 calculated structure factors can be provided by the authors upon request.

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## Results

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### Cation coordination

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123         The structure of hereroite contains 16 symmetrically independent Pb sites. Figure 2 shows the  
124 coordination of the Pb atoms with all the bonds shorter than 3.5 Å taken into account. The coordination  
125 configurations of the Pb atoms are variable and fall into two groups. Pb1, Pb2, Pb3, Pb5, Pb6, Pb7,  
126 Pb11, Pb13, and Pb16 sites are coordinated by the O<sup>2-</sup> anions in one coordination hemisphere. The  
127 number of the O atoms varies from three to four with Pb-O bonds in the range of 2.24-2.96 Å. In the  
128 opposite hemisphere, the above mentioned Pb atoms form from three to four long Pb-Cl bonds in range  
129 of 3.04-3.48 Å. Pb4, Pb8, Pb9, Pb10, Pb12, Pb14, and Pb15 have strongly asymmetrical coordinations  
130 by the O<sup>2-</sup> anions only. The observed distortion of Pb coordination polyhedra in hereroite is due to the  
131 stereoactive behaviour of the 6s<sup>2</sup> lone electron pairs. Recent study by Walsh et al. (2011) suggests also  
132 that distortion in Pb coordination depends on the electronic states of the coordinated anions and is the  
133 result of direct electronic interaction between the cation *s* and anion *p* orbitals. Our previous studies of  
134 Pb oxyhalides (Krivovichev and Burns 2001a, b, 2002, 2006; Siidra et al. 2007a, b, c; 2008a, b) support  
135 this conclusion by demonstration that halide ions and lone pairs on Pb<sup>2+</sup> cations associate in the same  
136 regions of the crystal structures. This phenomenon can be interpreted as a sign of the soft-soft attraction  
137 between halide ions and lone electron pairs known as halophilicity of the lone electron pairs.

138         The structure of hereroite contain two tetrahedrally coordinated sites. The As site is coordinated  
139 by four O anions with the <As-O> bond length equal to 1.70 Å. The AsO<sub>4</sub> tetrahedra are disordered  
140 with the O8 site split into the O8A (S.O.F. = 0.58(3)) and O8B (S.O.F. = 0.42(3)) sites. The second  
141 tetrahedral site, *T*, has mixed occupancy and the <*T*-O> bond length is 1.65 Å. The occupancy of this  
142 site can be written as [Si<sub>0.48</sub>As<sub>0.29</sub>V<sub>0.15</sub>Mo<sub>0.09</sub>], in accordance with the chemical data. Similar mixed-

143 occupied Si sites with the individual *T*-O distances comparable to those in hereroite occur in the  
144 structures of dixenite and saneroite (Nagashima and Armbruster, 2010).

145 The total number of oxygen sites is 21. The O1-O8 sites are bonded to As and T sites, being  
146 part of strong tetrahedral oxyanions. All other O atoms (O9-O20) are tetrahedrally coordinated by Pb  
147 atoms, which results in the formation of oxocentered OPb<sub>4</sub> tetrahedra (Fig. 3). The O9 site is partly  
148 occupied and plays the important role for charge-balance compensation between cationic and anionic  
149 parts of the structure. The <O-Pb> distances within the OPb<sub>4</sub> tetrahedra are in the range of 2.28-2.35 Å,  
150 which is in good agreement with the average value of 2.33 Å derived by Krivovichev et al. (1998).

151 There are five symmetrically independent Cl atoms coordinated by eight Pb atoms each, which  
152 is typical for Pb oxyhalides.

153 Bond-valence sums for the cations and anions in the structure of hereroite are listed in Table 2.  
154 They are in general agreement with the expected oxidation states, taking into account mixed occupancy  
155 for the T site and corresponding deviations of bond-valence sums for the O atoms of the TO<sub>4</sub> tetrahedra  
156 (O1, O2, O3, and O4).

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#### 158 Structure description

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160 The structure of hereroite (Fig. 3c) belongs to the 2:1 type and consists of alternating PbO-type  
161 layers and Cl sheets oriented parallel to the (010) plane. The PbO-type layer is a derivative of the [OPb]  
162 tetrahedral layer in litharge and can be obtained from the latter by removal of blocks of oxocentered  
163 tetrahedra. The AsO<sub>4</sub> and (Si,As,V,Mo)O<sub>4</sub> tetrahedral anions locate in cavities within the PbO-type  
164 layer. The formula of the layer can be written as [O<sub>20</sub>(O,□)Pb<sub>32</sub>] taking into account the partly occupied  
165 O9 site. The atomic structure of the layer including O9 atoms is shown in Figure 4a. Figure 4b shows  
166 the same layer viewed from the standpoint of oxocentered OPb<sub>4</sub> tetrahedra in polyhedral representation.  
167 Neglecting the low occupancy of the O9 site, the formula of the layer can be written as [O<sub>21</sub>Pb<sub>32</sub>]<sup>22+</sup>  
168 (Fig. 4c). There are two types of vacant cavities in the layer: a double-square sevenfold cavity and a

169 square fourfold cavity. The former is occupied by the  $\text{AsO}_4$  tetrahedron, whereas the latter is occupied  
170 by the mixed  $\text{TO}_4$  tetrahedron. The size of the cavities obviously correlates with the size of the  
171 tetrahedral anions: larger arsenate tetrahedra occupy larger cavities, whereas smaller tetrahedra occupy  
172 smaller cavities. The difference in size is rather small, but should not be neglected. This conclusion is  
173 supported by the observations of the structures of other Pb oxychlorides: the larger  $(\text{AsO}_4)^{3-}$  and  
174  $(\text{VO}_4)^{3-}$  anions in kombatite and sahlinite occupy sevenfold cavities, whilst smaller  $(\text{SO}_4)^{2-}$  anions  
175 occupy the fourfold cavities in symesite. As a result, the structure of the  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layer in hereroite  
176 may be described as a combination of the symesite and kombatite modules (Fig. 1d) organized into  
177 alternating diagonal stripes. Figure 4c shows the structure of the Cl sheet in hereroite, which is a  
178 vacancy-free square sheet of Cl ions typically observed in Pb oxyhalides.

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## Discussion

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182 Topological features of the O-Pb tetrahedral layers in lead oxyhalides can be described using  
183 the method of square lattices (Krivovichev et al. 2004, 2006; Siidra et al. 2011). Within this approach,  
184 each  $\text{OPb}_4$  tetrahedron is symbolized by a square. Each square is labeled by the number that  
185 corresponds to the number of the O site at the center of the tetrahedron. The  $[\text{OPb}]$  layer in litharge thus  
186 corresponds to the 2-dimensional (2-D) layer of squares (Table 1) that fill the 2-D plane without gaps  
187 and overlaps. In turn, tetrahedral layers in the PbO derivative structures correspond to two dimensional  
188 arrangements of black and white squares, where the latter symbolize vacancies. The topological  
189 function of a tetrahedron within the layer can be visualized by investigation of the local coordination of  
190 a given square by the adjacent squares, i.e., by all squares with which it has common points. The  
191 arrangement of black squares around the central square  $p$  is its first corona,  $c^1(p)$  (Krivovichev et al.  
192 1997a, b). The second corona,  $c^2(p)$ , is as a set of black squares that surround the first corona, and so  
193 on. Since the  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layer in hereroite consists of modules of the symesite- and kombatite-type



194 layers, these layers are analyzed first in order to understand topological relationships between the three  
195 minerals.

196 Figure 5a shows the 2-D array of squares that represents the arrangement of  $\text{OPb}_4$  tetrahedra  
197 within the  $[\text{O}_7\text{Pb}_{10}]^{6+}$  block in the structure of symesite. The following pairs of tetrahedra have identical  
198 first coronas:  $\text{O5Pb}_4$  and  $\text{O11Pb}_4$ ,  $\text{O6Pb}_4$  and  $\text{O10Pb}_4$ , and  $\text{O8Pb}_4$  and  $\text{O9Pb}_4$ . The  $\text{O7Pb}_4$  tetrahedron  
199 has a unique arrangement and therefore its topological function is unique, even from the viewpoint of  
200 the first-order analysis (Fig. 5c). To further investigate topological functions of different tetrahedra, one  
201 has to examine their higher-order configurations (Fig. 5e). The analysis shows that, within the above  
202 mentioned pairs, tetrahedra have the same topological functions, despite the fact that they are non-  
203 equivalent crystallographically. The topological symmetry of the  $[\text{O}_7\text{Pb}_{10}]^{6+}$  block in symesite is  
204 therefore lower than its crystallographic symmetry.

205 The 2-D array of oxocentered tetrahedra within the  $[\text{O}_9\text{Pb}_{14}]^{8+}$  layer in the structures of  
206 kombatite and sahlinite, and the schemes of the first coronas for all six symmetrically independent  
207 tetrahedra in the structure are shown in Figures 5b and d, respectively. The  $\text{O1Pb}_4$ - $\text{O3Pb}_4$  pair and the  
208  $\text{O7Pb}_4$ - $\text{O8Pb}_4$  pair have identical first coronas. However, their second coronas are different, and  
209 therefore all symmetry-independent  $\text{OPb}_4$  tetrahedra have different topological functions.

210 As the  $[\text{O}_{21}\text{Pb}_{32}]^{22+}$  layer in the structure of hereroite consists of both kombatite and symesite  
211 modules, it can be expected that its topological structure is more complex than that of the parent  
212 structures (Fig. 6). The analysis of square lattice shows that it is the case indeed. For instance, the O10-  
213 , O11-, O14-, O15-, O17-, and O19-centered  $\text{OPb}_4$  tetrahedra have the same first coronas. Analysis of  
214 the second-order configurations shows that O15- and O19-centered tetrahedra possess unique second  
215 coronas, whereas, for the rest the second coronas are identical. The O10-, O11-, O14-, and O17-  
216 centered  $\text{OPb}_4$  tetrahedra can be distinguished by their third coronas only, which means that their  
217 topological functions are unique. The O18- and O20-centered tetrahedra display even more complex  
218 behaviour, since they can be distinguished by their fourth coronas only. In general, all symmetry-  
219 independent  $\text{OPb}_4$  tetrahedra in hereroite are unique in their global topological properties.

220 The most striking feature of the structure of hereroite is the combination of the kombatite and  
221 symesite modules arranged within one structural unit. Recently, modular approach to the structures of  
222 complex inorganic phases attracted considerable attention (see, e.g. Ferraris et al. 2004) and proved to  
223 be the useful concept for rational design of novel families of inorganic compounds (Kabbour et al.  
224 2006). The structure of hereroite shows that this approach is also operational for complex Pb oxyhalide  
225 phases accommodating tetrahedral anions of slightly different size (e.g. silicate and arsenate). The  
226 observed dependence of the topology of the O-Pb layer upon the size of tetrahedral anionic species is  
227 remarkable and may be used for preparation of new layered inorganic compounds with novel structural  
228 architectures.

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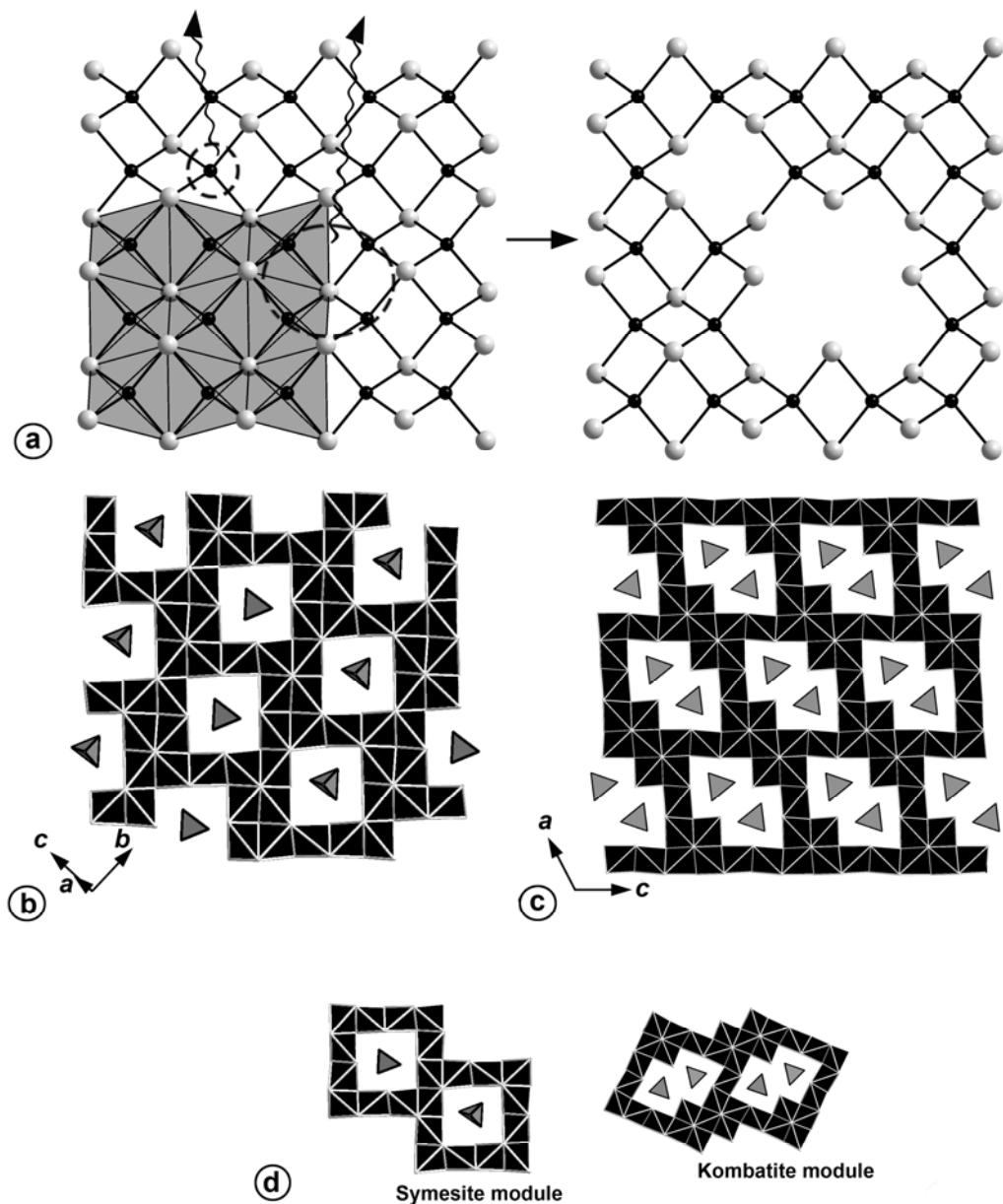
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## Figure Captions



374

375 Figure 1. The [OPb] layer in litharge and formation of vacancies due to the removal of Pb and

376 O atoms (a); the  $[O_7Pb_{10}]^{6+}$  layer in symesite shown as formed by edge-sharing  $OPb_4$

377 tetrahedra with the square shaped vacancies occupied by the  $SO_4$  anions (b); the  $[O_9Pb_{14}]^{10+}$

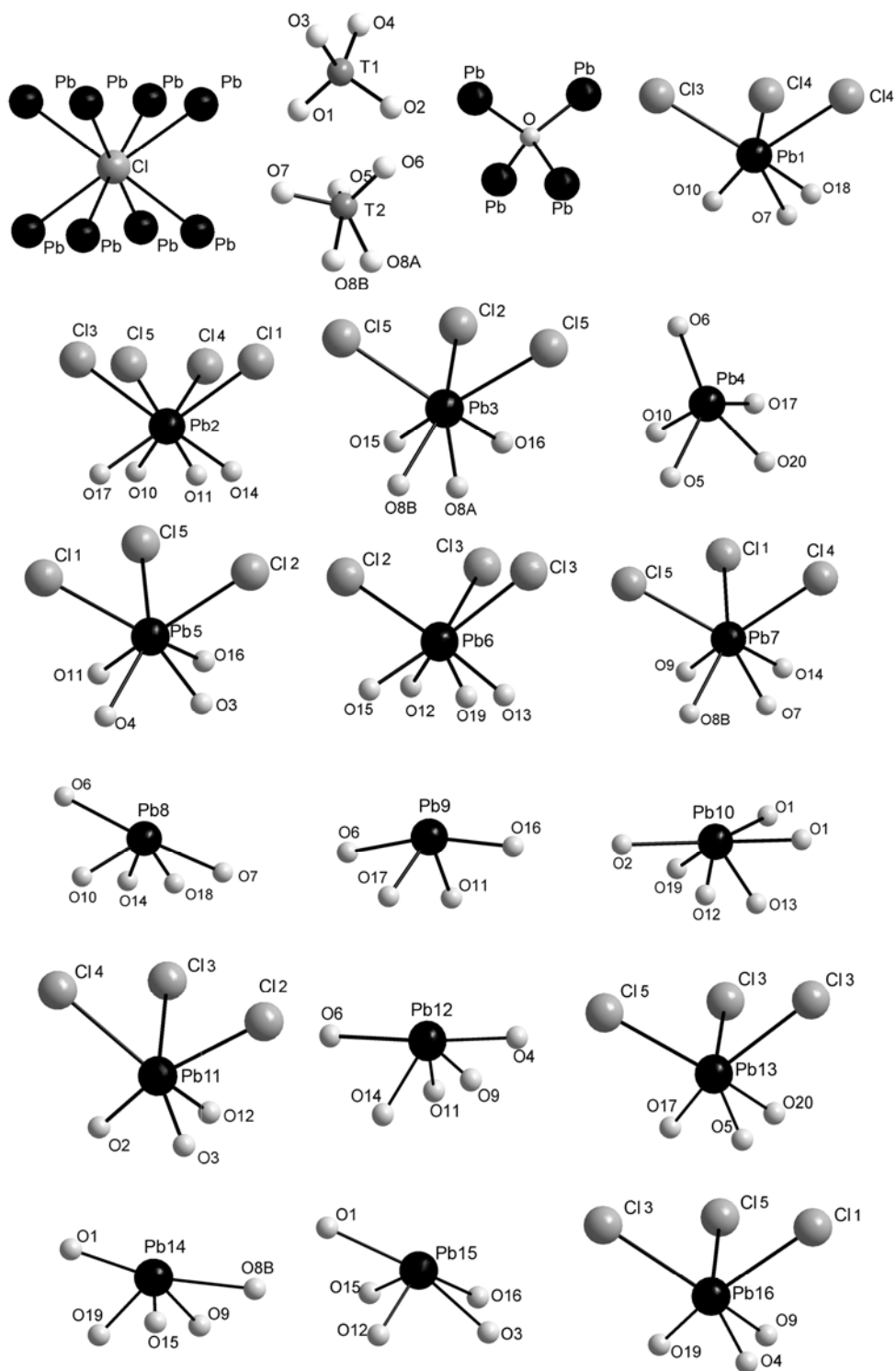
378 layer in kombatite and sahlinite with the double square vacancies occupied by the  $VO_4$  or

379  $AsO_4$  anions (c). Modules derived from the crystal structures of symesite and kombatite (d).

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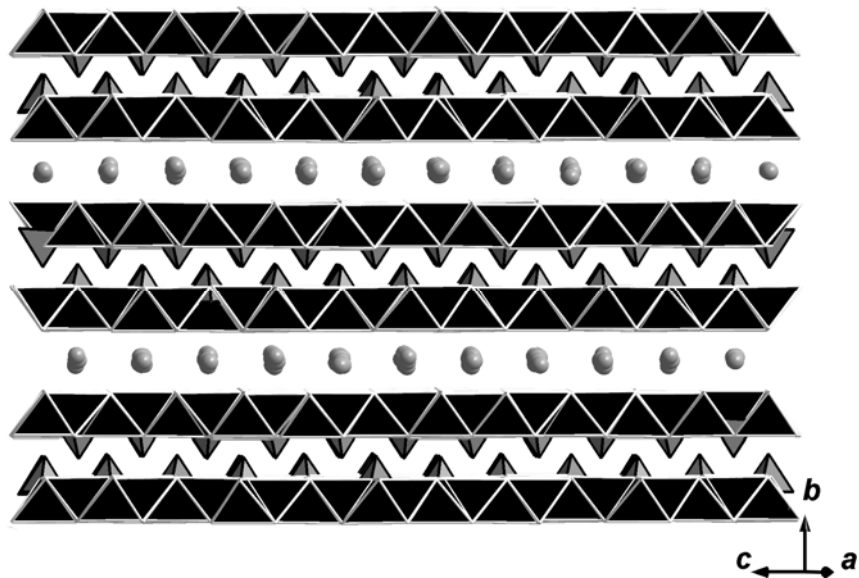


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383 Figure 2. Coordination of atoms in the crystal structure of hereroite.



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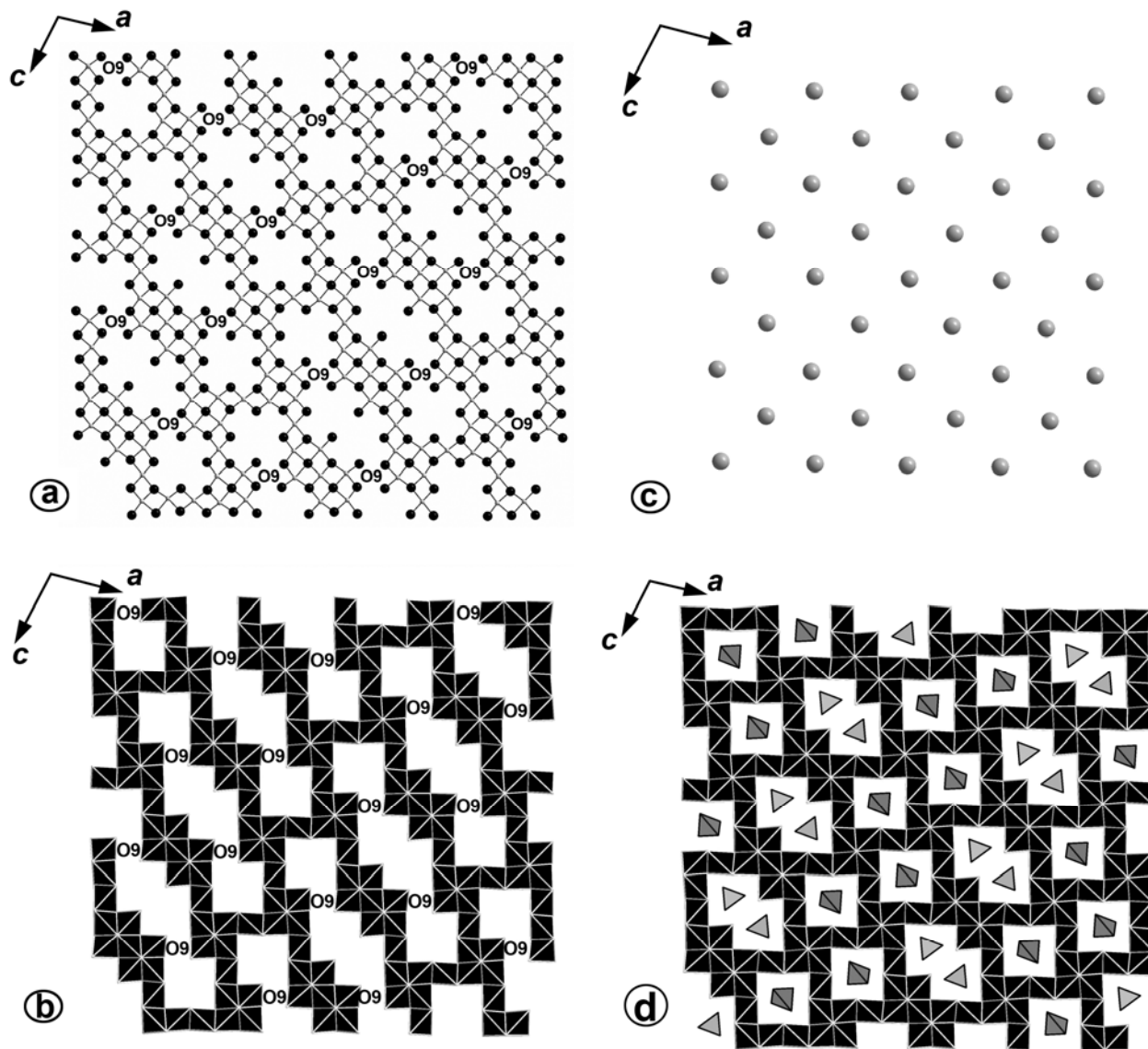
386 Figure 3. Projection of the crystal structure of hereroite along the extension of the PbO-type

387 layers. Legend: OPb<sub>4</sub> tetrahedra = dark, AsO<sub>4</sub> and TO<sub>4</sub> tetrahedra are grey, Cl atoms are

388 shown as spheres.

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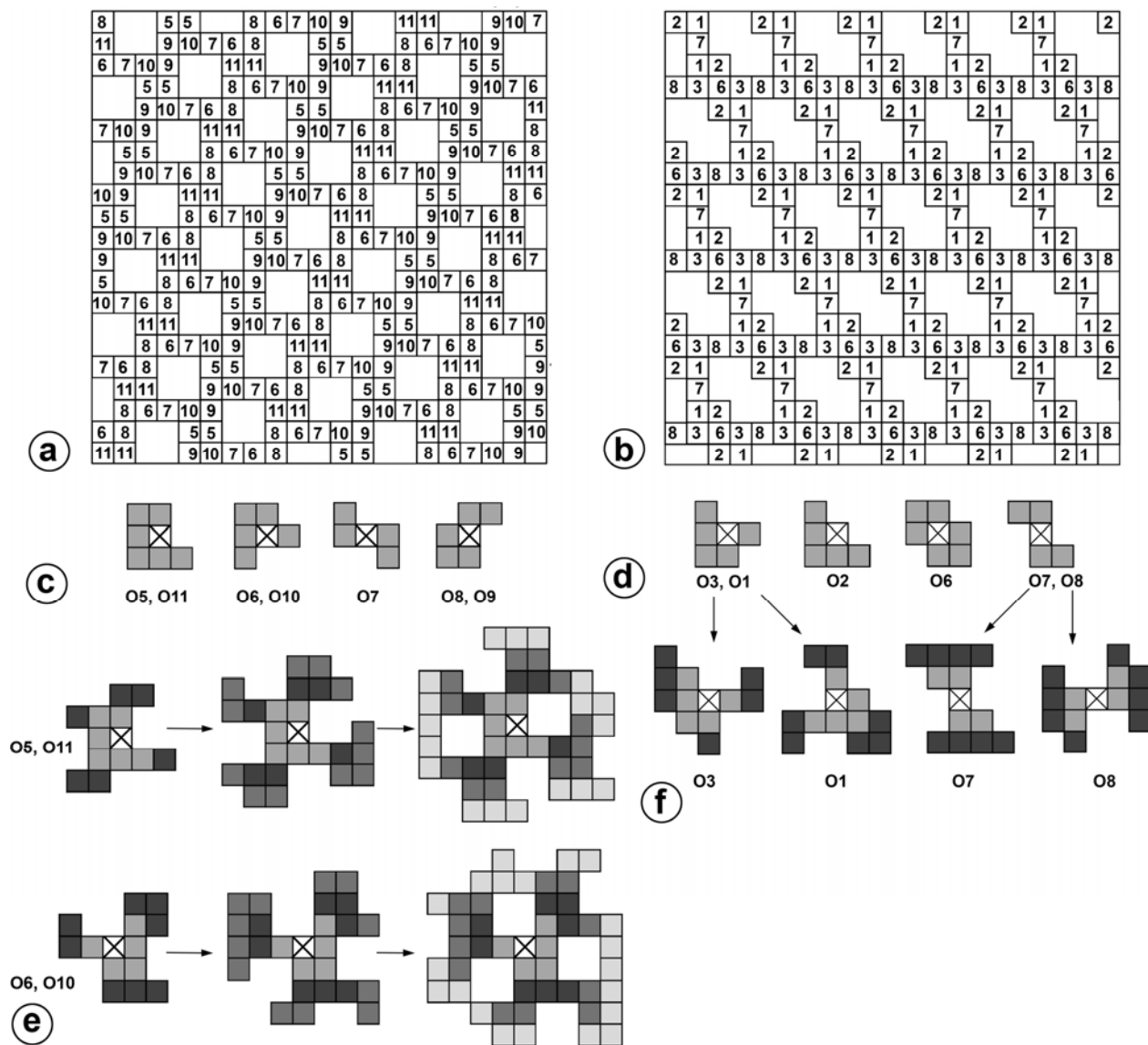
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393 Figure 4. Ball-and-stick (a) and polyhedral (b) representations of the PbO-type layer in the

394 crystal structure of hereroite. Under-occupied O9 site is shown. The layer of Cl<sup>-</sup> anions (c).395 General projection of the [O<sub>21</sub>Pb<sub>32</sub>]<sup>22+</sup> layer with tetrahedral anions (d).

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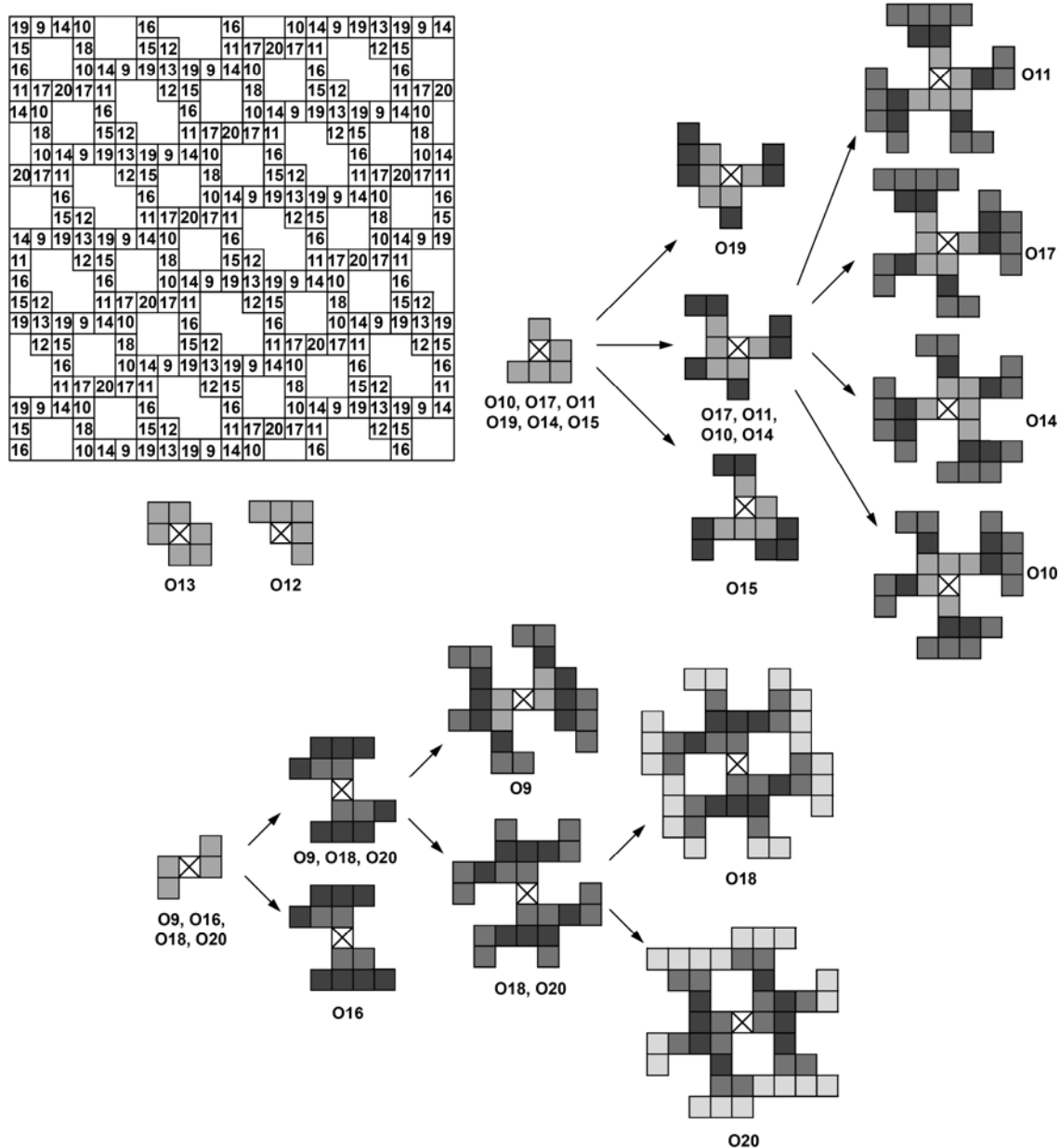
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399 Figure 5. Topological structure of the  $[O_7Pb_{10}]^{6+}$  layer in symesite (a) and the  $[O_9Pb_{14}]^{8+}$  layer  
 400 in komatite (b) examined using the method of square lattices: first coronas of the  $OPb_4$   
 401 tetrahedra (central tetrahedra are shown as crossed squares) in the structures of symesite  
 402 and komatite (c, d, respectively); second, third, and fourth coronas of the  $O_5Pb_4$ ,  $O_{11}Pb_4$ ,  
 403  $O_6Pb_4$ , and  $O_{10}Pb_4$  tetrahedra in symesite (e); second coronas of the  $O_1Pb_4$ ,  $O_3Pb_4$ ,  $O_7Pb_4$ ,  
 404 and  $O_8Pb_4$  tetrahedra in komatite (f).

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407 Figure 6. Topological structure of the  $[O_{21}Pb_{32}]^{22+}$  layer in hereroite examined using the  
 408 method of square lattices. See the text for details.

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414 **Table 1.** Crystallographic data for layered Pb oxychloride minerals with PbO-related structures

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Mineral name	Chemical formula	Space group	<i>a</i> [Å] / $\alpha$ [deg]	<i>b</i> [Å] / $\beta$ [deg]	<i>c</i> [Å] / $\gamma$ [deg]	Ref.
Rumseyite	(Pb <sub>2</sub> OF)Cl	<i>I4/mmm</i>	4.065 / 90	4.065 / 90	12.631 / 90	1
Blixite	Pb <sub>8</sub> O <sub>5</sub> (OH) <sub>2</sub> Cl <sub>4</sub>	<i>C2/c</i>	26.069 / 90	5.835 / 102.612	22.736 / 90	2
Symesite	Pb <sub>10</sub> (SO <sub>4</sub> )O <sub>7</sub> Cl <sub>4</sub> (H <sub>2</sub> O)	<i>B<math>\bar{1}</math></i>	19.727 / 82.21	8.796 / 78.08	13.631 / 100.04	3
Kombatite	Pb <sub>14</sub> (VO <sub>4</sub> ) <sub>2</sub> O <sub>9</sub> Cl <sub>4</sub>	<i>C2/c</i>	12.682 / 90	22.566 / 118.11	11.279	4
Sahlinite	Pb <sub>14</sub> (AsO <sub>4</sub> ) <sub>2</sub> O <sub>9</sub> Cl <sub>4</sub>	<i>C2/c</i>	12.704 / 90	22.576 / 118.37	11.287 / 90	5
Mereheadite	Pb <sub>47</sub> O <sub>24</sub> (OH) <sub>13</sub> Cl <sub>25</sub> (BO <sub>3</sub> ) <sub>2</sub> (CO <sub>3</sub> )	<i>Cm</i>	17.372 / 90	27.941 / 93.152	10.666 / 90	6
Hereroite	[Pb <sub>32</sub> O <sub>20</sub> (O, $\square$ )](AsO <sub>4</sub> ) <sub>2</sub> ((Si,As,V,Mo)O <sub>4</sub> ) <sub>2</sub> Cl <sub>10</sub>	<i>C2/c</i>	23.139 / 90	22.684 / 102.090(3)	12.389 / 90	7
Vladkrivovichevite	[Pb <sub>32</sub> O <sub>18</sub> ][Pb <sub>4</sub> Mn <sub>2</sub> O] Cl <sub>14</sub> (BO <sub>3</sub> ) <sub>8</sub> ·2H <sub>2</sub> O	<i>Pmmn</i>	12.759 / 90	27.169 / 90	11.515 / 90	8
Sundiusite*	Pb <sub>10</sub> O <sub>8</sub> (SO <sub>4</sub> )Cl <sub>2</sub>	<i>C2/m,</i> <i>Cm, C2</i>	24.67 / 90	3.781 / 100.07	11.881 / 90	9

416 \* structure unknown

417

418 References: (1) Turner et al. 2012b; (2) Krivovichev and Burns, 2006; (3) Welch et al. 2000; (4)

419 Cooper and Hawthorne 1994; (5) Bonaccorsi and Pasero 2003; (6) Krivovichev et al. 2009; (7) this

420 work; (8) Siidra et al. 2012; (9) Dunn and Rouse 1980.

421

422

423 **Table 2.** Crystallographic data and  
 424 refinement parameters for hereroite

<i>a</i> (Å)	23.139(4)
<i>b</i> (Å)	22.684(4)
<i>c</i> (Å)	12.389(2)
$\beta$ (°)	102.090(3)
<i>V</i> (Å <sup>3</sup> )	6358.8(18)
Space group	<i>C2/c</i>
<i>Z</i>	4
<i>D</i> <sub>calc</sub> (g/cm <sup>3</sup> )	8.144
$\mu$ (mm <sup>-1</sup> )	86.456
<i>F</i> <sub>000</sub>	12748.2
Crystal size (mm)	0.11×0.10×0.08
Radiation	MoK $\alpha$
<i>h</i> <sub>min</sub> , <i>h</i> <sub>max</sub>	-38, 33
<i>k</i> <sub>min</sub> , <i>k</i> <sub>max</sub>	-37, 37
<i>l</i> <sub>min</sub> , <i>l</i> <sub>max</sub>	-20, 20
$2\theta$ <sub>min</sub> , $2\theta$ <sub>max</sub>	1.27, 36.15
Total Ref.	56375
Unique Ref.	14483
Unique $ F_o  \geq 4\sigma_F$	6931
<i>R</i> <sub>1</sub>	0.054
<i>wR</i> <sub>2</sub>	0.108
<i>S</i>	1.002

425  
426**Table 3.** Atomic coordinates, displacement parameters ( $\text{\AA}^2$ ), bond-valence sums\* (valence units) and site-occupancy factors (SOFs) for hereroite

Atom	BVS	x	y	z	$U_{eq}$	$U_{11}$	$U_{22}$	$U_{33}$	$U_{23}$	$U_{13}$	$U_{12}$
Pb1	1.90	0.02370(3)	0.33013(3)	0.11179(5)	0.01240(13)	0.0127(3)	0.0114(3)	0.0122(3)	-0.0007(2)	0.0004(2)	0.0030(2)
Pb2	1.99	-0.13847(3)	0.33113(2)	-0.05572(5)	0.00938(10)	0.0116(2)	0.0066(2)	0.0097(2)	-0.0005(2)	0.00158(17)	-0.0008(2)
Pb3	2.17	0.24274(3)	0.16730(3)	0.00179(5)	0.01195(13)	0.0126(3)	0.0103(3)	0.0117(3)	0.0010(2)	-0.0002(2)	-0.0027(2)
Pb4	1.93	-0.03666(3)	0.55670(3)	0.37405(5)	0.01179(12)	0.0157(3)	0.0076(3)	0.0121(2)	-0.0009(2)	0.0032(2)	0.0037(2)
Pb5	1.88	-0.30232(3)	0.33470(3)	-0.22005(5)	0.01283(13)	0.0158(3)	0.0104(3)	0.0106(3)	0.0016(2)	-0.0010(2)	-0.0036(2)
Pb6	1.90	0.08073(2)	0.16427(2)	-0.17277(5)	0.00920(12)	0.0099(2)	0.0057(2)	0.0111(3)	0.0008(2)	0.00000(19)	-0.0001(2)
Pb7	1.95	-0.19191(3)	0.66864(3)	-0.28106(5)	0.01232(13)	0.0125(3)	0.0115(3)	0.0125(3)	-0.0048(2)	0.0016(2)	-0.0015(2)
Pb8	1.88	-0.07902(3)	0.44391(3)	0.15108(5)	0.01180(13)	0.0131(3)	0.0062(3)	0.0154(3)	-0.0033(2)	0.0014(2)	-0.0008(2)
Pb9	1.83	-0.19752(3)	0.44554(3)	-0.26142(5)	0.01285(13)	0.0136(3)	0.0076(3)	0.0175(3)	0.0046(2)	0.0035(2)	0.0017(2)
Pb10	1.86	-0.02290(3)	0.04875(3)	-0.11756(6)	0.01329(12)	0.0133(3)	0.0091(2)	0.0171(3)	0.0060(2)	0.0025(2)	-0.0008(2)
Pb11	1.87	0.03095(3)	0.16428(3)	0.11131(5)	0.01614(14)	0.0170(3)	0.0166(3)	0.0149(3)	-0.0063(3)	0.0034(2)	0.0028(3)
Pb12	1.91	-0.23979(3)	0.55461(3)	-0.48495(5)	0.01140(12)	0.0153(3)	0.0069(3)	0.0114(2)	0.0020(2)	0.0016(2)	-0.0016(2)
Pb13	1.83	-0.08201(3)	0.67123(3)	0.17201(5)	0.01177(12)	0.0133(3)	0.0097(3)	0.0125(3)	0.0035(2)	0.0030(2)	0.0019(2)
Pb14	1.96	0.18340(3)	0.05181(3)	-0.23282(5)	0.01399(13)	0.0170(3)	0.0115(3)	0.0126(3)	-0.0014(2)	0.0012(2)	0.0061(2)
Pb15	1.75	0.13716(3)	0.05425(3)	0.03687(5)	0.01445(13)	0.0137(3)	0.0092(3)	0.0205(3)	0.0069(2)	0.0036(2)	0.0025(2)
Pb16	1.78	0.14468(3)	0.16517(3)	-0.43715(5)	0.01309(12)	0.0142(3)	0.0117(3)	0.0146(3)	0.0041(2)	0.0056(2)	0.0026(2)
T1 <sup>s</sup>	3.88	0.09489(10)	-0.06941(11)	-0.15611(18)	0.0129(6)	0.0111(11)	0.0172(12)	0.0104(10)	0.0015(8)	0.0024(7)	-0.0016(9)
As1	5.09	-0.13500(8)	0.56799(8)	-0.05413(16)	0.0174(5)	0.0179(9)	0.0133(9)	0.0225(9)	0.0010(8)	0.0076(7)	0.0004(7)
Cl1	0.92	-¼	¼	0	0.0207(14)	0.022(3)	0.024(4)	0.017(3)	0.003(2)	0.007(2)	0.004(2)
Cl2	0.66	0.1401(1)	0.2363(2)	0.0553(4)	0.0204(8)	0.016(1)	0.027(2)	0.016(1)	-0.002(2)	0.0005(14)	0.001(1)
Cl3	0.78	-0.0270(1)	0.24850(18)	-0.1112(4)	0.0215(10)	0.019(2)	0.028(3)	0.018(2)	-0.0065(19)	0.0047(17)	-0.002(1)
Cl4	0.89	-0.0782(1)	0.25894(19)	0.1645(3)	0.0192(9)	0.021(2)	0.0171(19)	0.017(2)	0.0032(16)	-0.0002(17)	0.001(1)
Cl5	0.79	-0.3020(1)	0.75128(19)	-0.2273(4)	0.0200(10)	0.018(2)	0.027(3)	0.015(2)	0.0006(17)	0.0044(17)	0.001(1)
O1	1.76	0.0926(6)	0.0011(6)	-0.1634(10)	0.030(3)						
O2	1.60	0.0413(5)	-0.0931(6)	-0.0975(11)	0.031(3)						
O3	1.71	-0.4013(5)	0.3997(5)	-0.2770(9)	0.016(2)						
O4	1.75	-0.3410(5)	0.4083(6)	-0.0730(10)	0.024(3)						
O5	1.88	-0.0749(9)	0.5950(9)	0.0400(16)	0.077(6)						
O6	1.97	-0.1331(5)	0.4952(6)	-0.0590(10)	0.027(3)						
O7	1.98	-0.1185(9)	0.5952(11)	-0.1698(18)	0.099(8)						
O8A <sup>#</sup>	1.94	-0.1839(14)	0.5906(15)	0.015(3)	0.071(12)						
O8B <sup>£</sup>	1.66	-0.2059(16)	0.5928(17)	-0.115(3)	0.048(13)						
O9 <sup>¶</sup>	1.89	0.2278(5)	0.1153(5)	-0.3526(9)	0.008(3)						
O10	1.99	-0.0489(4)	0.3895(5)	0.0160(8)	0.014(2)						



O11	2.00	-0.2255(4)	0.3907(5)	-0.1289(8)	0.010(2)
O12	1.95	0.0560(4)	0.1105(5)	-0.0214(8)	0.012(2)
O13	2.10	0	0.1031(6)	- $\frac{1}{4}$	0.006(3)
O14	2.02	-0.1587(4)	0.3887(5)	0.0934(9)	0.014(2)
O15	2.01	0.1695(4)	0.1091(5)	-0.0956(8)	0.011(2)
O16	1.86	0.2221(5)	0.1126(5)	0.1404(9)	0.018(3)
O17	1.98	-0.1172(4)	0.3897(5)	-0.2119(9)	0.015(2)
O18	1.90	0	0.3853(7)	$\frac{1}{4}$	0.013(3)
O19	1.96	0.1077(4)	0.1034(5)	-0.3268(8)	0.012(2)
O20	1.94	0	0.6202(7)	$\frac{1}{4}$	0.015(3)

427 \* calculated using bond-valence parameters from Krivovichev and Brown (2001) for the  $\text{Pb}^{2+}$ -O bonds and from Brown and Altermatt (1985) for other  
 428 bonds (bond-valences for the T-O bonds calculated using the  $\text{Si}^{4+}$ -O bond-valence parameters)

429  $\text{\$ Si}_{0.48}\text{As}_{0.28}\text{V}_{0.16}\text{Mo}_{0.08}$  Note: assigned on the basis of chemical data.

430 # S.O.F. = 0.58(3)

431 £ S.O.F. = 0.42(3)

432 ¶ S.O.F. = 0.85(3)

433 **Table 4.** Selected bond lengths (Å) in the structure of hereroite

434

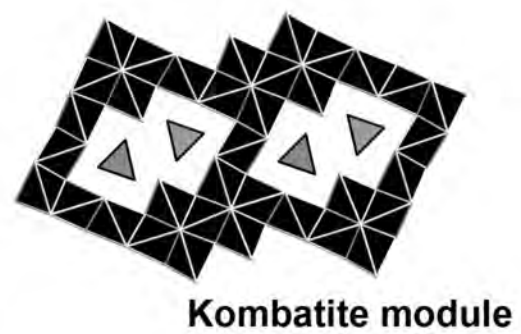
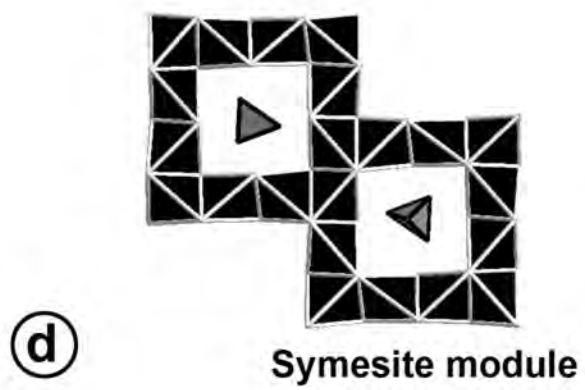
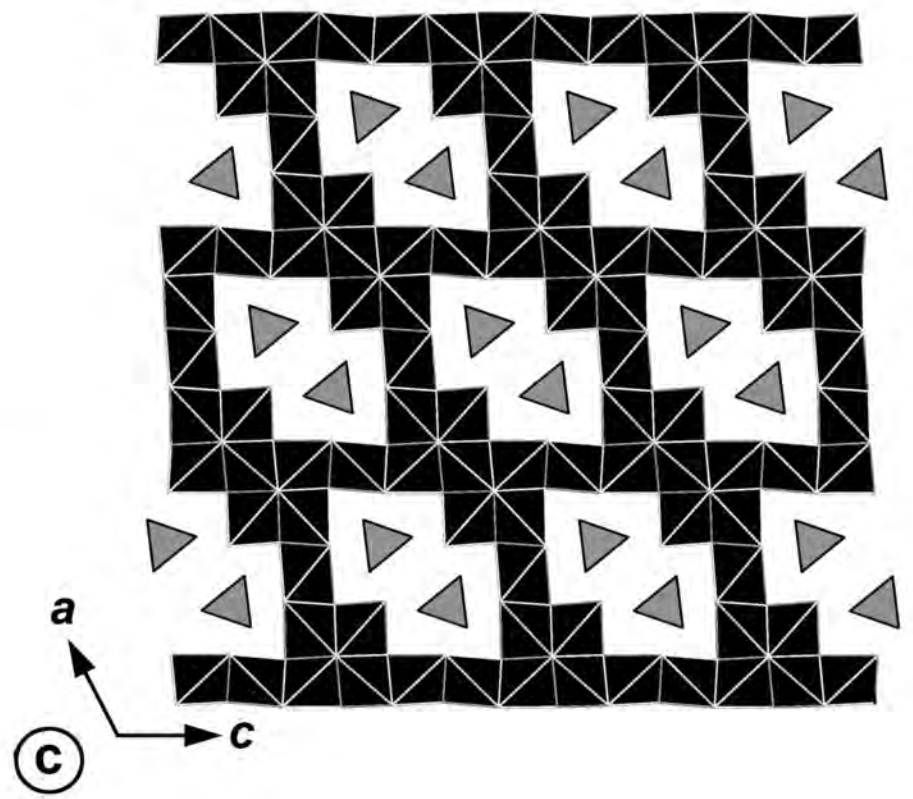
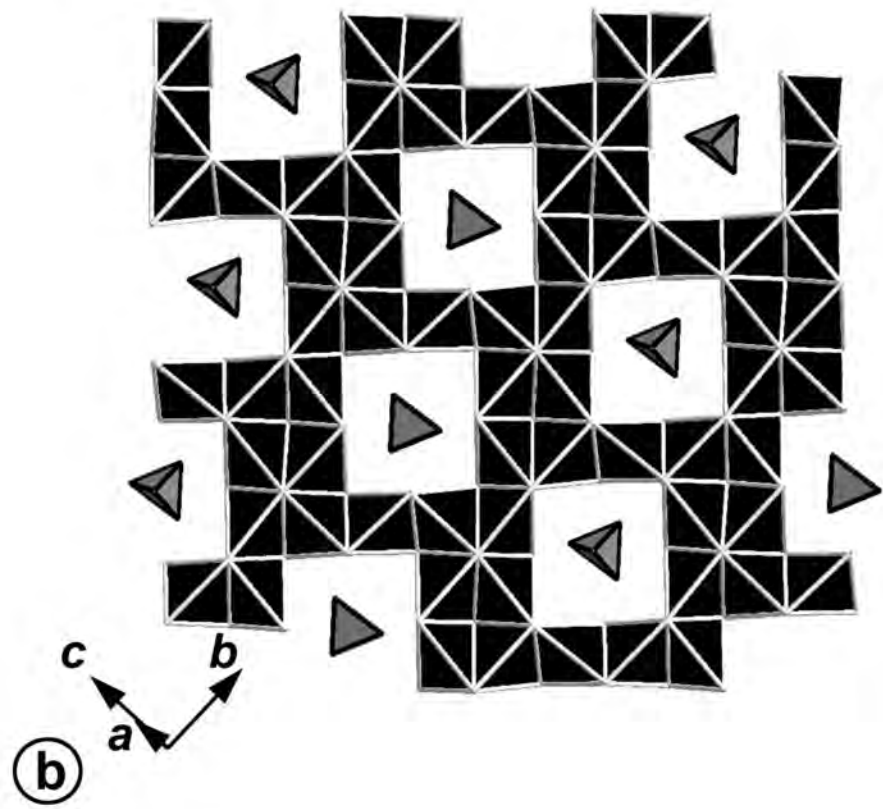
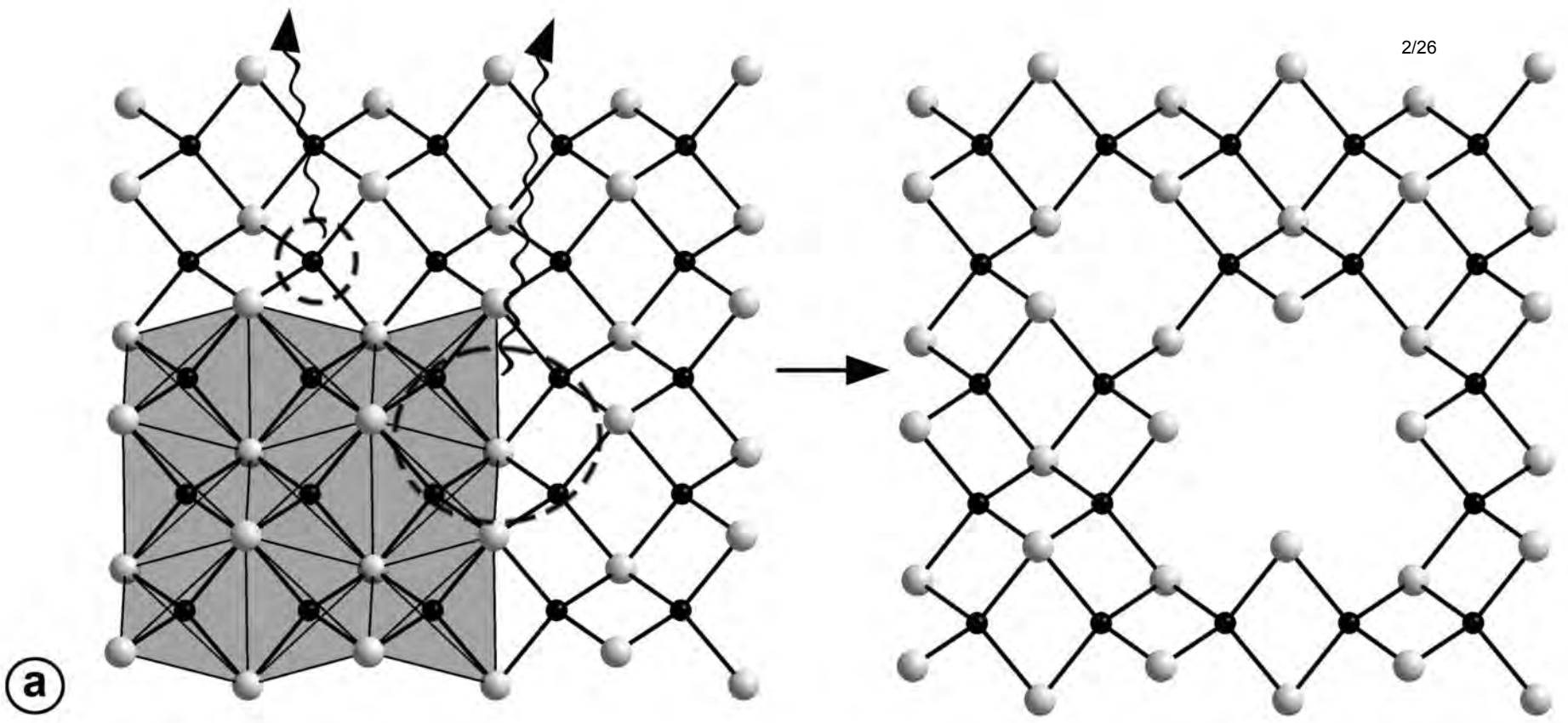
Pb1-O18	2.279(8)	Pb7-O9	2.239(11)	Pb14-O15	2.215(10)
Pb1-O10	2.281(10)	Pb7-O14	2.281(11)	Pb14-O19	2.223(10)
Pb1-O7	2.74(2)	Pb7-O7	2.56(2)	Pb14-O9	2.444(12)
Pb1-O5	2.96(2)	Pb7-O8B	2.75(4)	Pb14-O1	2.690(13)
Pb1-Cl4	3.040(5)	Pb7-Cl4	3.289(4)	Pb14-O8B	2.83(4)
Pb1-Cl4	3.228(4)	Pb7-Cl1	3.3232(7)		
Pb1-Cl3	3.329(4)	Pb7-Cl5	3.341(4)	Pb15-O12	2.260(10)
				Pb15-O15	2.306(11)
Pb2-O14	2.389(11)	Pb8-O14	2.220(10)	Pb15-O16	2.491(11)
Pb2-O11	2.437(10)	Pb8-O10	2.303(11)	Pb15-O1	2.757(12)
Pb2-O10	2.463(10)	Pb8-O18	2.382(8)	Pb15-O3	2.839(11)
Pb2-O17	2.480(11)	Pb8-O7	2.72(2)		
Pb2-Cl4	3.236(4)	Pb8-O6	2.886(12)	Pb16-O19	2.248(11)
Pb2-Cl4	3.295(4)			Pb16-O9	2.287(11)
Pb2-Cl1	3.3570(7)	Pb9-O17	2.227(10)	Pb16-O4	2.440(12)
Pb2-Cl5	3.373(5)	Pb9-O11	2.259(11)	Pb16-Cl5	3.279(4)
		Pb9-O16	2.393(11)	Pb16-Cl1	3.3231(7)
Pb3-O16	2.250(12)	Pb9-O6	2.866(12)	Pb16-Cl3	3.468(5)
Pb3-O15	2.284(10)	Pb9-O8A	2.96(3)		
Pb3-O8A	2.41(3)			T1-O1	1.60(1)
Pb3-O8B	2.66(4)	Pb10-O13	2.204(7)	T1-O2	1.65(1)
Pb3-Cl2	3.030(5)	Pb10-O19	2.302(10)	T1-O3	1.67(1)
Pb3-Cl5	3.390(4)	Pb10-O12	2.410(10)	T1-O4	1.69(1)
Pb3-Cl5	3.395(4)	Pb10-O2	2.961(13)		
		Pb10-O1	3.044(13)	As1-O8A	1.64(3)
Pb4-O10	2.207(11)	Pb10-O1	3.049(12)	As1-O6	1.65(1)
Pb4-O17	2.292(10)			As1-O7	1.68(2)
Pb4-O20	2.388(9)	Pb11-O12	2.220(11)	As1-O5	1.72(1)
Pb4-O5	2.721(19)	Pb11-O2	2.306(13)	As1-O8B	1.75(4)
Pb4-O6	2.798(13)	Pb11-O3	2.359(10)		
		Pb11-Cl2	3.205(5)		
Pb5-O16	2.269(12)	Pb11-Cl3	3.392(5)		
Pb5-O11	2.283(10)	Pb11-Cl4	3.482(4)		
Pb5-O3	2.691(10)				
Pb5-O4	2.756(13)	Pb12-O11	2.253(10)		
Pb5-Cl5	3.239(4)	Pb12-O14	2.313(10)		
Pb5-Cl2	3.250(4)	Pb12-O9	2.380(12)		
Pb5-Cl1	3.3476(7)	Pb12-O4	2.513(11)		
		Pb12-O6	3.026(13)		
Pb6-O13	2.363(8)				
Pb6-O12	2.404(10)	Pb13-O20	2.260(8)		
Pb6-O15	2.425(10)	Pb13-O17	2.265(11)		
Pb6-O19	2.536(10)	Pb13-O5	2.41(2)		
Pb6-Cl2	3.304(5)	Pb13-Cl3	3.273(5)		
Pb6-Cl3	3.311(4)	Pb13-Cl3	3.324(5)		
Pb6-Cl3	3.354(5)	Pb13-Cl5	3.396(4)		

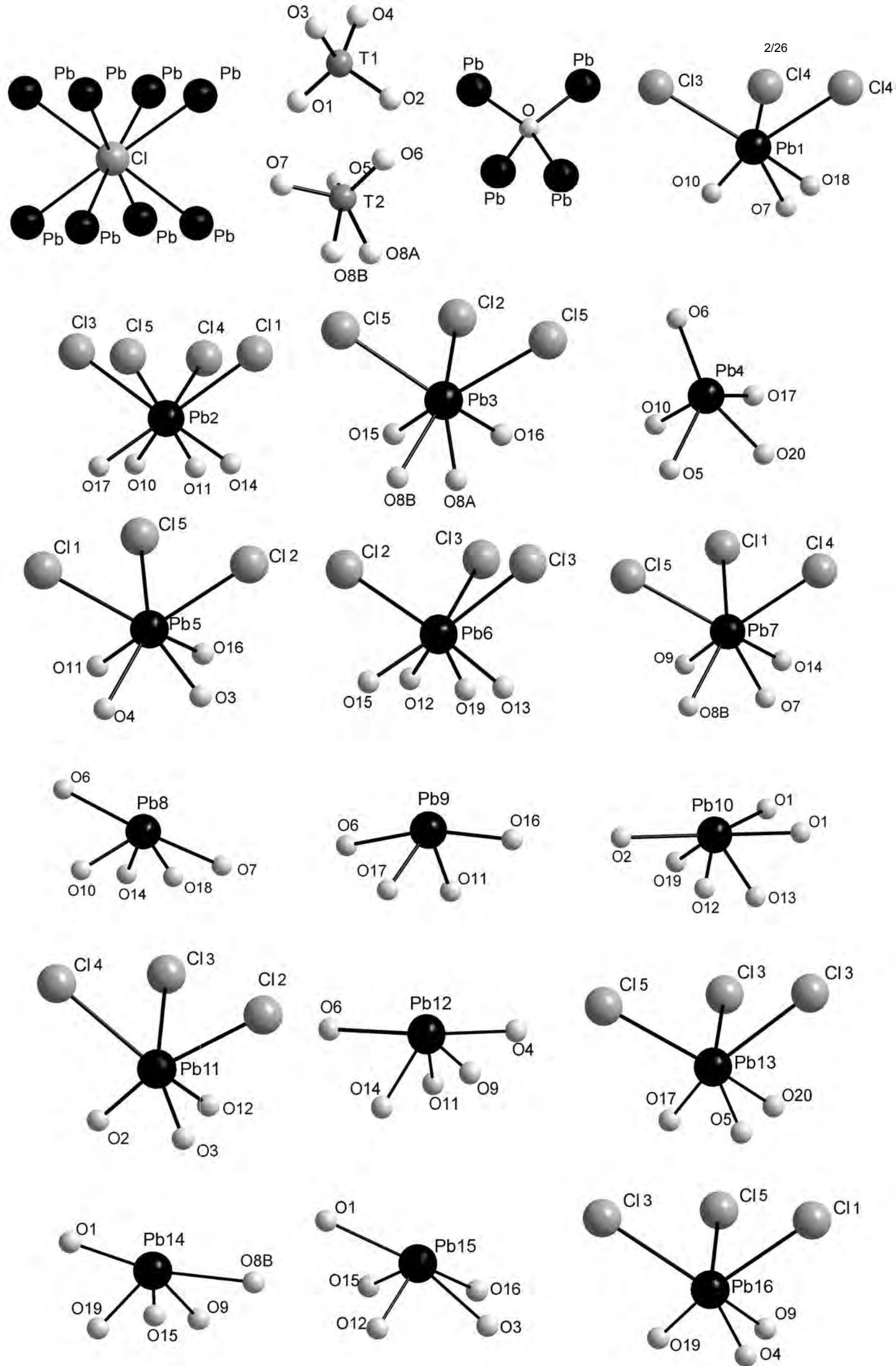
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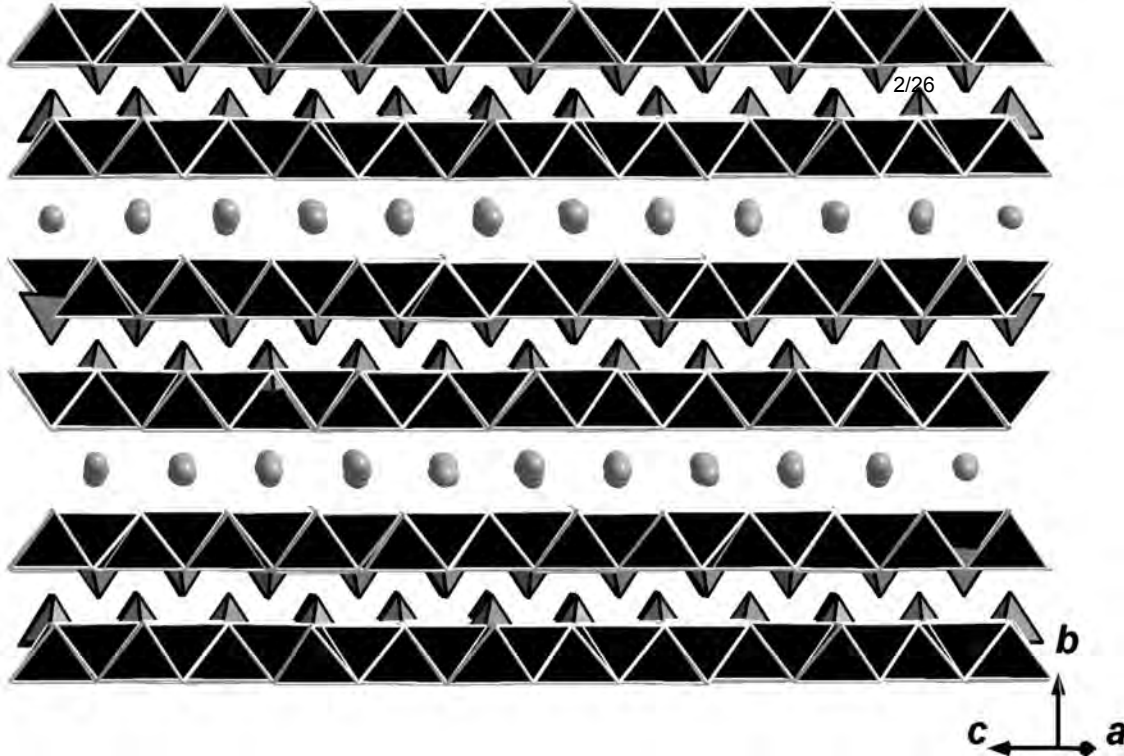
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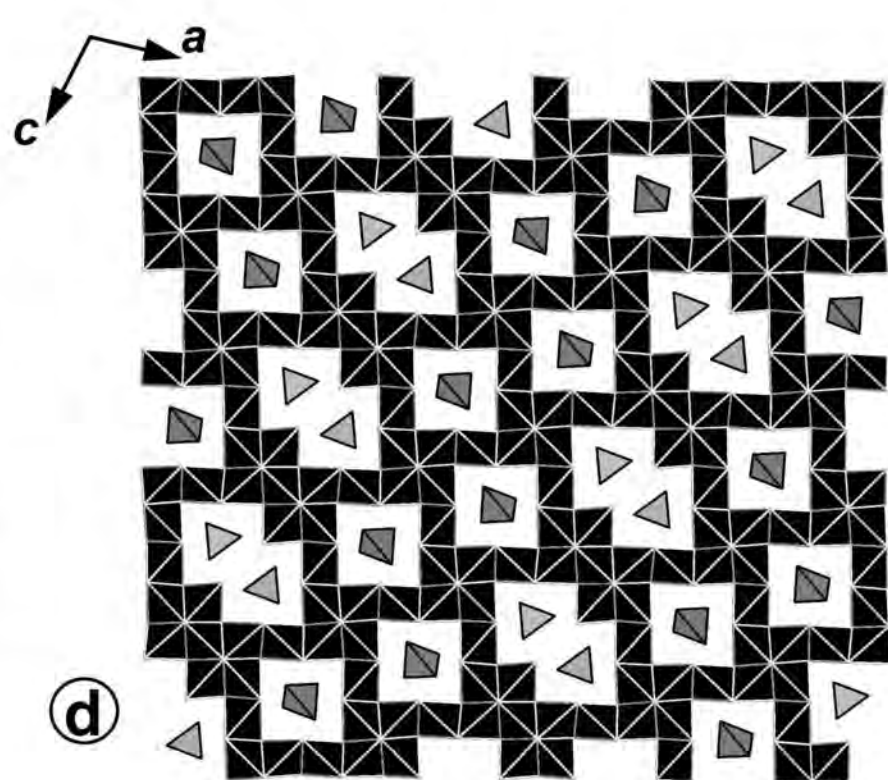
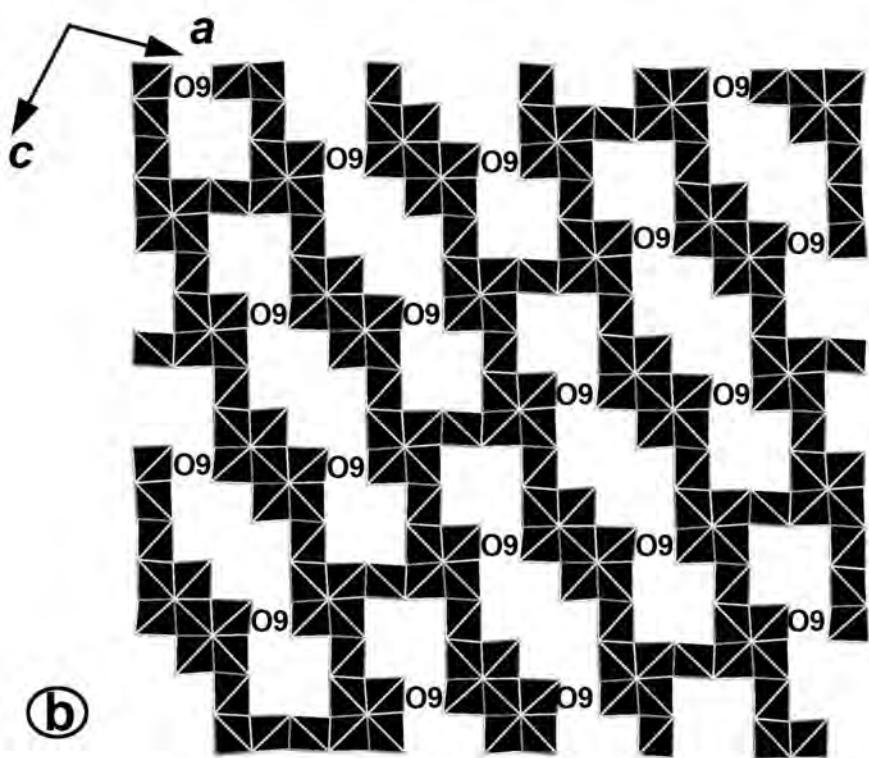
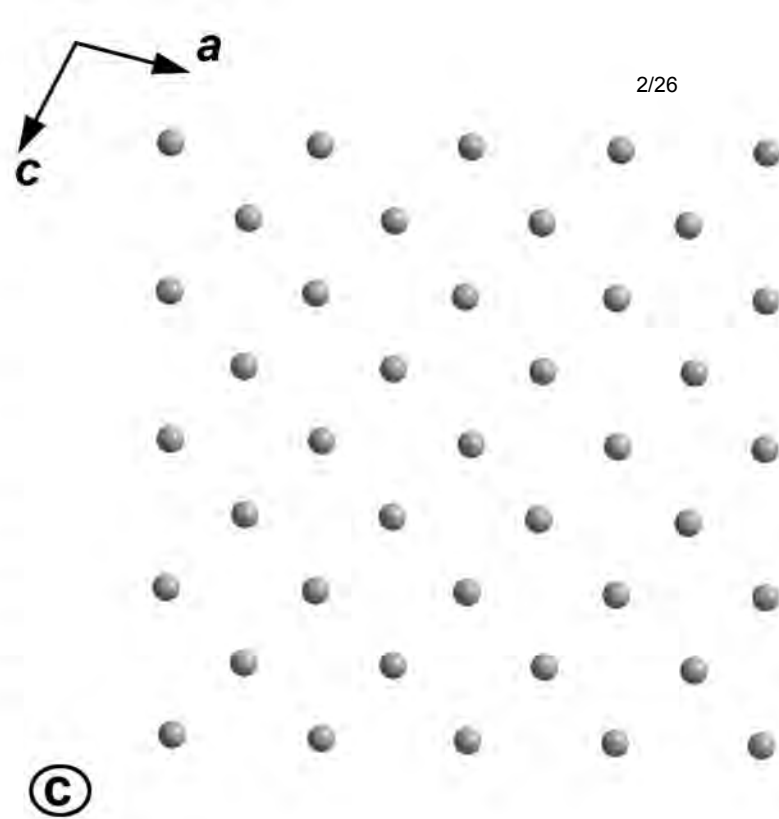
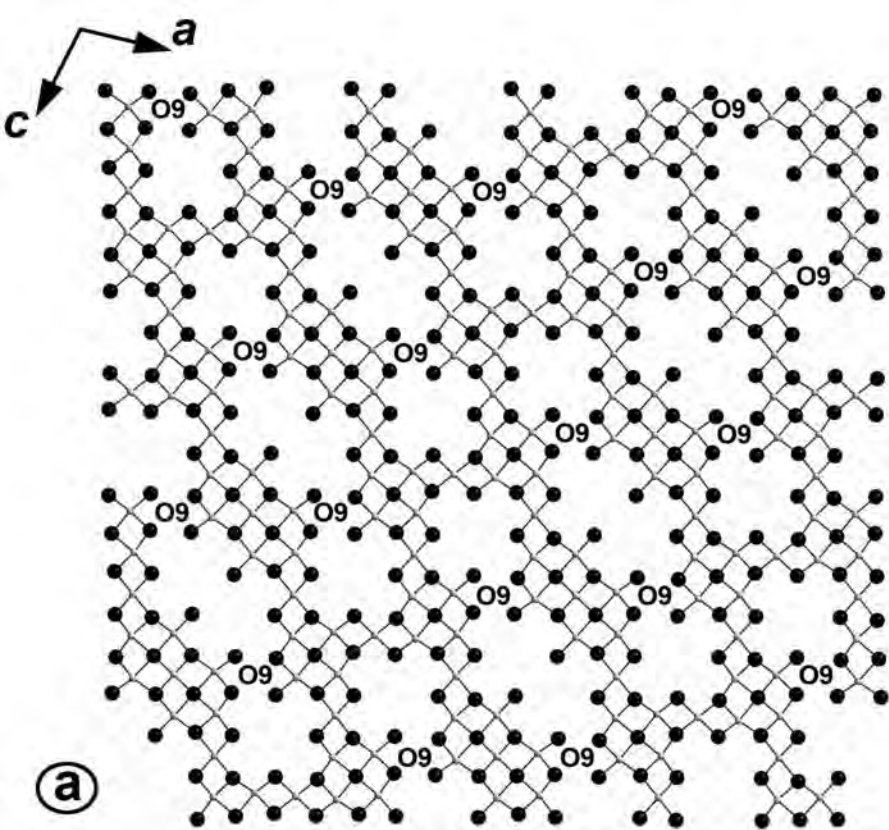
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