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"Revision 1"First principles study on the high-pressure phase transition and elasticity of KAISi₃O₈ hollandite

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ABSTRACT

11 In order to understand the fate of the host phase for potassium subducted into the 12 deep Earth's interior, we have studied the high-pressure stability and elastic 13 properties of KAlSi₃O₈ hollandite (K-hollandite) by means of the first principles 14 computation method. Based on experimental observations, the tetragonal K-15 hollandite I phase was found to undergo the ferroelastic second-order phase 16 transition to the monoclinic K-hollandite II phase at 14.9 GPa. This K-hollandite II 17 phase was mechanically stable up to 150 GPa (i.e., entirely in the Earth's lower 18 mantle), being consistent with previous studies. The Born's elastic stability criteria 19 indicate that the tetragonal mechanical instability occurs at similar pressure of 20 16.9 GPa with shear softening. This causes anomalous pressure dependence of the 21 wave velocities across the instability. Taking a Clapeyron slope of 7 MPa/K and a 22 temperature of 1800 K, the transition pressure becomes ~28 GPa corresponding to 23 about 770 km depth, which would be seismologically detectable and could be 24 comparable to seismic scatterers observed at the shallowest lower mantle. Next, we

- 25 studied the solid solution effect of sodium to K-hollandite, indicating that it is very
- 26 limited on the phase stability, though the Na-end member phase was found to be
- 27 metastable. Elasticity demonstrates strong anisotropy around 15 GPa due to the
- 28 ferroelastic nature.
- 29
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INTORODUCTION

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Long-lived radiogenic nuclides such as ⁴⁰K, ²³²Th, ²³⁵U, and ²³⁸U are essential to 34 understanding the Earth's long-term evolution, while short-lived such as ²⁶Al are 35 36 important only for the very early Earth (Urey 1955). It is believed that these elements 37 are distributed in the continental and oceanic crust and mantle and are concentrated 38 especially in the continental crust. Since recent studies have demonstrated that even 39 continental crust could be subducted into the Earth's deep interior despite of its 40 buoyancy as well as subducted slab (Scholl and von Huene 2007; Kawai et al. 2009), it 41 is important to understand the stability of the host phases of radiogenic nuclides such as ⁴⁰K. 42

Alkali feldspars in the KAlSi₃O₈-NaAlSi₃O₈ system are the most abundant group
of minerals in the Earth's crust. While alkali feldspars are minerals that can survive the
initial stage of subduction, hollandite-type aluminosilicates with the large square tunnels
formed by four double chains of edge-shared octahedra are a candidate for host phases
of large elements and incompatible large-ion lithophile elements (LILEs) such as Rb,
Ba Sr, K, Pb, La, Ce, and Th in the deep mantle. It is, therefore, essential to investigate
stability and elasticity of hollandite in the Earth's deep interior.

50 Potassium feldspar is one of the most abundant minerals in the continental crust. 51 Previous experimental studies on host phases of potassium show that KAlSi₃O₈ felspar 52 dissociates into an assemblage of K₂Si₄O₉ wadeite, Al₂SiO₅ kyanite, and SiO₂ coesite at 53 6-7 GPa and 1500 K, and the three phases further recombine into KAlSi₃O₈ hollandite 54 (hereafter called as K-hollandite I) at about 9 GPa and 1500 K (Urakawa et al. 1994; 55 Yagi et al. 1994). While experiments on K-rich alkali basalt using multi-anvil cell 56 implied a phase transition of K-hollandite I at 22.5 GPa and 2000 K because quenched 57 K-hollandite became Na- and Ca-richer (Wang and Takahashi 1999), some reported that

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58 K-hollandite I remains stable in the pressures range between 20 and 95 GPa at 900 K 59 (Tutti et al. 2001). On the other hand, a new high-pressure I2/m monoclinic structure of 60 KAlSi₃O₈ (K-hollandite II) from an I4/m tetragonal structure at 22 GPa and room 61 temperature was confirmed by in situ X-ray diffraction measurements (Sueda et al. 62 2004). Later, the phase transition was suggested to be second-order (ferroelastic) because of marginal volume change by high-pressure studies up to 32 GPa and room 63 64 temperature using laser-heated anvil cell (Ferroir et al. 2006) and then stability of K-65 hollandite II was investigated up to 128 GPa (Hirao et al. 2008). In contrast to the phase 66 stability, elasticity of the K-hollandite I and II phases was studied by theoretical studies 67 (Mookherjee and Steinle-Neumann 2009). However, these previous studies are 68 controversial in both the phase transition pressure and elasticity. Some studies reported 69 a second-order phase transition pressure of about 23 GPa (Caracas and Boffa Ballaran 70 2010; Deng et al. 2011), while the other suggested a first-order phase transition of 33 71 GPa (Mookherjee and Steinle-Neumann 2009). As for elasticity, one reported that the 72 K-hollandite I phase becomes mechanical unstable at 50 GPa (Mookherjee and Steinle-73 Neumann 2009), while another suggested a pressure of about 22 and 23 GPa (Caracas 74 and Boffa Ballaran 2010).

75 While hollandite is the host phase for large elements such as LILEs, potassium 76 could be replaced by sodium (Wang and Takahashi 1999). K_{0.8}Na_{0.2}AlSi₃O₈ hollandite I 77 was reported to undergo a phase transition to hollandite II at 17-18 GPa (Boffa Ballaran 78 et al. 2009). Then, K_{0.8}Na_{0.2}AlSi₃O₈ and KAlSi₃O₈ hollandite were theoretically 79 predicted to undergo phase transitions at 13 and 23 GPa, respectively, suggesting that 80 solid solution effect of sodium to K-hollandite be significant (Caracas and Boffa 81 Ballaran 2010). The high-pressure behaviour of the end member NaAlSi₃O₈ albite 82 composition has also been debatable. While some diamond-anvil cell experiments 83 reported that NaAlSi₃O₈ hollandite (Na-hollandite) could be stable at some pressure (Liu 84 1978; Tutti 2007), some multi-anvil experiments and theoretical studies suggested that

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- Na-hollandite be metastable phase (Yagi et al. 1994; Liu 2006; Akaogi et al. 2010;
- 86 Deng et al. 2010). In this paper, we have studied the high-pressure stability and
- 87 elasticity of the KAlSi₃O₈ hollandite, and solid solution effect of sodium to K-hollandite
- 88 by means of *ab initio* density functional computation methods.

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CRYSTAL STRUCTURE MODELS

91 The KAlSi₃O₈ hollandite I phase has a tetragonal crystal structure with the space 92 group I4/m, which has a large square tunnel formed by double chains of edge-shared 93 (Al,Si)O₆ octahedra, and K resides in the tunnel (Fig. 1a) (Ringwood et al. 1967; 94 Yamada et al. 1984). As the large square tunnels formed by four double chains supply a 95 space for large-ion elements such as K, Ba, Sr, Na and Pb, hollandite-type silicate are 96 considered as a possible repository of large alkali elements, especially potassium, in the 97 lower mantle. On the other hand, the K-hollandite II phase has a monoclinic crystal 98 structure with the space group I2/m, which has a distorted square formed by four double 99 chains (Sueda et al. 2004; Ferroir et al. 2006; Hirao et al. 2008; Mookherjee and Steinle-100 Neumann et al. 2009; Caracas and Boffa Ballaran 2010). There is, however, no 101 experimentally formal description in the crystallographic literature for the K-hollandite 102 II phase, while a theoretical study predicted the atomic positions satisfying the space 103 group *I2/m* monoclinic symmetry (Mookherjee and Steinle-Neumann et al. 2009).

The disorder in (Al,Si)O₆ octahedra in a single unit cell breaks the tetragonal symmetry. While a previous calculation using a single unit cell with 26 atoms breaks the symmetry (Deng et al. 2011), we therefore consider three supercells in order to maintain the tetragonal symmetry (Fig 1.b-d). First, a $1 \times 1 \times 2$ supercell consists of two layer of K₂Al₂Si₆O₁₆ along *c*-axis, where one layer is stacked along *c*-axis on another layer with a 90° rotation around the *c*-axis (Fig. 1b; model A). Next, a $2 \times 2 \times 1$ supercell

110	consists of four cells of K ₂ Al ₂ Si ₆ O ₁₆ in the <i>c</i> -plane, where one cell is adjacent to another
111	cell with a 90° rotation around the <i>c</i> -axis (Fig. 1c; model B). Finally, a $2 \times 2 \times 2$ supercell
112	consists of two layers of $K_8Al_8Si_{24}O_{64}$ model B, where one layer is stacked along <i>c</i> -axis
113	on another layer with a 90° rotation around the <i>c</i> -axis (Fig. 1d; model C). Although
114	conventional unit cells for model A, B, and C include 52, 104, and 208 atoms,
115	respectively, all the primitive unit cells include 52 atoms. Previous studies used a 104
116	atom 1×1×4 supercell in computation (Mookherjee and Steinle-Neumann et al. 2009)
117	but their supercell is exactly identical to our model A with translation symmetry
118	(personal communication with M. Mookherjee), and it is quite unclear why they chose
119	the reducible cell.

120 In order to investigate solid solution effect of sodium, we consider five $1 \times 1 \times 2$

supercells (model A) with Na content of 25 (Fig. 2a-b), 50 (Fig. 2c), 75 (Fig. 2d), 100

122 (Fig. 2e) mol%. The disorder effects are taken into account for Na content of 25 mol%.

123 We consider two configurations of Na. One maintains the tetragonal symmetry (Fig. 2a),

124 while another breaks the symmetry (Fig. 2b).

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FIRST PRINCIPLES COMPUTATION DETAILS

127 Our first principles calculations are based on the density functional theory 128 (Hohenberg and Kohn 1964; Kohn and Sham 1965) within the local density 129 approximation (LDA) (Ceperley and Alder 1980; Perdew and Zunger 1981). One 130 electron wavefunction satisfying the Kohn-Sham equation describes an electronic 131 ground state, which was expanded by using the plane wave basis set. Ionic core 132 potentials were approximated based on the pseudopotential method, which were 133 generated using the Vanderbilt's method (Vandelbilt 1990) for K, Na, and O and 134 Troullier and Martins' methods (Troullier and Martins 1991) for Al and Si. The electric

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135	configurations included in these pseudopotentials are $3s^23p^64s^1$ for K, $2s^22p^63s^1$ for Na,
136	$3s^23p^1$ for Al, $3s^23p^23d^0$ for Si, and $2s^22p^4$ for O. Most of these have been well tested in
137	our previous studies (Tsuchiya et al. 2004, 2005; Kawai and Tsuchiya 2010, 2012a,b).
138	The plane wave energy cutoff was set to 50 Ry, and the irreducible part of Brillouin
139	zone was sampled on $2 \times 2 \times 2$ for the supercells of the K-hollandite I and II phases (52
140	atoms as mentioned above) using the Monkhorst-Pack method (Monkhorst and Pack
141	1976). Effects of using the larger cutoff and k -points on the calculated properties were
142	found insignificant. The full elastic constant tensors of the class $4/m$ tetragonal and $2/m$
143	monoclinic structures have seven and thirteen independent components, respectively.
144	They were calculated using the stress-strain relation with applying strains of ± 0.01
145	(Kawai and Tsuchiya 2010, 2012a,b; Karki et al. 2001; Tsuchiya et al. 2004). We
146	confirmed that the linear relation was ensured enough for this strain range. All structural
147	parameters were fully relaxed to a static (0 K) configuration using the damped variable
148	cell shape constant pressure molecular dynamics technique (Wentzcovitch 1991) using
149	the PWSCF code until residual forces became less than 5.0×10^{-5} Ry/a.u. Computations
150	for the K-hollandite II were performed from the structure which was initially distorted
151	in the <i>c</i> -plane until the structure was fully relaxed.

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Results and discussion

153 In order to investigate the phase stability we calculated enthalpy for each supercell 154 model for both K-hollandite I and II phases (Fig. 3a). We found that model A is the 155 most stable, model C is the second for both K-hollandite I and II phases, although the 156 enthalpy difference between model A and C is little. The supercells were found to have 157 an enthalpy difference of about 30 kJ/mol between model A and B, which is produced 158 clearly due to the difference in the configurations of Si/Al. Si and Al locally generate 159 positive and negative charges, respectively. While Si is fairly randomly distributed in 160 model A, Al is aligned both in the *c*-plane and along the *c*-axis in the cell shown in

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model B. This causes an enthalpy difference between model A and B. It is a little
surprising that model C is about 3 kJ/mol more unstable than model A because model C
is more symmetric than model A. Hence, in this study we consider model A to be the Khollandite phase.

165 Next, we investigated the phase stability between the K-hollandite I and II phases. 166 The enthalpy of K-hollandite II phase deviates from that of K-hollandite I phase at 167 about 15 GPa and becomes lower than that of K-hollandite I phase above this pressure. 168 The enthalpies for the K-hollandite I and II phases asymptotically converge at about 15 169 GPa with decreasing pressure. This kind of behavior is a typical feature for a second-170 order phase transition, which is likely expected to have a positive Clapeyron slope. If 171 extrapolating experimental phase boundaries with a positive slope of $\sim 7MPa/K$ 172 (Nishiyama et al. 2005; Ferrior et al., 2006), we obtain phase transition pressures of 173 16~18 GPa at static temperature. This pressure is consistent well with our results. On 174 the other hand, a previous theoretical study (Mookherjee and Steinle-Neumann 2009) 175 reports that enthalpies for the K-hollandite I and II phases cross over one another at 176 about 33 GPa, indicating a typical feature for a first-order phase transition, even though 177 the asymptotic feature can be seen in their enthalpy of the K-hollandite II phase around 178 20 GPa. They seem to have failed in computation of the reference structure of the K-179 hollandite I phase. Meanwhile, a previous theoretical study determined the transition 180 pressure to be 23 GPa without any consideration of enthalpy difference between the K-181 hollandite I and II phases (Deng et al. 2011).

Pressure-volume relations calculated for the K-hollandite I and II phases are shown in Fig. 4. Their equations of state parameters (zero-pressure bulk modulus, B_0 , and its pressure derivative, B_0') determined by least-squares fitting to the third-order Birch-Murnaghan equation (Table 1) are $B_0 = 205.07$ GPa and $B_0' = 4.18$ for the Khollandite I phase and $B_0 = 192.26$ GPa and $B_0' = 4.16$ for the K-hollandite II phase. We

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187 can obtain good agreement with experimental data for the K-hollandite I phase (Ferroir 188 et al. 2006; Nishiyama et al. 2005; Zhang et al. 1993), although our computed volume 189 for the K-hollandite I phase is about 1.4 % smaller than experimental results due to no 190 thermal effect for a static condition (Fig. 4 and Table 1). The K-hollandite II is more 191 compressible than the K-hollandite I phase. Zero-pressure volume obtained by the 192 recent calculations within the generalized gradient approximation (GGA) (Deng et al. 193 2011) is larger than that computed in this study within LDA as usually seen. While the 194 zero-pressure volumes obtained in this study for both K-hollandite I and II phases match 195 well with each other, previous theoretical study shows difference in zero-pressure 196 volumes between K-hollandite I and II phases (Mookherjee and Steinle-Neumann 2009), 197 indicating that their computation is shy of convergence.

198 Computed lattice parameters for the K-hollandite I and II phases are shown in Fig. 199 5 as a function of pressure. These are in excellent agreement with the experimental data 200 for both the K-hollnadite I (Ferroir et al. 2006; Zhang et al. 1993) and II (Sueda et al. 201 2004; Ferroir et al. 2006; Hirao et al. 2008) phases. Although the bulk volume 202 compressibility does not differ significantly between the K-hollanidte I and II phases 203 (Fig. 4), the splitting of cell parameters (a and b) and the change of the angle γ start 204 rapidly at 20 GPa. The mean value (a+b)/2 and c do not change rapidly after the phase 205 transition as reported by experimental studies (Ferroir et al. 2006; Hirao et al. 2008). 206 Second-order phase transitions are critical phenomena. When we take a cell angle related to the monoclinic distortion, γ , as an order parameter, it can be described based 207 on the Landau theory (Landau and Lifshitz 1994) as $\gamma \bullet [(P-P_c)/P_c]^q$, where P_c and q208 209 are the critical pressure and the critical exponent related to the long-range order in the 210 system, respectively. Calculated pressure variation in γ was well fitted by this formula in the pressure range from 15 and 40 GPa and to provide $P_c = 14.9$ GPa and q = 0.463211 212 (inset of Fig. 5c). This critical exponent is found close to the ideal value for classical 213 second-order transitions of 1/2.

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214	Elastic constants of the K-hollandite I and II phases calculated are shown in Fig. 6
215	as a function of pressure from 0 to 150 GPa. Elastic constants of K-hollandite I and II
216	are almost the same up to 10 GPa. Note that the relation $C_{16} = -C_{26}$ is satisfied in the
217	class $4/m$ tetragonal structure (Nye 1985). This relation between the K-hollandite I and
218	II phases is seen for the monoclinic K-hollandite II below 10 GPa, indicating that the
219	monoclinic K-hollandite II spontaneously becomes tetragonal with decrease in pressure.
220	In contrast, the softening in C_{11} and C_{66} of the tetragonal K-hollandite I phase occurs
221	with increasing pressure over 40 GPa. This is consistent with previous theoretical
222	studies (Mookherjee and Steinle-Neumann 2009; Caracas and Boffa Ballaran 2010).
223	For $1 \le k \le n$, the <i>k</i> -th principal submatrix of matrix A is the $k \times k$ submatrix

formed from the first k rows and first k columns of matrix **A**. Its determinant is called as

the (leading) k-th principal minor (Gilbert 1991). From Sylvester's criterion, a real,

symmetric matrix is positive definite if and only if all its principle minors are positive.

227 Since the energy density must be positive definite quadratic form (Born and Huang

228 1954), all the six (leading) principal minors of elastic tensor C should be positive. For

the class 4/m tetragonal structure, principal minors are as follows:

$$C(6) = C_{66},$$

$$C(5,6) = C_{44}C_{66},$$

$$C(4,5,6) = C_{44}^2C_{66},$$

$$C(3,4,5,6) = C_{33}C_{44}^2C_{66},$$

$$C(2,3,4,5) = (C_{11}C_{33}C_{66} - C_{13}^2C_{66} - C_{16}^2C_{33})C_{44}^2$$

$$C(1,2,3,4,5,6) = [(C_{11}^2 - C_{12}^2)C_{33}C_{66} - 2(C_{11} - C_{12})C_{13}^2C_{66} + 4C_{13}^2C_{16}^2 - 2\{(C_{11} + C_{12})C_{33} - 2C_{13}^2\}C_{16}^2]C_{44}^2$$
(1)

- We computed all the values and found C(1,2,3,4,5,6) the mechanical stability condition
 for the tetragonal K-hollandite I phase. Hence, we computed a principal minor,
 C(1,2,3,4,5,6), in the pressure from 0 to 60 GPa and found that the mechanical
- instability occur around 16.9 GPa (Fig. 7a).

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The class 4/m tetragonal symmetry with seven independent elastic constants is also known as 4/m (TII) Laue symmetry (Farley et al. 1975; Blanchfield et al. 1982). When the components of the elastic tensor of a 4/m (TII) Laue symmetry crystal are transformed by rotation around *c*-axis through an angle

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$$\phi = \frac{1}{4} \tan^{-1} \left(\frac{4C_{16}}{C_{11} - C_{12} - 2C_{66}} \right),$$
 (2)

then the transformed elastic constant C_{16} becomes zero, the number of independent elastic constants is reduced from 7 to 6 and the elastic constant tensor C_{ij} makes the form of that of a *4/mmm* (TI) Laue symmetry crystal. For elasticity tensor to be positive definite, the requirements obtained in the same manners as noted above are as follows: $C_{11}-C_{12} > 0, (C_{11}+C_{12})C_{33}-2C_{13}^2 > 0, C_{44} > 0, and C_{66} > 0$ (3)

Since C_{11} - C_{12} > corresponds to the stability criterion for tetragonal stability for Khollandite I phase, we found that the mechanical instability occurs 16.9 GPa (Fig. 7b), which is equal to the above computed results. This is of course quite reasonable because this rotational transformation operation is just algebraic and should not affect the elasticity of material itself.

250 Previous theoretical study (Caracas and Boffa Ballaran 2010) shows that $C_{66} < 0$ 251 instead of C_{11} - $C_{12} < 0$ but we did not confirm $C_{66} < 0$ even using the values reported in 252 their paper. Instead, we found using their data tetragonal instability of C_{11} - C_{12} < 0 at 253 about 31 GPa for a 4/m (TII) Laue symmetry crystal with six independent elastic 254 constants and also found C(1,2,3,4,5,6) negative at almost the same pressure as 31 GPa 255 even for a 4/mmm (TI) Laue symmetry crystal with seven independent elastic constants, 256 although their data are too sparse in pressure to determine the critical pressure 257 accurately. The difference in critical pressures may depend on their adopted method. 258 They use the virtual crystal approximation, as it is called as *alchemical* approximation,

for the Al/Si disordered sites. This approximation allows us to reduce the cell size and, thereby, computational cost. On the other hand, since *alchemical* pseudopotential for the disordered sites seems to make crystal more symmetric, the K-hollandite I phase becomes more stable and the critical pressure increases. We suspect that *alchemical* pseudopotential causes this discrepancy in the transition pressure (de Gironcoli and Baroni 1992).

Taking the Voigt-Reuss-Hill averages of the calculated elastic constants, we have obtained the isotropic bulk (*B*) and shear moduli (*G*), and isotropic compressional (V_P), shear (V_S), and bulk sound velocities (V_{ϕ}) of the both K-hollandite I and II phases as a function of pressure (Fig. 8). While no significant anomaly is seen in the bulk modulus, a softening appears in the shear modulus associated with the K-hollandite I-II transition (Fig. 8b). The present calculations, therefore, show that the phase transition of Khollandite I phase at 15.7 GPa produces the shear elastic instability (Fig. 8b).

272 Next, the single crystal elastic wave velocities of both the K-hollandite I and II
273 phases are calculated. Taking a Fourier transformation over space and time of the elastic
274 equation of motion, we obtain Christoffel's equation:

$$275 \qquad \rho c^2 u_i = n_j n_l C_{ijkl} u_k \tag{4}$$

where ρ , *c*, and **u** are density, phase velocity and displacement vector respectively and $(n_x, n_y, n_z) = (\sin \theta \cos \phi, \sin \theta \sin \phi, \sin \theta)$ is the direction cosine normal to wave surface. For eq. 4 to be satisfied, it is necessary that

$$279 \quad \left|\rho c^2 \delta_{ik} - n_j n_l C_{ijkl}\right| = 0, \tag{5}$$

280 where δ_{ij} is the Kronecher's delta (Fryer and Frazer 1987). Solving this characteristic 281 equation, we can obtain seismic velocities for three quasi-waves such as quasi-P, quasi-

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282 S1 and quasi-S2 waves because polarizations may not be along the dynamic axis. The 283 azimuthal dependence of three phase velocities for the K-hollandite I phase at 0 and 20 284 GPa and the K-hollandite II phase at 20 and 100 GPa is shown in Fig. 9. While we do 285 not see changes of the anisotropic style, i.e., the fast and slow directions, for K-286 hollandite I phase between 0 and 20 GPa, quasi-S2 wave becomes very slow along 287 [110] suggesting mechanical instability. φ -dependence of the K-hollandite I and II 288 phase has periods of 180° and 90° due to the difference between tetragonal and 289 monoclinic systems. The fast and slow directions, in particular for the S waves, change 290 very much with pressure for the K-hollandite II phase, although the magnitude of 291 anisotropy decreases with increasing pressure. Next we define the azimuthal anisotropy 292 for the P and S waves by following relations as follows:

293
$$A_{P} = \frac{V_{P_{\text{max}}} - V_{P_{\text{min}}}}{V_{P}} \times 100; \quad A_{S} = \frac{V_{S_{\text{max}}} - V_{S_{\text{min}}}}{V_{S}} \times 100$$
(6)

where V_P and V_S are the isotropic aggregate velocities (Fig. 7). The K-hollandite becomes strongly anisotropic around 20 GPa (Fig. 10). In particular, the S wave

anisotropy diverges. This clearly reflects the ferroelastic nature of the phase transition.

297 Finally, the solid solution effect of sodium to K-hollandite is considered. In 298 order to investigate the phase stability, we compute enthalpy for each symmetric 299 supercell for Na content of 25 (Fig. 2a), 50 (Fig. 2c), 75 (Fig. 2d), 100 (Fig. 2e) mol% 300 for both K-hollandite I and II phases (Fig. 3d). Significant solid solution effects to phase 301 stability cannot be found via enthalpy comparison. We also compute the mechanical 302 stability conditions for each supercell and find that the instability occurs in the pressure 303 range from 16 to 18 GPa for all the Na content (Fig. 7). This also suggests that solid 304 solution effect is insignificant. This pressure is very consistent with previous 305 experimental results which reported that K_{0.8}Na_{0.2}AlSi₃O₈ hollandite undergo phase 306 transition to hollandite II at 17-18 GPa [14]. On the other hand, previous theoretical

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307	studies predicted that $K_{0.8}Na_{0.2}AlSi_3O_8$ and $KAlSi_3O_8$ hollandites undergo phase
308	transitions at 13 and 23 GPa, respectively, and the solid solution effect of sodium to K-
309	hollandite is significant (Caracas and Boffa Ballaran 2010). As noted above, we suspect
310	that alchemical pseudopotential causes this discrepancy in the transition pressure (de
311	Gironcoli and Baroni 1992). We also compute enthalpy differences between
312	$K_{0.75}Na_{0.25}AlSi_3O_8$ hollandite I and $KAlSi_3O_8$ hollandite I and between
313	$K_{0.75}Na_{0.25}AlSi_3O_8$ hollandite II and KAlSi_3O_8 hollandite II. Since these values are
314	coincident within 0.001 kJ/mol both before and after the phase transition. This suggests
315	that the solid solution effects of sodium to both K-hollandite I and II are almost the
316	same in contrast to previous experimental results (Wang and Takahashi 1999). This may
317	be caused by other effects such as the solid solution effects of sodium to other phases in
318	their sample.

319 Then, the disorder effect in (K,Na) in the tunnel (Fig. 2a-b) is studied. Although 320 the supercell without the tetragonal symmetry (Fig. 2b) becomes spontaneously strained 321 with increasing pressure, this pressure of about 17 GPa is rarely different from the 322 pressure at which the supercell with the tetragonal symmetry (Fig. 2a) undergoes phase 323 transition to monoclinic hollandite II (Fig. 3b). The supercell without the symmetry (Fig. 324 2b) is about 2.5 kJ/mol less stable than with the symmetry (Fig. 2a). Since this enthalpy 325 difference is very small, both configurations with and without tetragonal configurations 326 will be randomly distributed under high-*P*,*T* condition in the mantle. In order to 327 understand random configuration effects to the phase transition between K-hollandite I 328 and II, we investigate the pressure dependence of a cell-parameter, γ , for configurations 329 both with and without tetragonal symmetry (Fig. 5d). Although a configuration without 330 tetragonal symmetry (green line) has a distorted structure in all the pressure range, 331 distortion is very small up to 15 GPa and γ without tetragonal symmetry increases in the 332 same manner as that with tetragonal symmetry (blue line). This suggests that even for a 333 configuration without tetragonal symmetry phase transition occurs at the same pressure

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334	as with tetragonal symmetry. Hence, random distribution of structure with and without
335	tetragonal symmetry would not influence phase transition pressure. Also, although it
336	may be monoclinic locally, due to random distributions K-hollandite should be
337	tetragonal up to the phase transition pressure.
338	Next, phase stability in NaAlSi $_{3}O_{8}$ system is considered. We calculate enthalpy
339	for an assemblage of jadeite (NaAlSi ₂ O ₆) and stishovite (SiO ₂), an assemblage of a CF-
340	type phase (NaAlSiO ₄) and stishovite ($2SiO_2$), and NaAlSi ₃ O ₈ hollandite I and II, and

- have found that an assemblage of jadeite and stishovite is stable up to 18 GPa and an
- 342 assemblage of a CF-type phase and stishovite is stable over 18 GPa. This is consistent
- 343 with previous theoretical studies (Deng et al. 2010; Kawai and Tsuchiya 2010, 2012c),
- although the transition pressure of 18 GPa (Kawai and Tsuchiya 2010, 2012c) is
- different from that of 33.6 GPa (Deng et al. 2010). The former matches well with
- 346 experimental results (Akaogi et al. 2010). Our calculation suggests that the Na-
- hollandite be metastable phase as previous studies noted.

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GEOPHYSICAL IMPLICATIONS

350 Sediment and Archaean granites become less dense near the Earth's surface but 351 contain about 20 volumetric% of the K-hollandite I phase above 9 GPa in the upper mantle, respectively (Irifune et al. 1994; Wu et al. 2009). Taking a temperature of 1300 352 353 (cold geotherm) 1800 K (normal geotherm) into account, the mechanical instability 354 (Fig. 7b) of the K-hollandite I occurs at 24.8 and 28.3 GPa, with a Clapeyron slope of 7 355 MPa/K (Nishiyama et al. 2005), corresponding to the depth of 692 and 771 km, 356 respectively. In depth range between the depth at which mechanical stability starts and 357 the depth at which phase transition is completed, the velocity changes due to both 358 softening of shear modulus of the K-hollandite I phase and phase transition between the

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- K-hollandite I and II phase can be observed. This is consistent with evidence on the
 seismic scatterers observed in the shallowest lower mantle (Kaneshima 2009),
 suggesting that buoyant granitic material can be subducted into the deep mantle and be
 carried to another gravitationally stable depth range other than the Earth's surface
 (Kawai et al. 2009).
- 364 Since the K-hollandite is stable entirely in the lower mantle, they can be host 365 phase for large elements in the Earth's deep interior. The K-hollandite controls a 366 significant proportion of the whole-rock budget of incompatible elements. Recycling of 367 continental sediments including the K-hollandite subducted into the mantle transition 368 zone can explain the geochemical trace element abundance in EM-I type ocean-island 369 basalts (Rapp et al. 2008). Also, geochemical studies on potassium-rich lamproites from 370 Gaussberg, Antarctica (Murphy et al. 2002) suggested that the source of lamproites 371 should be effectively isolated for 2-3 Gyr in the mantle. Since the granitic materials 372 including the K-hollandite are gravitationally stable in the mantle transition zone 373 (Kawai et al. 2009), the mantle transition zone could be reservoir of incompatible 374 LILEs, solving the terrestrial Pb-isotope paradox (Murphy et al. 2002; Kawai et al. 375 2010).

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CONCLUSIONS

Using theoretical methods, first we examined the phase stabilities of the Khollandite I and II phases and confirmed that the K-hollandite I phase becomes mechanically unstable at 15.7 GPa and undergoes the ferroelastic (second order) phase transition to the K-hollandite II phase at about 15.7 GPa and that the K-hollandite II is stable entirely in the lower mantle and can be the host phase for geochemically important large elements. Next, we investigated elasticity of both the K-hollandite

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- 384 phases and found softening of shear modulus and velocity changes due to phase
- transition, consistent with seismic observation. Finally, we studied the solid solution
- 386 effect of sodium to K-hollandite and found that this effect to phase stability is very
- 387 small in contrast to previous studies and that Na-hollandite is metastable phase.

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- 543 Fig. 1 (a) Crystal structures of the KAlSi $_3O_8$ hollandite phases. Deep purple,
- 544 light blue, dark blue and red spheres represent K, Al, Si, and O atoms,
- respectively. (b) 1×1×2 tetragonal supercell (model A), (c) 2×2×1 tetragonal
- 546 supercell (model B), and (d) 2×2×2 tetragonal supercell (model C). (See
- 547 supplemental data for details of atomic positions for each model.)
- 548 Fig. 2 (a) Crystal structures of the (K,Na)AlSi₃O₈ hollandite phases. Deep purple,
- 549 yellow, light blue, dark blue and red spheres represent K, Na, Al, Si, and O
- atoms, respectively. 1×1×2 supercell with Na content of 25 mol% with (e; conf
- A) and without (f; conf B) the tetragonal symmetry. 1×1×2 supercell with Na
- 552 content of 50 (g), 75 (h), and 100 (i) mol% with tetragonal symmetry. (See
- supplemental data for details of atomic positions for each model.)
- 554 Fig. 3. (a) The calculated enthalpy difference of the K-hollandite I and II phases 555 for three supercell models relative to model A of the K-hollandite I phase (see 556 Fig. 1). The phase transition pressure from K-hollandite I to II is determined to 557 be 15 GPa by fitting the cell angle y (see text for details). (b) The calculated 558 enthalpy differences of the hollandite with Na content of 25 mol% between conf 559 B and tetragonal conf A (red) and between conf B and monoclinic conf A (blue). 560 (c) The enthalpy differences among an assemblage of jadeite (NaAlSi₂ O_6) and 561 stishovite (SiO₂), an assemblage of a CF-type phase (NaAlSiO₄) and stishovite 562 (2SiO₂), and NaAlSi₃O₈ hollandite I and II. The enthalpy of hollandite II is taken 563 as a reference. (d) The enthalpy differences of the K-hollandite I and II phases 564 with Na content of 0, 25, 50, 75, and 100 mol% relative to K-hollandite I phase 565 with each Na content.
- 566 Fig. 4. Volumes calculated for model A of the K-hollandite I (circles) and II
- 567 (triangles) phases. Fitting curves are computed using a third-order Birch-
- 568 Murnaghan equation of state. We also show experimental results for the K-

- 569 hollandite I (circles) (Ferroir et al. 2006; Zhang et al. 1993) and II (triangles)
- 570 (Sueda et al. 2004; Ferroir et al. 2006; Hirao et al. 2008) in addition to
- 571 theoretical results by Mookherjee and Steinle-Neumann (2009).
- 572 Fig. 5. Lattice parameters (a and b in a; c in b; γ in c) calculated for model A of
- 573 the K-hollandite I (circles) and II phases (triangles). We also show experimental
- 574 results for the K-hollandite I (open circles) (Ferroir et al. 2006; Zhang et al.
- 575 1993) and II (open triangles) (Sueda et al. 2004; Ferroir et al. 2006; Hirao et al.
- 576 2008) phases. (inset of c) calculated cell angle, γ , (red triangle) with a fit (green
- 577 line) (see text for details). (d) Calculated lattice parameters, γ, for K-hollandite I
- 578 (red line) and II phases (blue line) with tetragonal symmetry and K-hollandite
- 579 without tetragonal symmetry (green line) with Na content of 25 mol%.
- 580 Fig. 6. Elastic constants calculated as a function of pressure for model A of the
- 581 K-hollandite I (0-60 GPa) and II (0-150 GPa) phases. (a)-(c) show longitudinal,
- 582 off-diagonal, and shear components for the tetragonal K-hollandite I and
- 583 monoclinic K-hollandite II phases.
- 584 Fig. 7. Mechanical stability criterion computed for the K-hollandite I phase with
- 585 Na content of 0, 25, 50, 75, and 100 mol%. (a) C(1,2,3,4,5,6) for the class 4/m
- tetragonal structure. (b) C_{11} - C_{12} for the class 4/mmm tetragonal structure.
- Fig. 8. (a) Aggregate bulk and shear moduli of the K-hollandite I (red) and II (blue) phases. (b) Longitudinal, bulk and shear wave velocities and densities of the K-hollandite I and II phases. The vertical bold lines represent the calculated transition pressure of hollandites.
- 591 Fig. 9. Velocities of quasi-waves ($qV_{P,} qV_{S1}$, and qV_{S2}), where qS1 is the faster 592 of the two split shear waves, as a function of the propagation direction for the K-

- 593 hollandite I phase at 0 and 20 GPa and the K-hollandite II phase at 20 and 100
- 594 GPa.
- 595 Fig. 10. Pressure dependence of the elastic anisotropy of the K-hollandite I and
- 596 II phases.
- 597 **Table 1** Equation of state parameters for the K-hollandite I and II phases
- 598 calculated at the static temperature. Results of previous theoretical and
- 599 experimental studies are also shown.

	V ₀ (cm ³ /mol)	B ₀ (GPa)	B' ₀
Hollandite-I (calc:this study)	140.70	205.07	4.18
Hollandite-I (calc) ^a	137.09	225.0	4.3
Hollandite-I (calc) ^b	147.43	174.0	3.99
Hollandite-I (exp) °	142.3	180	4.0
d d	143.1	183	4.0
Hollandite-I (exp) ^e	142.7	201	4.0
Hollandite-II (calc:this study)	140.7	192	4.16
Hollandite-II (calc) ^a	137.3	221	3.9
Hollandite-II (calc) ^b	147.4	168	4.0

- 600 ^a Mookherjee and Steinle-Neumann (2009)
- 601 ^b Deng et al. (2011)
- 602 ^c Zhang et al. (1993)
- 603 ^d Nishiyama et al. (2005)
- 604 ^e Ferroir et al. (2006)

- 606 **Table 2** Elastic constants (c_{ij}), aggregate bulk (B) and shear (G) moduli of the
- 607 K-hollandite I and II phases calculated at 0 GPa. We also show results of
- 608 previous theoretical studies.

	Hollandite-I	Hollandite-I	Hollandite-I	Hollandite-II	Hollandite-
	(This study)	a	b	(This study)	ll ^a
C ₁₁	290	342	303	290	382
(GPa)					
C ₂₂	-	-	-	290	395
C ₃₃	530	568	517	530	575
C ₄₄	154	165	175	154	177
C ₅₅	-	-	-	154	182
C ₆₆	112	129	139	112	48
C ₁₂	181	186	202	181	132
C ₁₃	103	118	123	103	140
C ₂₃	-	-		103	107
C ₁₆	-42	-	44	-42	-16
C ₂₆	-	-		42	4
C ₃₆	-	-		-0.1	38
C ₄₅	-	-		0.0	13

В	207	230	224	204	220
G	111	124	143 (117)*	110	134
VP	9.46	9.87	10.19	9.45	9.92
(km/s)			(9.74)*		
Vs	5.28	5.53	5.98 (5.41)*	5.28	5.75
V_{ϕ}	7.23	7.53	7.49	7.23	7.37

^a Mookherjee and Steinle-Neumann (2009)
 ^b Caracas and Boffa Ballaran (2010)
 * recomputed using the data by Caracas and Boffa Ballaran (2010)















C(1,2,3,4,5,6) (GPa⁶)





