

1 **REVISION 2**

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3 **Oxy-dravite, Na(Al<sub>2</sub>Mg)(Al<sub>5</sub>Mg)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, a new mineral**  
4 **species of the tourmaline supergroup**

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17 **ABSTRACT**

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19 Oxy-dravite, Na(Al<sub>2</sub>Mg)(Al<sub>5</sub>Mg)(Si<sub>6</sub>O<sub>18</sub>)(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, is a new mineral of the  
20 tourmaline supergroup. The holotype specimen originates from the locality of Osarara (Narok  
21 district, Kenya) and occurs in quartz-muscovite schist. Crystals of oxy-dravite are dark red,  
22 partially translucent with a vitreous luster, a pink streak and conchoidal fracture. It has a Mohs  
23 hardness of approximately 7, and a calculated density of 3.073 g/cm<sup>3</sup>. In plane polarized light,  
24 oxy-dravite is pleochroic (O = orange and E = pink) and uniaxial negative: ω = 1.650(5), ε =  
25 1.620(5). Oxy-dravite is rhombohedral, space group *R3m*, with the unit-cell parameters *a* =  
26 15.9273(2) and *c* = 7.2001(1) Å, *V* = 1581.81(4) Å<sup>3</sup>, *Z* = 3. Chemical characterization based on  
27 electron microprobe analysis, single-crystal structure refinement, Mössbauer and optical  
28 spectroscopy, resulted in the unit formula:

29 
$$^X(\text{Na}_{0.83}\square_{0.15}\text{K}_{0.02})_{\Sigma 1.00} ^Y(\text{Al}_{1.34}\text{Fe}^{3+}_{0.58}\text{Mg}_{1.03}\text{Fe}^{2+}_{0.03}\text{Ti}_{0.02})_{\Sigma 3.00} ^Z(\text{Al}_{4.95}\text{Mg}_{1.03}\text{Fe}^{3+}_{0.02})_{\Sigma 6.00}$$
  
30 
$$^T(\text{Si}_{5.98}\text{Al}_{0.02}\text{O}_{18})_{\Sigma 6.00} ^B(\text{BO}_3)_3 ^V(\text{OH})_3 ^W(\text{O}_{0.76}\text{OH}_{0.24})_{\Sigma 1.00}.$$

31 While the end-member formula of oxy-dravite may be formalized as  
32 Na<sup>Y</sup>(Al<sub>3</sub>)<sup>Z</sup>(Al<sub>4</sub>Mg<sub>2</sub>)Si<sub>6</sub>O<sub>18</sub>(BO<sub>3</sub>)<sub>3</sub>(OH)<sub>3</sub>O, the most representative ideal structural formula is

33  $\text{Na}^Y(\text{Al}_2\text{Mg})^Z(\text{Al}_5\text{Mg})\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ . The difference between these two formulae is  
34 solely in Al-Mg order-disorder, i.e., there is no difference in chemical composition. Although  
35 the Mg-Al disorder over the *Y* and *Z* sites is controlled by the short-range bond-valence  
36 requirements of  $\text{O}^{2-}$  at the O1 ( $\equiv$  W) site, the amount of Mg at the *Z* site is a function of the  
37 degree of cation size mismatch at *Z*.

38 The crystal structure of oxy-dravite was refined to statistical index *R*1 of 1.17% using  
39 1586 equivalent reflections collected with MoK $\alpha$  X-radiation. Oxy-dravite is chemically  
40 related to dravite (and fluor-dravite),  $\text{NaMg}_3\text{Al}_6\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3(\text{OH},\text{F})$ , by the heterovalent  
41 substitution  $\text{Al}^{3+} + \text{O}^{2-} \rightarrow \text{Mg}^{2+} + (\text{OH},\text{F})^{1-}$ .

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## INTRODUCTION

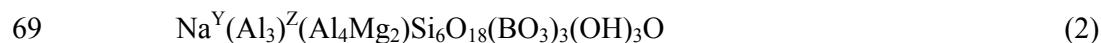
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46 The tourmaline supergroup minerals are widespread, occurring in a wide variety of  
47 sedimentary, igneous and metamorphic rocks. They are known as valuable indicator minerals  
48 that can provide information on the compositional evolution of their host rocks, chiefly due to  
49 their ability to incorporate a large number of elements (e.g., Novák et al. 2004; Agrosi et al.  
50 2006; Lussier et al. 2011a; Novák et al. 2011; van Hinsberg et al. 2011). However, the  
51 chemical composition of tourmalines is also strongly controlled by various crystal-structural  
52 constraints (e.g., Hawthorne 1996, 2002a; Bosi 2010, 2011; Henry and Dutrow 2011; Skogby  
53 et al. 2012) as well as by temperature (e.g., van Hinsberg and Schumacher 2011). Tourmaline  
54 supergroup minerals are complex borosilicates and their crystal structure and crystal chemistry  
55 have been widely studied (e.g., Foit 1989; Hawthorne and Henry 1999; Bosi and Lucchesi  
56 2007; Lussier et al. 2008; Bosi et al. 2010; Lussier et al. 2011b). In accordance with Henry et  
57 al. (2011), the general formula of tourmaline may be written as:  $\text{XY}_3\text{Z}_6\text{T}_6\text{O}_{18}(\text{BO}_3)_3\text{V}_3\text{W}$ ,  
58 where X ( $\equiv$   $^{[9]}X$ ) =  $\text{Na}^+$ ,  $\text{K}^+$ ,  $\text{Ca}^{2+}$ ,  $\square$  (= vacancy); Y ( $\equiv$   $^{[6]}Y$ ) =  $\text{Al}^{3+}$ ,  $\text{Fe}^{3+}$ ,  $\text{Cr}^{3+}$ ,  $\text{V}^{3+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ ,  
59  $\text{Mn}^{2+}$ ,  $\text{Li}^+$ ; Z ( $\equiv$   $^{[6]}Z$ ) =  $\text{Al}^{3+}$ ,  $\text{Fe}^{3+}$ ,  $\text{Cr}^{3+}$ ,  $\text{V}^{3+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ ; T ( $\equiv$   $^{[4]}T$ ) =  $\text{Si}^{4+}$ ,  $\text{Al}^{3+}$ ,  $\text{B}^{3+}$ ; B ( $\equiv$   $^{[3]}B$ ) =  
60  $\text{B}^{3+}$ ; W ( $\equiv$   $^{[3]}O1$ ) =  $\text{OH}^{1-}$ ,  $\text{F}^{1-}$ ,  $\text{O}^{2-}$ ; V ( $\equiv$   $^{[3]}O3$ ) =  $\text{OH}^{1-}$ ,  $\text{O}^{2-}$  and where, for example, T  
61 represents a group of cations ( $\text{Si}^{4+}$ ,  $\text{Al}^{3+}$ ,  $\text{B}^{3+}$ ) accommodated at the [4]-coordinated *T* sites.  
62 The dominance of these ions at one or more sites of the structure gives rise to a range of  
63 distinct mineral species (Henry et al. 2011).

64 The name oxy-dravite was first proposed by Hawthorne and Henry (1999) for the  
65 hypothetical formula:



67 Successively the name was confirmed in the approved classification scheme of Henry et al.  
68 (2011), but with a new ideal formula:



70 Although these two structural formulas are chemically equivalent, formula (1) shows a  
71 disordering of Mg over the Y and Z sites, whereas formula (2) shows an ordering of Mg at the  
72 Z site. As these two structural formulae show long-range disordered (1) and long-range  
73 ordered site populations (2) for the bulk composition  $\text{NaMg}_2\text{Al}_7\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ , we  
74 define (1) as the disordered formula and (2) as the ordered formula. The name oxy-dravite has  
75 been previously used for some time in the literature despite the lack of a formal definition of  
76 the mineral (e.g., Žáček et al. 2000; Novák et al. 2004; Bosi et al. 2010).

77 The new species as well as the name have been approved by the Commission on New  
78 Minerals, Nomenclature and Classification of the International Mineralogical Association  
79 (IMA 2012-004a). The holotype specimen of oxy-dravite is deposited in the collections of the  
80 Museum of Mineralogy, Earth Sciences Department, Sapienza University of Rome, Italy,  
81 catalogue number 33066. A formal description of the new species oxy-dravite is presented  
82 here, including a full characterization of its physical, chemical and structural attributes.

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#### 85 **OCCURRENCE, APPEARANCE AND PHYSICAL AND OPTICAL PROPERTIES**

86

87 The holotype specimen originates from the locality of Osarara (Narok district, Kenya),  
88 and occurs in quartz-muscovite schist (Dunn et al. 1975). The crystal is inclusion-free and  
89 occurs as an euhedral crystal approximately  $7 \times 7 \times 15$  mm in size. It is dark red in color, with  
90 pink streak, partially translucent and has a vitreous luster (Fig. 1). It is brittle and shows  
91 conchoidal fracture. The Mohs hardness is approximately 7 and the calculated density is  $3.073$   
92  $\text{g/cm}^3$ . In transmitted light, oxy-dravite is pleochroic with O = orange and E = pink. Oxy-  
93 dravite is uniaxial negative with refractive indices, measured by the immersion method using  
94 white light from a tungsten source, of  $\omega = 1.650(5)$ ,  $\varepsilon = 1.620(5)$ . The mean index of  
95 refraction, density and chemical composition lead to a compatibility index of 0.020, classed as  
96 excellent (Mandarino 1976, 1981).

97 It is worth pointing out that the dark red bulk color as well as the pleochroism observed  
98 is most likely caused by relatively minor concentrations of  $\text{Fe}^{3+}$  present in the mineral (Taran  
99 and Rossman 2002).

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## EXPERIMENTAL METHODS

103

### 104 **Single-crystal structural refinement (SREF)**

105 A representative fragment of the type specimen was selected for X-ray diffraction  
106 measurements on a Bruker KAPPA APEX-II single-crystal diffractometer, at Sapienza  
107 University of Rome (Earth Sciences Department), equipped with a CCD area detector ( $6.2 \times$   
108  $6.2 \text{ cm}^2$  active detection area,  $512 \times 512$  pixels) and a graphite crystal monochromator, using  
109  $\text{MoK}\alpha$  radiation from a fine-focus sealed X-ray tube. The sample-to-detector distance was 4  
110 cm. A total of 3265 exposures (step =  $0.2^\circ$ , time/step = 20 s) covering the full reciprocal  
111 sphere with a high redundancy of about 8 was used. Final unit-cell parameters were refined by  
112 means of the Bruker AXS SAINT program using reflections with  $I > 10 \sigma(I)$  in the range  $8^\circ <$   
113  $2\theta < 69^\circ$ . The intensity data were processed and corrected for Lorentz, polarization, and  
114 background effects with the APEX2 software program of Bruker AXS. The data were  
115 corrected for absorption using the multi-scan method (SADABS). The absorption correction  
116 led to a significant improvement in  $R_{\text{int}}$  (from 0.0363 to 0.0266). No violations of  $R3m$   
117 symmetry were noted.

118 Structural refinement was carried out with the SHELXL-97 program (Sheldrick 2008).  
119 Starting coordinates were taken from Bosi and Lucchesi (2004). Variable parameters that were  
120 refined include: scale factor, extinction coefficient, atomic coordinates, site scattering values  
121 and atomic displacement factors. To obtain the best values of statistical indexes ( $R1$ ,  $wR2$ ), a  
122 fully ionized O scattering curve was used, whereas neutral scattering curves were used for the  
123 other atoms. In detail, the X site was modeled using Na scattering factors. The occupancy of  
124 the Y site was obtained considering the presence of Al vs. Fe. The Z, T and B sites were  
125 modeled, respectively, with Al, Si and B scattering factors and with a fixed occupancy of 1,  
126 because refinement with unconstrained occupancies showed no significant deviations from  
127 this value. Three full-matrix refinement cycles with isotropic displacement parameters for all  
128 atoms were followed by anisotropic cycles until convergence was attained. No significant  
129 correlations over a value of 0.7 between the parameters were observed at the end of

130 refinement. Table 1 lists crystal data, data collection information and refinement details; Table  
131 2 gives the fractional atomic coordinates and site occupancies; Table 3 gives the displacement  
132 parameters; Table 4 gives selected bond distances.

133

### 134 **X-ray powder diffraction**

135 The X-ray powder-diffraction pattern for the oxy-dravite sample was collected using a  
136 Panalytical X'pert powder diffractometer equipped with an X'celerator silicon-strip detector.  
137 The range 5-80° (2 $\theta$ ) was scanned with a step-size of 0.017° during 30 minutes using a sample  
138 spinner with the sample mounted on a background-free holder. The diffraction data (in Å for  
139 CuK,  $\lambda_1 = 1.54060$  Å), corrected using Si as an internal standard, are listed in Table 5. Unit-  
140 cell parameters from the powder data were refined using the program UnitCell (Holland and  
141 Redfern 1997):  $a = 15.919(1)$  Å,  $c = 7.200(1)$  Å,  $V = 1580.3(1)$  Å<sup>3</sup>.

142

### 143 **Electron Microprobe analysis (EMPA)**

144 Electron microprobe analyses of the crystal used for X-ray diffraction refinements were  
145 obtained by using wavelength-dispersive spectrometry with a Cameca SX50 instrument at the  
146 “Istituto di Geologia Ambientale e Geoingegneria (Rome, Italy), CNR”, operating at an  
147 accelerating potential of 15 kV and a sample current of 15 nA (5  $\mu$ m beam diameter). Minerals  
148 and synthetic compounds were used as standards: wollastonite (SiK $\alpha$ , CaK $\alpha$ ), magnetite  
149 (FeK $\alpha$ ), rutile (TiK $\alpha$ ), corundum (AlK $\alpha$ ), fluorophlogopite (FK $\alpha$ ), periclase (MgK $\alpha$ ), jadeite  
150 (NaK $\alpha$ ), K-feldspar (KK $\alpha$ ), sphalerite (ZnK $\alpha$ ), along with metallic Cr, V, Mn and Cu. The  
151 PAP matrix correction procedure (Pouchou and Pichoir 1991) was applied to reduce the raw  
152 data. The results, which are summarized in Table 5, represent mean values of 10 spot analyses.  
153 In accordance with the documented very low concentration of Li in dravitic samples (e.g.,  
154 Henry et al. 2011), the Li<sub>2</sub>O content was assumed to be insignificant. Manganese, Cr, V, Zn,  
155 Cu, Ca and F were found to be below their respective detection limits varying between 0.03  
156 wt. % and 0.05 wt. %.

157

### 158 **Mössbauer spectroscopy (MS)**

159 The oxidation state of Fe was determined by Mössbauer spectroscopy at room  
160 temperature using a conventional spectrometer system operating in constant-acceleration  
161 mode. In order to spare the holotype material, the absorber was prepared by filling a small

162 quantity of ground material in a 1-mm hole in a lead plate, and the spectrum then acquired  
163 using a closely positioned  $^{57}\text{Co}$  point-source in rhodium matrix with a nominal activity of 10  
164 mCi. The spectrum was calibrated against  $\alpha\text{-Fe}$  foil and folded before fitting using the MDA  
165 software by Jernberg and Sundqvist (1983). The resultant spectrum (Fig. 2) shows a relatively  
166 broad central absorption doublet with weak shoulder features. On a sample of similar  
167 composition from Osarara, Narok district, Kenya (NMNH#126030), Mattson and Rossman  
168 (1984) obtained a comparable spectrum and interpreted the broad bands as being caused by  
169 relaxation effects. The spectrum of the holotype specimen was fitted with two bands assigned  
170 to  $\text{Fe}^{3+}$  and one band assigned to  $\text{Fe}^{2+}$ , resulting in an  $\text{Fe}^{3+}/\Sigma\text{Fe}$  ratio of approximately 0.96.

171

### 172 **Optical spectroscopy (OAS)**

173 Polarized room temperature optical absorption spectra in the  $\epsilon$  and  $\omega$  directions were  
174 recorded on a 31  $\mu\text{m}$  thick crystal section in the UV/VIS to NIR spectral range (330-1500 nm)  
175 with a Zeiss MPM800 microscope spectrometer (cf. Halenius et al. 2011). The obtained  
176 spectra (not shown here) are very similar to those reported by Mattson and Rossman (1984)  
177 and Taran and Rossman (2002) for the NMNH#126030 sample, showing three absorption  
178 bands at 486, 540 and 545 nm related to  $\text{Fe}^{3+}\text{-Fe}^{3+}$  pair transitions and two  $\text{Fe}^{2+}$  bands around  
179 700 and 1100 nm (Taran and Rossman 2002).

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181

## 182 **RESULTS**

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### 184 **Determination of atomic proportions**

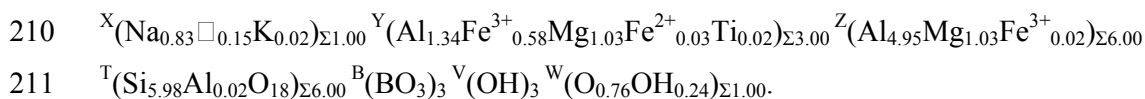
185 In agreement with the structural refinement results, the boron content was assumed to  
186 be stoichiometric in the sample of oxy-dravite ( $B = 3.00$  atoms per formula unit, apfu). In fact,  
187 both the site-scattering results and the bond lengths of  $B$  and  $T$  are consistent with the  $B$  site  
188 fully occupied by boron, but along with it being absent at the  $T$  site. The OH content can then  
189 be calculated by charge balance with the assumption ( $T + Y + Z = 15.00$ ). The atomic  
190 proportions were calculated on this assumption (Table 5). The excellent match between the  
191 number of electrons per formula unit (epfu) derived from chemical and structural analysis  
192 supports this procedure: 231.8 epfu vs. 232.8 epfu, respectively.

193

194 **Site populations**

195 Anion site populations were apportioned following the protocols of Grice and Ercit  
196 (1993) and Henry et al. (2011): the O3 site (V position in the general formula) is occupied by  
197 OH and O<sup>2-</sup>, while the O1 site (W position in the general formula) is occupied by O<sup>2-</sup> and OH.  
198 The cation distribution at the T, Y and Z sites was optimized by using a quadratic program (for  
199 details, see Bosi and Lucchesi 2004) to minimize the residuals between calculated and  
200 observed data (based on the chemical and structural analysis). Site scattering values,  
201 octahedral and tetrahedral mean bond distances (i.e., <Y-O>, <Z-O> and <T-O>) were  
202 calculated as the linear contribution of each cation multiplied by its specific bond distance  
203 (Table 6). More details about the specific distances derived from the ionic radii are found in  
204 Bosi and Lucchesi (2007). The robustness of this approach was confirmed by another  
205 optimization procedure (Wright et al. 2000) which led to very similar cation distributions in  
206 the present sample (Table 6). This result represents another example of convergence of these  
207 two procedures to similar solutions for tourmaline (i.e., Bosi and Lucchesi 2007; Filip et al.  
208 2012; Bosi et al. 2012; Bosi et al. 2013).

209 The empirical structural formula of oxy-dravite is:



212 The results from bond-valence analysis are also consistent with the proposed structural  
213 formula. Bond valence calculations, using the formula and bond-valence parameters from  
214 Brown and Altermatt (1985), are reported in Table 7. In particular, note that the value of bond  
215 valence sum incident at the O1 site (1.53 valence units) implies that O1 ( $\equiv$  W) is only partially  
216 occupied by an OH group (expected value close to 1). This is consistent with O<sup>2-</sup> being  
217 dominant at the O1 site (expected value close to 2). This finding can be verified by means of  
218 the empirical relation reported in the CNMNC-approved proposal IMA 2012-004a by Bosi  
219 and Skogby, in which the OH content at O1 in tourmaline can be estimated by considering the  
220 bond valence sum at O1 and the F content according to:  $^W\text{OH} = [2 - \text{BVS}(\text{O1}) - 0.20 - \text{F}]$ .  
221 The resulting value (0.27 apfu) is very close to the OH value obtained by stoichiometry (0.24  
222 apfu). It is noteworthy that the very small amounts of Fe<sup>3+</sup> at the Z site (0.02 apfu) optimized  
223 in the structural formula are in good agreement with the presence of Fe<sup>3+</sup> at Z observed in the  
224 polarized optical absorption spectrum of oxy-dravite samples from Osarara, with a band at 540

225 nm (**E** parallel to the **c** axis) assigned to electronic  $Z\text{Fe}^{3+}-Z\text{Fe}^{3+}$  pair transitions (Mattson and  
226 Rossman 1984; Taran and Rossman 2002).

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228

229

## DISCUSSION

230

### 231 **Red tourmalines**

232 Although not previously classified as oxy-dravite, unusual red  $\text{Fe}^{3+}$ -rich tourmalines  
233 from Osarara (Narok district, Kenya) have been previously studied by different techniques, in  
234 order to define the chemistry, the structure and the interactions between and among ions in the  
235 atomic sites (Dunn et al. 1975, EMPA; Mattson and Rossman 1984, EMPA, MS and OAS;  
236 Hawthorne et al. 1993, EMPA and SREF; Taran and Rossmann 2002, OAS). However, some  
237 differences occur with respect to the data obtained for the holotype specimen: for example, the  
238 unit-cell parameters are  $a = 15.9273(2)$  and  $c = 7.2001(1)$  Å in our sample, and  $a = 15.947(2)$   
239 and  $c = 7.214(1)$  Å in Hawthorne et al. (1993). Moreover, although the structural formula  
240 reported by Hawthorne et al. (1993) shows a cation site populations quite similar to that of our  
241 sample, the anion site populations at O3 and O1 were not specified, since the atomic  
242 proportions were calculated by the assumption  $V^{+W}(\text{OH}+\text{F}) = 4.00$  apfu, instead of the more  
243 appropriate assumption  $(T + Y + Z) = 15.00$  apfu. The latter assumption implies: 1) no  
244 vacancies at the octahedrally-coordinated sites in line with the crystal structure information  
245 and 2) a site population at O1 in line with the bond-valence sum value incident at the O1 site  
246 ( $\sim 1.5$  valence units). As a result, the anion population is  $W(\text{O}_{0.69}\text{OH}_{0.31})$ . It is noteworthy that  
247 the anion population is consistent with the short-range bond-valence constrains around O1,  
248 thus implying that the Mg-Al disorder over the *Y* and *Z* sites should be coupled to the  
249 incorporation of  $\text{O}^{2-}$  at the O1 site (e.g., Taylor et al. 1995; Hawthorne 1996).

250

### 251 **Nomenclature**

252 The empirical structural formula of the sample examined in this study indicates this  
253 tourmaline-supergroup mineral is best classed in the alkali group, oxy-subgroup 3 (Henry et al.  
254 2011) with Na dominant at the X position of the general formula, oxygen dominant at the W  
255 position with  $\text{O}^{2-} > \text{OH}$ , and Al as the dominant cation at the Y and Z positions. Disregarding  
256 the minor constituents (K,  $\text{Fe}^{2+}$  and Ti and  $^{\text{T}}\text{Al}$ ) and replacing ( $\text{ } + \text{OH}^-$ ) by ( $\text{Na}^+ + \text{O}^{2-}$ ) and



257  $\text{Fe}^{3+}$  by Al, the empirical structural formula of the studied tourmaline can be approximated as  
258  $\text{Na}^{\text{Y}}(\text{Al}_2\text{Mg})^{\text{Z}}(\text{Al}_5\text{Mg})\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ , that is, one having a disordered formula (1). As  
259 this composition has multiple cations at more than one site (i.e., Al and Mg), it is not in  
260 accordance with the characteristics of an ordered end-member formula as defined by  
261 Hawthorne (2002b).

262

### 263 **End-member formula**

264 The empirical formula presented above can be rearranged to the two possible end-  
265 members:  $\text{Na}^{\text{Y}}(\text{AlMg}_2)^{\text{Z}}(\text{Al}_6)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$  and  
266  $\text{Na}^{\text{Y}}(\text{Al}_3)^{\text{Z}}(\text{Al}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ .

267 In accordance with the short-range bond-valence requirements at O1 (e.g., Hawthorne  
268 1996), however, the end-member with Mg ordered at Y would suggest the existence of an  
269 unstable short-range arrangement  $^{\text{Y}}(\text{Al} + 2\text{Mg})\text{-}^{\text{W}}(\text{O}^{2-})$  in oxy-dravite, whereas the end-  
270 member with Mg ordered at Z is both consistent with the occurrence of the stable short-range  
271 arrangement  $^{\text{Y}}(3\text{Al})\text{-}^{\text{W}}(\text{O}^{2-})$  and with the characteristics of an end-member composition  
272 (Hawthorne 2002b). Consequently, the  $\text{Na}^{\text{Y}}(\text{Al}_3)^{\text{Z}}(\text{Al}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$  formula is  
273 preferred as the end-member oxy-dravite composition. Finally, note that both the disordered  
274 formula (1) and the empirical formula are consistent with the stable short-range arrangement  
275  $^{\text{Y}}(2\text{Al} + \text{Mg})\text{-}^{\text{W}}(\text{O}^{2-})$  (Hawthorne 1996).

276

### 277 **Name and relation to other species**

278 By analogy to the relationship between oxy-schorl and schorl (Bačík et al. 2013), oxy-  
279 chromium-dravite and chromium-dravite (Bosi et al. 2012), oxy-vanadium-dravite and  
280 “vanadium-dravite” (Bosi et al. 2013), the name oxy-dravite is given in relation to dravite. The  
281 prefix *oxy* represents the heterovalent substitution  $\text{Al}^{3+} + \text{O}^{2-} \rightarrow \text{Mg}^{2+} + (\text{OH})^{1-}$  relative to the  
282 root composition of dravite. Replacing (OH) by F in the above heterovalent substitution, also  
283 suggests a relationship between oxy-dravite and fluor-dravite, the latter being the fluor-  
284 equivalent of dravite (Clark et al. 2011). Comparative data for oxy-dravite, fluor-dravite and  
285 dravite are given in Table 8. Ideally, oxy-dravite is related to oxy-schorl, oxy-chromium-  
286 dravite, oxy-vanadium-dravite and povondraite through the homovalent substitution of  $\text{Mg}^{2+}$   
287 for  $\text{Fe}^{2+}$ ,  $\text{Al}^{3+}$  for  $\text{Cr}^{3+}$ ,  $\text{Al}^{3+}$  for  $\text{V}^{3+}$  and  $\text{Al}^{3+}$  for  $\text{Fe}^{3+}$ , respectively (Table 9).

288

### 289 **Mg-Al disorder**

290 It must be stressed that the difference between the ordered and disordered formula  
291 being considered for oxy-dravite wholly relates Mg-Al order-disorder and does not impact on  
292 the resulting chemical composition. The two atoms of Mg can be disordered over the Y and Z  
293 sites, or fully ordered at Z as occurs in the other oxy(O1)-tourmalines such as oxy-chromium-  
294 dravite and oxy-vanadium-dravite (Table 9). In this regard, several empirical structural  
295 formulae of oxy-dravite reported in the literature approach  
296  $\text{Na}^Y(\text{Al}_2\text{Mg})^Z(\text{Al}_5\text{Mg})\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$  (e.g., Bloodaxe et al. 1999; Bosi and Lucchesi,  
297 2004, 2007; Bosi et al. 2010) rather than the end-member proposed herein. The occurrence of  
298 Mg at the Z site and Al at the Y site is commonly coupled to the occurrence of  $\text{O}^{2-}$  at the O1  
299 site, controlled by the relationship  $2^Y\text{Mg}^{2+} + {}^Z\text{Al}^{3+} + {}^W(\text{OH})^{1-} \leftrightarrow 2^Y\text{Al}^{3+} + {}^Z\text{Mg}^{2+} + {}^W\text{O}^{2-}$   
300 (Hawthorne 1996). This relationship has been generalized as  $3^Y\text{Mg}^{2+} + 2^Z\text{Al}^{3+} + {}^W(\text{OH})^{1-} \leftrightarrow$   
301  $3^Y\text{Al}^{3+} + 2^Z\text{Mg}^{2+} + {}^W\text{O}^{2-}$  by Henry *et al.* (2011) on the basis of findings of Bosi and Lucchesi  
302 (2007), namely that a maximum of 2 divalent cations (i.e.,  $\text{Mg}^{2+} + \text{Fe}^{2+}$ ) per formula unit can  
303 occur at the Z site. The latter authors, however, based their findings on Al-poor tourmalines, in  
304 which the Al is substituted by ions with larger ionic radius (e.g.,  $\text{Fe}^{3+}$  and Cr). Actually, the  
305 amounts of Mg at Z are a product of the mismatch in size of cations occupying the Z-site; in  
306 fact, this mismatch is larger when involving Mg-Al rather than Mg-Cr<sup>3+</sup> or Mg-V<sup>3+</sup> as the  
307 ionic radii of  ${}^Z\text{Cr}^{3+}$  (0.613 Å) or  ${}^Z\text{V}^{3+}$  (0.653 Å) are much larger than that of  ${}^Z\text{Al}$  (0.543 Å)  
308 (Bosi and Lucchesi 2007). Figure 3 displays the increase of  $\Sigma\text{R}^{2+}$  at Z (such as Mg and  $\text{Fe}^{2+}$ ) as  
309 a function of replacement of  $\Sigma\text{Al}$  at Y and Z by larger cations like  $\text{V}^{3+}$ ,  $\text{Cr}^{3+}$  and  $\text{Fe}^{3+}$  in  
310 tourmalines, showing that the amounts of Mg at the Z site is dependent on the  $\text{R}^{3+}$ -cation size.  
311 These observations strongly support the concept that Mg is distributed over the Y and Z sites in  
312 Al-rich oxy(O1)-tourmalines. Note that although the occurrence of  $\text{Fe}^{2+}$  at Z seems to be  
313 limited to  $\text{Fe}^{3+}$ -bearing Al-rich tourmalines (e.g., Filip et al 2012), Mg is observed as to be the  
314 only divalent cation occurring at the Z site in Al-poor tourmalines (e.g., Bosi et al 2004, 2012,  
315 2013).

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#### CONCLUDING REMARKS

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320 Some important characteristics of oxy-dravite are:

321 1) the end-member formula is:  $\text{Na}^Y(\text{Al}_3)^Z(\text{Al}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ ;

- 322 2) the most representative ideal formula is:  $\text{Na}^Y(\text{Al}_2\text{Mg})^Z(\text{Al}_5\text{Mg})\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$ ;  
323 3) the Mg-Al disorder over the Y and Z sites is controlled by the short-range bond-valence  
324 requirements of  $\text{O}^{2-}$  at O1;  
325 4) the amount of Mg at the Z site is a function of the degree of cation size mismatch at Z.

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329

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475 **FIGURE 1.** Photos of the holotype fragment of oxy-dravite in reflected (a) and transmitted (b)  
476 light.

477 **FIGURE 3.** Room-temperature Mössbauer spectrum of oxy-dravite, fitted with two doublets  
478 (dotted lines) assigned to  $\text{Fe}^{3+}$  (centroid shift: 0.35 mm/s; quadrupole splittings:  
479 0.76 and 1.60 mm/s), and one doublet (solid line) assigned to  $\text{Fe}^{2+}$  (centroid shift:  
480 1.00 mm/s; quadrupole splitting: 2.43 mm/s, relative to  $\alpha\text{-Fe}$ ). The obtained  
481  $\text{Fe}^{3+}/\Sigma\text{Fe}$  area ratio is 0.96, with an estimated error of  $\pm 0.02$ . Thick line denotes  
482 summed spectrum.

483 **FIGURE 3.** Variation in  $\Sigma\text{Al}$  at Y and Z as a function of  $\Sigma\text{R}^{2+}$ -cations (Mg,  $\text{Fe}^{2+}$ , Co, Ni) at Z.  
484 The green arrow represents the substitution  $(\text{Cr} + \text{V}^{3+} + \text{Fe}^{3+}) \rightarrow \text{Al}$ , the red arrow  
485 represents the substitution  $2^{\text{Y}}\text{Al}^{3+} + {}^{\text{Z}}\text{Mg}^{2+} + {}^{\text{W}}\text{O}^{2-} \rightarrow 2^{\text{Y}}\text{Mg}^{2+} + {}^{\text{Z}}\text{Al}^{3+} + {}^{\text{W}}(\text{OH})^{1-}$ .  
486 Red circle represents the oxy-dravite from the present study. Filled circles  
487 represent 139 samples: uvite from Razmanova et al. (1983), 129 tourmalines of  
488 different compositions from the dataset of Bosi and Lucchesi (2007), dravite from  
489 Bosi (2008), 2 oxy-dravite samples plus 2 dravite samples from Bosi et al. (2010),  
490 oxy-chromium-dravite from Bosi et al. (2012), oxy-vanadium-dravite from Bosi et



491 al. (2013), 2 oxy-schorl samples from Bačík et al. (2013). Black crosses represent  
492 Co-rich and Ni-rich tourmalines from Rozhdestvenskayaa et al. (2012). Red  
493 squares represent the ideal dravite, ideal oxy-dravite with ordered (O-) and  
494 disordered (D-) formula (see text).

Table 1. Single-crystal X-ray diffraction data details for oxy-dravite

Crystal sizes (mm)	0.20 × 0.25 × 0.30
<i>a</i> (Å)	15.9273(2)
<i>c</i> (Å)	7.2001(1)
<i>V</i> (Å <sup>3</sup> )	1581.81(4)
Range for data collection, 2θ (°)	5 - 69
Reciprocal space range <i>hkl</i>	-22 ≤ <i>h</i> ≤ 25 -25 ≤ <i>k</i> ≤ 25 -11 ≤ <i>l</i> ≤ 11
Total number of frames	3265
Set of measured reflections	7124
Unique reflections, <i>R</i> <sub>int</sub> (%)	1586, 1.42
Completeness (%)	99.9
Redundancy	9
Absorption correction method	SADABS
Refinement method	Full-matrix least-squares on <i>F</i> <sup>2</sup>
Structural refinement program	SHELXL-97
Extinction coefficient	0.0046(2)
Flack parameter	0.03(3)
<i>wR</i> <sub>2</sub> (%)	3.33
<i>R</i> <sub>1</sub> (%) all data	1.17
<i>R</i> <sub>1</sub> (%) for <i>I</i> > 2σ( <i>I</i> )	1.16
GooF	1.140
Largest diff. peak and hole (±e <sup>-</sup> /Å <sup>3</sup> )	0.26 and -0.28

Notes: *R*<sub>int</sub> = merging residual value; *R*<sub>1</sub> = discrepancy index, calculated from *F*-data; *wR*<sub>2</sub> = weighted discrepancy index, calculated from *F*<sup>2</sup>-data; GooF = goodness of fit; Diff. Peaks = maximum and minimum residual electron density. Radiation, MoKα = 0.71073 Å. Data collection temperature = 293 K. Space group *R*3*m*; *Z* = 3.

**TABLE 2.** Fractional atom coordinates and site occupancy for oxy-dravite ( $\text{\AA}^2$ ).

Site	$x/a$	$y/b$	$z/c$	Site occupancy
<i>X</i>	0	0	0.22939(18)	Na <sub>0.914(5)</sub>
<i>Y</i>	0.12223(2)	0.061116(11)	0.63657(5)	Al <sub>0.826(2)</sub> Fe <sub>0.174(2)</sub>
<i>Z</i>	0.297683(16)	0.261493(17)	0.61014(4)	Al <sub>1.00</sub>
<i>B</i>	0.10989(4)	0.21978(8)	0.45430(16)	B <sub>1.00</sub>
<i>T</i>	0.191532(13)	0.189631(14)	0	Si <sub>1.00</sub>
O1	0	0	0.7698(2)	O <sub>1.00</sub>
O2	0.06076(3)	0.12151(6)	0.48724(12)	O <sub>1.00</sub>
O3	0.26112(7)	0.13056(4)	0.50994(12)	O <sub>1.00</sub>
O4	0.09356(3)	0.18712(6)	0.07152(12)	O <sub>1.00</sub>
O5	0.18445(6)	0.09223(3)	0.09252(12)	O <sub>1.00</sub>
O6	0.19447(4)	0.18433(4)	0.77729(9)	O <sub>1.00</sub>
O7	0.28509(4)	0.28474(4)	0.07754(8)	O <sub>1.00</sub>
O8	0.20920(4)	0.27005(4)	0.43936(9)	O <sub>1.00</sub>
H3	0.2453(15)	0.1227(7)	0.396(3)	H <sub>1.00</sub>

**TABLE 3.** Displacement parameters ( $\text{\AA}^2$ ) for oxy-dravite

Site	$U^{11}$	$U^{22}$	$U^{33}$	$U^{23}$	$U^{13}$	$U^{12}$	$U_{\text{eq}}/U_{\text{iso}}^*$
X	0.0217(4)	0.0217(4)	0.0235(6)	0	0	0.0108(2)	0.0223(3)
Y	0.00895(14)	0.00675(11)	0.01145(14)	-0.00158(4)	-0.00316(8)	0.00448(7)	0.00881(9)
Z	0.00531(9)	0.00536(9)	0.00583(9)	0.00039(7)	-0.00008(7)	0.00262(7)	0.00552(5)
B	0.0054(3)	0.0060(4)	0.0076(4)	0.0005(3)	0.00027(16)	0.0030(2)	0.00626(17)
T	0.00458(8)	0.00429(8)	0.00590(8)	-0.00061(6)	-0.00054(6)	0.00226(6)	0.00490(5)
O1	0.0127(4)	0.0127(4)	0.0075(6)	0	0	0.00637(19)	0.0110(2)
O2	0.0112(3)	0.0042(3)	0.0098(4)	0.0009(2)	0.00044(12)	0.00210(15)	0.00918(15)
O3	0.0221(4)	0.0124(2)	0.0049(3)	0.00040(14)	0.0008(3)	0.0110(2)	0.01206(15)
O4	0.0070(2)	0.0140(4)	0.0094(3)	-0.0015(3)	-0.00075(13)	0.00699(18)	0.00936(14)
O5	0.0148(4)	0.0066(2)	0.0083(3)	0.00080(13)	0.0016(3)	0.00740(18)	0.00900(13)
O6	0.0090(2)	0.0089(2)	0.0054(2)	-0.00140(16)	-0.00081(16)	0.00505(18)	0.00751(10)
O7	0.0064(2)	0.0053(2)	0.0084(2)	-0.00106(17)	-0.00179(18)	0.00047(17)	0.00776(9)
O8	0.0039(2)	0.0088(2)	0.0140(2)	0.00227(18)	0.00051(18)	0.00275(19)	0.00909(10)
H3							0.018*

Notes: Equivalent ( $U_{\text{eq}}$ ) and isotropic ( $U_{\text{iso}}$ ) displacement parameters; H-atom was constrained to have a  $U_{\text{iso}}$  1.5 times the  $U_{\text{eq}}$  value of the O3 oxygen.

**TABLE 4.** Selected bond distances (Å) for oxy-dravite.

<i>B</i> -O2	1.3760(14)	<i>Y</i> -O1	1.9397(8)
<i>B</i> -O8 <sup>A</sup> (× 2)	1.3740(8)	<i>Y</i> -O2 <sup>B</sup> (× 2)	1.9955(6)
< <i>B</i> -O>	1.375	<i>Y</i> -O3	2.1216(10)
		<i>Y</i> -O6 <sup>C</sup> (× 2)	1.9858(6)
<i>T</i> -O4	1.6246(3)	< <i>Y</i> -O>	2.004
<i>T</i> -O5	1.6396(4)		
<i>T</i> -O7	1.6030(6)	<i>Z</i> -O3	1.9985(5)
<i>T</i> *-O6	1.6077(6)	<i>Z</i> -O6	1.9082(6)
< <i>T</i> -O>	1.619	<i>Z</i> -O8 <sup>E</sup>	1.8993(6)
		<i>Z</i> -O7 <sup>E</sup>	1.9073(6)
<i>X</i> -O2 <sup>B,F</sup> (× 3)	2.5012(12)	<i>Z</i> -O7 <sup>D</sup>	1.9572(6)
<i>X</i> -O4 <sup>B,F</sup> (× 3)	2.8201(10)	<i>Z</i> -O8	1.9258(6)
<i>X</i> -O5 <sup>B,F</sup> (× 3)	2.7284(10)	< <i>Z</i> -O>	1.933
< <i>X</i> -O>	2.683		
		O3-H	0.85(2)

Notes: Standard uncertainty in parentheses. Superscript letters: A = (*y* - *x*, *y*, *z*); B = (*y* - *x*, -*x*, *z*); C = (*x*, *x* - *y*, *z*); D = (*y* - *x* + 1/3, -*x* + 2/3, *z* + 2/3); E = (-*y* + 2/3, *x* - *y* + 1/3, *z* + 1/3); F = (-*y*, *x* - *y*, *z*). Transformations relate coordinates to those of Table 2.

\* Positioned in adjacent unit cell.

**TABLE 5.** X-ray powder diffraction data for oxy-dravite

$I_{(meas)}$ (%)	$d_{(meas)}$ (Å)	$d_{(cal)}$ (Å)	$h$	$k$	$l$
44	6.377	6.383	1	0	1
25	4.978	4.979	0	2	1
14	4.598	4.595	3	0	0
67	4.222	4.221	2	1	1
64	3.983	3.980	2	2	0
84	3.483	3.484	0	1	2
10	3.379	3.377	1	3	1
9	3.011	3.008	4	1	0
100	2.963	2.962	1	2	2
7	2.898	2.896	3	2	1
7	2.622	2.621	3	1	2
68	2.576	2.575	0	5	1
19	2.400	2.400	0	0	3
14	2.377	2.376	2	3	2
14	2.343	2.342	5	1	1
8	2.190	2.189	5	0	2
9	2.163	2.162	4	3	1
13	2.128	2.128	0	3	3
15	2.056	2.055	2	2	3
35	2.041	2.040	1	5	2
6	2.019	2.018	1	6	1
52	1.915	1.918	3	4	2
7	1.876	1.876	1	4	3
6	1.849	1.848	6	2	1
5	1.785	1.785	1	0	4
18	1.660	1.660	0	6	3
25	1.633	1.640	2	7	1
11	1.593	1.592	5	5	0
2	1.528	1.532	9	0	0
14	1.507	1.507	0	5	4
8	1.504	1.504	8	2	0
4	1.481	1.481	2	4	4
14	1.456	1.456	5	1	4
6	1.432	1.430	7	4	0
6	1.418	1.417	6	5	1
11	1.410	1.410	4	3	4
6	1.329	1.329	3	5	4
5	1.310	1.309	10	1	0
9	1.276	1.277	5	0	5

Notes:  $I_{(meas)}$  = measured intensity,  $d_{(meas)}$  = measured interplanar spacing;  $d_{(calc)}$  = calculated interplanar spacing;  $hkl$  = reflection indices. Estimated errors in  $d_{(meas)}$ -spacing range from 0.01 Å for large  $d$ -values to 0.001 Å for small  $d$ -values.

**TABLE 6.** Chemical composition of oxy-dravite

	Weight %		Apfu
SiO <sub>2</sub>	37.01(10)	Si	5.98(5)
TiO <sub>2</sub>	0.14(3)	Ti <sup>4+</sup>	0.017(4)
B <sub>2</sub> O <sub>3</sub> <sup>*</sup>	10.76	B	3.00
Al <sub>2</sub> O <sub>3</sub>	33.11(20)	Al	6.31(5)
Fe <sub>2</sub> O <sub>3</sub> <sup>†</sup>	5.00	Fe <sup>3+</sup>	0.61(1)
FeO <sup>†</sup>	0.19	Fe <sup>2+</sup>	0.025(3)
MgO	8.56(8)	Mg	2.06(2)
Na <sub>2</sub> O	2.65(3)	Na	0.83(1)
K <sub>2</sub> O	0.10(1)	K	0.021(1)
H <sub>2</sub> O <sup>*</sup>	2.65	OH	3.24
Total	100.58		

*Notes:* number of ions calculated on basis of 31 (O, OH, F).  
Uncertainties for oxides (in brackets) are standard deviation of 10 spots. B<sub>2</sub>O<sub>3</sub> and H<sub>2</sub>O uncertainty assumed at 5%. Standard uncertainty for ions was calculated by error-propagation theory.

\* Calculated by stoichiometry.

† Determined by Mössbauer spectroscopy, (FeO)<sub>EMPA</sub> = 4.69(9).

**TABLE 7.** Site populations (*apfu*), site scattering factors (*epfu*) and mean bond lengths (Å) for oxy-dravite

Site	Site population	Site scattering		Mean bond length	
		refined	calculated	refined	calculated*
X	0.83 Na + 0.02 K + 0.15 □	10.05(5)	9.53		
Y	1.34 Al + 1.03 Mg + 0.58 Fe <sup>3+</sup> + 0.03 Fe <sup>2+</sup> + 0.02 Ti (1.20 Al + 1.16 Mg + 0.60 Fe <sup>3+</sup> + 0.03 Fe <sup>2+</sup> + 0.02 Ti)**	45.8(1)	46.0	2.004	2.002
Z	4.94 Al + 1.03 Mg + 0.02 Fe <sup>3+</sup> (5.11 Al + 0.90 Mg)**	78 <sup>†</sup>	77.3	1.933	1.931
T	5.98 Si + 0.02 Al (6.00 Si)**	84 <sup>†</sup>	84	1.619	1.619
B	3.00 B	15 <sup>†</sup>	15	1.375	1.374
O3	3.00 (OH)	24 <sup>†</sup>	24		
O1	0.76 O + 0.24 (OH)	9 <sup>†</sup>	9		

Notes: O2, O4...O8 sites are fully populated by O<sup>2-</sup>; apfu = atoms per formula unit; epfu = electrons per formula unit.

\* Calculated from the ionic radii (Bosi and Lucchesi 2007).

\*\* Site populations optimized by the procedure of Wright et al. (2000).

<sup>†</sup>Fixed in the final stages of refinement.



**TABLE 8.** Bond valence calculations (valence unit) for oxy-dravite

Site	X	Y	Z	T	B	BVS
O1		0.51 <sup>x3</sup> →				1.53
O2	0.15 <sup>x3</sup> ↓	0.44 <sup>x2</sup> ↓→			0.99	2.01
O3		0.31	0.40 <sup>x2</sup> →			1.11
O4	0.07 <sup>x3</sup> ↓			1.00 <sup>x2</sup> →		2.06
O5	0.08 <sup>x3</sup> ↓			0.96 <sup>x2</sup> →		2.00
O6		0.45 <sup>x2</sup> ↓	0.51	1.04		2.00
O7			0.51	1.06		2.02
			0.45			
O8			0.49		0.99 <sup>x2</sup> ↓	2.00
			0.52			
BVS	0.91	2.59	2.87	4.06	2.97	
MFV*	0.85	2.65	2.83	4.00	3.00	

\* MFV = mean formal valence from site populations

**TABLE 9.** Comparative data for oxy-dravite, fluor-dravite and dravite.

	Oxy-dravite Na(Al <sub>2</sub> Mg)(Al <sub>5</sub> Mg)Si <sub>6</sub> O <sub>18</sub> (BO <sub>3</sub> ) <sub>3</sub> (OH) <sub>3</sub> O	Fluor-dravite NaMg <sub>3</sub> Al <sub>6</sub> Si <sub>6</sub> O <sub>18</sub> (BO <sub>3</sub> ) <sub>3</sub> (OH) <sub>3</sub> F	Dravite NaMg <sub>3</sub> Al <sub>6</sub> Si <sub>6</sub> O <sub>18</sub> (BO <sub>3</sub> ) <sub>3</sub> (OH) <sub>3</sub> (OH)
<i>a</i> (Å)	15.9273(2)	15.955(3)	15.96(2)
<i>c</i> (Å)	7.2001(1)	7.153(2)	7.21(2)
<i>V</i> (Å <sup>3</sup> )	1581.81(4)	1576.9(6)	1590.5
Space group	<i>R3m</i>	<i>R3m</i>	<i>R3m</i>
Optic sign	Uniaxial (–)	Uniaxial (–)	Uniaxial (–)
$\omega$	1.650(5)	1.645(2)	1.634 - 1.661
$\epsilon$	1.620(5)	1.621(2)	1.612 - 1.632
Colour	Dark red	Blackish brown	Pale brown to dark-brown to brownish-black, dark-yellow, blue
Pleochroism	O = orange E = pink	O = pale yellow-brown E = colorless	O = pale yellow E = colorless, yellowish, greenish, brownish
Reference	This work	Clark et al. (2011)	<a href="http://www.mindat.org">www.mindat.org</a>

**TABLE 10.** Selected tourmaline formulae.

Name	Representative formula	Substitution related to oxy-dravite
Oxy-dravite	$\text{Na}^{\text{Y}}(\text{Al}_2\text{Mg})^{\text{Z}}(\text{Al}_5\text{Mg})\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$	
Oxy-schorl	$\text{Na}^{\text{Y}}(\text{AlFe}^{2+}_2)^{\text{Z}}(\text{Al}_6)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$	$\text{Fe}^{2+} \rightarrow \text{Mg}^{2+}$
Oxy-chromium-dravite	$\text{Na}^{\text{Y}}(\text{Cr}_3)^{\text{Z}}(\text{Cr}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$	$\text{Cr}^{3+} \rightarrow \text{Al}^{3+}$
Oxy-vanadium-dravite	$\text{Na}^{\text{Y}}(\text{V}_3)^{\text{Z}}(\text{V}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$	$\text{V}^{3+} \rightarrow \text{Al}^{3+}$
Povondravite	$\text{Na}^{\text{Y}}(\text{Fe}^{3+}_3)^{\text{Z}}(\text{Fe}^{3+}_4\text{Mg}_2)\text{Si}_6\text{O}_{18}(\text{BO}_3)_3(\text{OH})_3\text{O}$	$\text{Fe}^{3+} \rightarrow \text{Al}^{3+}$

Figure 1a



Figure 1b

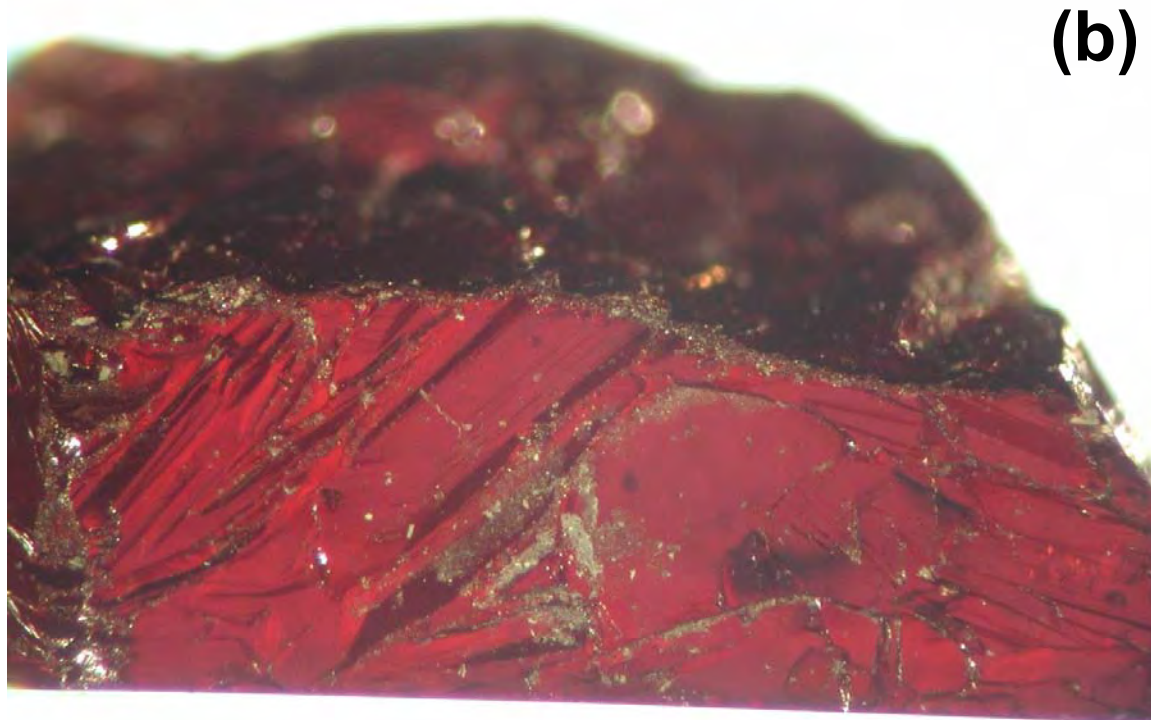


Figure 2

