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3	Calibration of zircon as a Raman spectroscopic pressure sensor to high temperatures
4	and application to water-silicate melt systems
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14	Abstract - The shifts in wavenumber of the $v_3(SiO_4)$ (~1008 cm ⁻¹) Raman band of fully
15	crystalline synthetic zircon with changing pressure (P) and temperature (T) were calibrated
16	for application as a Raman spectroscopic pressure sensor in optical cells to about 1000 °C and
17	10 GPa. The relationship between wavenumber (v) of this band and T from 22 to 950 °C is
18	described by the equation $v[\text{cm}^{-1}] = 7.54 \cdot 10^{-9} \cdot T^3 - 1.61 \cdot 10^{-5} \cdot T^2 - 2.89 \cdot 10^{-2} \cdot T + 1008.9$, where T
19	is given in [°C]. The pressure dependence is nearly linear over the studied range in P . At ~25
20	°C, the $\partial v/\partial P$ slope to 6.6 GPa is 5.69 cm ⁻¹ /GPa, and that to 2 GPa is 5.77 cm ⁻¹ /GPa. The
21	$\partial v/\partial P$ slope does not significantly change with temperature, as determined from experiments
22	conducted along isotherms up to 700 °C. Therefore, this pressure sensor has the advantage
23	that a constant $\partial v / \partial P$ slope of 5.8±0.1 cm ⁻¹ /GPa can be applied in experiments to pressures of
24	at least about 6.6 GPa without introducing a significant error. The pressure sensor was tested
25	to determine isochores in experiments with $H_2O+Na_2Si_3O_7$ and $H_2O+NaAlSi_3O_8$ fluids to 803
26	°C and 1.65 GPa. These pressures were compared to pressures calculated from the EoS of

27 H₂O based on the measured vapor dissolution or ice melting temperature for the same 28 experiment. Pressures determined from the zircon sensor in runs in which NaAlSi₃O₈ melt 29 dissolved in aqueous fluid were close to or lower than the pressure calculated from the EoS of H₂O using the vapor dissolution or ice melting temperature. In experiments with 30 31 H₂O+Na₂O+SiO₂ fluids, however, the pressure obtained from the Raman spectrum of zircon 32 was often significantly higher than that estimated from the EoS of H_2O . This suggests that the 33 pressures along some critical curves of water-silicate melt pseudobinary systems should be 34 revised.

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36 Keywords: zircon, Raman spectroscopy, temperature, pressure sensor, diamond anvil cell

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38 **1. Introduction**

39 Experiments using diamond-anvil cells have been pivotal in many studies of minerals and 40 fluids at high pressure (P) and temperature (T) because these cells permit not only optical 41 observation of the sample but also *in situ* measurement of physical and chemical properties by 42 "photon-in – photon-out" techniques. The pressure in the sample chamber of such cells must 43 be determined indirectly, which is commonly done by measurement of the liquid-vapor 44 homogenization temperature of the fluid pressure medium with application of an appropriate 45 equation of state or using the shift in wavenumber of a Raman or fluorescence line of a 46 calibrant inside the sample chamber. However, fluorescence intensities decrease rapidly with increasing temperature, such that Raman lines are the often better option for pressure 47 determination at elevated temperatures, particularly above 400 °C. The ${}^{5}D_{0}-{}^{7}F_{0}$ fluorescence 48 line of SrB₄O₇:Sm²⁺ is still detectable at 627 °C, but this pressure sensor is quite soluble in 49 aqueous fluids (Datchi et al. 2007). The 464 cm⁻¹ Raman line of quartz is often applied to 50 determine pressure because of its large shift with P ($\partial v/\partial P \sim 9$ cm⁻¹/GPa) and relatively small 51 shift with T ($\partial v/\partial T \sim -0.014$ cm⁻¹/deg) (Schmidt and Ziemann 2000). However, its use is 52

53 limited by phase transitions and high solubility of quartz in many fluids and melts. The 54 wavenumber difference between two Raman lines of berlinite, AlPO₄, is more sensitive to pressure than that of the 464 cm⁻¹ Raman line of quartz, and its determination does not require 55 56 a high-resolution Raman spectrometer, but berlinite reacts readily with aqueous fluids 57 (Watenphul and Schmidt 2012). In situations in which quartz or berlinite are not suitable for 58 use as spectroscopic pressure sensors, zircon is one alternative because it is stable over a large 59 range in P, T, and fluid composition (e.g., Watson and Harrison 1983). Other inert Raman spectroscopic calibrants such as ¹³C diamond or cubic boron nitride (*c*-BN) display only small 60 wavenumber shifts with pressure of ~ 2.83 cm⁻¹/GPa (first-order Raman line, Schiferl et al. 61 1997; Datchi et al. 2007), or ~3.27 cm⁻¹/GPa (TO mode, Datchi and Canny 2004), and also 62 fairly large and nonlinear shifts in the wavenumber with temperature. This makes the 63 64 application of ¹³C diamond or *c*-BN more difficult at relatively low pressures because of limitations in the attainable accuracy and precision in determining wavenumbers and actual 65 sample temperatures. As described in the following sections, zircon offers the advantage over 66

67 these phases that the shift in the wavenumber of its most intense Raman line with pressure is 68 about twice as large as those of *c*-BN and 13 C-diamond.

The four most intense Raman bands of zircon (Fig. 1) at wavenumbers $>300 \text{ cm}^{-1}$ have been 69 70 assigned to internal vibrations of the SiO₄ tetrahedra. Specific assignments include $v_3(SiO_4)$ (antisymmetric stretching mode) to the band at 1008 cm⁻¹, $v_1(SiO_4)$ (symmetric stretching 71 mode) to the band at 975 cm⁻¹, $v_2(SiO_4)$ (symmetric bending mode) to the band at 439 cm⁻¹, 72 73 and $v_4(SiO_4)$ (antisymmetric bending mode) or an external lattice mode to the band at 356 cm⁻¹ ¹ (Dawson et al. 1971; Syme et al. 1977; Nasdala et al. 2003). Although there is a large 74 75 number of Raman spectroscopic studies on zircon, very few data are available on its Raman 76 band parameters as functions of pressure and temperature. The effect of pressure on the 77 positions of these bands and on the linewidth of $v_3(SiO_4)$ has previously been characterized at 78 ambient temperature for fully crystalline and radiation-damaged zircon (Knittle and Williams 79 1993, Nasdala et al. 2008), and that of temperature on positions and linewidths at -183 °C and 80 22 °C by Syme et al. (1977). To our knowledge, such data are lacking for elevated 81 temperatures. In addition, there is significant disagreement on the $\partial v/\partial P$ slope of the v₃(SiO₄) 82 band, i.e., 5.9 cm⁻¹/GPa (Nasdala et al. 2008, synthetic zircon) vs. 4.8±0.2 cm⁻¹/GPa (Knittle 83 and Williams 1993, natural zircon with little radiation damage). 84 In this study, we calibrated zircon as Raman spectroscopic pressure sensor for experiments 85 using diamond-anvil or other optical cells, particularly for experiments at high temperatures 86 and with fluid compositions in which the spectrosopic pressure sensors quartz and berlinite 87 are readily dissolved and can no longer be applied. For this application, we focused on the 88 shift in the wavenumber of the $v_3(SiO_4)$ band with pressure and temperature, because it has 89 the largest $\partial v/\partial P$ slope of Raman bands of zircon (Knittle and Williams 1993). In addition, 90 this band is very intense in most crystallographic orientations of the zircon relative to the direction of polarization of the laser (its intensity is only zero at $E \parallel c$ and the incident beam \perp 91 92 {110}, Dawson et al. 1971). Additional objectives were to test if fully crystalline zircon 93 persists in H₂O+Na₂O+SiO₂ \pm Al₂O₃ fluids at high P and T, and to compare pressures 94 calculated from EoS of H₂O to pressures determined from Raman spectra of zircon in 95 experiments with H_2O + silicate melts.

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97 2. Methods

The zircon sample was synthesized in an 1 atm furnace from 500 mg of oxide mixture (10.8 mass% ZrO₂, 26.8 mass% SiO₂, and 62.4 mass% PbO as flux) and seeds of natural crystalline Hf-bearing zircon in a Pt crucible using a temperature ramp from 1430 to 1350 °C at 1 deg/h with subsequent quench, followed by treatment with 48% hydrofluoric acid to remove excess quartz. Electron microprobe analyses gave the following mean composition (mass%) SiO₂ 32.73 \pm 0.36, HfO₂ 1.48 \pm 0.31, ZrO₂ 66.40 \pm 0.28 (46 analyses). This corresponds to a composition of Zr_{0.987}Hf_{0.013}[SiO₄]. Lead was present in low concentrations in some analyses, but was mostly below the detection limit. The backscattered electron images of the sampleverified its compositional homogeneity.

107 All spectra recorded for calibration of the zircon Raman spectrum were acquired using a 108 HORIBA Jobin Yvon LabRAM HR800 UV-Vis Raman spectrometer (gratings 1800 109 lines/mm, focal length 800 mm, confocal pinhole aperture set at 100 µm to reduce the background from the diamond anvils, excitation by 488 nm line of an Ar⁺ laser), equipped 110 111 with a CCD-detector (1024×256 pixel) and an Olympus microscope. The unpolarized 112 spectra were obtained in backscattering geometry using $20 \times$ objectives with long working 113 distances of 25 mm or 20.5 mm to avoid damage to the objectives from heat transfer during 114 experiments at high temperature. For each spectrum, the positions of the studied Raman bands of zircon and of the pressure calibrant quartz (464 cm⁻¹ quartz Raman line) were calibrated 115 116 based on the plasma lines from the Ar^+ laser at 1057.25 cm⁻¹ (514.5319 nm) and 737.49 cm⁻¹ (506.2036 nm), or 380.40 cm⁻¹ (497.2157 nm), 351.70 cm-1 (496.5073 nm), and 221.60 cm⁻¹ 117 118 (493.3206 nm). This was achieved by removal of the interference filter to record the plasma 119 lines and the Raman bands in the same spectrum, without a change in the spectrometer 120 settings. The use of the 488 nm line for excitation had the advantage that the plasma lines at 1057.25 cm⁻¹ (514.5319 nm) and 737.49 cm⁻¹ (506.2036 nm) and the $v_1(SiO_4)$ and $v_3(SiO_4)$ 121 122 Raman bands of zircon could be recorded in the same spectral window. The spectral resolution was about 0.9 ± 0.1 cm⁻¹ based on the line width of the recorded plasma lines. All 123 124 spectra from the calibration runs were processed using the software package PeakFit v4.11 125 from SYSTAT Software Inc. The Raman lines of quartz were fitted with the asymmetric 126 model Pearson IV. The Raman lines of zircon and the plasma lines from the laser were fitted 127 with the symmetric Gauss+Lorentz area function using a linear baseline correction in the 128 selected section. The range in the wavenumber of the sections used to fit the $v_1(SiO_4)$ and 129 v_3 (SiO₄) Raman bands of zircon is given in Tab. 1. This resulted in good fits with 130 insignificant or small nonstatistical residuals. The accuracy in the determined peak positions

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is about ± 0.1 to ± 0.2 cm⁻¹. No temperature reduction for the Bose-Einstein occupation factor was performed before fitting the spectrum profile. Note that application of other functions to fit the Raman bands of zircon gives normally very similar results for the wavenumber, but this is not always the case for the line width. In addition, deviations in wavenumber due to different applied functions for band fitting have no significant effect on determined pressure if the shift in the wavenumber relative to that at a reference *P* and *T* is measured.

137 Raman spectra of zircon as a function of temperature at atmospheric pressure were recorded 138 using a Linkam TS1000 heating stage. The temperature was calibrated based on optical 139 observation of melting of a small halite crystal (800.7 °C at 0.1 MPa) near the center of a Pt 140 crucible, which itself was close to the center of heated ceramic crucible of the stage. 141 Subsequently, a small zircon chip was then placed at the same location at which halite melting 142 had been observed. This was done to minimize potential effects of temperature gradients. 143 Raman spectra were then obtained at various temperatures between 22 to 950 °C with a 144 NIKON M Plan SLWD 20 \times objective (numerical aperture 0.35). The output laser power set 145 via the laser controller was 200 mW, which corresponded to a measured laser power of 34.0 146 mW "downstream" of the objective. These measurements were verified by additional 147 measurements at temperatures between 25 and 741 °C using an externally-heated Bassett-type 148 hydrothermal diamond-anvil cell (HDAC) (Bassett et al. 1993), the design of which had been 149 modified to improve the accuracy of temperature measurements. This cell was equipped with 150 type Ia ultra-low fluorescence and ultra-low birefringence grade diamond anvils (culet 151 diameter 0.9 mm). The temperature in the sample chamber of the cell was measured using K-152 type (Ni/NiCr) thermocouples attached to the diamond anvils. These thermocouples were 153 calibrated using the α - β transition temperature of quartz (574 °C, 0.1 MPa) and the triple 154 point of H_2O (0.01 °C, 0.6 kPa). The power input to the resistive heaters was controlled using 155 Eurotherm[®] 2408 temperature controllers, which held the set temperature within \pm 0.2 °C. 156 Accuracy and reproducibility of the temperature measurements were about ± 0.1 to ± 0.2 °C

157 at ~20 °C and usually better than \pm 1 °C at temperatures between 100 and 600 °C (Schmidt 158 2009). Other conditions for recording the Raman spectra of zircon and the plasma lines were 159 the same as those during the run using the Linkam TS1000 heating stage, except that the 160 output laser power set via the laser controller was 300 mW.

161 The same HDAC was used to obtain Raman spectra of zircon as a function of pressure up to 162 1.95 GPa at ambient temperature. Zircon and quartz crystals were loaded into the sample 163 chamber (a hole with an initial diameter of $\sim 400 \ \mu m$ in a Re gasket that separated the 164 diamond anvils) together with distilled water as pressure medium. The pressure was increased 165 incrementally by stepwise compression of the sample chamber. The pressure was determined 166 by recording Raman spectra of quartz immediately before and after that of zircon and using 167 the average of the pressures obtained from the wavenumber shift of the 464 cm⁻¹ Raman line 168 of quartz based on Eq. (2) of Schmidt and Ziemann (2000). A NIKON M Plan SLWD 20× 169 objective (numerical aperture 0.35) was used, and the output laser power set via the laser 170 controller was 300 mW, which corresponded to approximately 30 mW in the sample 171 chamber. Another series of experiments at about 23 °C and pressures up to 6.6 GPa was 172 carried out using a megabar-type piston-cylinder diamond-anvil cell (DAC) (Mao and 173 Hemley 1998) with type IIa low fluorescence diamond anvils with culet size of 400 µm. 174 Chips of zircon and quartz and annealed ruby spheres were loaded into the sample chamber, 175 which consisted of a hole of 200 µm diameter in a pre-indented stainless steel gasket. A 176 16:3:1 mixture of methanol, ethanol, and distilled water was used as pressure medium to 177 ensure hydrostatic conditions. Similarly to the HDAC experiments as function of pressure at 178 ambient temperature, the pressure was incrementally increased by stepwise compression of 179 the sample chamber, and spectra of zircon, quartz, and ruby were collected for each pressure 180 step. An OLYMPUS SLMPlan N 20× objective (numerical aperture 0.25) was used. The 181 output laser power set via the laser controller was 206 mW, which corresponded to a 182 measured laser power of 40.4 mW "downstream" of the objective. For recording the strong R_1 183 fluorescence line of ruby, the laser power was decreased by a factor of 1000 by inserting a 184 filter. This also ensured that no laser heating of the ruby spheres occurred. Watenphul and 185 Schmidt (2012) showed that pressures determined from the calibrated wavenumber shift of the 464 cm⁻¹ Raman line of quartz using equation (2) of Schmidt and Ziemann (2000) remain 186 187 in good agreement to pressures of at least 10 GPa with pressures obtained from the shift of the 188 R_1 ruby fluorescence line using pressure dependencies determined by Piermarini et al. (1975). Therefore, the pressure for this run series are from the shift of the 464 cm⁻¹ Raman line of 189 quartz, with verification by the pressures from the shift of the R₁ fluorescence line of ruby 190 191 (using the calibration by Piermarini et al. 1975). The deviations of these pressures from those 192 determined from quartz remained small, but increased from approximately -0.01 GPa at the 193 lowest pressure to approximately +0.25 GPa at the highest pressure of this series.

194 For the experiments at simultaneously elevated pressures and temperatures, we used the same 195 HDAC as for the run at 0.1 MPa to 741 °C and for the series at 27 °C to 1.95 GPa, except that 196 the upper anvil was replaced by a synthetic type IIa ultra-low fluorescence grade diamond 197 anvil (culet diameter 0.9 mm). The measured temperature was calibrated as described for the 198 run at 0.1 MPa to 741 °C. Chips of the zircon sample and natural quartz, and distilled water 199 were loaded into the HDAC sample chamber, for which an iridium gasket was used to ensure 200 that the gasket material remained inert during the experiment at high temperature and 201 pressure. The sample chamber was sealed and heated to 600 °C to clean the chamber and the 202 loaded crystals by reaction of possible fluorescence-causing organic contaminants with water 203 to form methane (Schmidt 2009). After cooling, the sample chamber was opened, and fresh 204 distilled water was loaded. Before it was sealed again, an air bubble was allowed to grow by 205 controlled leaking until the desired approximate initial fluid density was attained. The sample 206 was then studied by visual observation and Raman spectroscopy at different temperatures and 207 pressures along a near isochoric pressure-temperature path. Then, the fluid density was 208 changed by adjusting of the liquid-vapor ratio, and data along the next near isochoric path

209 were collected. Data at three to four different pressures were obtained for each isotherm (300, 210 400, 500, 600, and 700 °C) and for one experiment at 200 °C. The pressure at elevated 211 temperature was determined from the measured liquid-vapor homogenization temperature 212 after isochoric cooling and the IAPWS 1995 equation of state of H₂O (Wagner and Pruß 213 2002). At temperatures up to 500 °C, the wavenumber shift of the 464 cm⁻¹ Raman line of 214 quartz (Schmidt and Ziemann 2000; equations (2) and (3) therein) was used for additional 215 pressure determination. At these conditions, the concentration of dissolved silica is relatively 216 small, and the pressure from both methods was in good agreement as expected, and as already 217 noted by Schmidt and Watenphul (2010) for an experiment with quartz and water to 600 °C, 218 920 MPa. The resolution (random uncertainty) of the pressure determination from the wavenumber shift of the 464 cm⁻¹ Raman line of guartz was about 25 MPa (Schmidt 2009). 219 During the runs, the cell was flushed with 99% Ar + 1% H_2 gas to prevent oxidation of 220 221 metallic parts. The objective was an OLYMPUS SLMPlan N 20× (numerical aperture 0.25), 222 and the output laser power set via the laser controller 206 mW (40.4 mW measured in the 223 focal point of the objective). Additional data at 200 °C, 360 MPa and 300 °C, 304 MPa were 224 obtained from a run using spectroscopic grade ethanol as pressure medium, for which the pressure was determined from the wavenumber shift of the 464 cm⁻¹ Raman line of guartz 225 226 (Schmidt and Ziemann, 2000) and emission lines from a Ne lamp were used to calibrate the 227 wavenumber.

In all run series, reference Raman spectra of the solids in the sample chambers of the heating stage or pressure cell (diamond-anvil cell) were acquired at ~0.1 MPa and ambient temperature at the start of the series. In the runs using a diamond-anvil cell, trapping of a small air bubble when the sample chamber was sealed ensured the low reference pressure. Additional reference spectra were acquired in these experiments at the end of the series, after pressure release. 234 The zircon pressure sensor was tested during HDAC experiments on $H_2O+NaAlSi_3O_8$ and 235 H₂O+Na₂Si₃O₇ fluids. Details of those studies will be reported elsewhere, whereas here we 236 will only discuss the outcomes of using the zircon pressure sensor. Note that the same zircon 237 sample was used in all of these experiments, and the bulk composition was determined using 238 the method described by Manning et al. (2008). Specifically, the mass of the doubly polished 239 silicate glass piece loaded into the sample chamber was calculated from its determined 240 volume and its known density. The mass of the added water was calculated from the volume 241 of the sample chamber, less the volumes of glass and zircon chips, and its density as obtained 242 from the liquid-vapor homogenization temperature upon first heating. All volumes were 243 obtained by optical micrometry. The Raman spectra of the test experiments were recorded 244 using HORIBA Jobin Yvon LabRAM HR800 Raman spectrometers (gratings 1800 lines/mm, focal length 800 mm, excitation by the 514.53 nm line of an Ar⁺ laser or by the 472.937 nm 245 246 line from a diode laser), and the wavenumber was calibrated using emission lines from a Ne 247 lamp.

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249 **3. Results**

Table 1 lists the data for the wavenumbers and the linewidth (FWHM = full width at half 250 251 maximum) of the $v_3(SiO_4)$ Raman band of zircon as function of temperature and pressure. The 252 temperature dependence of the position of the $v_3(SiO_4)$ Raman band at 0.1 MPa is shown in 253 Figure 2. Two different heating stages were used, a Linkam TS1000 and a HDAC. There is 254 good agreement between both series, although temperature gradients in the sample chambers 255 were certainly different. Because of the larger range in temperature, the data obtained using 256 the Linkam TS1000 were used to fit the relationship between wavenumber v of the $v_3(SiO_4)$ 257 Raman band of fully crystalline zircon and temperature T at 0.1 MPa

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$$v[\text{cm}^{-1}] = 7.54(88)*10^{-9}*T^3 - 1.61(13)*10^{-5}*T^2 - 2.89(5)*10^{-2}*T + 1008.90(6)$$
 (1)

259 where $22 < T [^{\circ}C] < 950$. Numbers in parentheses denote 1σ errors in last significant digits. 260 The deviation of wavenumbers calculated from Eqn (1) and data from the run using the Linkam TS1000 stage is between -0.12 cm⁻¹ and 0.11 cm⁻¹, mostly $\leq \pm 0.05$ cm⁻¹. The 261 deviation of wavenumbers calculated from Eqn (1) and data from the other experiments at 0.1 262 MPa using the HDAC or DAC ranges from -0.45 cm⁻¹ to 0.34 cm⁻¹. 263 Within the uncertainty of the data, the $v_3(SiO_4)$ (~1008 cm⁻¹ at ambient temperature) Raman 264 265 band of synthetic zircon displayed a linear change in wavenumber with changing pressure (Fig. 3). At 23 °C, the obtained $\partial v/\partial P$ slope of the v₃(SiO₄) band is 5.69 cm⁻¹/GPa based on 266 267 the data from the DAC experiments to the maximum pressure of 6.6 GPa. Slight nonlinearity may be inferred from the increase in that slope to 5.76 cm⁻¹/GPa if only the data to 2.52 GPa 268 are included in the linear fit. This is confirmed by the slope of 5.77 cm^{-1} /GPa from the HDAC 269 270 series to 1.98 GPa at 27 °C (Fig. 3). Figure 4 shows that the $\partial v/\partial P$ slope does not significantly 271 change with T, as determined for the 700 and 600 °C isotherms ($\partial v/\partial P \sim 5.81$ and 5.89 cm⁻ ¹/GPa). The $\partial v/\partial P$ slopes along other isotherms (300, 400, and 500 °C) do not differ by much 272 273 from these values, but are less reliable because of the smaller range in pressure and are 274 therefore given in parentheses in Fig. 4.

In this study, we focused on the shift in the wavenumber of the of the $v_3(SiO_4)$ (~1008 cm⁻¹) 275 276 Raman band. Therefore, results for the linewidth of the $v_3(SiO_4)$ band and results for other 277 Raman bands are only briefly summarized here. In the following, these Raman bands v_i will 278 be denoted by their approximate position at ambient conditions, i.e., v_{975} , v_{439} , v_{356} , v_{224} , v_{214} , 279 and v_{202} . Their positions and widths as function of temperature and pressure are listed in the 280 electronic appendix (Table 2). Additionally, Fig. 5 shows a series of Raman spectra recorded 281 at various temperatures and constant pressure (0.1 MPa). At 0.1 MPa, the v_{439} band displays the largest increase in the full width at half maximum with temperature from ~5 cm⁻¹ at 22 °C 282 to ~24 cm⁻¹ at 900 °C. The increases in the linewidths of the Si–O stretching bands v_{1008} and 283 v_{975} with temperature at 0.1 MPa are moderate, from ~2 cm⁻¹ at 22 °C to ~13 cm⁻¹ at 900 °C. 284

285	The FWHM of the v_{356} band increases from ~4 cm ⁻¹ at 22 °C to ~18 cm ⁻¹ at 900 °C, that of the
286	v_{224} band from ~3 cm ⁻¹ at 22 °C to ~12 cm ⁻¹ at 900 °C, that of the v_{214} band from ~3 cm ⁻¹ at 22
287	°C to ~14 cm ⁻¹ at 900 °C, and the FWHM of the v_{202} band shows the smallest increase from
288	~1 cm ⁻¹ at 22 °C to ~4 cm ⁻¹ at 900 °C. The temperature dependencies of the wavenumbers of
289	of the Si–O stretching bands v_{1008} and v_{975} at 0.1 MPa are distinctly nonlinear, whereas those
290	of the v_{439} , v_{356} , v_{224} , v_{214} , and v_{202} bands are close to linear. In the following, the $\partial v/\partial T$ slopes
291	at 22 and 700 °C are reported in the case of nonlinear behavior to enable comparison between
292	different Raman bands. The ν_{1008} band shows the largest wavenumber decrease with
293	temperature at 0.1 MPa. The first derivative of Eqn (1) gives -0.0296 cm ⁻¹ /deg at 22 $^{\circ}$ C and -
294	0.0404 cm ⁻¹ /deg at 700 °C. The $\partial v/\partial T$ slopes at 0.1 MPa of the other bands are -0.0194 cm ⁻¹
295	$^{1}/deg$ at 22 °C and -0.0240 cm ⁻¹ /deg at 700 °C for the v ₉₇₅ , -0.0166 cm ⁻¹ /deg for v ₄₃₉ , -0.0187
296	cm ⁻¹ /deg for v_{356} (-0.0147 cm ⁻¹ /deg at 22 °C and -0.0196 cm ⁻¹ /deg at 700 °C), -0.0056 cm ⁻¹
297	$^{1}/deg$ for v ₂₂₄ , -0.0105 cm ⁻¹ /deg for v ₂₁₄ (-0.0092 cm ⁻¹ /deg at 22 °C and -0.0112 cm ⁻¹ /deg at
298	700 °C), and -0.0034 cm $^{-1}/deg$ for the ν_{202} band (-0.0022 cm $^{-1}/deg$ at 22 °C and -0.0042 cm $^{-1}$
299	1 /deg at 700 °C). The isothermal change in the wavenumber with pressure is nearly constant
300	for all studied Raman bands. At 23 °C, the obtained $\partial v/\partial P$ slopes for the data to 6.6 GPa are
301	5.69 cm ⁻¹ /GPa for the $v_{1008} = v_3(SiO_4)$ band (Fig. 3), 5.16 cm ⁻¹ /GPa for v_{975} , 1.45 cm ⁻¹ /GPa for
302	v_{439} , 4.56 cm ⁻¹ /GPa for v_{356} , 0.22 cm ⁻¹ /GPa for v_{224} , 1.45 cm ⁻¹ /GPa for v_{214} , and -0.45 cm ⁻¹ /GPa for v_{214}
303	¹ /GPa for the v_{202} band. As determined for the v_{1008} band, the $\partial v/\partial P$ slopes of the v_{975} , v_{356} ,
304	v_{224} , v_{214} , and v_{202} bands at 23°C do likewise not differ significantly from the $\partial v/\partial P$ slopes
305	along other isotherms (500, 600, and 700 °C). The $\partial v/\partial P$ slope of the v ₄₃₉ band may be
306	dependent on temperature because the slopes along the 300, 400, 500, 600, and 700 $^\circ\mathrm{C}$
307	isotherms are slightly but consistently smaller than that at 23 °C. Observed changes in the
308	linewidths of the v_{975} , v_{439} , v_{356} , v_{224} , v_{214} , and v_{202} Raman bands with pressure along isotherms
309	are probably mostly insignificant (Tab. 1 and electronic appendix (Tab. 2)).

310 The zircon pressure sensor was applied to determine isochores for aqueous fluids with 6, 35, 311 or 37 mass% NaAlSi₃O₈ (Figs. 6a-c), and 11, 26, or 61 mass% Na₂Si₃O₇ (Figs. 6d-f) to 803 312 °C and 1.65 GPa. This was done along cooling P-T paths, for which the sample chamber 313 volume usually remains constant (Bassett et al. 1993; Schmidt and Ziemann 2000). In the case 314 of the NaAlSi₃O₈ dissolution experiments in water, the difference between the isochores from 315 the zircon sensor and the isochores calculated from the liquid-vapor homogenization 316 temperature using the EoS of H_2O (Wagner and Pruß 2002) was small or insignificant to 317 about 700 °C. Above this temperature, the pressure from the zircon pressure sensor increased 318 less with temperature than the pressure from the EoS of H_2O in the experiments with the bulk compositions H₂O + 35 mass% NaAlSi₃O₈ and H₂O + 37 mass% NaAlSi₃O₈ (Figs. 6b-c). For 319 320 experiments with $H_2O + Na_2Si_3O_7$, the dP/dT slopes of the isochores obtained from zircon 321 were about the same or steeper than those calculated from homogenization temperature and 322 EoS of H₂O (Figs. 6d-f). The largest observed difference between both pressures was \sim 520 323 MPa at 500 °C in the experiment with $H_2O + 61 \text{ mass}\% \text{ Na}_2\text{Si}_3O_7$ (Fig. 6f). None of the 324 calibration and few of the test experiments with H₂O+SiO₂±Na₂O±Al₂O₃ fluids pointed to a 325 high zircon solubility. Fairly high zircon solubilities were indicated in an experiment with 326 $H_2O + 80 \text{ mass}\% \text{ Na}_2\text{Si}_3O_7$, in which ~0.26 vol% zircon dissolved completely upon heating to 327 600 °C at ~500 MPa, a run with a H₂O+6 mol% Na₂O+7 mol% SiO₂ fluid, in which ~0.3 328 vol% zircon dissolved completely at 600 °C and ~900 MPa, and a run with a H₂O+6 mol% 329 Na₂O+4 mol% SiO₂ fluid, in which ~ 0.89 vol% zircon dissolved completely at 600 °C and 330 ~1400 MPa. Furthermore, there was no optical and Raman spectroscopic detection of 331 formation of secondary Zr-oxides or Zr-silicates in any of these runs.

332

333 4. Discussion and Conclusions

There is close agreement in the obtained $\partial v/\partial P$ slopes of the v₃(SiO₄) band for different isotherms (5.76 cm⁻¹/GPa at 23 °C to 2.52 GPa, 5.77 cm⁻¹/GPa at 27 °C to 1.98 GPa, 5.89 cm⁻¹ ¹/GPa at 600 °C to 1.05 GPa, and 5.81 cm⁻¹/GPa at 700 °C to 1.24 GPa), although the pressures were determined using different techniques. The results indicate that the shift in the position of the v₃(SiO₄) band with changing pressure is close to constant ($\partial v/\partial P = 5.8\pm0.1$ cm⁻¹/GPa) and independent of temperature over the whole range in pressures and temperatures studied here. The possible slightly nonlinear behavior was not taken into account because it is within the error of the linear approximation of ±0.1 cm⁻¹/GPa, which corresponds to ±0.12 GPa at 7 GPa.

343 This nearly linear, fairly large, and temperature-independent shift of the $v_3(SiO_4)$ band with 344 pressure is only one argument for the suitability of crystalline zircon as pressure sensor for 345 experiments using optical cells at high P and T. Zircon is of particular advantage in situations 346 where other phases cannot be used as pressure sensors because of phase transitions or high 347 solubility. Diamond-anvil cell experiments at ambient temperature showed that synthetic 348 zircon transforms into reidite at 19.7 GPa (Van Westrenen et al. 2004). This pressure should 349 be higher than the equilibrium transition pressure (Knittle and Williams 1993). The transition 350 pressure determined by Liu (1979) is 12 GPa at 1027 °C. Ono et al. (2004) found a positive 351 $\partial P/\partial T$ slope of the phase boundary between zircon and reidite (8.33 GPa at 827 °C, 8.67 GPa 352 at 1027 °C). This indicates that zircon should be applicable as Raman spectroscopic pressure 353 sensor to pressures of ~ 10 GPa even at elevated temperature or, in other words, over most of 354 the range in P and T relevant for experiments on aqueous fluids at subduction zone 355 conditions.

However, the wavenumber of the $v_3(SiO_4)$ Raman band also depends on the degree of radiation-induced disorder and the crystallite size of recrystallized zircon (e.g., Nasdala et al. 1995, Zhang et al. 2000; Geisler and Pidgeon 2002) and on the composition, e.g., it increases substantially with the Hf concentration (~0.12 cm⁻¹/mol% HfSiO₄, Hoskin and Rodgers 1996; 0.13 cm⁻¹/mol% HfSiO₄, own data) and decreases with addition of Th (e.g., Syme et al., 1977). The Zr–Hf ratio in the zircon used in this study was very close to the estimated atomic 362 ratio of the crustal abundances of Zr and Hf of about 70:1, and natural zircons do usually not 363 deviate much from this ratio (Finch and Hanchar 2003, and references therein). Therefore, 364 most natural zircons with little radiation-damage can also be used as Raman spectroscopic 365 pressure sensor. Systematic errors in pressure determination, e.g., induced by the composition 366 of natural zircon samples, can be eliminated by measuring the shift in the wavenumber of the 367 $v_3(SiO_4)$ Raman band as wavenumber difference Δv relative to the position of that band at ambient temperature and low (near atmospheric) pressure in the same sample. It needs to be 368 369 emphasized that the actual sample temperature should be well known for accurate 370 determination of P using this sensor because of the temperature sensitivity of the wavenumber $(\partial v/\partial T \text{ is } \sim -0.03 \text{ cm}^{-1}/\text{deg at } 22 \text{ °C and } \sim -0.04 \text{ cm}^{-1}/\text{deg at } T \ge 500 \text{ °C})$. This temperature 371 372 sensitivity may be the largest source of error in the application of zircon as Raman 373 spectroscopic pressure sensor. The largest deviations of wavenumbers from experiments at 0.1 MPa using the HDAC (Tab. 1) and wavenumbers calculated from Eqn (1), which is based 374 on the data from experiments using the Linkam TS1000 heating stage, were 0.16 cm⁻¹ and -375 376 0.45 cm^{-1} , respectively. These deviations in wavenumber correspond to deviations in pressure 377 of +28 and -78 MPa, but include small systematic errors related to temperature measurement 378 and calibration. Deviations in the wavenumber of data at ambient temperature range between -0.08 to 0.34 cm⁻¹. We therefore consider \pm 50 MPa to be a reasonable estimate of the 379 380 reproducibility and random error of this technique.

An additional advantage of using zircon as a pressure sensor is its chemical durability at most conditions. As an amphoteric element, zirconium is more soluble in acidic or basic fluids than at near neutral pH. However, the solubility of crystalline zircon in H₂O+HCl is rather moderate, $\sim 10^{-2} m_{Zr}$ at 500 °C in $\sim 7 m$ HCl (Schmidt et al. 2006). In aqueous fluids, zircon is stable below quartz saturation, but not at very low silica activities (e.g., Newton et al. 2005). The solubility of zircon in hydrous peraluminous andesitic to granitic melts is low, ~ 100 to 1330 ppm Zr at 750 to 1020 °C and 120 to 600 MPa (Watson and Harrison 1983). At 700– 388 800 °C, 200 MPa, the solubility of zircon is below 100 ppm Zr in felsic peraluminous hydrous 389 melts in the system SiO₂-Al₂O₃-Na₂O-K₂O, but increases strongly and linearly as function of 390 increasing alkali/aluminum ratio, with 3.9 mass% Zr dissolved up to at 391 $(Na_2O+K_2O)/Al_2O_3=2.0$ (Watson 1979). In our study, the highest congruent zircon solubility 392 was observed in a test experiment, in which 0.89 vol% zircon dissolved completely at 600 °C 393 and ~1400 MPa in a highly alkaline aqueous fluid ($H_2O+6 \text{ mol}\% \text{ Na}_2O+4 \text{ mol}\% \text{ SiO}_2$).

394 An obvious application of zircon as Raman spectroscopic pressure sensor are studies of 395 hydrous silicate melts or other systems for which α -quartz is of limited applicability due to 396 phase transitions and high solubility. It has long been suspected that pressures in such 397 experiments may not be well approximated by the EoS of H₂O due to the considerable excess 398 volume of mixing in such systems (e.g., Mysen and Wheeler 2000). However, until recently 399 pressures estimated from the EoS of H_2O had to be assumed in studies of hydrous silicate 400 melts, because an appropriate spectroscopic pressure sensor for use in such experiments was 401 lacking (e.g., Shen and Keppler 1997; Bureau and Keppler 1999; Manning et al. 2008; 402 Borchert et al. 2009, 2010). The isochores determined in this study in experiments on 403 mixtures of aqueous fluids and silicate melts using the zircon sensor demonstrate that high 404 mutual solubility of fluid and melt can have a significant effect on the P-V-T properties. 405 Application of the EoS of H_2O (Wagner and Pruß 2002) appears to yield good or reasonable 406 approximations of the actual pressure for low concentrations of dissolved NaAlSi₃O₈. If more 407 than about 10 mass% NaAlSi₃O₈ was dissolved in the aqueous fluid, the isochores deviated 408 towards lower pressure from those calculated from the EoS of H₂O (Figs. 6b-c). The situation 409 is different for $H_2O + Na_2Si_3O_7$ mixtures, for which the pressure from the EoS of H_2O was 410 often much lower than the actual pressure, particularly at high concentrations of dissolved 411 sodium silicate (Fig. 6f) and at relatively high bulk densities (Figs. 6d and 6f). The results of 412 these tests suggest that it may be worthwile to revise phase equilibria involving aqueous fluid

413	and silicate melt, e.g., to redetermine the P - T location of critical curves in pseudobinary
414	systems such as those studied by Shen and Keppler (1997) and Bureau and Keppler (1999).
415	
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519

520 Figure captions

521

Fig. 1. Raman spectrum of synthetic fully crystalline zircon (near endmember composition $Zr_{0.987}Hf_{0.013}[SiO_4]$) at 22 °C, 0.1 MPa, with band assignments. Arrows indicate the direction of the shift in the wavenumber of these bands with temperature and pressure.

525

Fig. 2. Position of the $v_3(SiO_4)$ Raman band of zircon vs. temperature at 0.1 MPa. Errors are smaller than the symbol size. The solid line represents the fit of the data measured using the Linkam TS1000. The given equation is valid for $22 \le T [°C] \le 950$.

529

Fig. 3. Position of the $v_3(SiO_4)$ Raman band of zircon vs. pressure at ~22 °C (DAC experiment) and ~27 °C (HDAC experiment). Errors are smaller than the symbol size. Solid lines represent linear fits of the data. Numbers refer to the slopes of these linear equations. Details are given in the text.

534

Fig. 4. Position of the $v_3(SiO_4)$ Raman band of zircon vs. pressure along isotherms from 27 to 700 °C (HDAC experiments, pressure determined from EoS of H₂O (Wagner and Pruß 2002) and/or quartz as Raman spectroscopic pressure sensor (Schmidt and Ziemann 2000)). Values for the $\partial v/\partial P$ slopes are obtained from linear fits (solid and dashed lines). If these values are

- 539 more uncertain because of the more limited pressure range, the lines are dashed and the 540 numbers for the $\partial v / \partial P$ slopes are given in parentheses.
- 541

Fig. 5. Series of Raman spectra of synthetic fully crystalline zircon $(Zr_{0.987}Hf_{0.013}[SiO_4])$ recorded at various temperatures and 0.1MPa using the Linkam TS1000. The spectra are shifted along the intensity axis, but are not normalized to the acquisition time or to the temperature based on the Bose-Einstein occupation factor. * = plasma lines from laser.

546

547 Fig. 6. Isochores determined from the zircon pressure sensor (thick grey lines and symbols,

548 diamonds = individual pressure determinations) using the calibration from this study, and

549 isochores (thick black lines) and densities from the EoS of H₂O (Wagner and Pruß 2002).

550 Dashed lines – isochores for the same bulk composition, but at a lower density. Also shown

are *P*-*T* conditions of phase transitions (vapor dissolution or ice I melting, and silicate melt

dissolution (circles), qz = quartz. The estimated uncertainty in pressure from the $v_3(SiO_4)$

553 Raman band of zircon is \pm 50 MPa (about symbol size). (a) H₂O + 6 mass % NaAlSi₃O₈, (b)

554 $H_2O + 35 \text{ mass } \% \text{ NaAlSi}_3O_8$, (c) $H_2O + 37 \text{ mass } \% \text{ NaAlSi}_3O_8$, (d) $H_2O + 11 \text{ mass } \%$

555 Na₂Si₃O₇, (e) H₂O + 26 mass % Na₂Si₃O₇, (f) H₂O + 61 mass % Na₂Si₃O₇.

556

557 Table 1. Experimentally determined wavenumber and linewidth of the $v_3(SiO_4)$ Raman band

558	of fully cryst	talline	synthetic	zircon	(Zr _{0.987} Hf _{0.013} [SiO ₄])	as	functions	of temperature	e and
559	pressure.								

Т	Р	$v_3(SiO_4)$	$\Gamma v_3(SiO_4)$	section	$T_{\rm LV=L}$	experimental
[°C]	[MPa]	$[cm^{-1}]$	$[\text{cm}^{-1}]$	$[cm^{-1}]$	[°C]	details
22	0.1	1008.15	2.3	950-1050		TS1000
50	0.1	1007.43	2.5	950-1050		TS1000
100	0.1	1005.92	2.7	950-1050		TS1000
150	0.1	1004.34	3.1	950-1050		TS1000
200	0.1	1002.56	3.6	925-1075		TS1000
250	0.1	1000.83	4.0	925-1075		TS1000
300	0.1	998.96	4.5	875-1100		TS1000
350	0.1	997.16	5.0	875-1100		TS1000
400	0.1	995.17	5.6	875-1100		TS1000

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450	0.1	993.29	6.2	875-1100	TS1000
500	0.1	991.30	6.8	875-1100	TS1000
550	0.1	989.34	7.5	875-1100	TS1000
600	0.1	987.37	8.2	875-1100	TS1000
650	0.1	985.36	8.8	875-1100	TS1000
700	0.1	983.43	9.5	875-1100	TS1000
750	0.1	981.47	10.5	850-1100	TS1000
800	0.1	979.31	11.1	825-1125	TS1000
850	0.1	977.37	12.0	825-1125	TS1000
900	0.1	975.25	12.9	825-1125	TS1000
950	0.1	973.34	13.9	825-1125	TS1000
25.5	0.1	1008.23	2.7	950-1050	HDAC
25.3	0.1	1008.04	2.7	950-1050	HDAC
98.8	0.1	1006.13	3.1	950-1050	HDAC
98.8	0.1	1005.91	3.3	950-1050	HDAC
197.6	0.1	1002.56	4.0	925-1075	HDAC
197.6	0.1	1002.92	3.7	925-1075	HDAC
296.4	0.1	999.32	4.6	875-1100	HDAC
395.2	0.1	995.87	5.6	875-1100	HDAC
494	0.1	991.59	6.8	875-1100	HDAC
592.9	0.1	987.95	8.1	875-1100	HDAC
691.6	0.1	984.16	9.5	875-1100	HDAC
741	0.1	982.11	10.2	850-1100	HDAC
26.8	0.1	1007.95	2.6	950-1050	HDAC,W
26.5	46	1008.35	2.5	950-1050	HDAC,W,Q
26.5	88	1008.47	2.6	950-1050	HDAC,W,Q
26.2	173	1009.08	2.6	950-1056	HDAC,W,Q
26.3	265	1009.67	2.7	950-1056	HDAC,W,Q
26.6	440	1010.61	2.6	950-1056	HDAC,W,Q
26.5	606	1011.55	2.6	955-1056	HDAC,W,Q
26.6	874	1013.00	2.6	955-1056	HDAC,W,Q
26.6	1100	1014.46	2.9	960-1056	HDAC,W,Q
26.7	1264	1015.32	2.6	960-1056	HDAC,W,Q
26.7	1439	1016.48	2.8	960-1056	HDAC,W,Q
27.0	1303	1015.75	2.6	960-1056	HDAC,W,Q
26.9	1682	1017.87	2.7	960-1056	HDAC,W,Q
26.8	1682	1017.83	2.7	960-1056	HDAC,W,Q
27.0	1948	1019.23	2.8	960-1056	HDAC,W,Q
26.9	1948	1019.31	2.8	960-1056	HDAC,W,Q
26.7	1978	1019.35	2.7	950-1075	HDAC,W,Q
24.2	0.1	1007.85	2.3	950-1075	DAC,MEW
23.1	41	1008.04	2.3	950-1075	DAC,MEW,Q
22.8	241	1009.33	2.4	950-1075	DAC,MEW,Q
22.9	1217	1014.76	2.3	950-1075	DAC,MEW,Q
23.0	1503	1016.49	2.3	950-1075	DAC,MEW,Q
23.1	2516	1022.39	2.2	950-1075	DAC,MEW,Q
23.2	3258	1026.43	2.2	950-1080	DAC,MEW,Q
23.2	3953	1030.34	2.2	950-1085	DAC,MEW,Q
23.2	4403	1032.87	2.3	950-1085	DAC,MEW,Q
23.2	4903	1035.89	2.2	950–1085	DAC,MEW,Q

23.2	5579	1039.83	2.2	950-1090		DAC,MEW,Q
23.3	6126	1042.87	2.2	950-1090		DAC,MEW,Q
23.4	6645	1045.29	2.4	950-1090		DAC,MEW,Q
23.4	44	1008.09	2.2	950-1075		DAC,MEW,Q
21.8	0.1	1008.19	2.2	950-1050		HDAC,W
200	120	1003.39	3.5	900-1075	131.7	HDAC,W,EoS
200	360	1004.42	3.6	950-1040		HDAC,E,Q
300	9	999.22	4.5	875-1100	$P_{ m V}$	HDAC,W,EoS
300	149	1000.00	4.5	875-1100	206.8	HDAC,W,EoS
300	310	1000.83	4.4	875-1100	131.7	HDAC,W,EoS
300	304	1000.71	4.5	940-1050		HDAC,E,Q
400	108	995.88	5.5	875-1100	303.9	HDAC,W,EoS
400	309	997.22	5.5	875-1100	206.8	HDAC,W,EoS
400	502	998.17	5.5	875-1100	131.7	HDAC,W,EoS
500	213	992.67	6.7	875-1100	303.9	HDAC,W,EoS
500	466	994.17	6.7	875-1100	206.8	HDAC,W,EoS
500	854	996.42	6.6	875-1100	71.4	HDAC,W,EoS
500	652	995.40	6.6	875-1100	144.0	HDAC,W,EoS
600	316	989.30	8.1	875-1100	303.9	HDAC,W,EoS
600	618	991.01	8.0	875-1100	206.8	HDAC,W,EoS
600	1052	993.53	8.0	875-1100	71.4	HDAC,W,EoS
600	829	992.37	8.0	875-1100	144.0	HDAC,W,EoS
700	437	985.79	9.5	875-1100	298.1	HDAC,W,EoS
700	765	987.88	9.6	875-1100	206.8	HDAC,W,EoS
700	1242	990.59	9.4	900-1100	71.4	HDAC,W,EoS

Notes: T = temperature; P = pressure; $v_3(SiO_4)$ = determined wavenumber of the $v_3(SiO_4)$ Raman band of zircon; $\Gamma v_3(SiO_4)$ = full width at half maximum of the $v_3(SiO_4)$ Raman band of zircon (not corrected for spectral resolution); section [cm⁻¹] = section of the spectrum used to fit the $v_1(SiO_4)$ and $v_3(SiO_4)$ Raman bands (plasma lines were omitted and fitted individually); $T_{LV=L}$ = liquid-vapor homogenization temperature of the aqueous pressure medium after isochoric cooling. P_V = experiment at vapor pressure. TS1000 = Linkam TS1000 heating stage, HDAC, = hydrothermal diamond-anvil cell, DAC = piston-cylinder diamond-anvil cell. Pressure medium: W = water; MEW = 16:3:1 mixture of methanol, ethanol, and distilled water; E = ethanol. Pressure determination technique: Q = wavenumber shift of the 464 cm⁻¹ Raman line of quartz using equation (2) of Schmidt and Ziemann (2000); EoS = liquid-vapor homogenization temperature after isochoric cooling and equation of state of H₂O (Wagner and Pruß 2002).

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Fig. 1

















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Fig. 6